

A satellite view of Earth's clouds, showing a dense, swirling pattern of white and light blue clouds against a darker blue background. The clouds are concentrated in the center and spread out towards the edges, creating a complex, textured appearance.

**Numerical Weather Prediction
Parametrization of Subgrid Physical Processes**

Clouds (2)
Sub-grid Cloud Cover
(or “Sub-grid heterogeneity of cloud and humidity”)

Richard Forbes

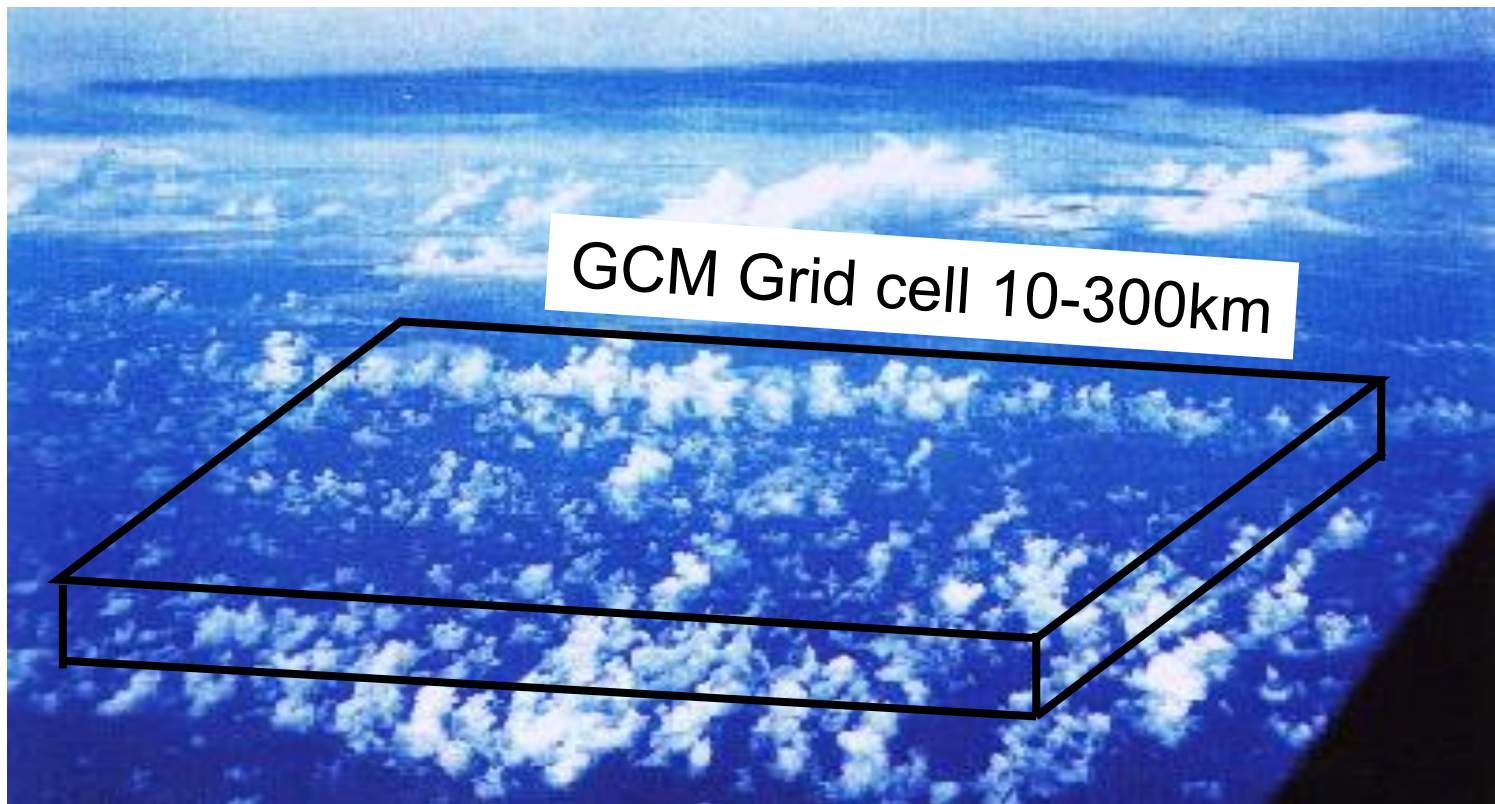
(With thanks to Adrian Tompkins
and Christian Jakob)

forbes@ecmwf.int

Clouds in GCMs: Representing sub-grid heterogeneity



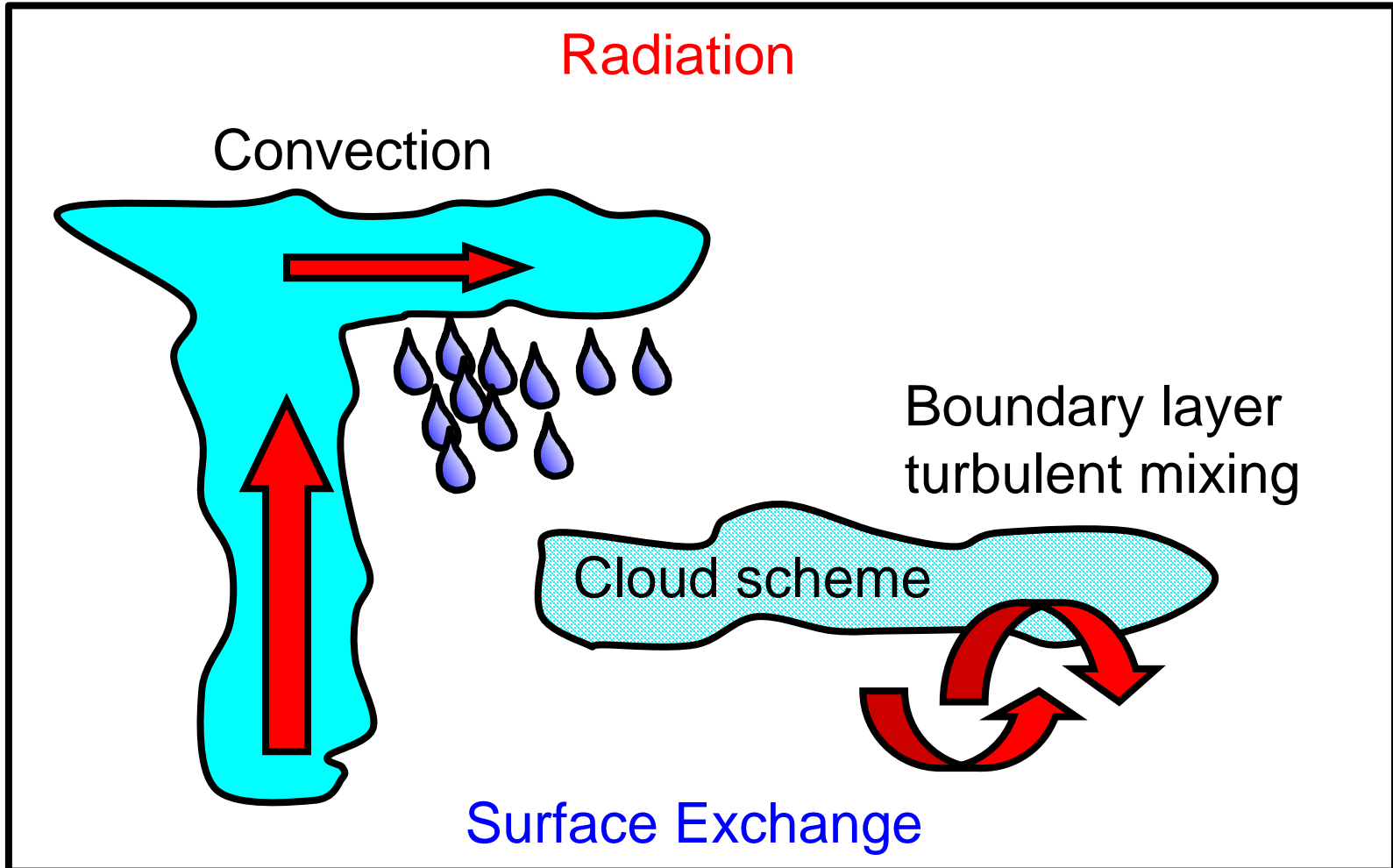
Many of the observed clouds and especially the processes within them are of **subgrid-scale size** (both horizontally and vertically)



Clouds in GCMs: Representing sub-grid heterogeneity



Many heterogeneity assumptions across the model parametrizations...



Why represent heterogeneity?

