Model Physics

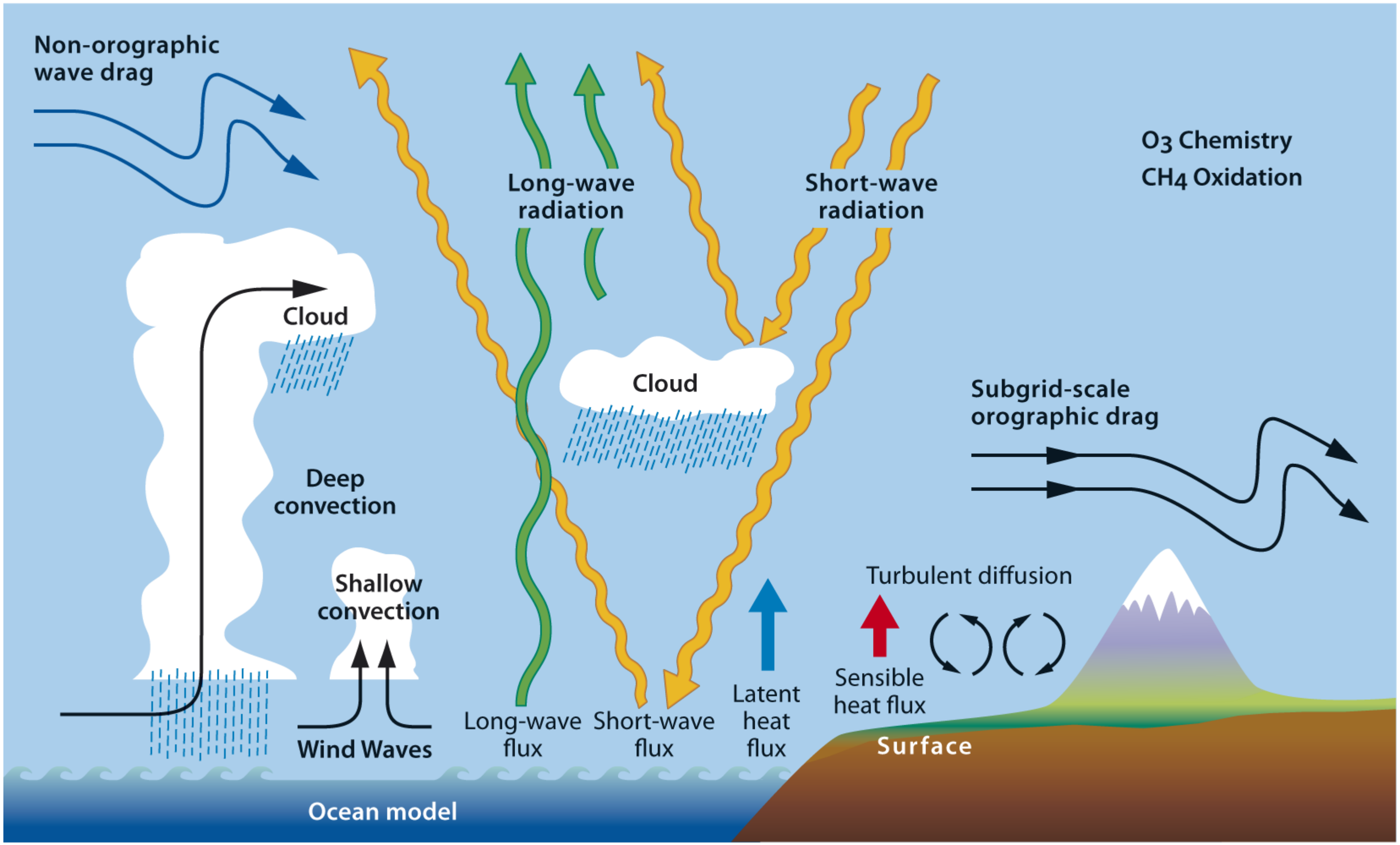
- A few **basics**
- **High resolution**
- A few problems
- A few **products**

for the Model Section: Peter Bechtold

http://www.ecmwf.int/en/learning/education-material/introductory-lectures-nwp

Parameterized processes in the ECMWF model

from the surface to the stratosphere



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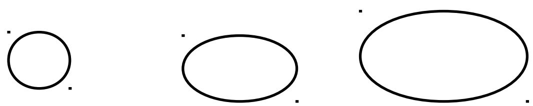
Precipitation JJA: Sensitivity to Model Formulation
Seasonal integrations

GPCP JJA 1990-2006



15
13
11
9
7
5
3
1

33R1(old vdiff)-33R1

10
4
2
0.5
-0.5
-2
-4
-10

33R1:2008 -GPCP

10
4
2
0.5
-0.5
-2
-4
-10

33R1(old radiation)-33R1

10
4
2
0.5
-0.5
-2
-4
-10

33R1(old convection)-33R1

10
4
2
0.5
-0.5
-2
-4
-10

33R1(old soil hydrology)-33R1

10
4
2
0.5
-0.5
-2
-4
-10

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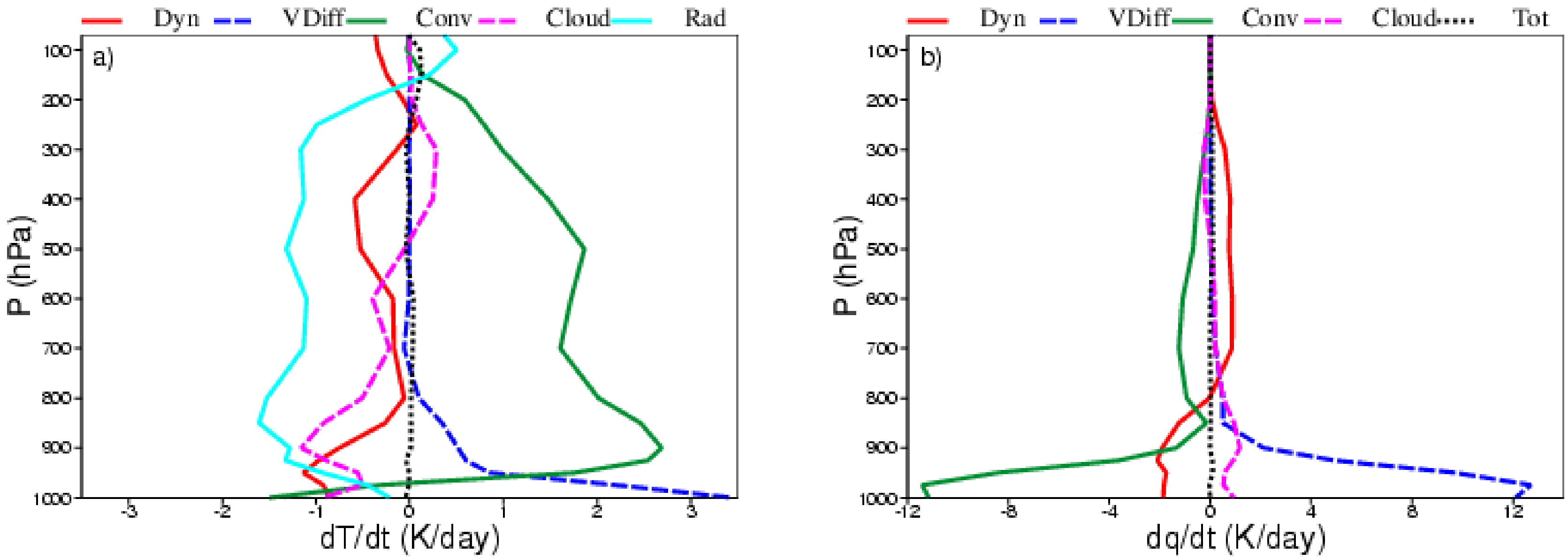
Model Tendencies - Tropics



For Temperature, above the boundary layer, there is roughly an equilibrium Radiation-Convection, but Dynamics and Clouds also important, whereas for moisture there is roughly an equilibrium between dynamical transport (moistening) and convective drying. - *Global Budgets are very* similar

All processes are important, nevertheless the driving force for atmospheric dynamics and convection is the radiation

The weather and thermal equilibria: exercises

- Suppose we have a series of fine day with an anticyclone, the temperature above the boundary-layer barely changes, Why?

$$\frac{d\theta}{dt} \approx 0 \quad \Rightarrow w\frac{d\theta}{dz} = \frac{d\theta}{dt}\Big|_{rad} = -\frac{2K}{86400s} \Rightarrow$$

w ~ -0.5 cm/s subsidence

~0.5 K/100 m

- But what happens when it is raining 100 mm/day ?

$$\int_{surf}^{10km} Cp\frac{dT}{dt}\rho_{air}\, dz = L_v\rho_{water}\, Pr(m/s)$$

$$c_p = 1005\, J/kg/K; \quad \rho_{water} = 1000 kg/m3; \quad L_v = 2.5x10^6 \text{ J/kg}$$

$$\mathrm{Pr} = 100\frac{mm}{day} = 1.147\, m/s\, x10^{-6}$$

100 mm/day precipitation heats the atmospheric column by 2867 W/m2 or by 25 K/day on average. This heating must be compensated by uplifting of w ~ 10 cm/s ➔ heavy precip/convection requires large-scale perturbation.

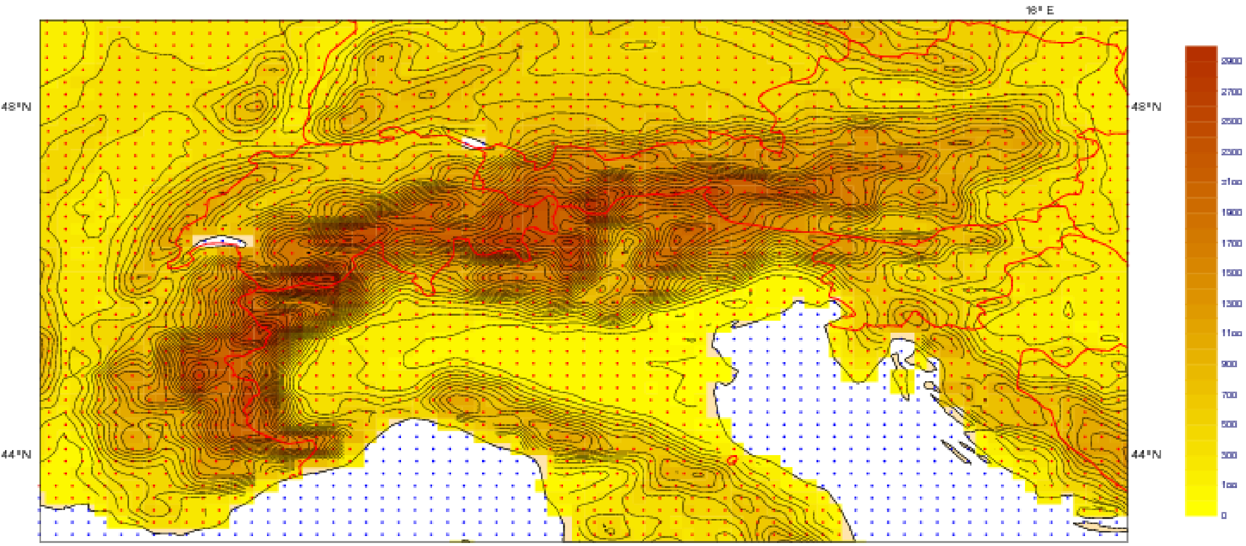
The 2016 horizontal resolution upgrade:

The Grids and effects from improved Numerics

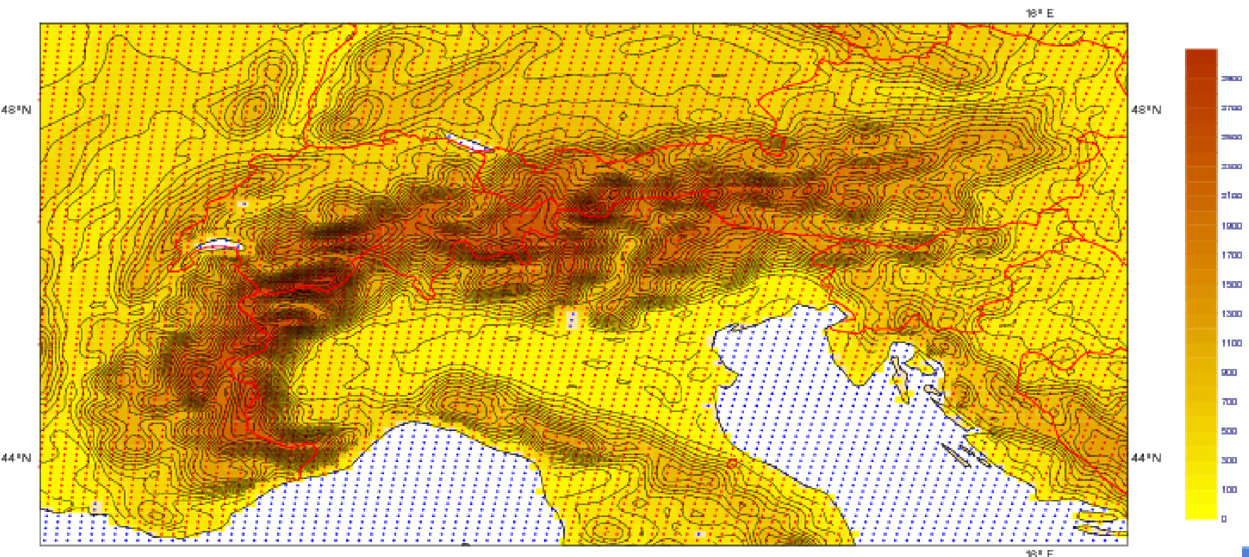
ECMWF

From Tl1279 (16 km) to TCo1279 (9 km)

