Using stochastic physics to represent model uncertainty

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Using stochastic physics to represent model uncertainty

- Why represent model error in an ensemble forecast?
- What are the sources of model uncertainty?
- How do we represent model uncertainty?
 - 2 stochastic physics schemes in the IFS
- Impact of stochastic physics schemes in the IFS:
 - Medium-range ensemble (ENS)
 - Seasonal forecast (S4)
- Towards process-level simulation of model uncertainty

Ensemble reliability

• In a reliable ensemble, ensemble spread is a predictor of ensemble error



- Ensemble member
- Ensemble mean
- Observation

i.e. averaged over many ensemble forecasts,

 $e(\bar{x}) \approx \sigma(x)$

For a thorough discussion of this relationship:



Martin Leutbecher's lectures

Ensemble reliability

• In an under-dispersive ensemble,

 $e(\bar{x}) \gg \sigma(x)$



- Ensemble member
- Ensemble mean
- Observation

and ensemble spread does not provide a good estimate of error.

What happens when the ensemble includes no representation of model uncertainty?



What happens with no representation of model uncertainties?

Ensemble standard deviation ("Spread")



TL399/255, resolution change at D15, 20 members

For details of skill measures:

Martin Leutbecher's lectures

What happens with no representation of model uncertainties?

Probabilistic skill (CRPS)



TL399/255, resolution change at D15, 20 members

For details of skill measures:

Martin Leutbecher's lectures

Model uncertainty: where does it come from?

• Atmospheric processes parametrised in the model:





Model error: where does it come from?

- Other parametrised atmospheric processes?
 - Surface coupling
 - Radiation-aerosol interactions
 - ...
- Other sources:
 - Dynamics / numerics
 - Coupled system: land-surface / oceans / sea-ice
- Other sources: processes not captured by the underlying model?
 - Atmosphere exhibits upscale propagation of kinetic energy (KE)
 - Occurs at ALL scales: no concept of "resolved" and "unresolved" scales
 - How can the model represent upscale KE transfer from unresolved to resolved scales?

Model uncertainty: how to simulate it?

- What do errors due to model uncertainty look like?
- Can we characterize them: relative size and timescales associated with different sources?
- How can we represent them?
- Multi-model ensembles
- Multi-physics ensembles
- Perturbed parameter ensembles
- "Stochastic parametrisations"

Stochastic physics schemes in IFS

- IFS ensemble forecasts (ENS and S4) include 2 model uncertainty schemes:
 - Stochastically perturbed parametrisation tendencies (SPPT) scheme
 - Stochastic kinetic energy backscatter (SKEB) scheme
- SPPT scheme: simulates uncertainty due to sub-grid parametrisations
- SKEB scheme: parametrises missing and uncertain upscale transfer of KE

Stochastically Perturbed Parametrisation Tendencies (SPPT) scheme

- Initially implemented in IFS, 1998 (Buizza et al., 1999); revised in 2009:
- Simulates model uncertainty due to physics parameterisations by
 - taking the net tendencies from the physics parametrisations:

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\boldsymbol{X} = \begin{bmatrix} X_U, X_V, X_T, X_Q \end{bmatrix}
coming from radiation schemes
gravity wave drag
vertical mixing
convection
cloud physics
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• and perturbing with multiplicative noise $r \in [-1, +1]$ as:

$$\mathbf{X}' = (1 + \mu r)\mathbf{X}$$

where $\mu \in [0,1]$ tapers the perturbations to zero near the surface & in the stratosphere.

Shutts et al. (2011, ECMWF Newsletter); Palmer et al., (2009, ECMWF Tech. Memo.)

- 2D random pattern in spectral space:
- First-order auto-regressive [AR(1)] process for evolving spectral coefficients \hat{r} $\hat{r}(t + \Delta t) = \phi \hat{r}(t) + \rho \eta(t)$

where $\phi = \exp(-\Delta t/\tau)$ controls the correlation over timestep Δt ;

and spatial correlations (Gaussian) for each wavenumber define ho for random numbers, η

- Resulting pattern mapped into grid-point space *r*:
- clipped such that $r \in [-1, +1]$
- same pattern is applied to T, q, u, v
- applied at all model levels to preserve vertical structures**
- ***Except*: tapered to zero at model top/bottom, to avoid:
 - instabilities due to perturbations in the boundary layer;
 - perturbing stratospheric tendencies dominated by well-constrained clear-skies radiation

- 2D random pattern, *r*:
- Time-correlations: AR(1)
- Spatial-correlations: Gaussian
- Clipped such that $r \in [-1, +1]$
- Applied at all model levels to preserve vertical structures**