Important scales of cloud cover & reflectance

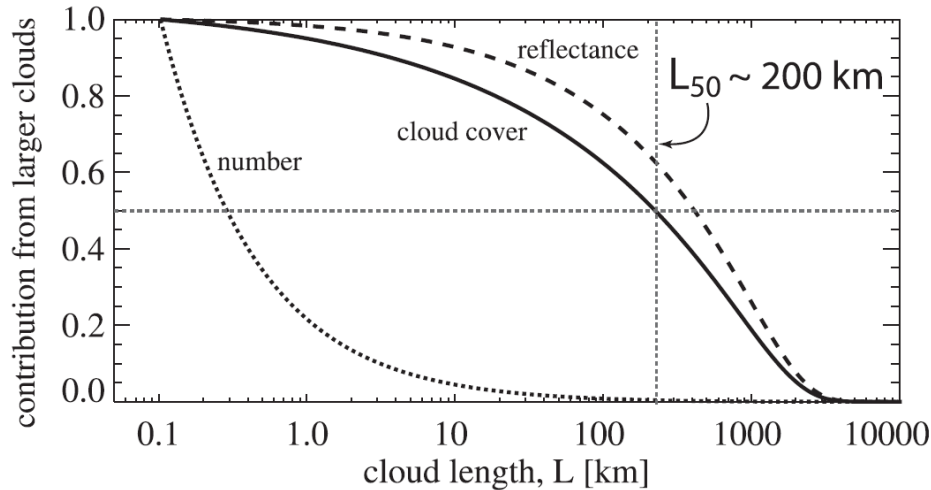


Fig 6. Contribution to global cloud cover (solid), number (dotted) and visible reflectance (dashed) from clouds with chord lengths greater than L (based on MODIS, aircraft and NWP data).

(from Wood and Field 2011, JClimate)

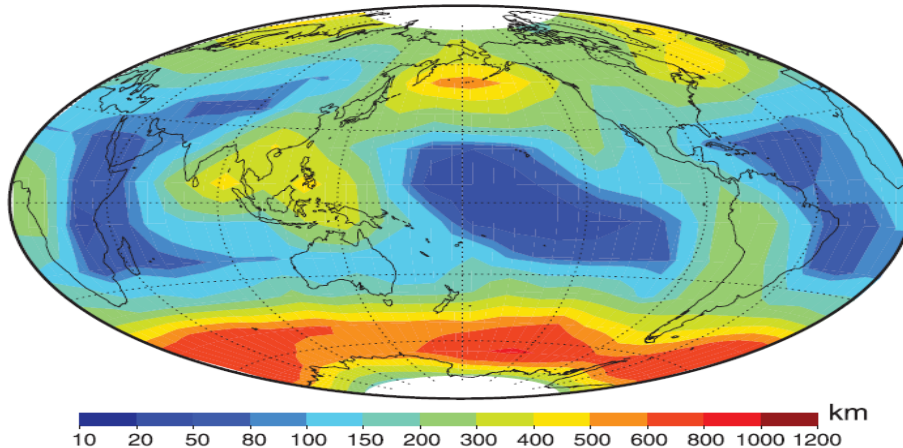


Fig 8. Map of the cloud size for which 50% of cloud cover comes from larger clouds (from 2 years of MODIS data)

15% of global cloud cover comes from clouds smaller than 10 km
(smaller scales dominate over subtropical ocean)

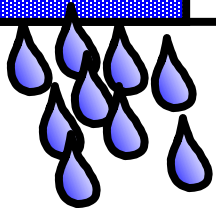
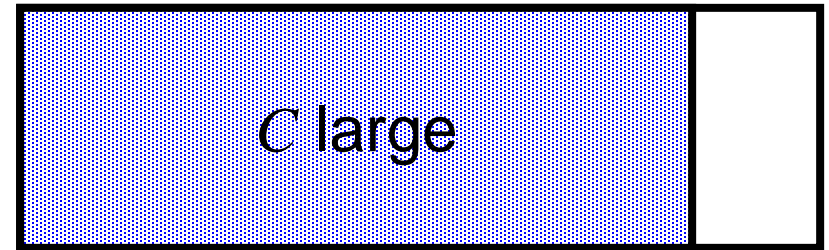
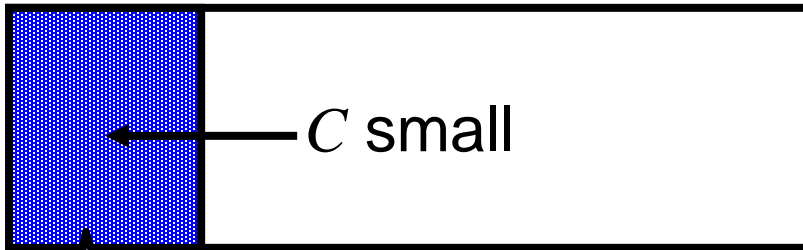
Why represent heterogeneity?

Important for microphysics



Imagine a cloud with condensate mass q_l and cloud fraction C
The in-cloud mass mixing ratio is q_l/C

GCM grid box



precipitation not equal in each case since
cloud-to-rain autoconversion is **nonlinear**

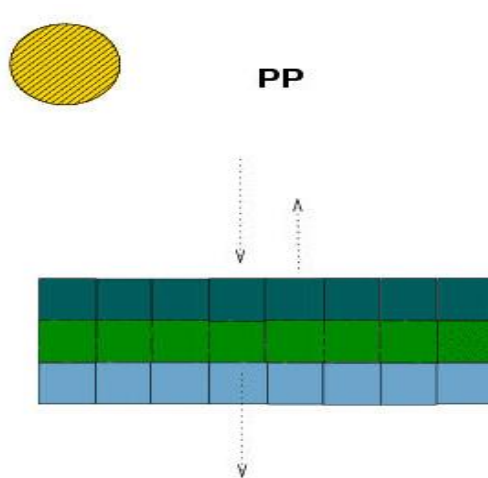
- Complex microphysics perhaps a wasted effort if assessment of cloud fraction C is poor!
- In addition, in-cloud condensate heterogeneity should also be represented, i.e. not all the cloud is precipitating?

Why represent heterogeneity?

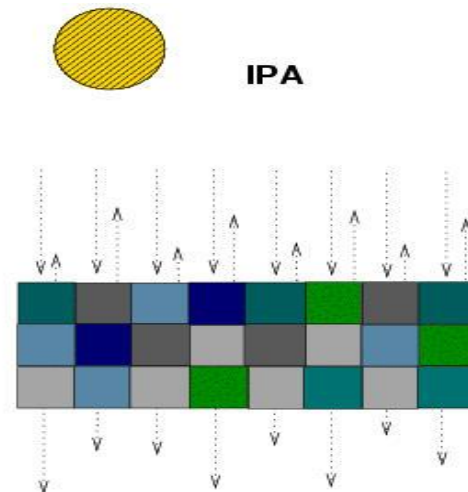


Important for radiation

- Assuming homogeneity can lead to biased radiative calculations (e.g. Cahalan et al. 1994, Barker et al 1996).
- Monte Carlo Independent Column Approximation, for example, can treat the inhomogeneity of in-cloud condensate and vertical overlap in a consistent way between the cloud and radiation schemes



Traditional approach
(homogeneous)



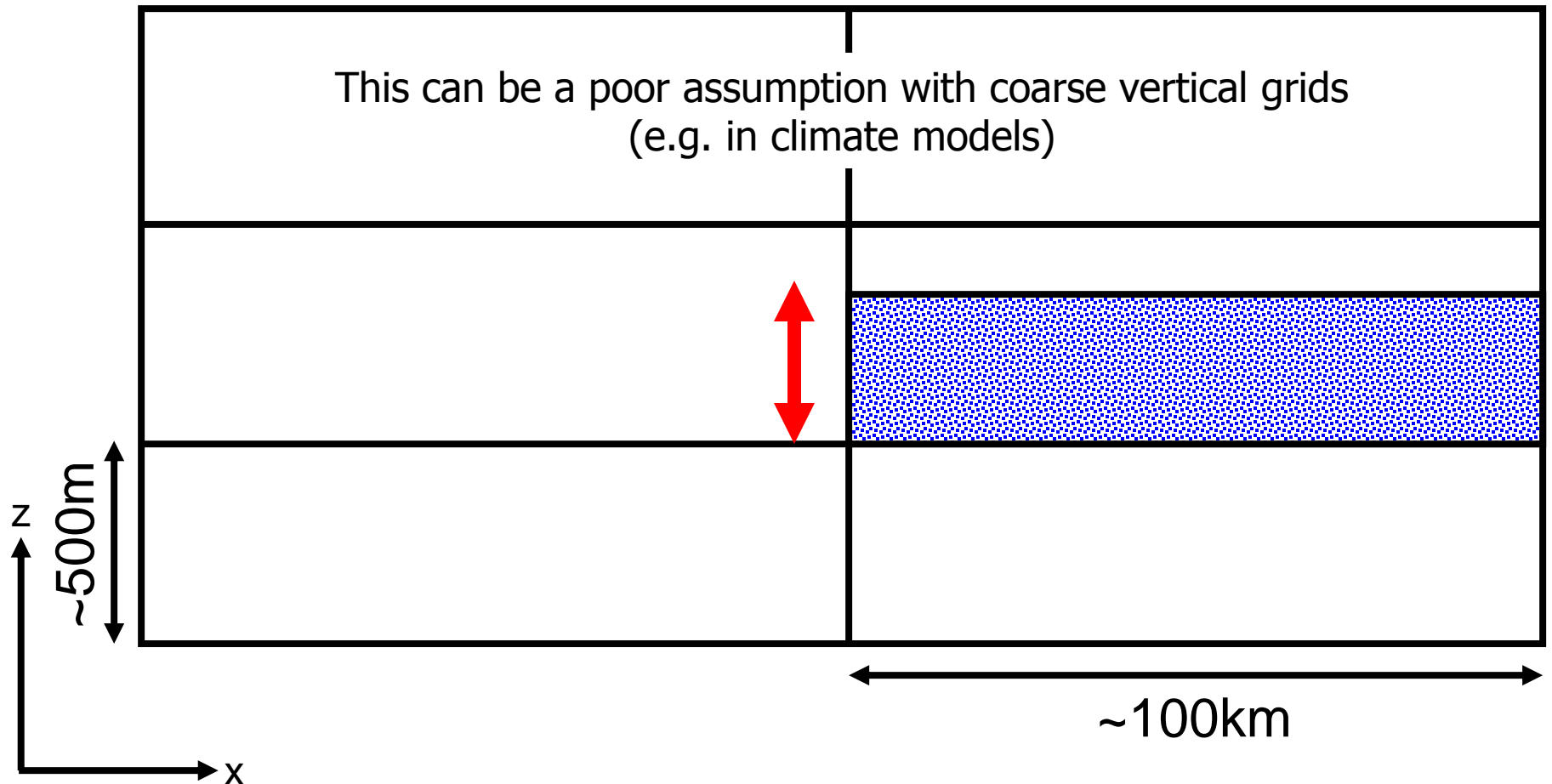
Independent Column
Approximation, e.g.
MCICA