OROGRAPHY, GRID POINTS AND LAND_SEA MASK FOR N640 ORIGINAL GRID
orography shaded (height in m), land grid points (red), sea grid points (blue)

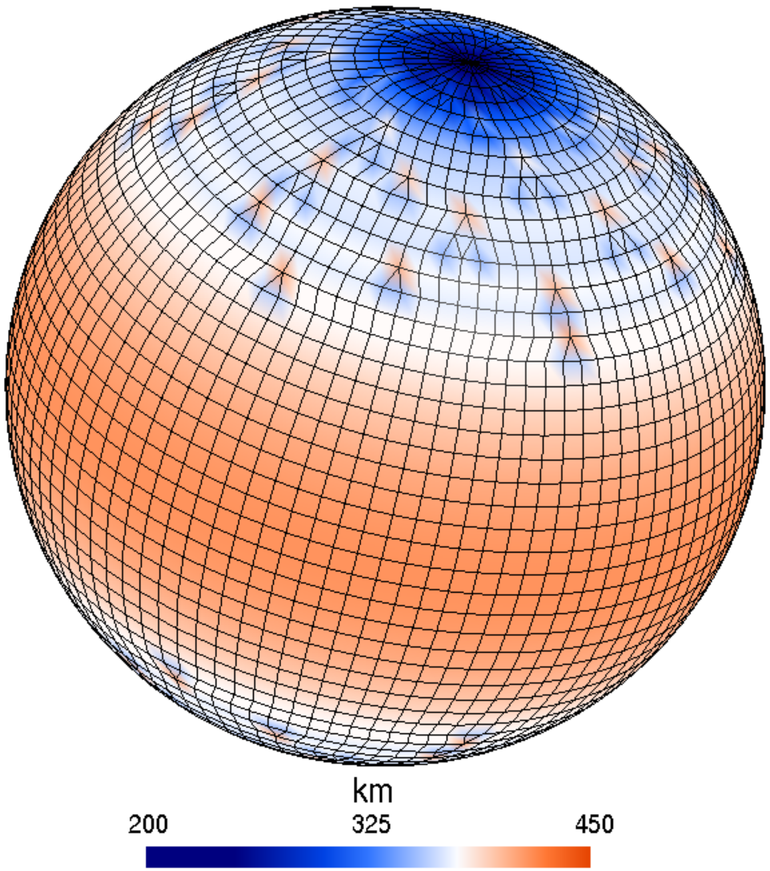


OROGRAPHY, GRID POINTS AND LAND_SEA MASK FOR O1280 OCTAHEDRAL GRID
orography shaded (height in m), land grid points (red), sea grid points (blue)

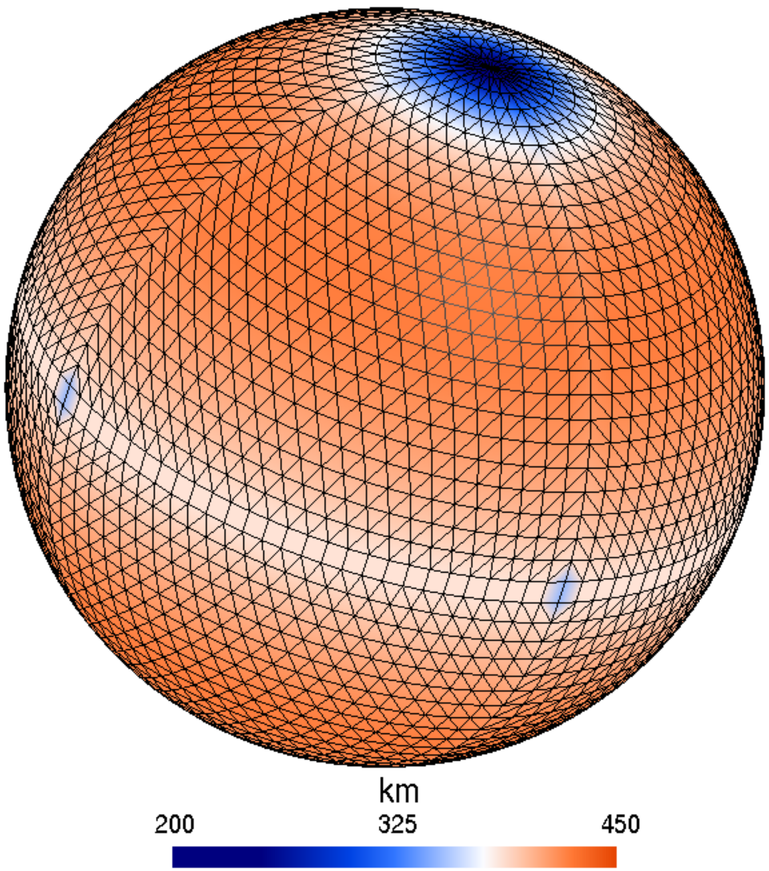


- Same max number of waves on the sphere=1279
- Less spectral smoothing applied to TC1279 orography than in Tl1279
- In the linear=Tl grid 2 grid-points represent one wave, while in the cubic=TC grid, a wave is represented by 4 grid-points =>much more accurate
- note that most computations are done in grid-point space
- The TC Gaussian grid is further reduced to a TC octahedral to save grid points

A new grid ….
and a more uniform resolution, ~9 km over Europe



N24 reduced Gaussian grid

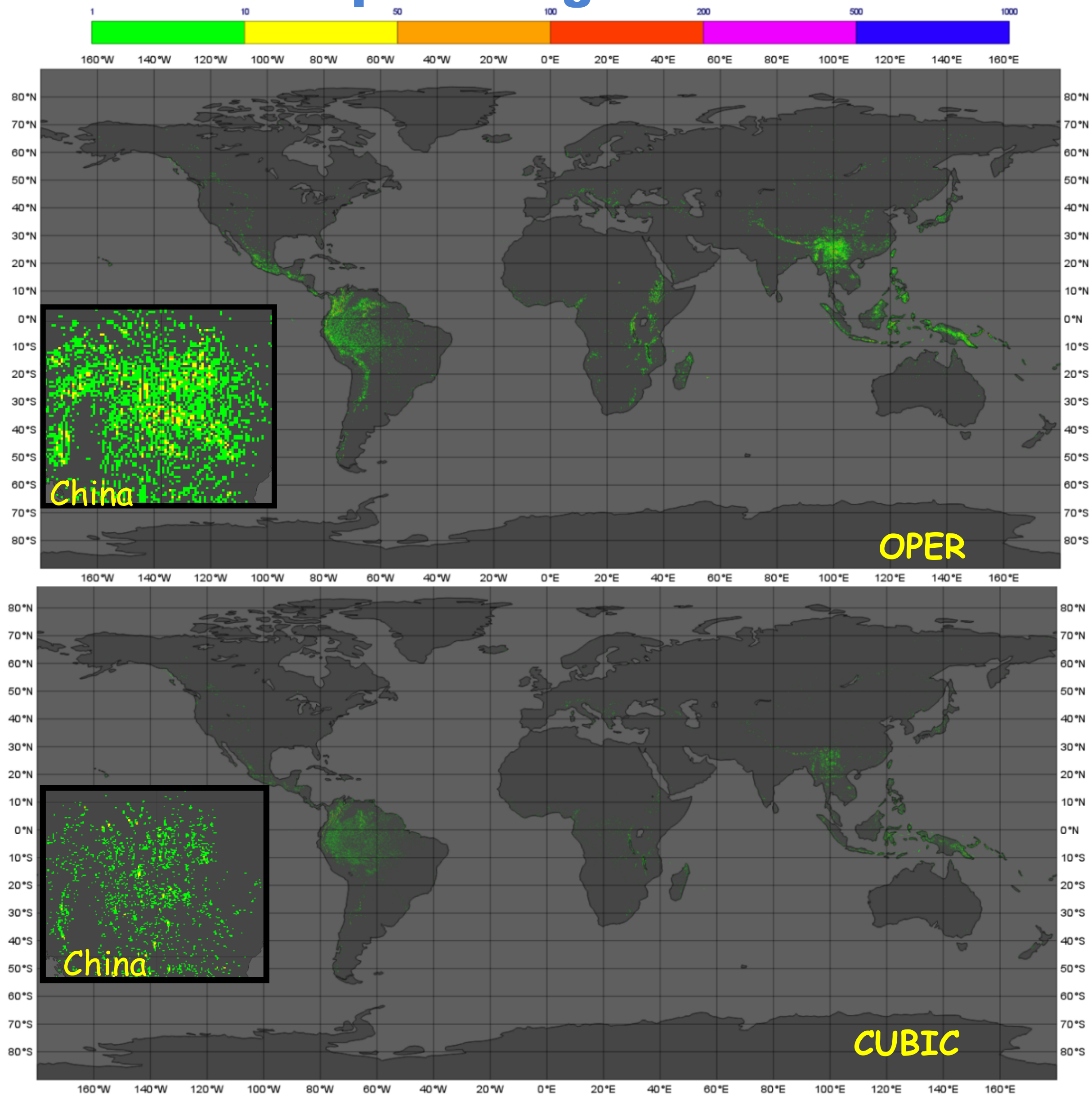


N24 octahedral Gaussian grid

Improvements: ....

Strong reduction of spurious grid-scale rainfall events (LSP)

Frequency of rain events >20mm/6h



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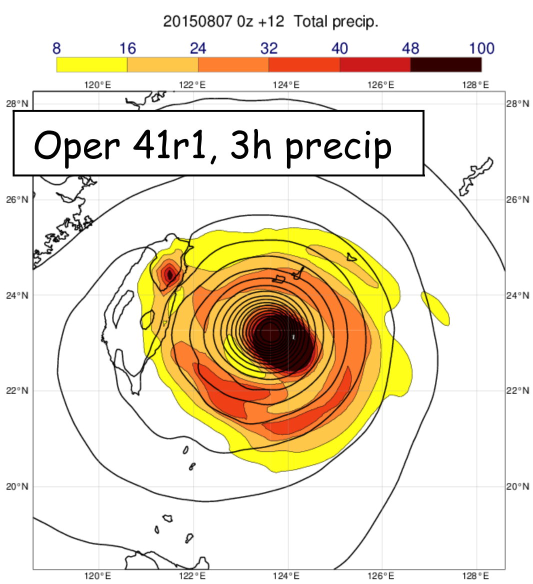
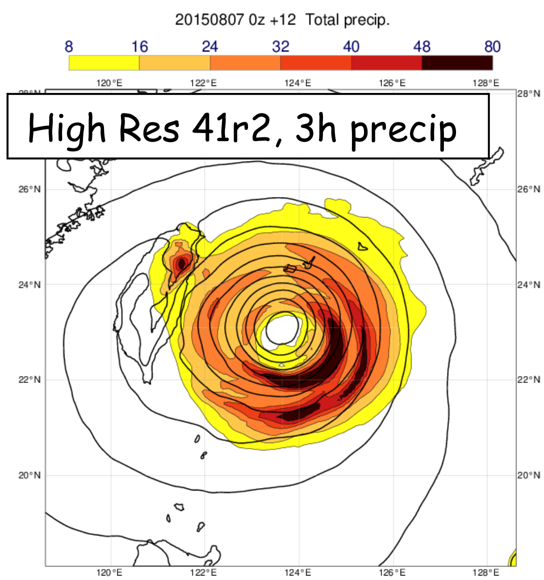
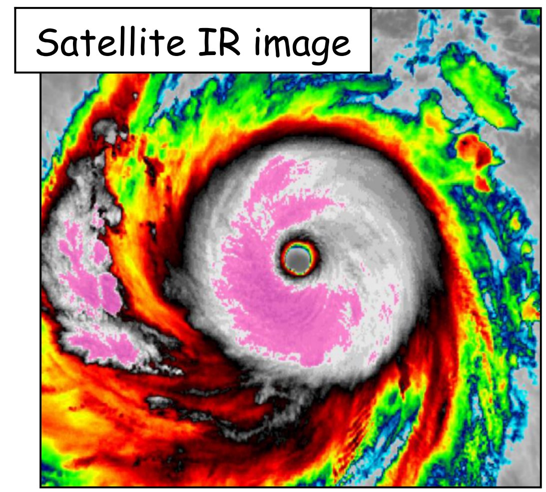
Improvements: Numerics

- Instability in Numerics due to departure point calculation in the semi-Lagrangian advection, leading to unrealistic tropical cyclone structures