***Except*: tapered to zero at model top/bottom



3 correlation scales:

i)	6 hours,	500 km,	$\sigma = 0.52$
ii)	3 days,	1 000 km,	$\sigma = 0.18$
iii)	30 days,	2 000 km,	$\sigma = 0.06$



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Stochastic Kinetic Energy Backscatter (SKEB) scheme

- Introduced into IFS, 2010:
- Attempts to simulate a process otherwise absent from the model –

upscale transfer of energy from sub-grid scales

• Represents backscatter of Kinetic Energy (KE) by adding perturbations to U and V via a forcing term to the streamfunction:

$$F_{\varphi} = \left(b_R D\right)^{1/2} F^*$$

where

D is an estimate of the (smoothed) total (local) dissipation rate due to the model,

 $b_{\rm R}$ is the backscatter ratio – a scaling factor,

 F^* is a 3D evolving random pattern field.

Shutts et al. (2011, ECMWF Newsletter); Palmer et al., (2009, ECMWF Tech. Memo.); Shutts (2005, QJRMS); Berner et al. (2009, JAS)

SKEB scheme

$$F_{\varphi} = \left(b_R D\right)^{1/2} F^*$$

- 3D random pattern field F^* :
 - First-order auto-regressive [AR(1)] process for evolving F^* $F^*(t + \Delta t) = \phi F^*(t) + \rho \eta(t)$

where $\phi = \exp(-\Delta t/\tau)$ controls the correlation over timestep Δt ;

and spatial correlations (power law) for wavenumbers define ρ for random numbers, η

- vertical space-(de)correlations: random phase shift of η between levels

SKEB perturbations

$$F_{\varphi} = \left(b_R D\right)^{1/2} F^*$$

- *D* is an estimate of sub-grid scale production of KE, and includes:
- D_{num} = numerical dissipation from
 - explicit horizontal diffusion (bi-harmonic, ∇^2); and
 - estimate due to semi-Lagrangian interpolation error
- D_{con} = estimated KE generated by updraughts and detrainment within sub-grid deep convection
- Note: as of the resolution upgrade (32 -> 19 km) in March 2016:
 - New numerical diffusion operator is no longer consistent with the biharmonic diffusion assumed by SKEB (for D_{num}) => numerical dissipation contribution has been deactivated

How are the perturbation patterns determined?

• Characteristics of errors due to model uncertainty cannot easily be determined from observations:

- uncertain processes small-scale (space and time)
- lack of observational coverage

- Can attempt to use models: coarse-graining studies (e.g. Shutts and Palmer, 2007)
 - take high-resolution model simulations as "truth"
 - average model fields and tendencies (or streamfunction) to a grid-resolution typical of the forecast model
 - compare the contribution of "sub-grid" scales in the coarse-grained simulation with parametrisations in the forecast model
 - coarse-graining studies were used to justify and inform scales in SPPT and SKEB

IFS ensembles: ENS and System 4 (S4)

- **ENS** = ensemble prediction system for
 - medium-range forecasts (up to 15 days) and
 - monthly forecasts (up to 32 days) [Frederic Vitart's lecture]
- **S4** = seasonal forecasting system [Tim Stockdale's lecture]
 - up to 7 months
- Both forecast systems include representations of model uncertainty via SPPT and SKEB

• ENS:

- 1 control forecast + 50 perturbed members
- T_{co}639 (~19 km) resolution to day 15; T_{co}319 (~32 km) days 45
- 91 vertical levels, up to 0.01hPa

Impact of SPPT and SKEB in ENS

Ensemble standard deviation ("Spread")





Impact of SPPT and SKEB in ENS Probabilistic skill (CRPS)



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Impact of SPPT and SKEB in ENS

- Adding SPPT + SKEB perturbations:
 - increases ensemble "spread" (= ensemble standard deviation),
 - i.e. ensemble members describe greater region of the parameter space
 - some reduced ensemble mean errors
 - SPPT has a much greater impact than SKEB
- In the extra-tropics:
 - Experiments: perturbations in days 0-5 contribute most effect
- In the tropics:
 - Experiments: effect of perturbations rapidly lost at all times

Impact of SPPT and SKEB in S4

- System 4 (S4), November 2011: introduction of (revised) SPPT and SKEB
- Operational configuration:
 - T255 (~80 km), 91 vertical levels (up to 0.01 hPa)
 - Coupled ocean model: NEMOv3.0, 1 degree (~110 km), 42 vertical levels
 - 51 members
 - Initialised on 1st of each month
 - Forecast lead times: to 7 months
- Recent work with S4 to assess impact of stochastic schemes
- For longer time-scales, consider impact in terms of:
 - Noise-induced drift, i.e. change in model mean
 - Noise-activated regime transition, e.g. Pacific-N. American region regimes