Macroscale Issues of Parameterization



VERTICAL COVERAGE

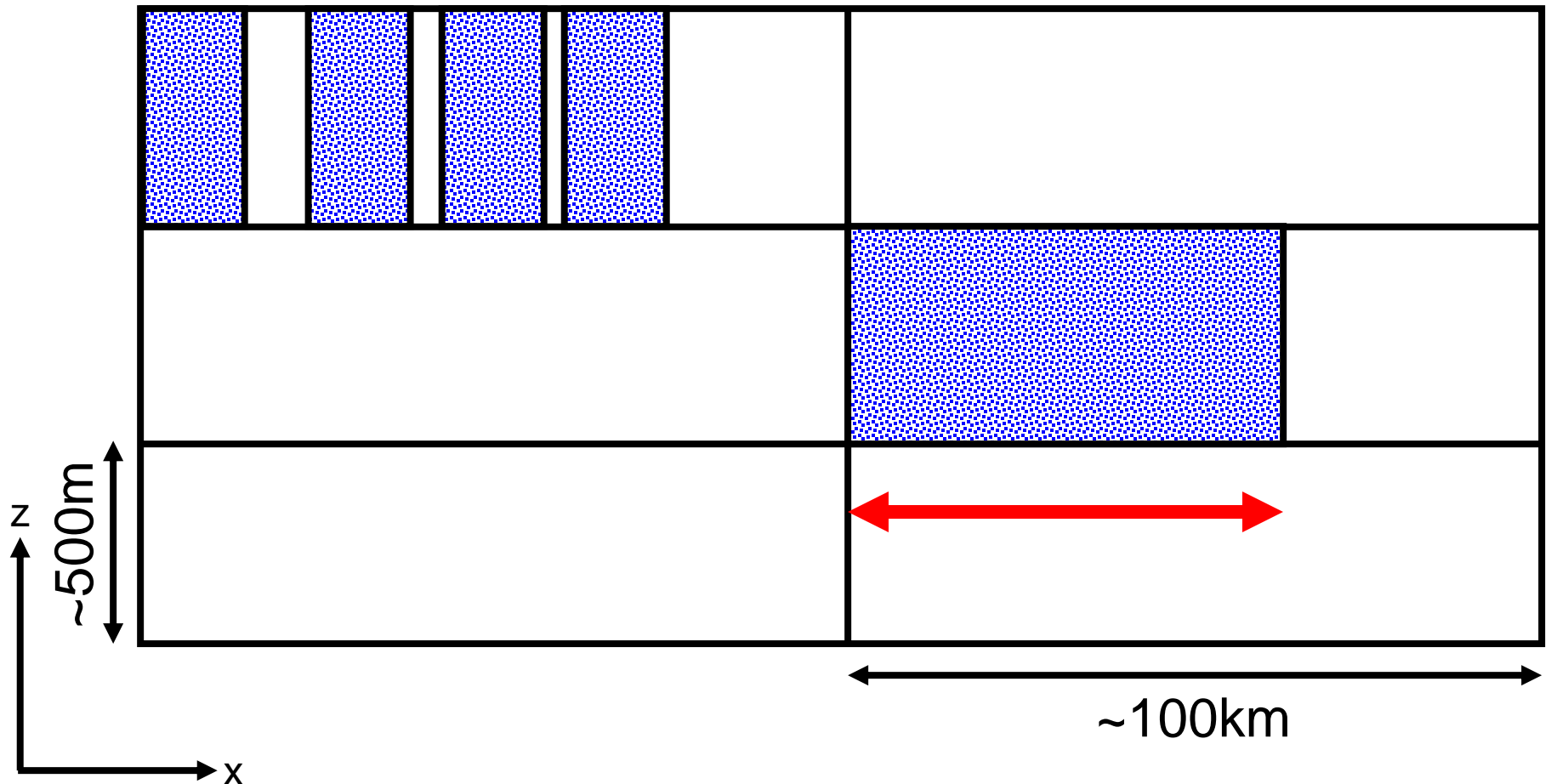
Most models assume that this is 1



Macroscale Issues of Parameterization



HORIZONTAL COVERAGE, C
Spatial arrangement ?