Tropical Cyclone Soudelor
Aug 2015







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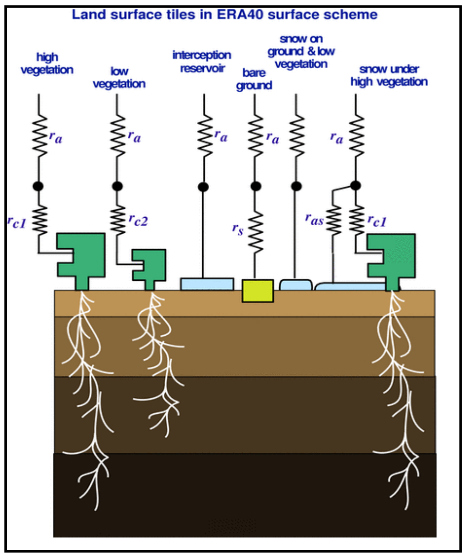
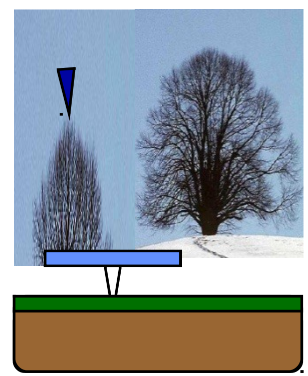
Physical processes: Surface temperatures wind and snow

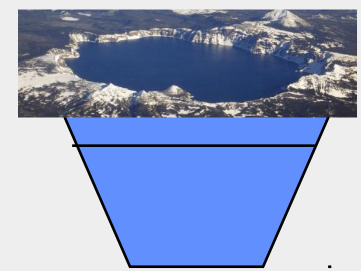
Land surface model evolution

2000/06	2007/11	2009/03	2009 &	2010	2015

- **TESSEL**
- **Hydrology**-TESSEL
- **new SNOW**
- **FLAKE**

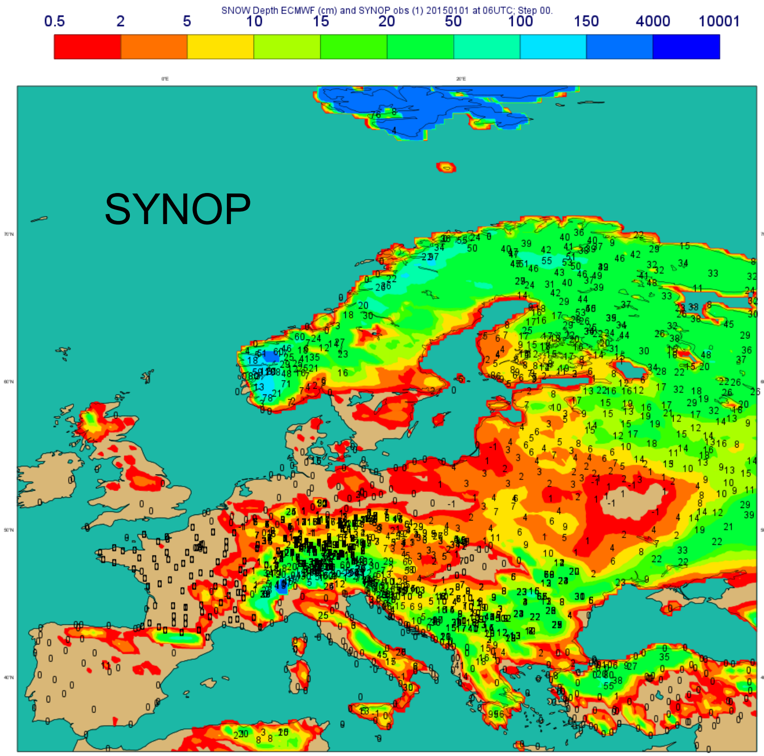






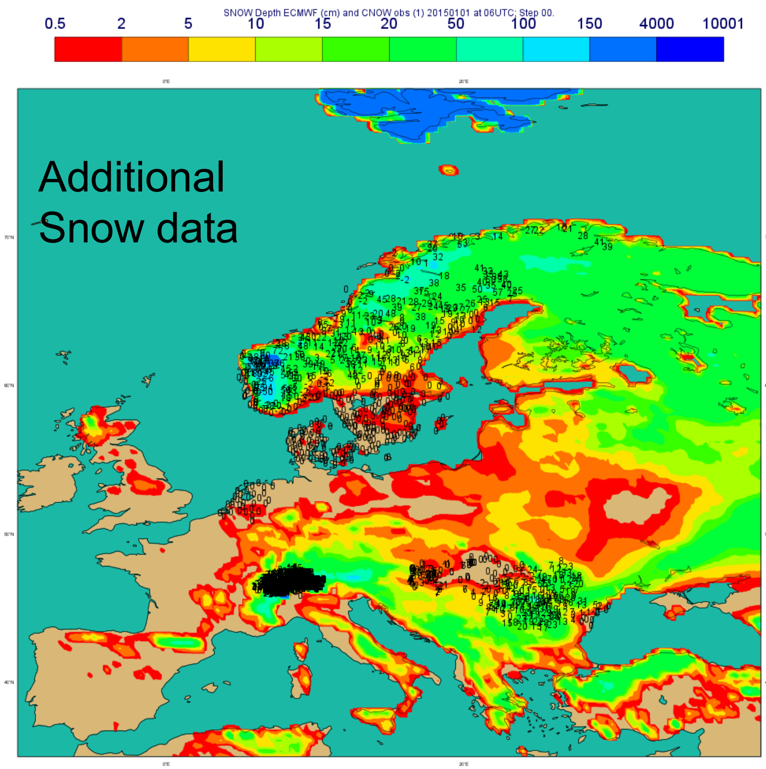
Snow Observations
Snow SYNOP and National Network data

Snow Depth (cm)



Available on the GTS (Global Telecommunication System)

2015 01 01 at 06UTC



Additional data from national networks (7 countries):

Sweden (>300), Romania(78), The Netherlands (33), Denmark (43), Hungary (61), Norway (183), Switzerland (332).

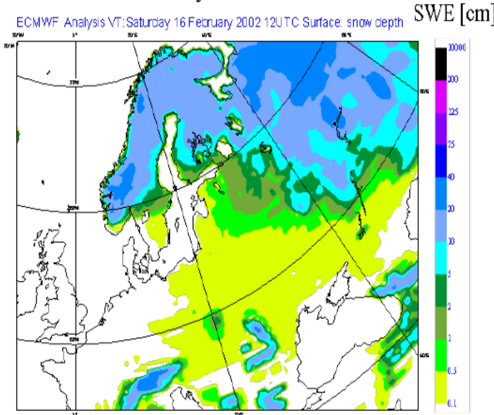
→**Dedicated BUFR (2011)**

(de Rosnay et al. ECMWF Res. Memo, R48.3/PdR/1139, 2011)

Snow analysis uses Synop and Satellite Obs

MODIS 16/02/2002

Analysis based on SYNOP



Snow extent is overestimated in the analysis when it is based on SYNOP data only

However, satellite only gives snow cover!
And the big change in 2014 was the way satellite data is used, i.e it is assimilated with large observation error, also if
FG =no snow, Sat=snow   => Sat snow≈5 cm

Fc errors (scores) very sensitive to snow (analysis)

See also ECMWF Newsletter no 143, article pp 26-31, Spring 2015

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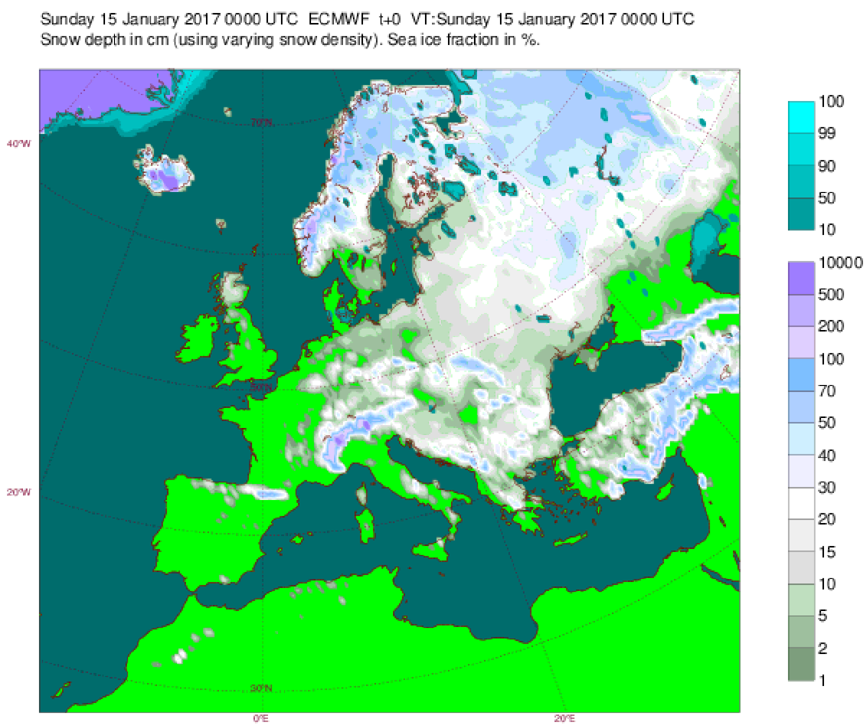
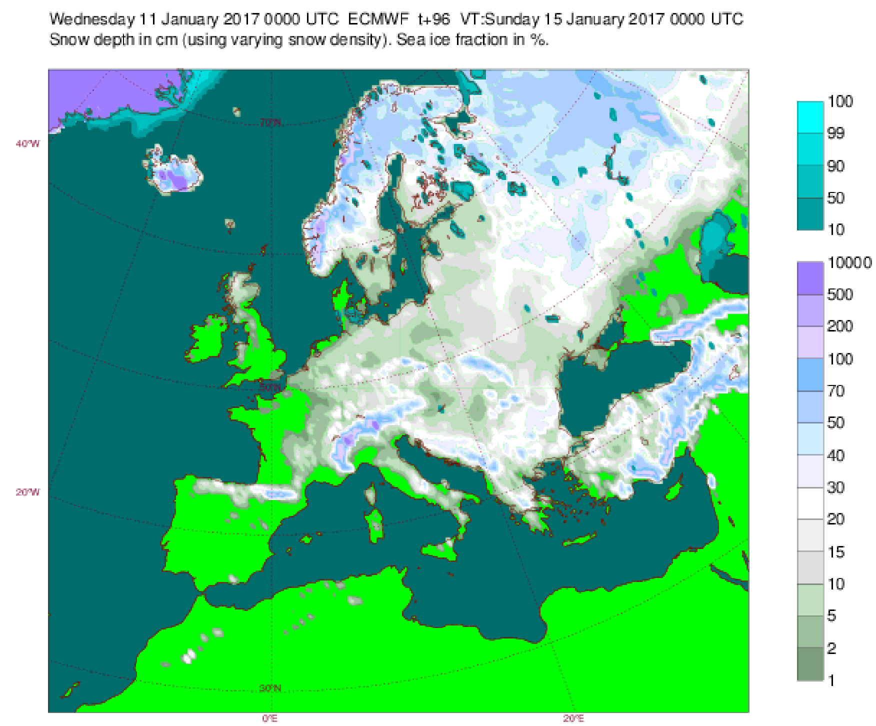
Archived prognostic snow related quantities

- **Snow depth (water equivalent), Sd => actual depth=Sd*(Rl=1000)/Rsn**

  below 10 cm snow depth, snow cover becomes fractional

- **Snow density (typically factor 10 lower than water-> 1 mm precip~1 cm snow), Rsn** (mixture old/new snow, wind compression)
- **Snow temperature, Tsn**
- **Snow albedo, Asn**

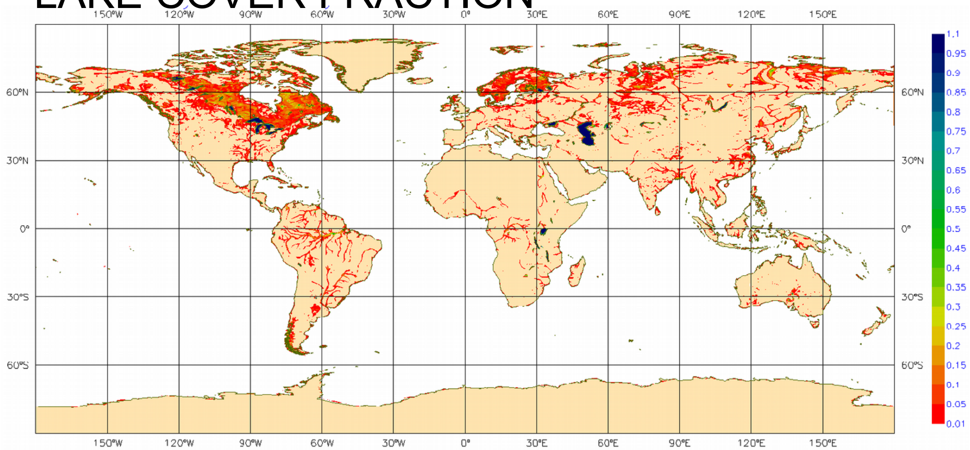




http://www.ecmwf.int/en/forecasts/charts/medium/snow-depth-and-sea-ice

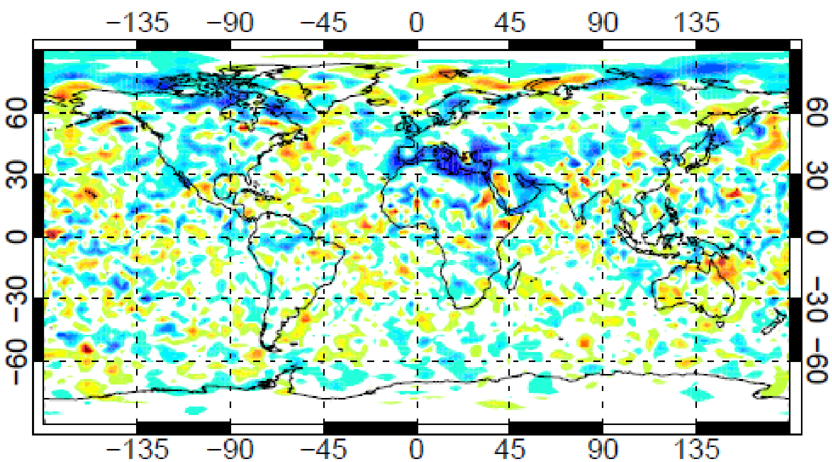
Impact of water bodies in IFS version June 2015