Impact of SPPT and SKEB in S4

- Recent work with S4 to assess impact of stochastic schemes:
 - Hindcast period: 1981-2010
 - Start dates: May, Aug & Nov
 - Ensemble size: 51
 - Verification of forecasts to lead times: 4-7 months
- Considers impact of SPPT + SKEB on:
 - Systematic errors
 - Madden-Julian Oscillation (MJO) statistics
 - ENSO forecast quality
 - Circulation regimes over the Pacific-North American region [Franco Molteni's lecture]

Impact of SPPT and SKEB in S4: Systematic errors

- Activating SPPT + SKEB reduces some biases:
 - Outgoing longwave radiation (OLR)
 - Total cloud cover
 - Total precipitation
 - Zonal winds (850 hPa)
- Greatest improvements in the tropics: reduces overly active tropical convection
- SPPT is responsible for most of the difference

See Weisheimer et al. (2014, Phil. Trans. R. Soc. A)

Impact of SPPT and SKEB in S4: Madden Julian Oscillation



*stochphys*OFF

Wheeler and Hendon Index:

Projection of daily data on 2 dominant combined EOFs

of OLR, u200 and u850 over 15°N-15°S

Weisheimer et al. (2014, Phil. Trans. R. Soc. A)



Impact of SPPT & SKEB in S4: Increased amplitude of MJO events



Impact of SPPT & SKEB in S4: ENSO forecast quality - Niño4 SSTs





System 4 has:

- Reduced forecast errors
- Increased ensemble spread
- Improved correlation

Impact of SPPT & SKEB in S4: Pacific North America (PNA) circulation regimes



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Stochastic physics: summary

- Errors due to model uncertainty arise from unresolved and misrepresented processes
 - finite-resolution of a discrete numerical model
 - parametrisations use simplified, bulk methods to represent complex, multi-scale sub-grid processes
- Difficult to characterise sources of model uncertainty due to lack of observations
- Without representing model uncertainty, ensemble forecasts are under-dispersive
- Stochastic methods for representing model uncertainty improve ensemble reliability
- ECMWF ensembles include 2 stochastic physics schemes:
 - SPPT: represents uncertainty due to sub-grid physics parameterisations
 - SKEB: simulates upscale transfer of kinetic energy from unresolved scales
- Medium-range: increased ensemble spread, greater probabilistic skill
- Seasonal: reduction in biases; better representation of MJO, ENSO, PNA regimes

Stochastic physics: brief outlook for IFS

Towards process-level model uncertainty representation



- **Aim**: to improve the physical consistency
- Generate flux perturbations at the top of atmosphere (TOA) and surface that are consistent with tendency perturbations within the atmospheric column
- Conservation of water
- Remove ad hoc tapering in boundary layer and stratosphere
- Include multi-variate aspects of uncertainties

Stochastic physics: brief outlook for IFS

Towards process-level model uncertainty representation



Approach:

Stochastically Perturbed Parametrisations (SPP) (Ollinaho et al., submitted QJ, 2016)

- Embed stochasticity within IFS parametrisations
- Perturb parameters/variables directly
- Specify spatial/temporal correlations
- Target uncertainties that matter (level of uncertainty and impact)
- Require that stochastic schemes converge to deterministic schemes in limit of vanishing variance

Stochastically Perturbed Parametrisations (SPP) scheme

Towards process-level model uncertainty representation

Stochastic perturbations are applied to unperturbed parameters / variables in the physics parametrisations, $\hat{\xi}_j$:

$$\xi_j = \hat{\xi}_j \exp(\Psi_j)$$

where

 $\Psi_j \sim \mathcal{N}(\mu_j, \sigma_j^2)$

Development started with parameter perturbations to target cloudy-skies radiation

Now includes parameters/variables from:

- Turbulent diffusion and subgrid orography
- Cloud and large-scale precipitation
- Radiation
- Convection



(Ollinaho et al., submitted QJ, 2016)

Stochastically Perturbed Parametrisations (SPP) scheme

Towards process-level model uncertainty representation



- Standard deviation of 0-3h Temperature tendency
- SPP induces larger (smaller) tendency perturbations within (above) the boundary layer than SPPT
- Correlations between SPP and SPPT standard deviations are small at early lead times => two schemes are generating different perturbation structures

Based on 6 boreal winter cases;

Unit (top panels): K/3h



References & reading

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