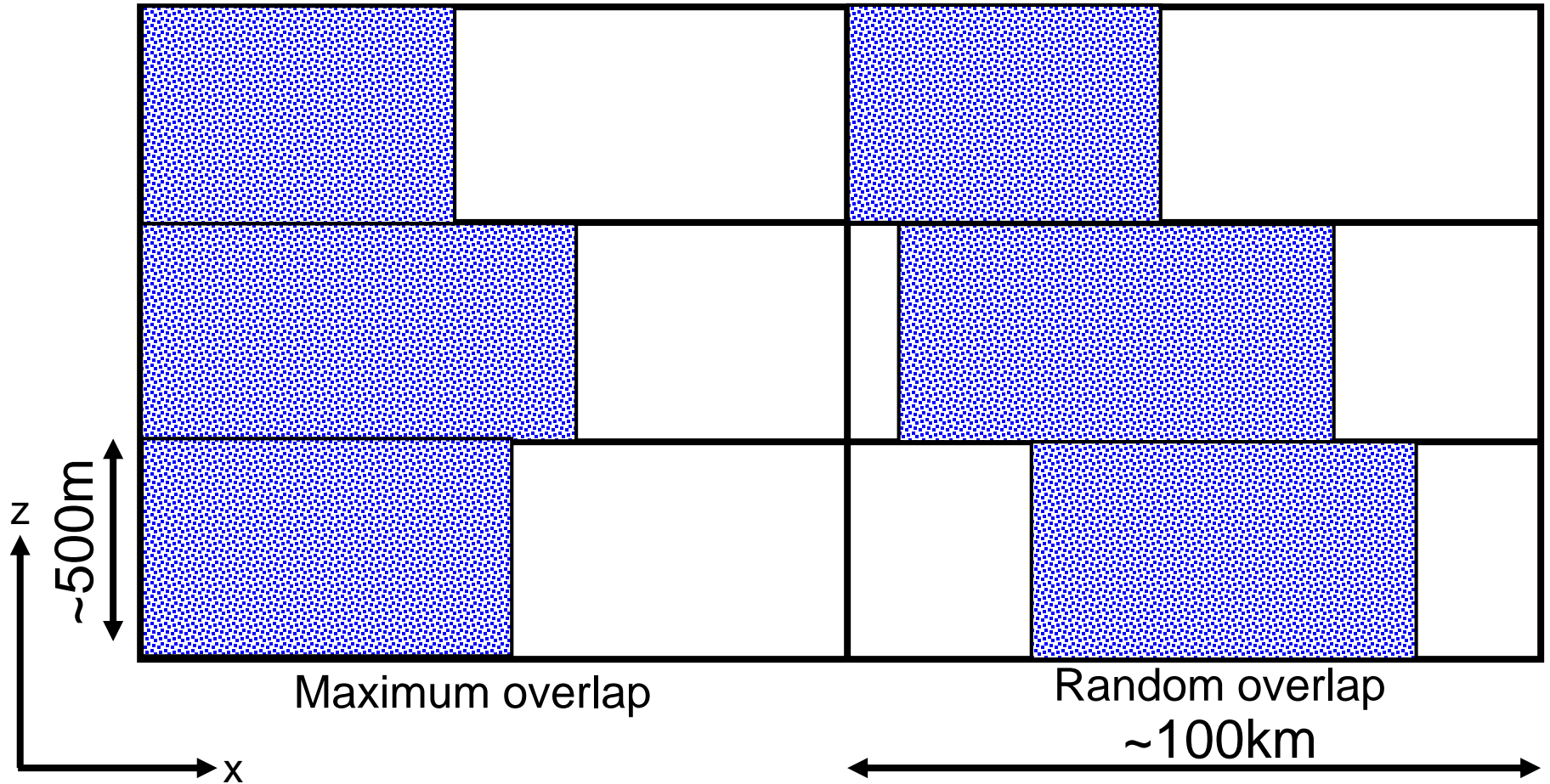


Macroscale Issues of Parameterization



Vertical overlap of cloud

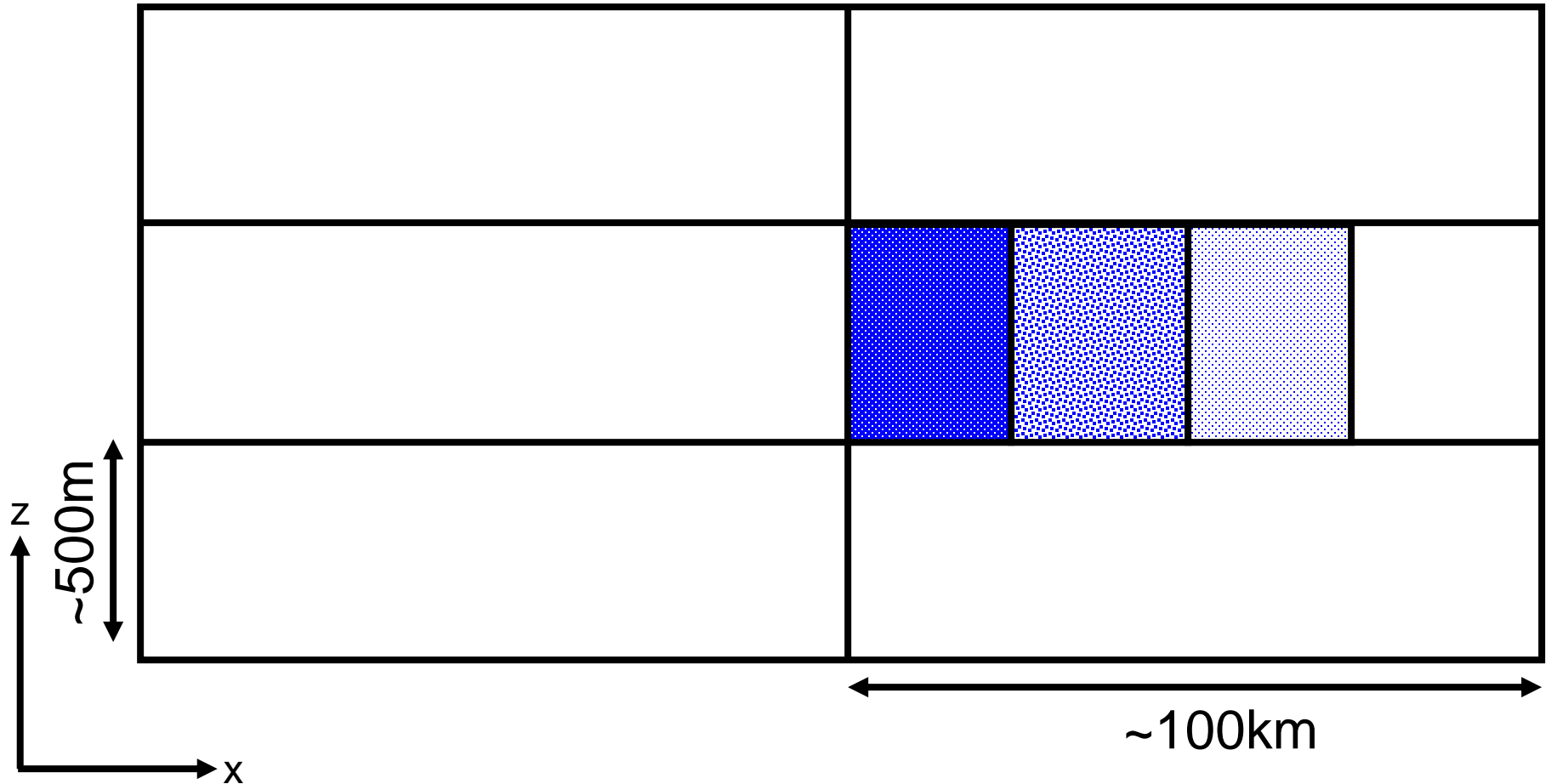
Important for radiation and microphysics interaction



Macroscale Issues of Parameterization



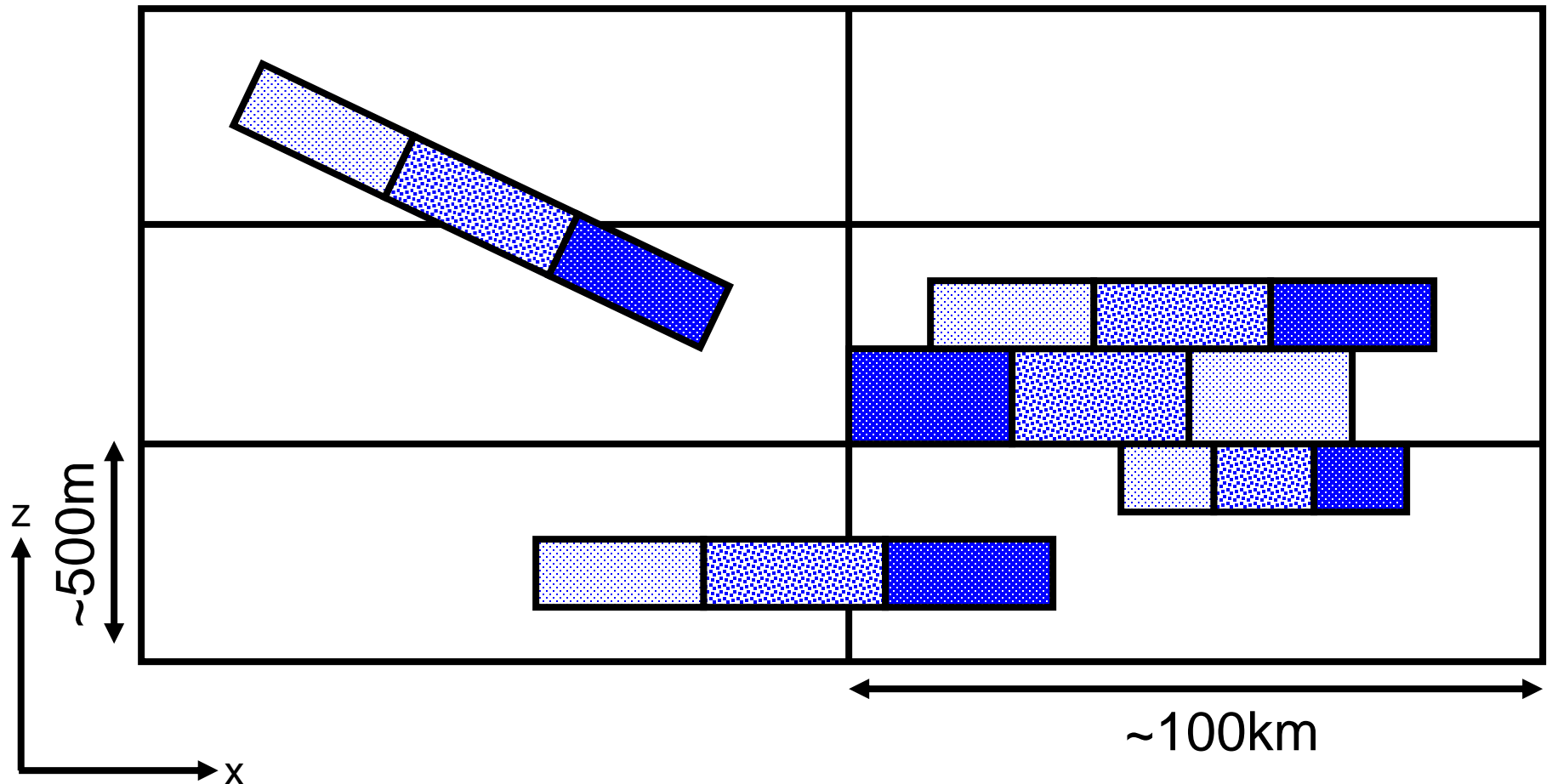
In-cloud inhomogeneity
in terms of cloud water, particle size/number



Macroscale Issues of Parameterization



Just these issues can become very complex!!!





First: Some assumptions!