LAKE COVER FRACTION



T+48; 1000hPa

Summer experiment 15−Jun−2013 to 5−Jul−2013



Forecast of 2m temperature are improved in proximity of lakes and coastal areas

Why also coastal areas, these are not Lakes ?!...... cause before if land-sea mask>0.5 then only land point……… but doesn't solve T2m coastal problem for Norway

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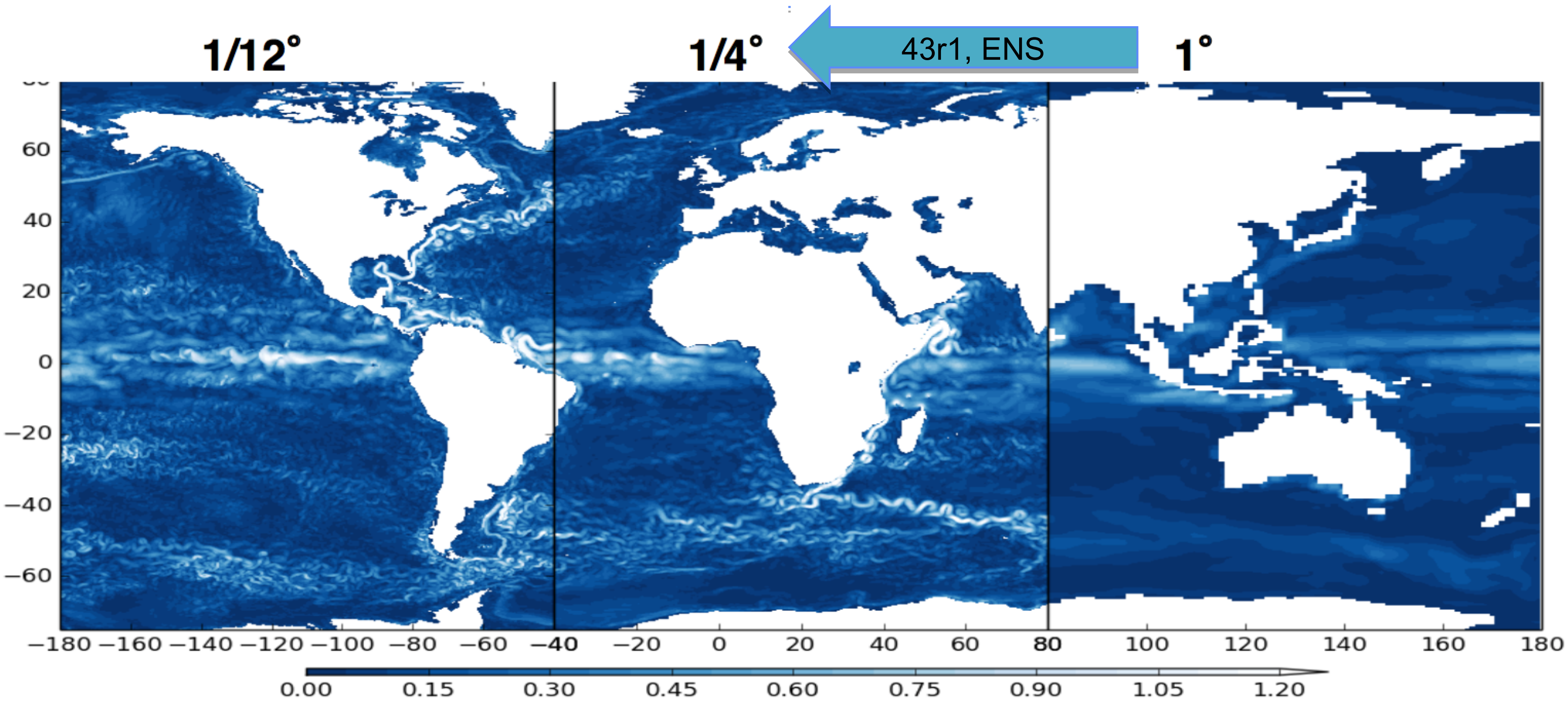
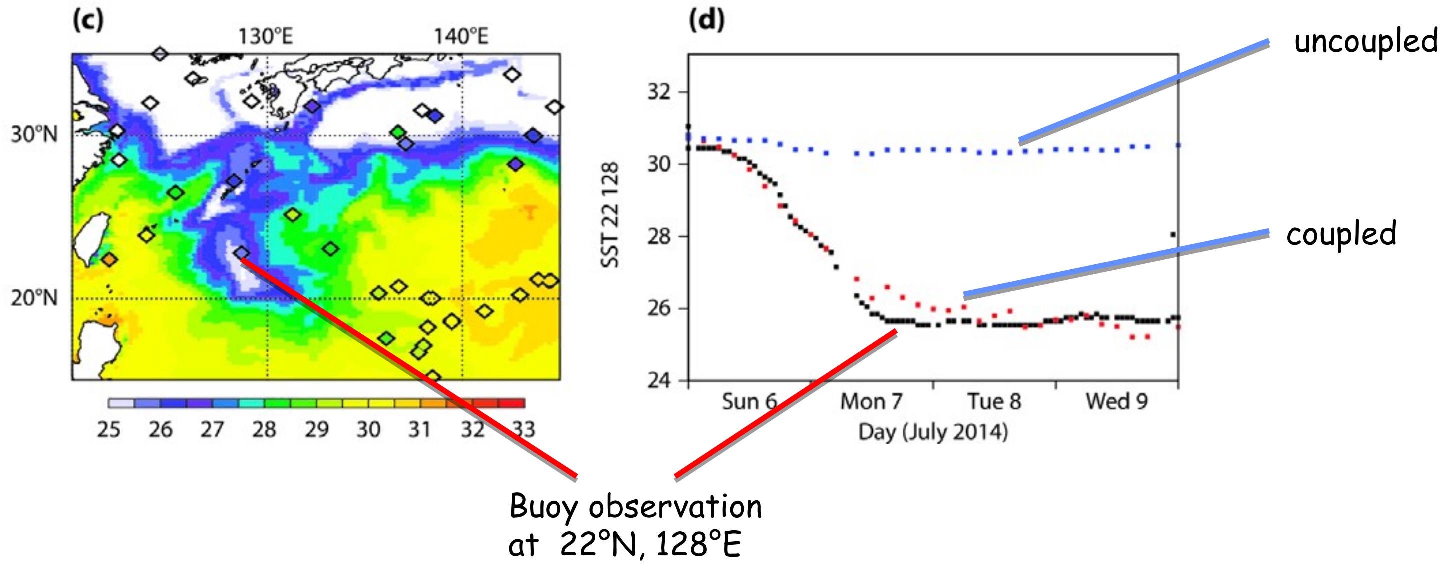
Ocean surface currents at various resolutions



Eddy resolving | Eddy permitting | Eddy parameterising

Coupled ocean vs uncoupled simulation

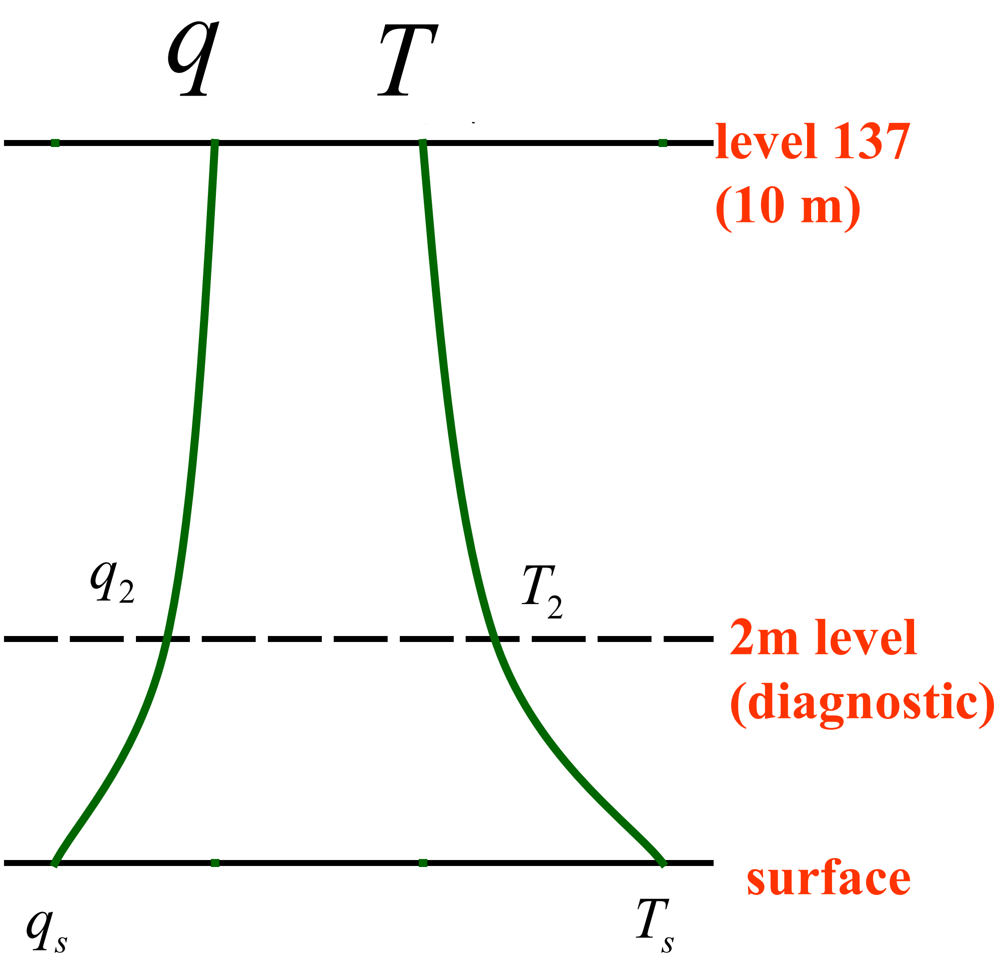
Tropical cyclone *Neoguri* with TCo1279 (HRES)



4-day forecast SSTs from the coupled forecast initialised at 0UTC on 6 July 2014 at the location of a buoy with approximate position 22°N, 128°E.

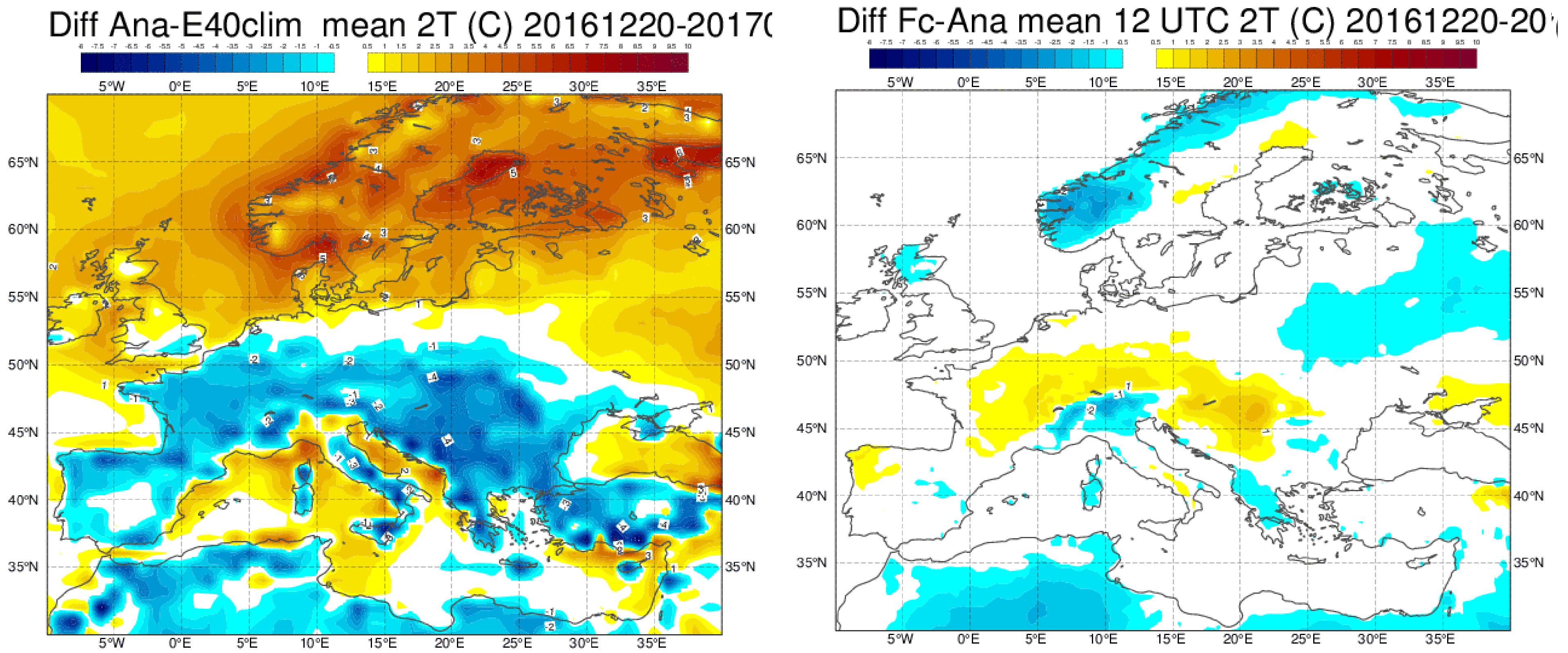
(Rodwell et al, ECMWF Technical Report 759, 2015)

T and q interpolation to the 2m level



- q_s and T_s are determined by the land surface scheme or by SST.
- Main purpose of land surface scheme is to provide correct area averaged fluxes of heat and moisture.
- Land surface scheme considers different sub-areas (tiles) but effect on screen level variables is not accounted for yet.

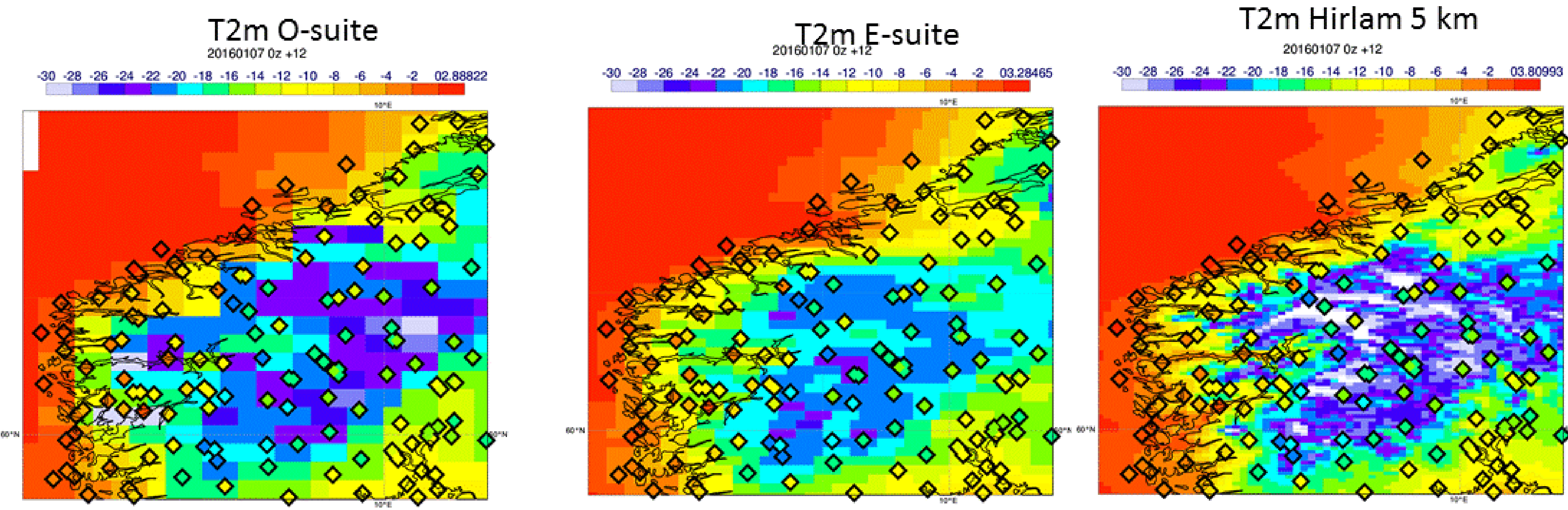
T2m mean errors (K) 20 Dec 2016- 25. Jan 2017



land mask applied (contour interval 0.5 K, start at +- 0.5 K)

Temperature negative error reduction in 41r2 resolution upgrade:

Coastal T-errors reduced through approximate radiation updates in space and time



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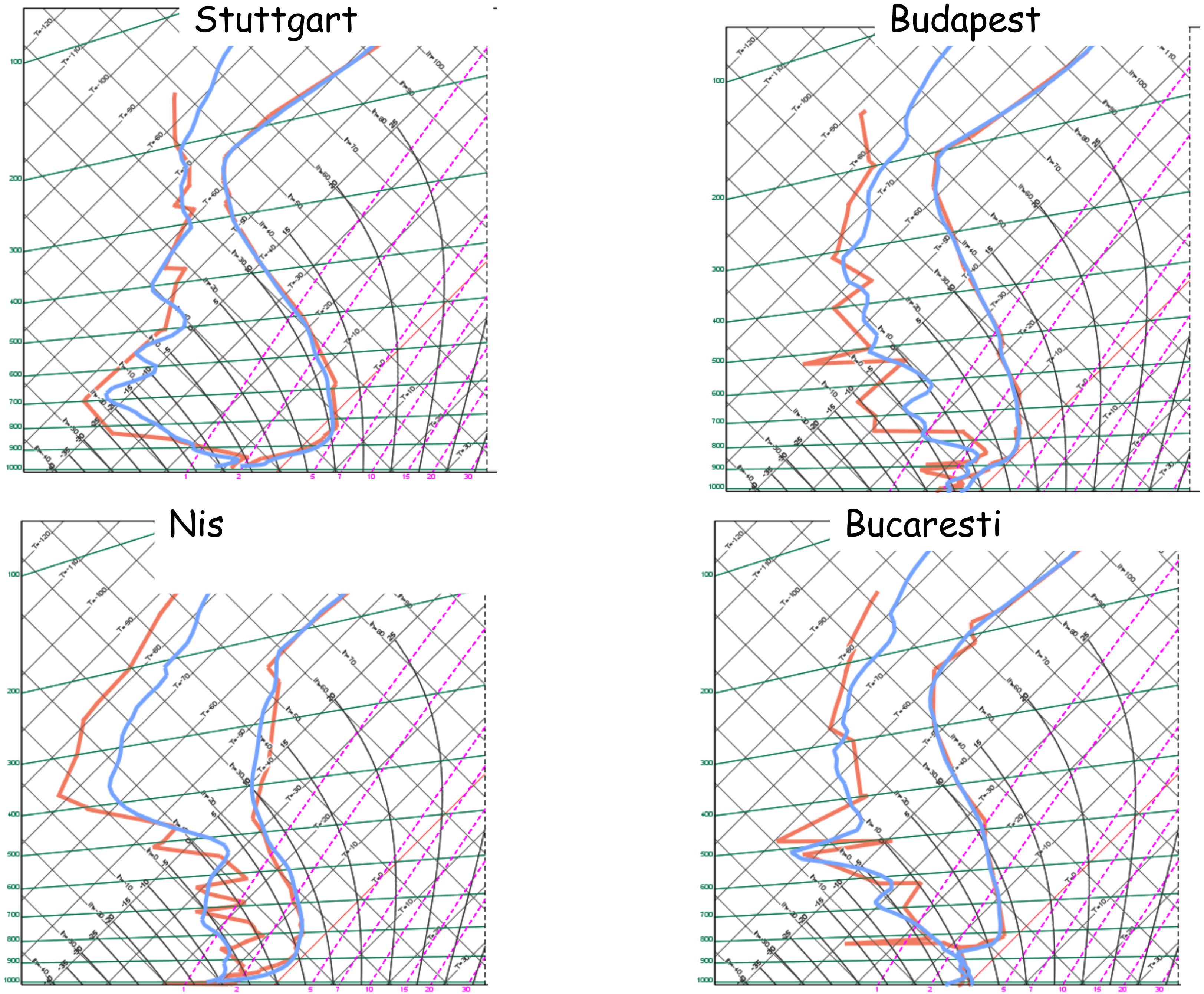
Example of 22 Jan 2017 00 t+24 Fc vs Sondes



Summary of wintertime 2m T errors

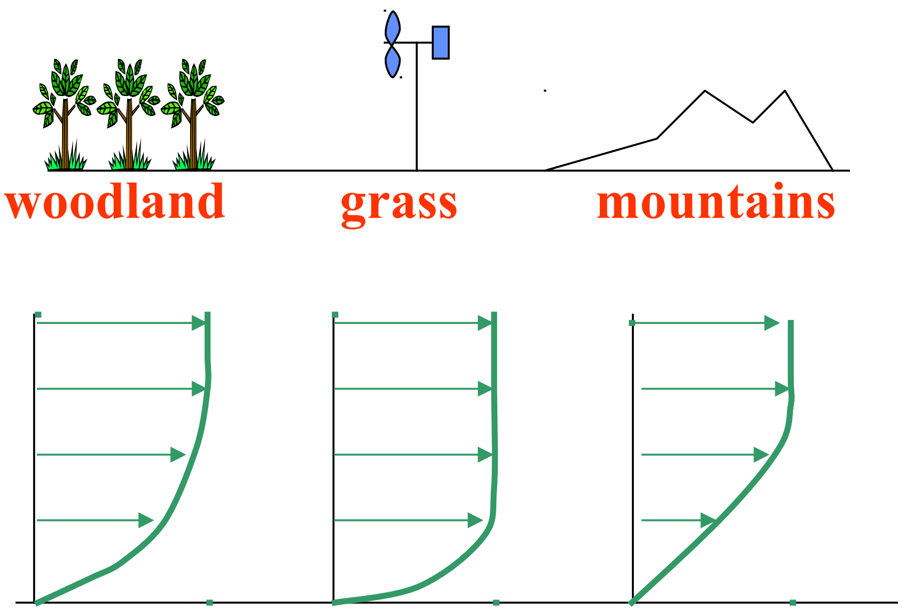
Overall not bad, mean error < 0.5 K, improved over previous years but still

- Regional differences, now mainly too cold, particular night-time problem, especially apparent over orography

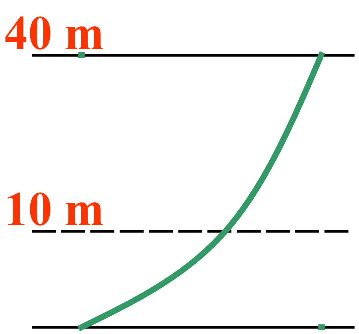
Various possible reasons: coupling (coefficient) with ground heat flux, error in lake temperatures (not frozen), stable boundary-layer mixing, low-level clouds, snow

- Overestimation of summertime night temperatures (coupling with soil, vegetation not shown) … should have been partly addressed (to be seen)

10 m wind



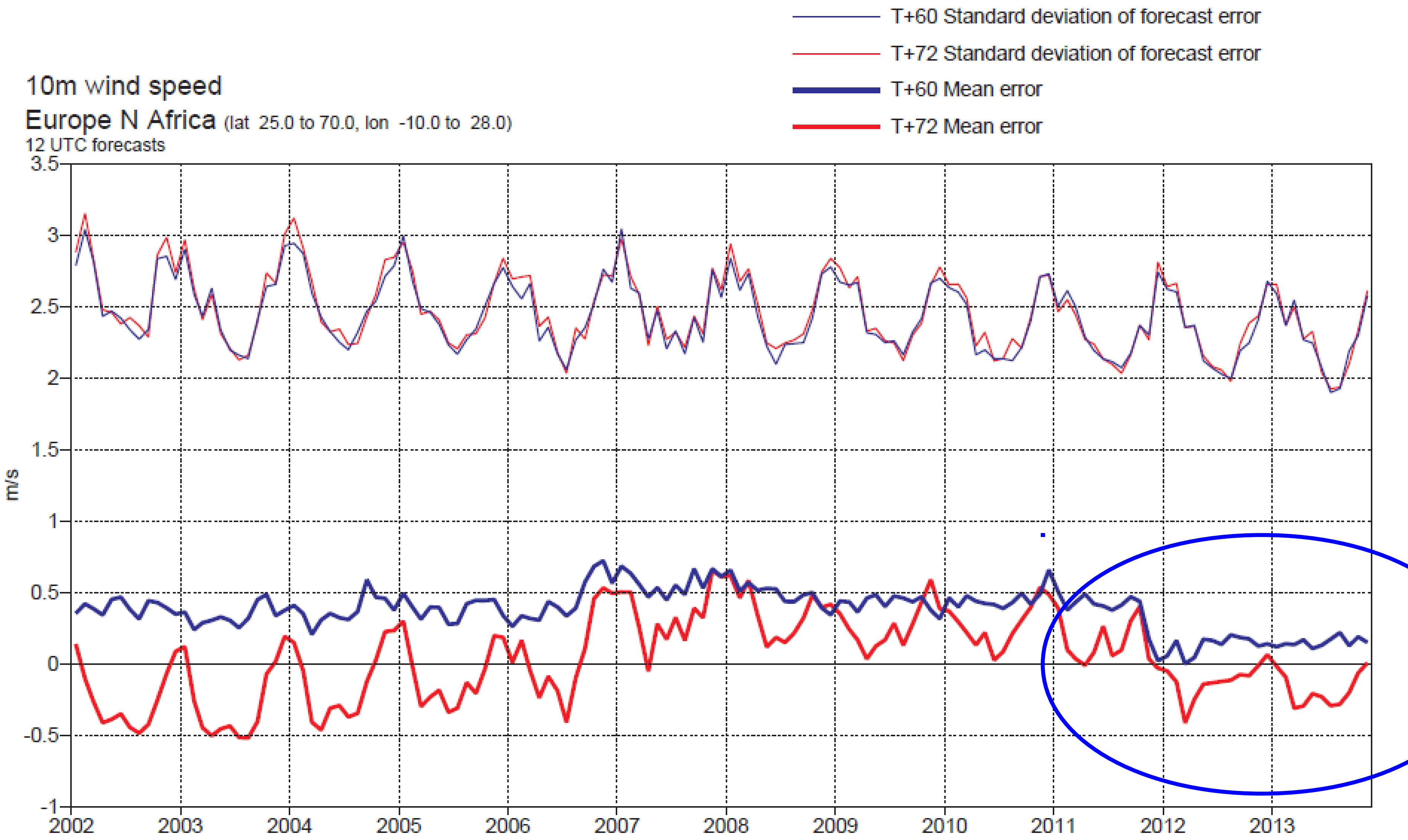
- Local wind depends strongly on local exposure.
- ECMWF model has roughness length parametrisation to obtain realistic "area averaged" surface drag.
- Resulting wind is low over land because rough elements dominate.