q_v = water vapour mixing ratio

q_c = cloud water (liquid/ice) mixing ratio

q_s = saturation mixing ratio = $F(T,p)$

q_t = total water (vapour+cloud) mixing ratio

RH = relative humidity = q_v / q_s

1. Local criterion for formation of cloud: $q_t > q_s$

This assumes that no supersaturation can exist

2. Condensation process is fast (cf. GCM timestep)

$$q_v = q_s$$

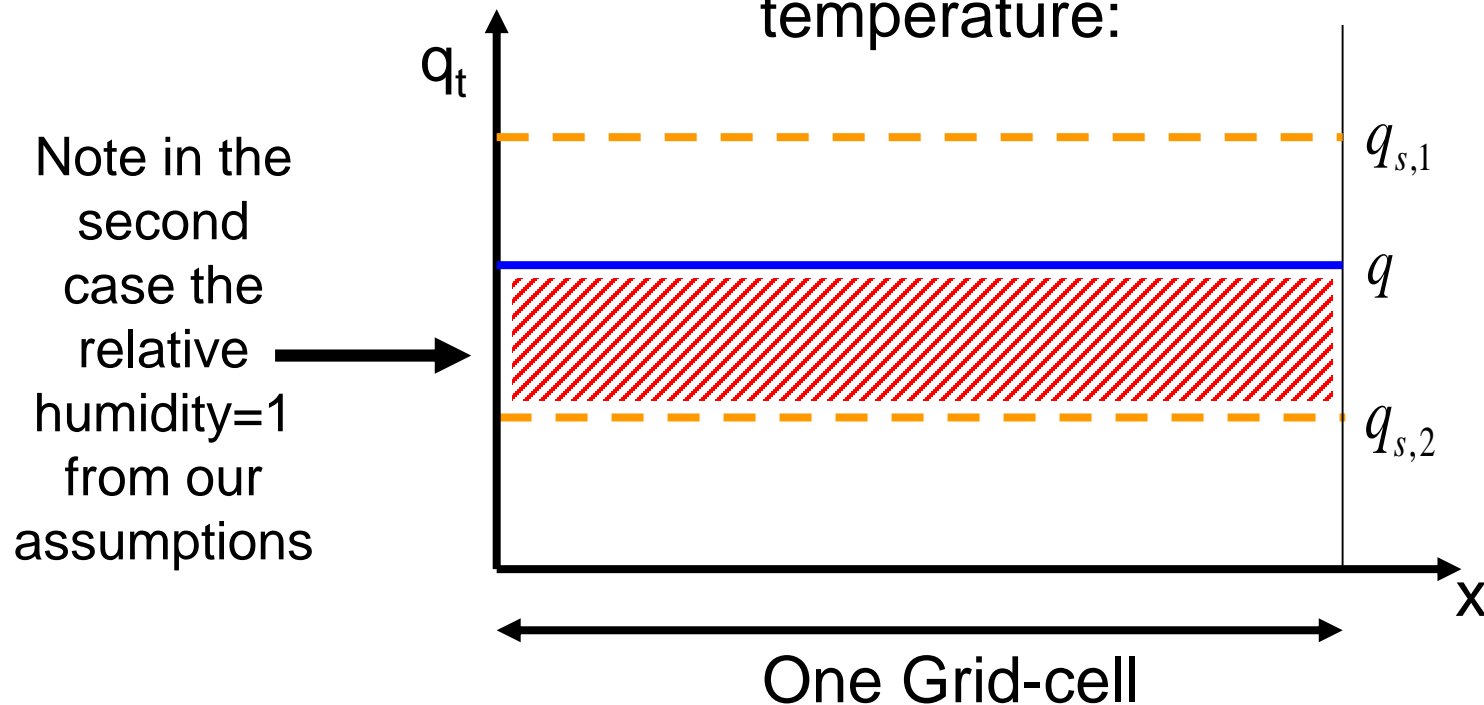
$$q_c = q_t - q_s$$

!!Both of these assumptions less applicable in ice clouds!!



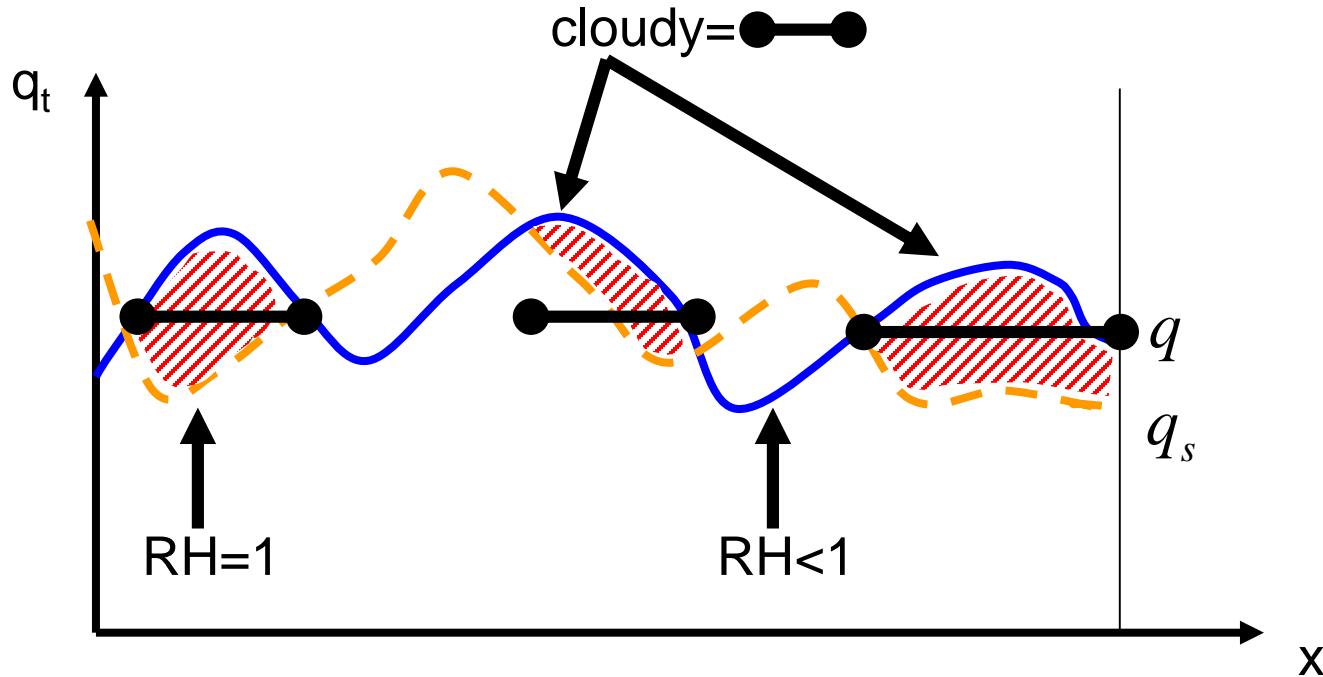
Partial cloud cover

Homogeneous distribution
of water vapour and
temperature:



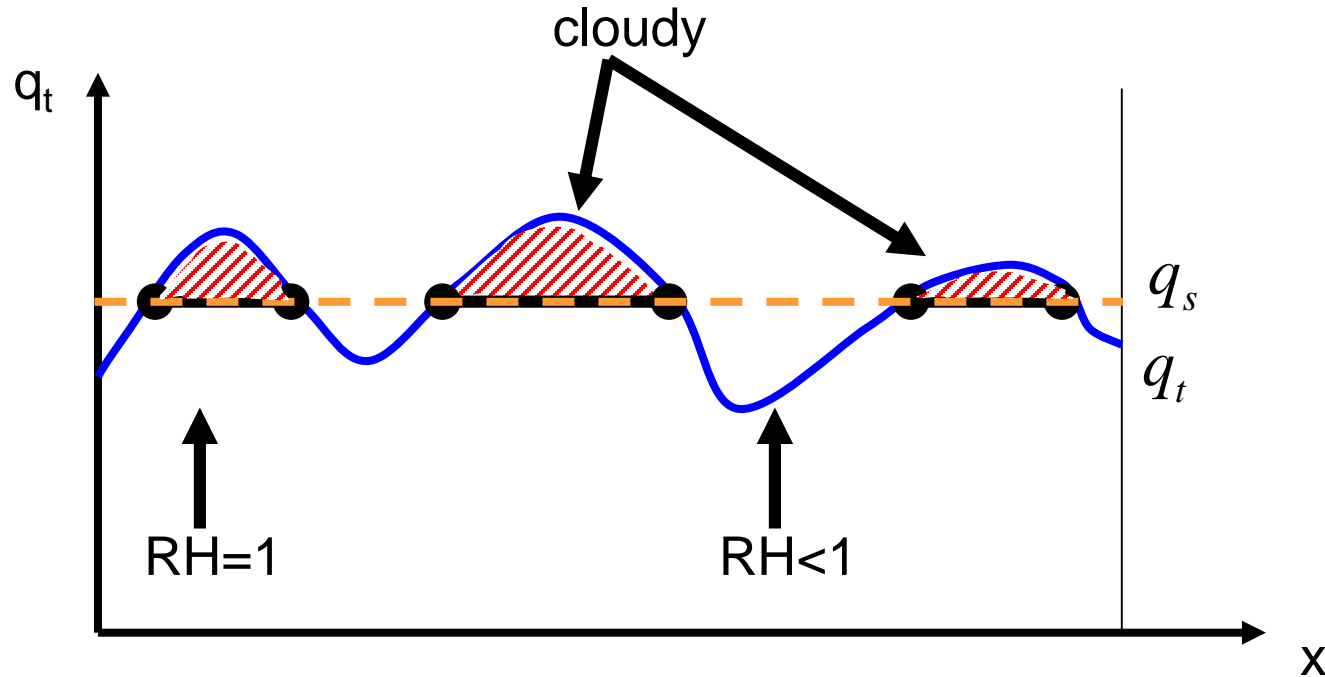
Partial coverage of a grid-box with clouds is only possible if there is an inhomogeneous distribution of temperature and/or humidity.