Post-processing of wind at 10 m

- Post-processed 10 m wind interpolates wind from 40 m (was 75 m before Nov. 2011)) assuming roughness length for grassland.
- Note: this exposure correction is only a partial correction to account for local effects (which tend to be more complex).

Changes to the roughness length table (Nov 2011)



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Wind Gusts: what is it ?

WMO definition:

Gusts are defined as wind extremes observed by anemometer. A 3 second running average is applied to the data. The report practice is such that gusts are reported as extremes over the previous hour, or the previous 3 or 6 hours.

The mean wind is reported as a 10 min average which is the last 10-minute interval of the hour; it should be comparable with instant output of the model 10 m wind, as it can be interpreted as some space and/or time average.

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Wind Gusts in the IFS

Gusts are computed by adding a turbulence component and a convective component to the mean wind:

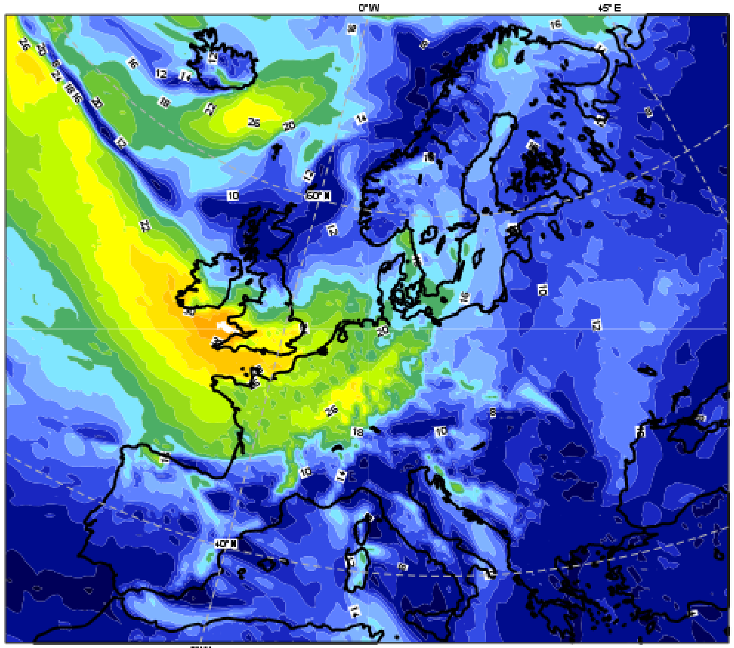
$$U_{gust} = U_{10} + 7.71\,U_{*}\,f(z/L) + \underbrace{0.6\max(0, U_{850} - U_{925})}_{deep\ convection}$$

where $\mathbf{U_{10}}$ is the 10m wind speed (obtained as wind speed at first model level, or interpolated down from 40m level), $\mathbf{U_{*}}$ is the friction velocity – itself obtained from the wind speed at the first model level, and **L** is a stability parameter.

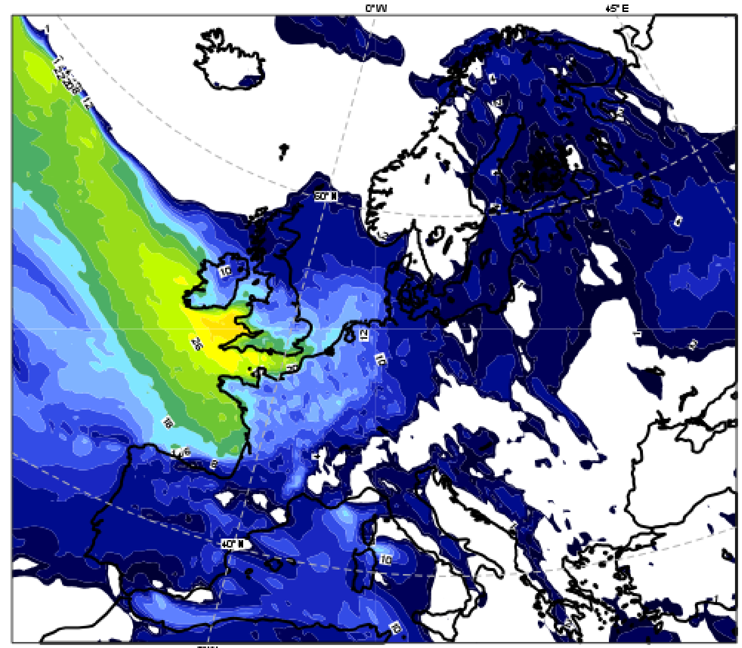
The convective contribution is set proport. to the wind shear between model levels corresponding to 850 hPa and 950hpa, respectively.

Wind gusts 8 Feb 2016 12 UTC

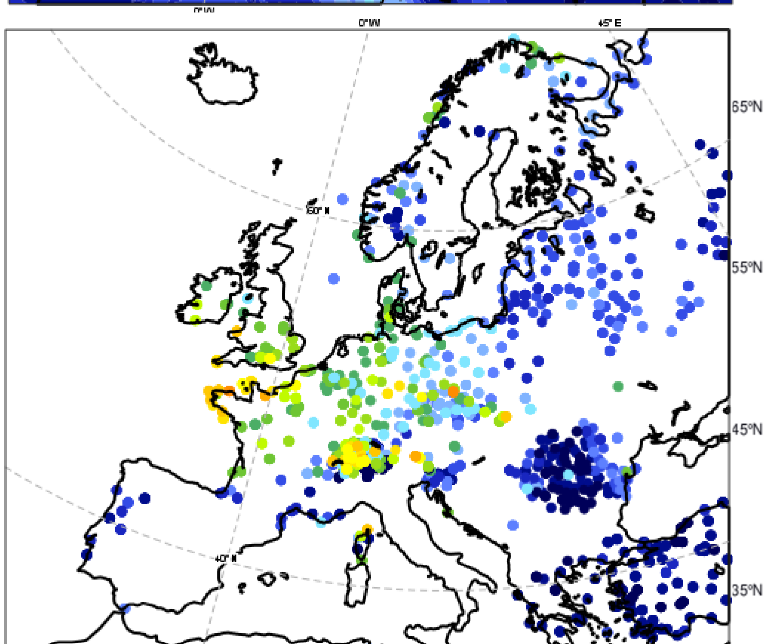
Gust

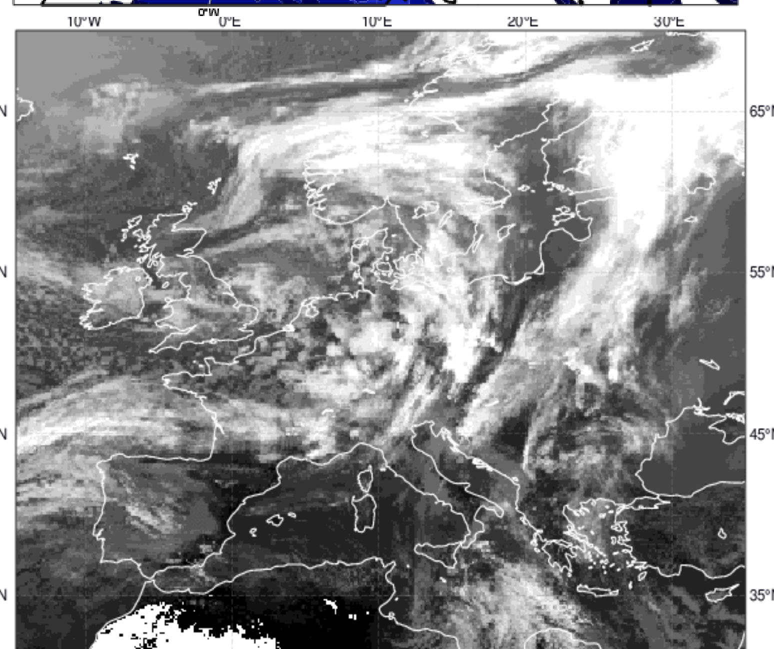


Wind speed



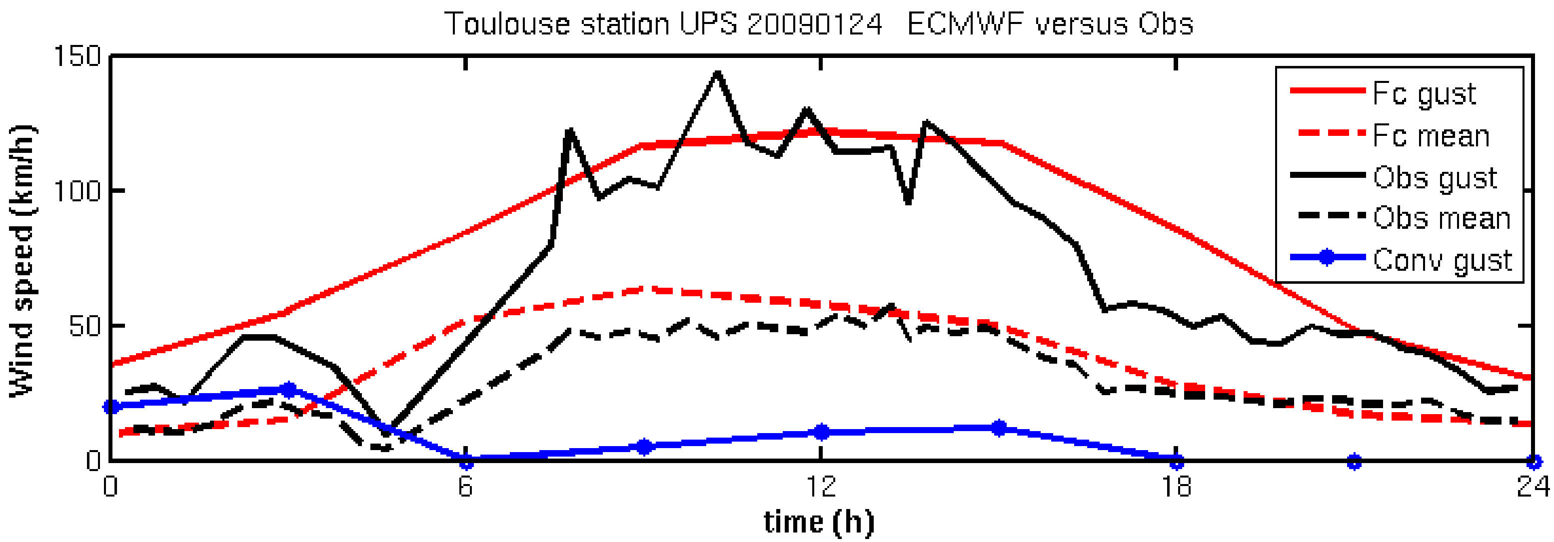
Obs





Wind gusts

Time series against anemometer 24 January 2009 (storm Klaus)



Observed mean wind speed (dashed black line) and maximum wind speed (solid black line) for 24 January 2009 at a meteorological station at Toulouse University, France (courtesy Jean-Luc Attié and Pierre Durand), together with corresponding 3-hourly forecast values (red lines) from the operational deterministic forecast from 23 January 12 UTC. The blue line denotes the convective contribution to the gusts.

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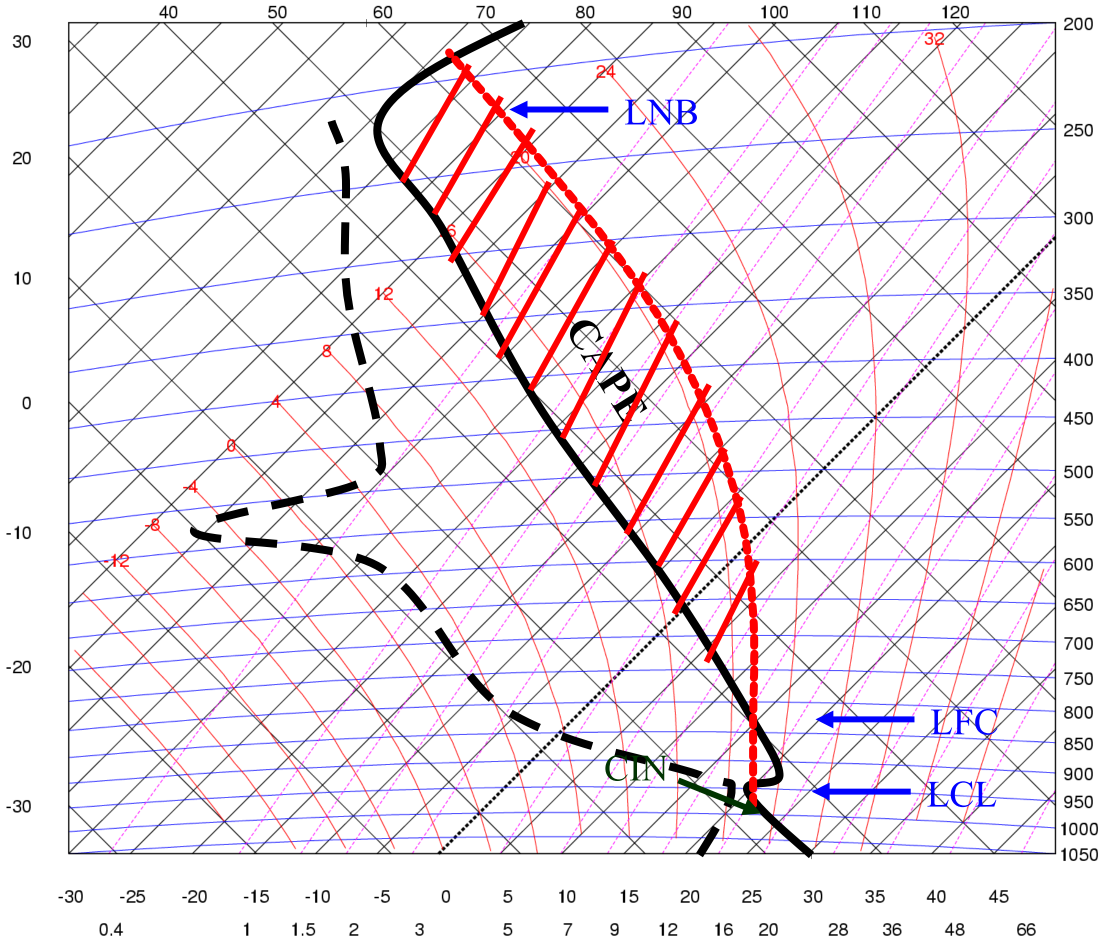
Physical processes: Summer and winter convection

Parcel convective In(stability): CAPE (CIN)

Idealised Profile



$$CAPE \approx \int_{base}^{top} g \frac{T_{cld} - T_{env}}{T_{env}} dz$$

In Thermodynamic diagram use T to compute CAPE, otherwise use virtual temperature T_v instead

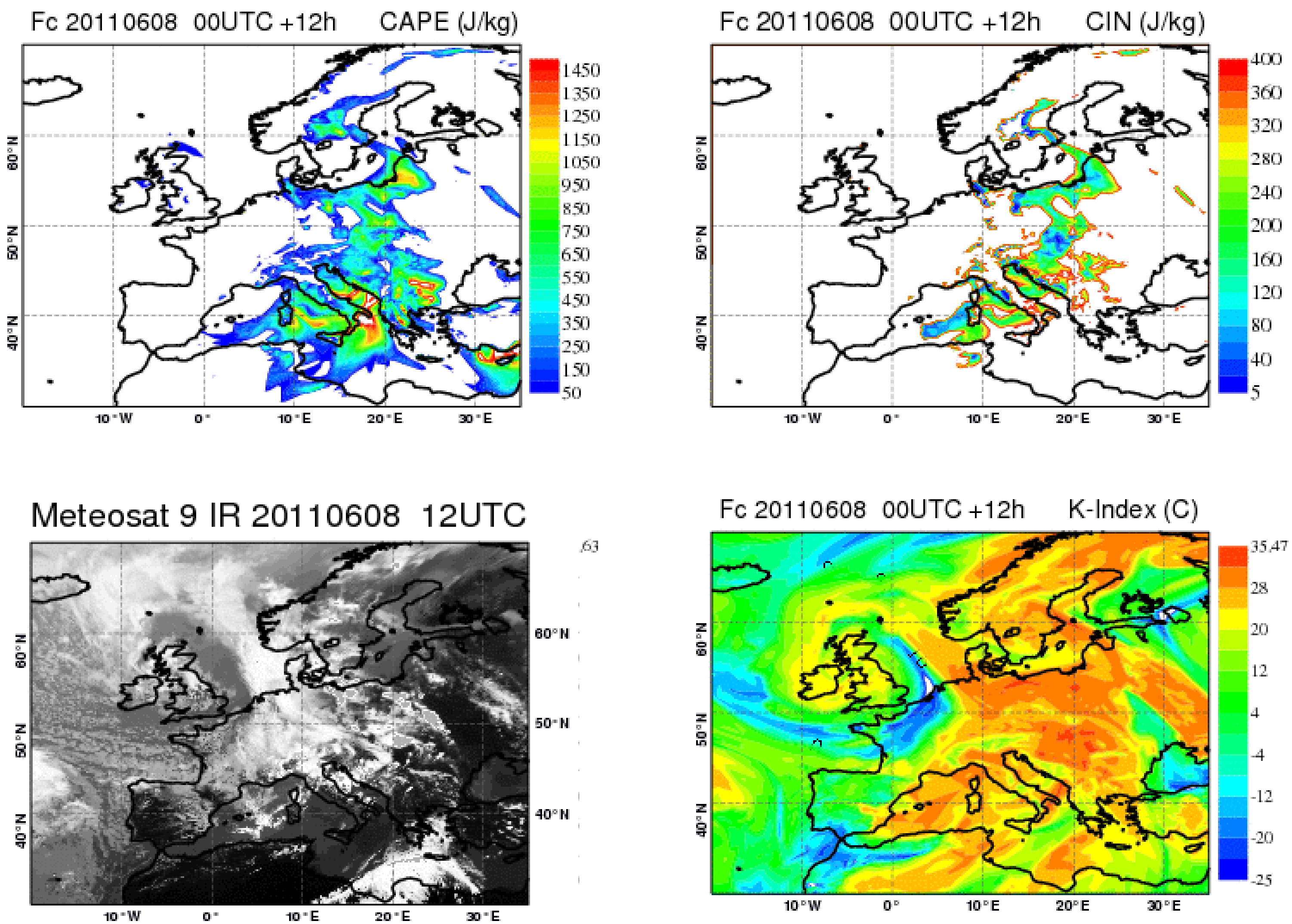
$$\frac{dw}{dt} = w\frac{dw}{dz} = \frac{1}{2}\frac{dw^2}{dz} \approx g\frac{T'}{\overline{T}}$$

$$w^2(z) = 2\int_0^z g\frac{T'}{\overline{T}} dz = 2 \cdot CAPE$$

Maximum updraught velocity (vertical velocity in cloud) $w = \sqrt{2 \cdot CAPE}$

In the IFS convection parameterization the amount of CAPE determines the intensity of convection (rainfall) – the computation of CAPE depends on the specified entrainment and the departure level of the air parcel (LCL=lifting condensation level, LFC=level of free convection, LNB=level of neutral buoyancy)

Convective Indices
requested by Member States (User Meeting June 2011)



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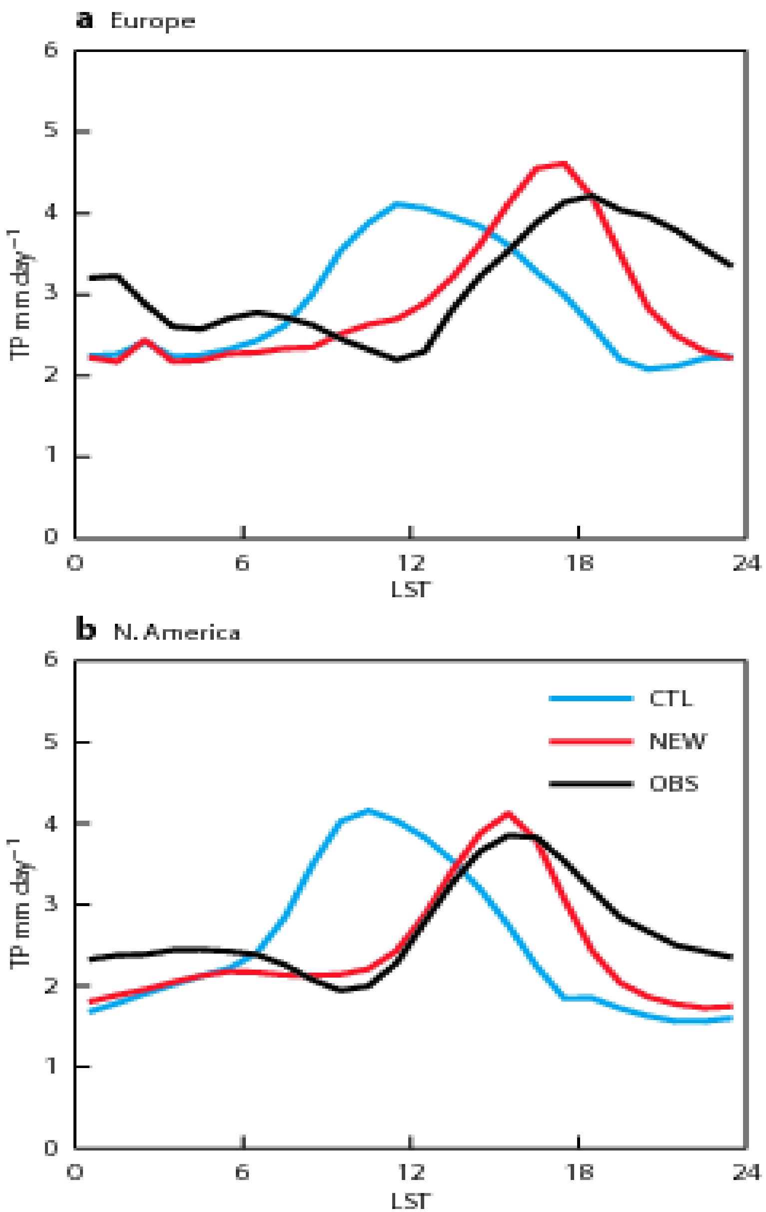
Diurnal cycle: realistic since Nov 2013



JJA 2011-2012 hourly rainfall composite against Radar

See ECMWF Newsletter No 136 Summer 2013
Bechtold et al., 2014, J. Atmos. Sci.

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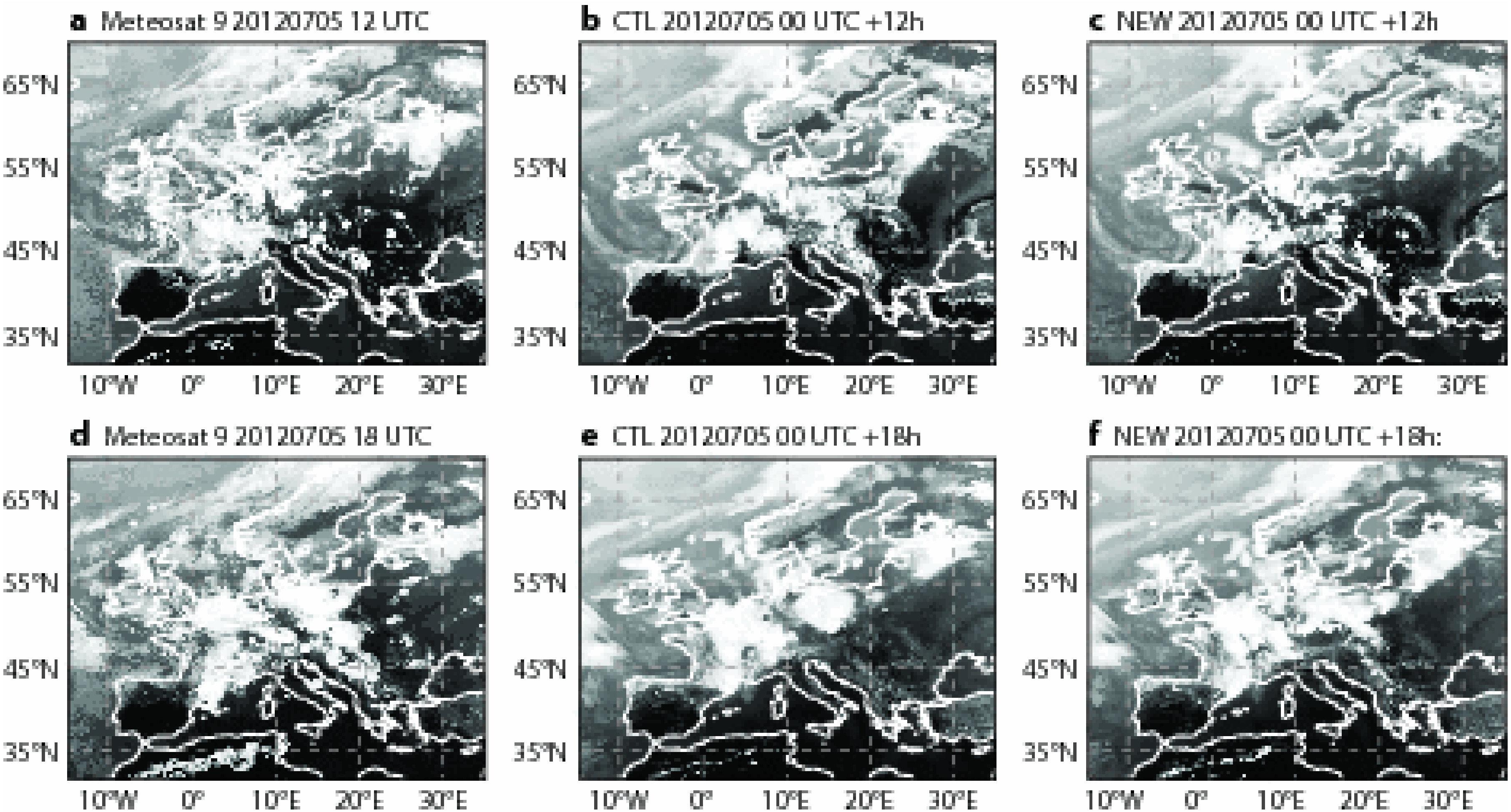
Diurnal cycle: Impact on weather forecasts