Heterogeneous Distribution of T and q



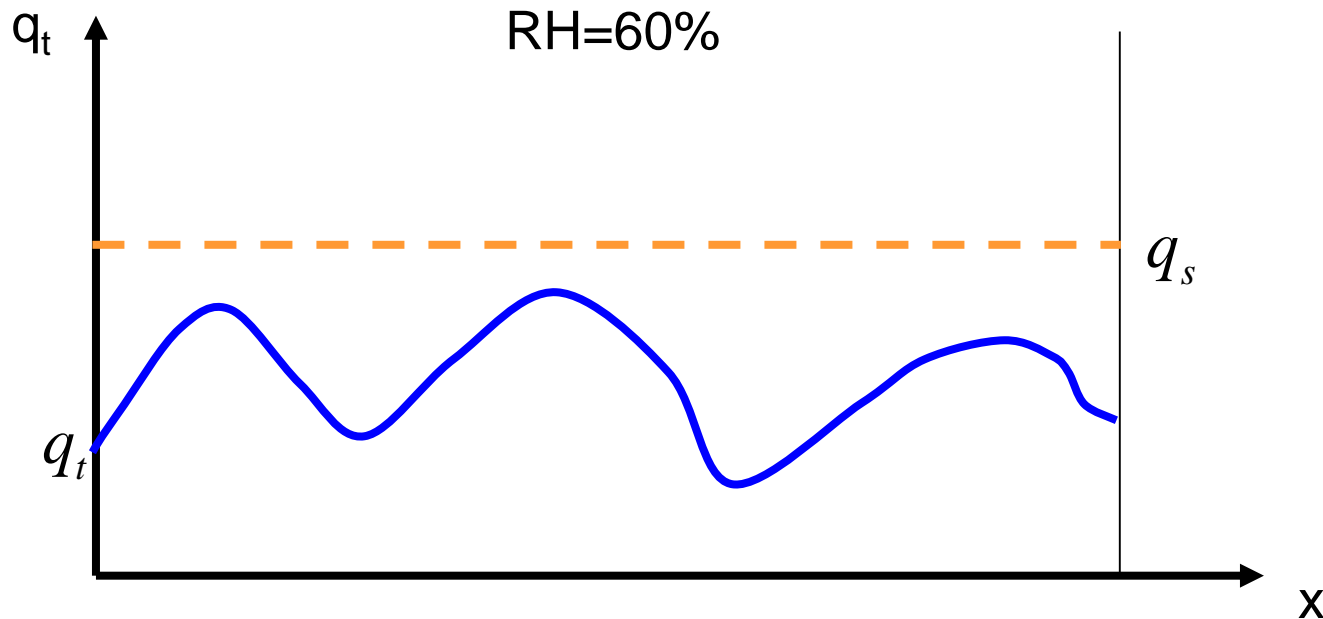
Another implication of the above is that clouds must exist before the grid-mean relative humidity reaches 1.

Heterogeneous Distribution of q only



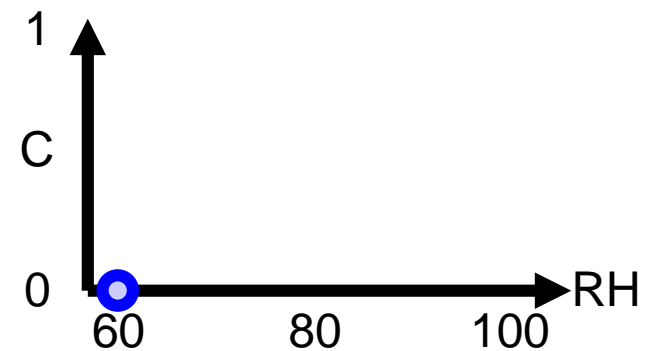
- The interpretation does not change much if we only consider humidity variability
- Throughout this talk I will neglect temperature variability
- Analysis of observations and model data indicates humidity fluctuations are more important *most* of the time.

Simple Diagnostic Cloud Schemes: Relative Humidity Schemes

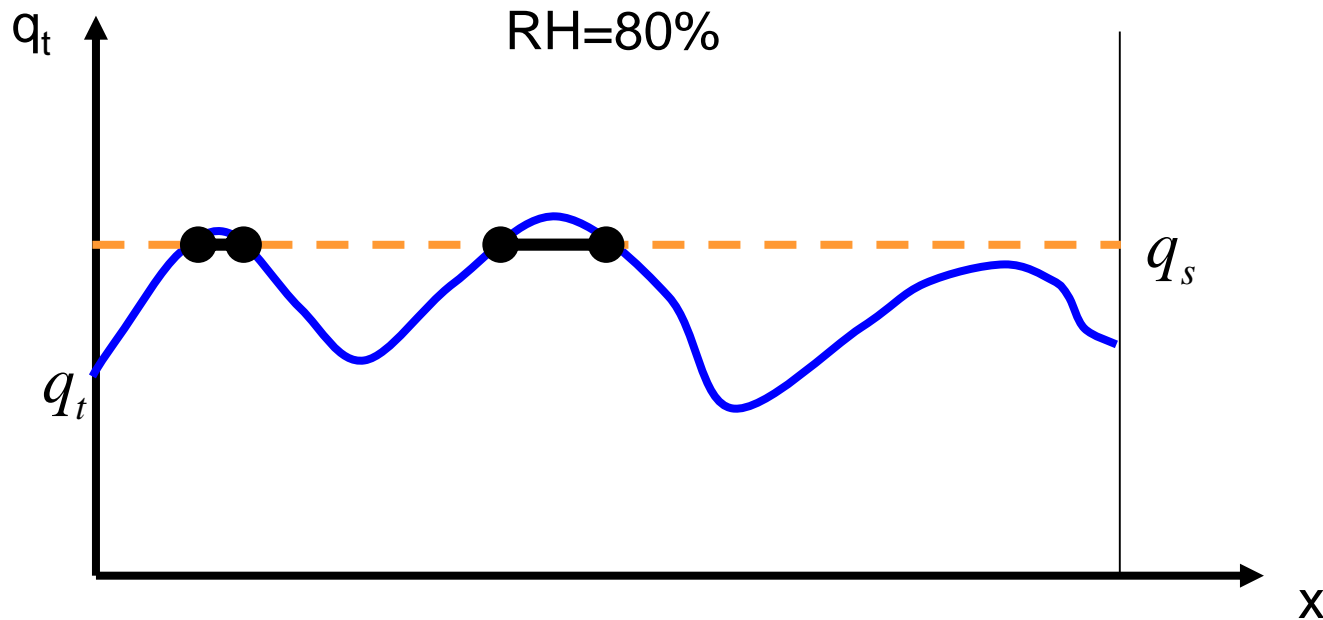


Take a grid cell with a certain (fixed) distribution of total water.

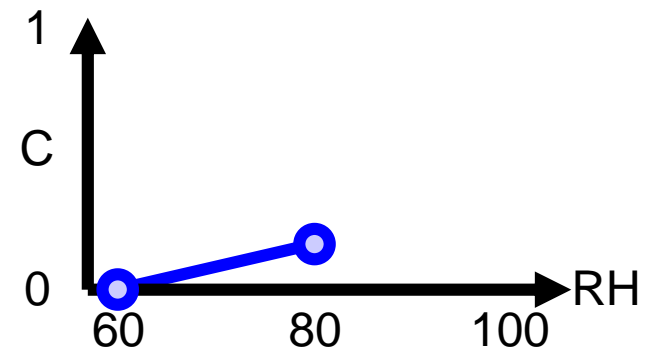
At low mean RH, the cloud cover is zero, since even the moistest part of the grid cell is subsaturated



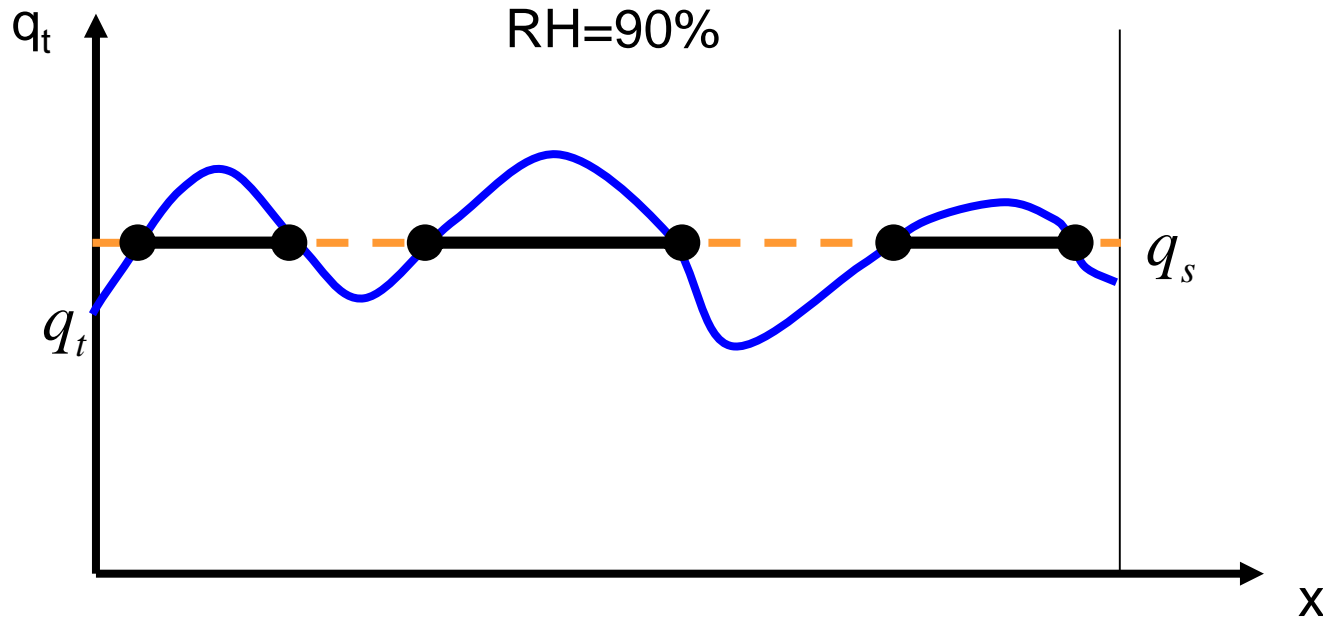
Simple Diagnostic Cloud Schemes: Relative Humidity Schemes



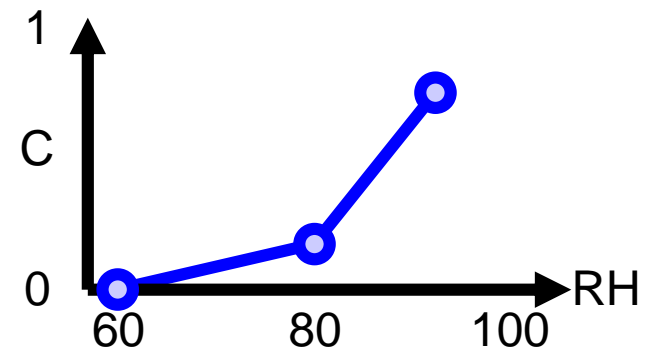
Add water vapour to the gridcell,
the moistest part of the cell
become saturated and cloud
forms. The cloud cover is low.