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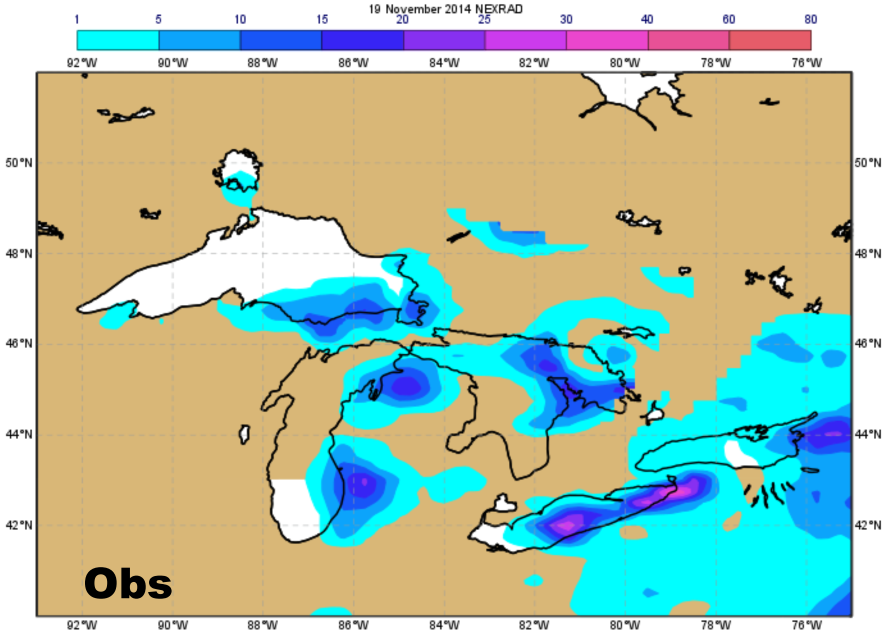
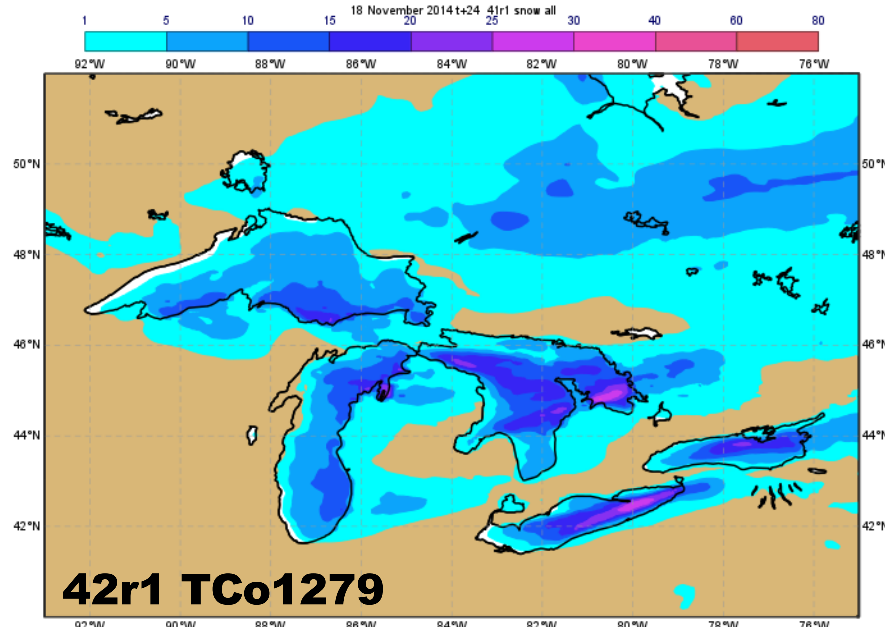
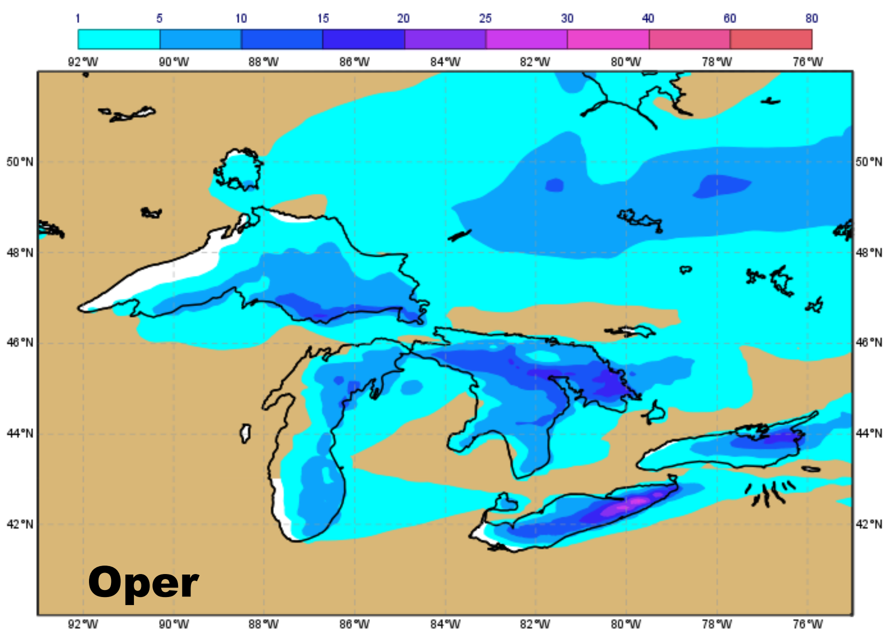
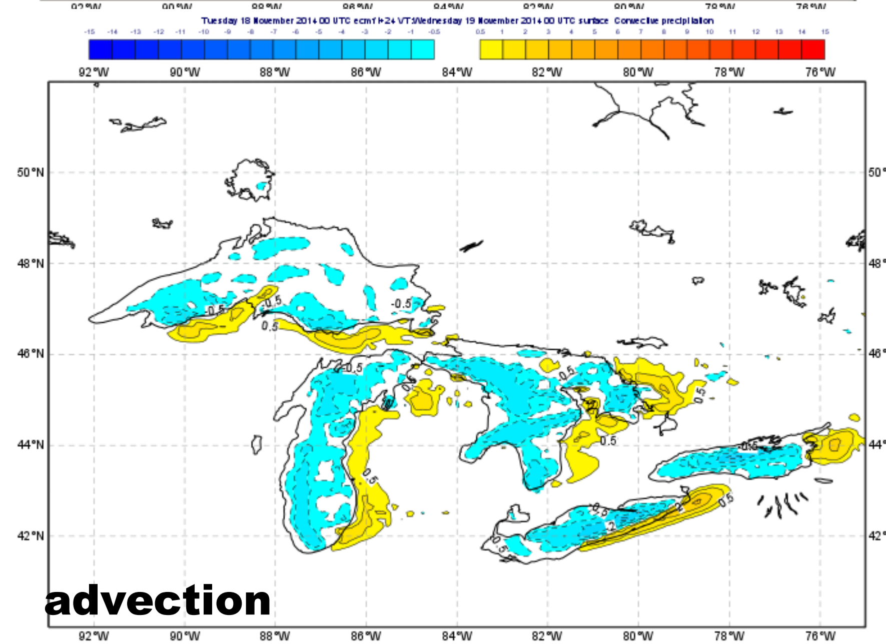
Winter convection: snow showers









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Summary: issues for improvement

- T2m winter can still be difficult: stable boundary-layer, coupling with surface (ground, lakes) and low-level clouds
- Still underestimation of convective night-time precip and some overestimation of light precipitation (drizzle)
- Inland penetration of (convective) showers and convective organisation improved but can still be improved
- Too strong Indian and SE Asian Summer Monsoon (some positive effect from new aerosol climatology in 2017)
- Melting of fresh snow on ground somewhat too slow
- and for long-range forecasts the coupling between the stratosphere and the troposphere

New products and things coming up in 2017

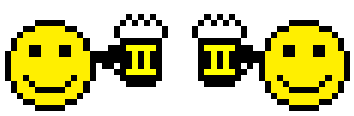
- New products: Ceiling (m), convective cloud top height (m), height of 0 and 1 Deg C wet bulb temperature, direct beam surface radiation

- New radiation scheme and Aerosol climatology -> improved (reduced precipitation) Indian summer monsoon

- Revised mixed phase for microphysics in convection

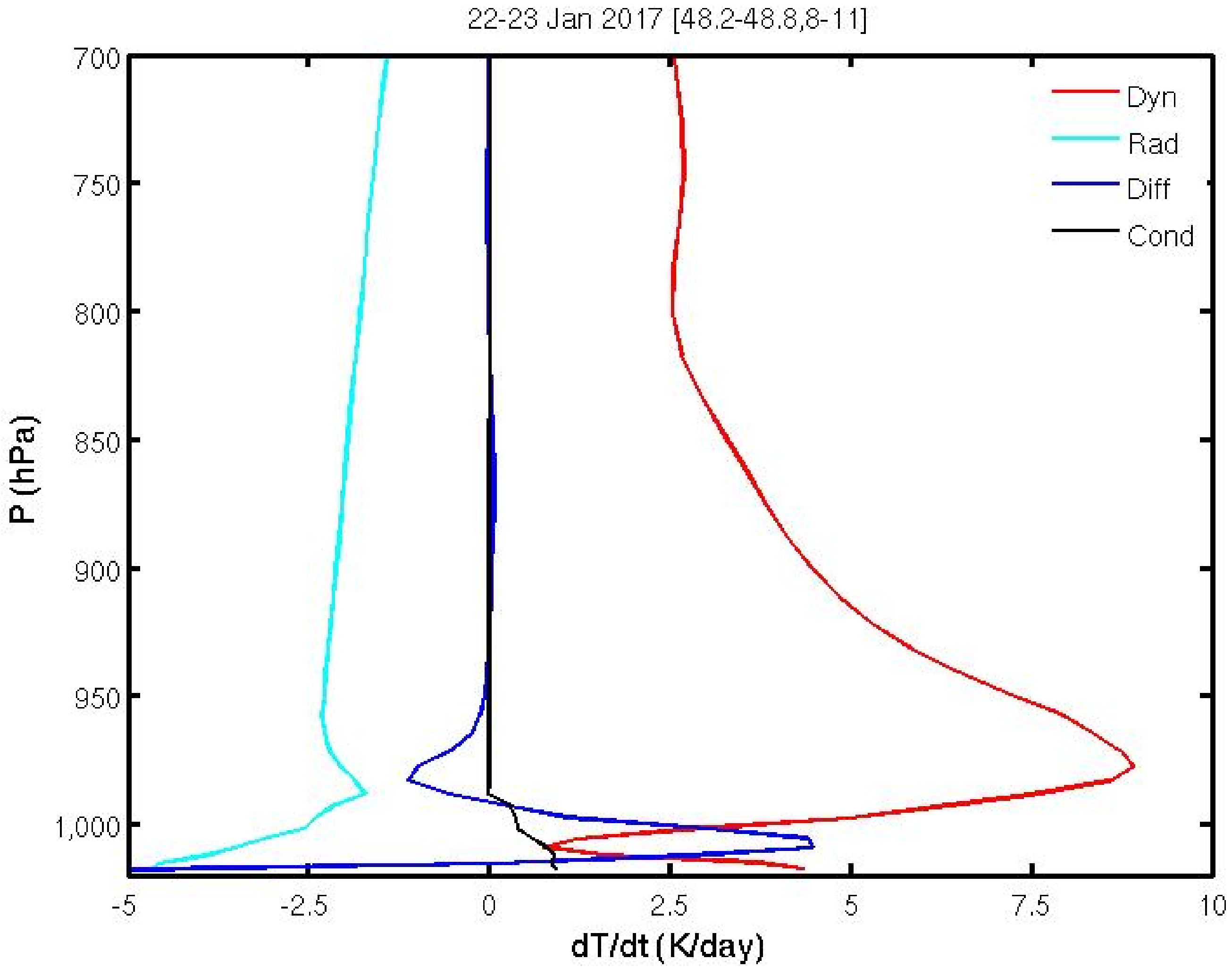
- Possible: coupling of HRES from t=0

- Possible: revised warm phase microphysics and revised boundary-layer clouds – shallow convection



ECMWF

Model tendencies during an inversion situation



ECMWF