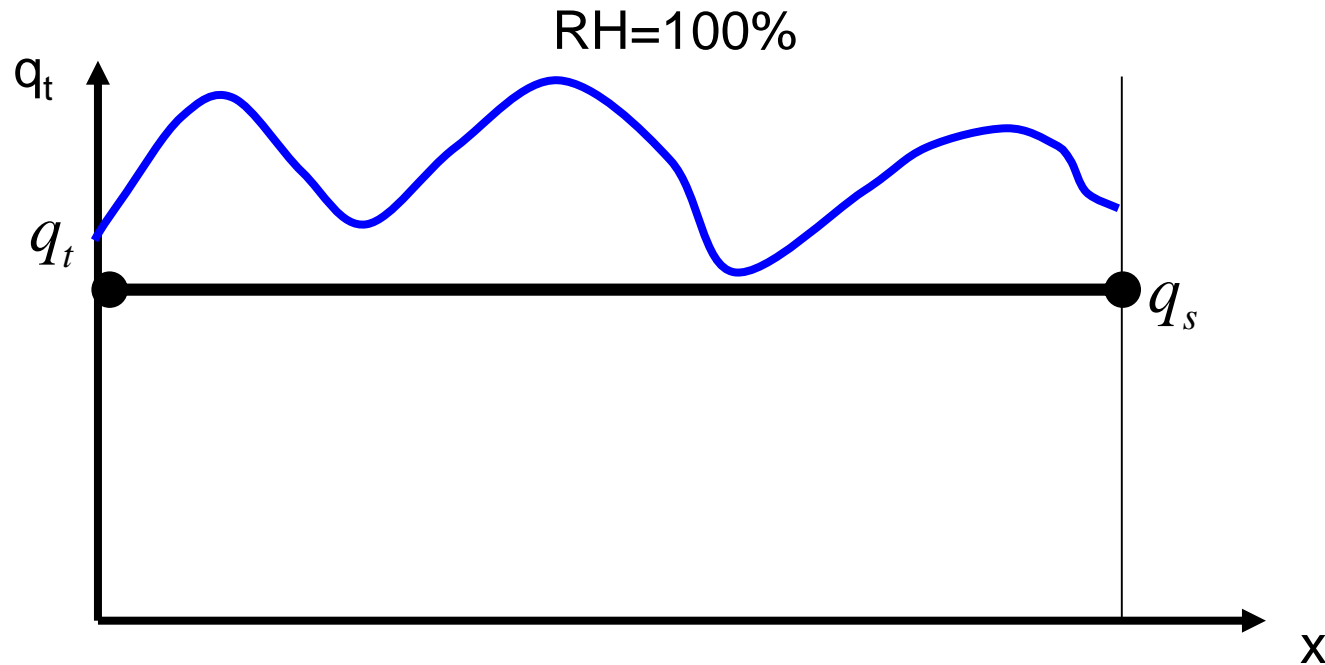
Simple Diagnostic Cloud Schemes: Relative Humidity Schemes



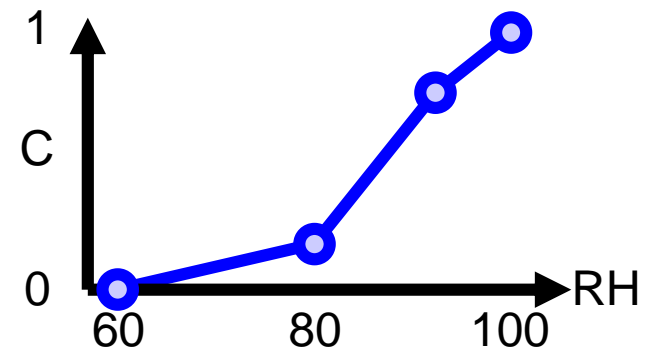
Further increases in RH
increase the cloud cover



Simple Diagnostic Cloud Schemes: Relative Humidity Schemes



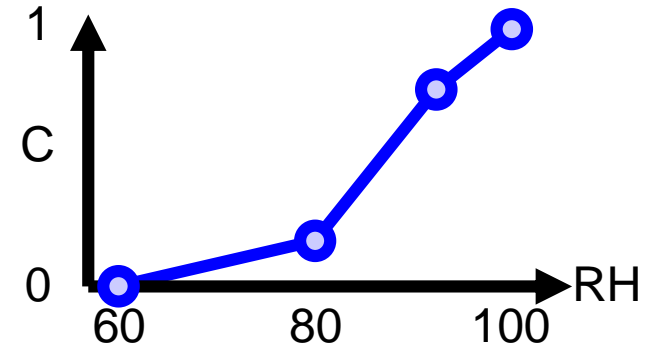
- The grid cell becomes overcast when $RH=100\%$, due to lack of supersaturation
- Diagnostic RH-based parametrization $C = f(RH)$



Diagnostic Relative Humidity Schemes



- Many schemes, from the 1970s onwards, based cloud cover on the relative humidity (RH)
- e.g. Sundqvist et al. MWR 1989:



$$C = 1 - \sqrt{\frac{1 - RH}{1 - RH_{crit}}}$$

 Remember this for later!

RH_{crit} = critical relative humidity at which cloud assumed to form
(= function of height, typical value is 60-80%)

Diagnostic Relative Humidity Schemes



- Since these schemes form cloud when $RH < 100\%$, they implicitly assume subgrid-scale variability for total water, q_t , (and/or temperature, T).
- However, the actual PDF (the shape) for these quantities and their variance (width) are often not known.
- They are of the form: “*Given a RH of X% in nature, the mean distribution of q_t is such that, on average, we expect a cloud cover of Y%.*”

Diagnostic Relative Humidity Schemes



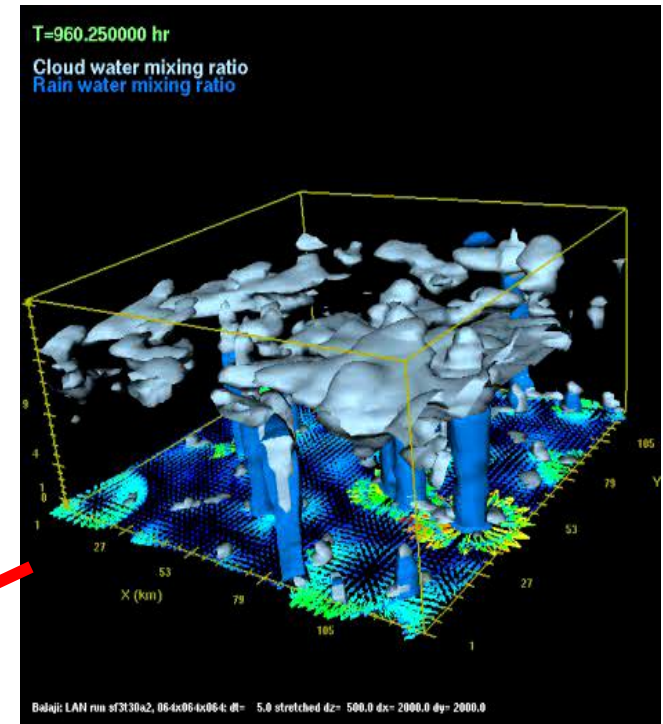
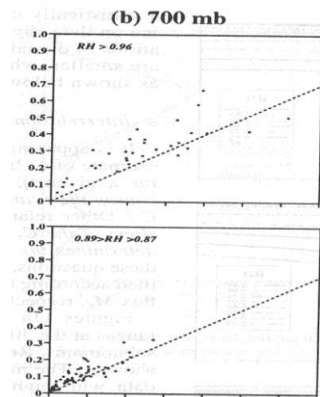
- Advantages:
 - Better than homogeneous assumption, since clouds can form before grids reach saturation.
- Disadvantages:
 - Cloud cover not well coupled to other processes.
 - In reality, different cloud types with different coverage can exist with same relative humidity. This can not be represented.
- Can we do better?

Diagnostic Relative Humidity Schemes



- Could add further predictors
- E.g: Xu and Randall (1996) sampled cloud scenes from a 2D cloud resolving model to derive an empirical relationship with two predictors:

$$C = F(RH, q_c)$$



- More predictors, more degrees of freedom = flexible
- But still do not know the form of the PDF (is model valid? representative for all situations?)
- Can we do better?

Diagnostic Relative Humidity Schemes



- Another example is the scheme of Slingo, operational at ECMWF until 1995.
- This scheme also adds dependence on vertical velocities
- Use different empirical relations for different cloud types, e.g., middle level clouds:

$$C_m = \begin{cases} 0 & \omega \geq 0 \\ C_m^* \omega / \omega_{crit} & \omega_{crit} \leq \omega < 0 \\ C_m^* & \omega < \omega_{crit} \end{cases} \quad C_m^* = \left[\max \left(\frac{RH - RH_{crit}}{1 - RH_{crit}}, 0 \right) \right]^2$$

Relationships seem Ad-hoc? Can we do better?

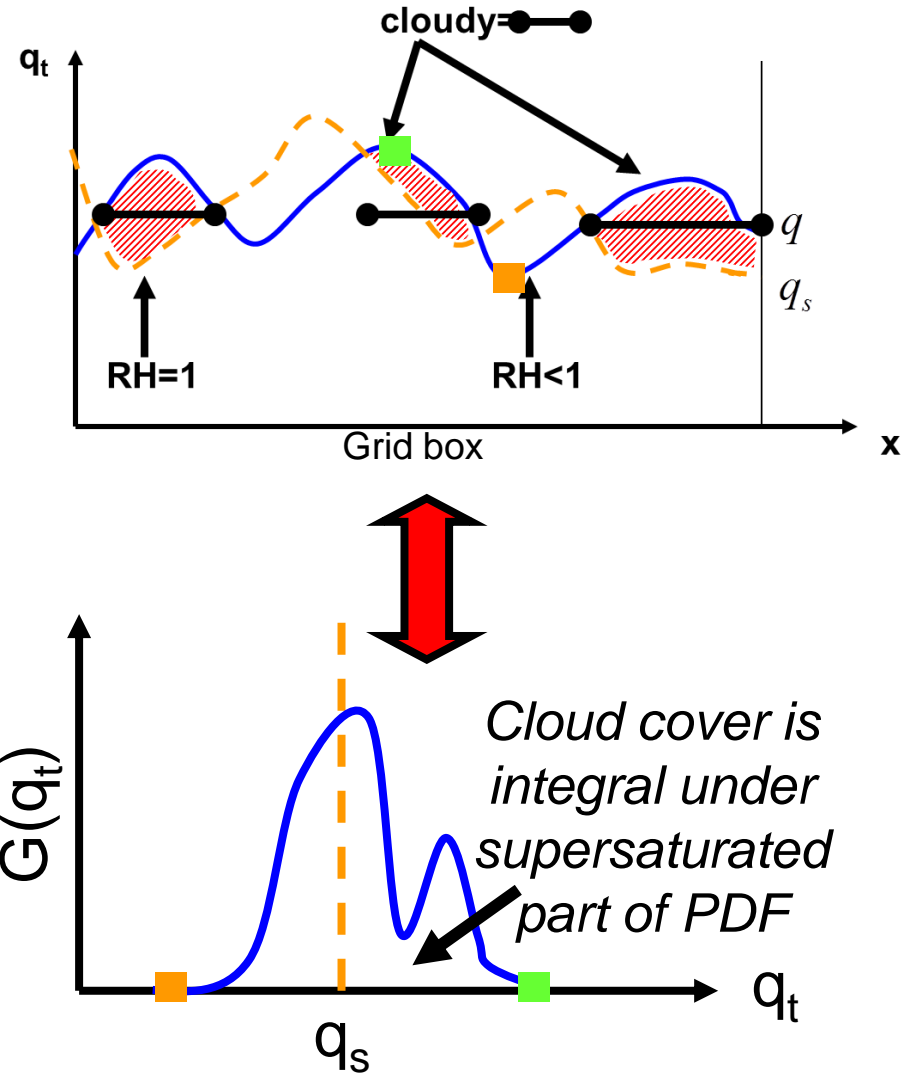


Statistical PDF Schemes

- Statistical schemes explicitly specify the probability density function (PDF), G , for the total water q_t (and sometimes also temperature)

$$C = \int_{q_s}^{\infty} G(q_t) dq_t$$

$$q_c = \int_{q_s}^{\infty} (q_t - q_s) G(q_t) dq_t$$



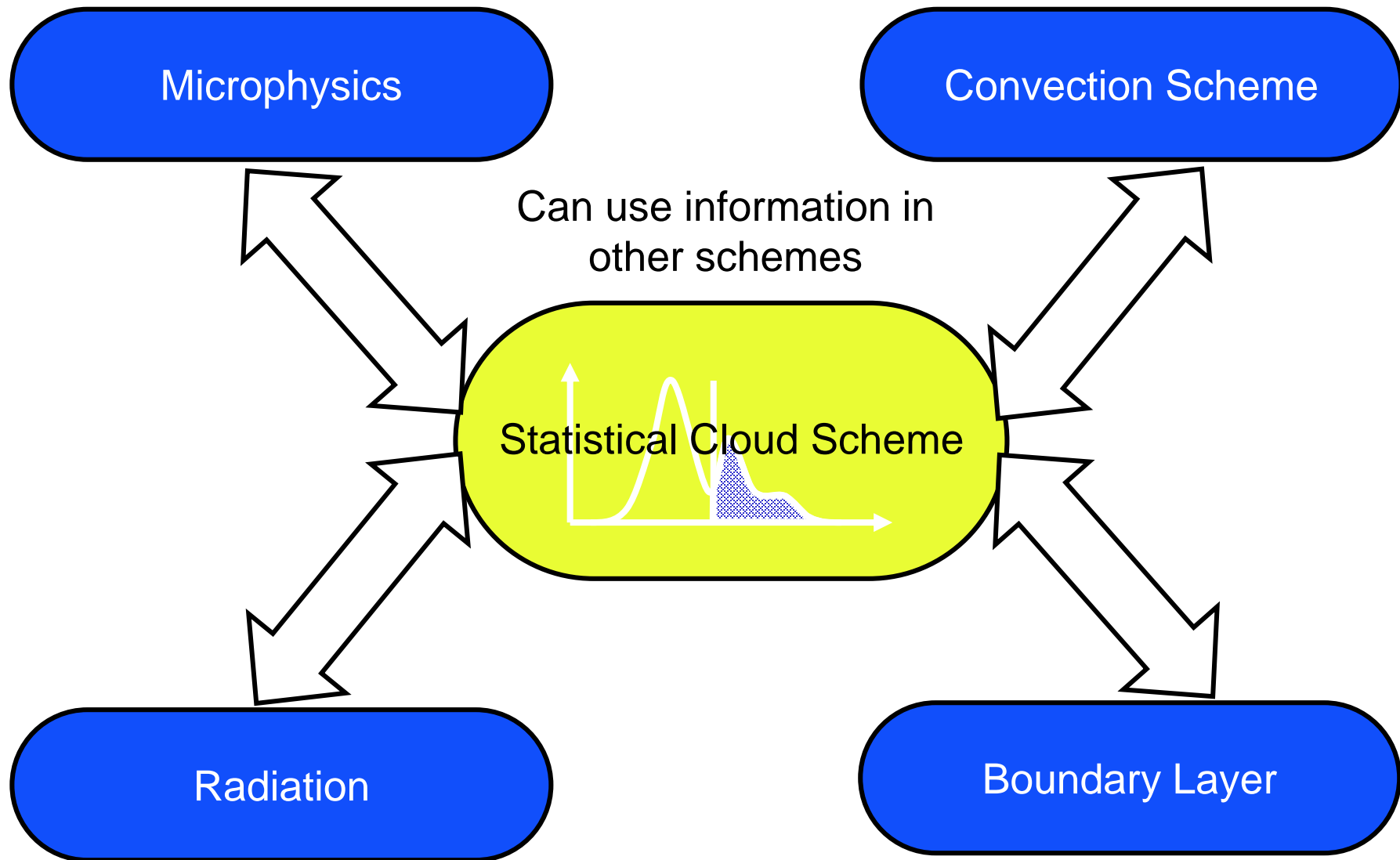
Statistical PDF Schemes



- Knowing the PDF has advantages:
 - Information concerning subgrid fluctuations of humidity and cloud condensate is available (for all parametrizations) , e.g.
 - More accurate calculation of radiative fluxes
 - Unbiased calculation of microphysical processes
 - Use of underlying PDF means cloud variables (condensate, cloud fraction) are always self-consistent.
 - Physically-based. Can evaluate with observations.

(Note, location of clouds within grid cell is still not known)

Statistical PDF scheme: Consistency across parametrizations



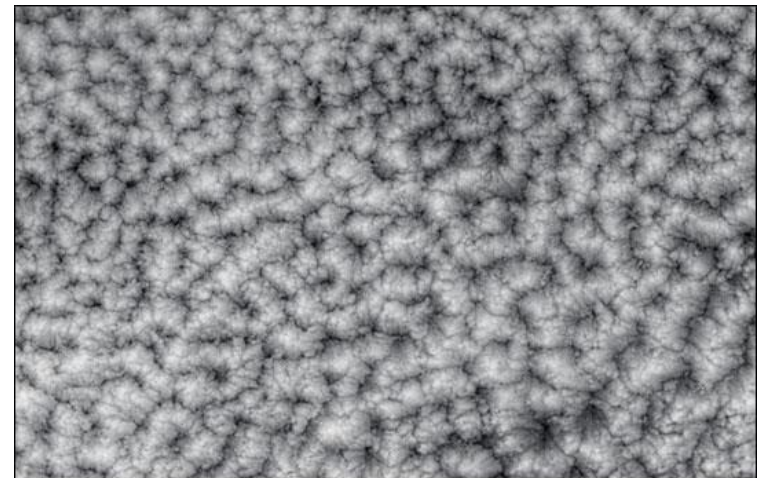
Building a statistical cloud scheme

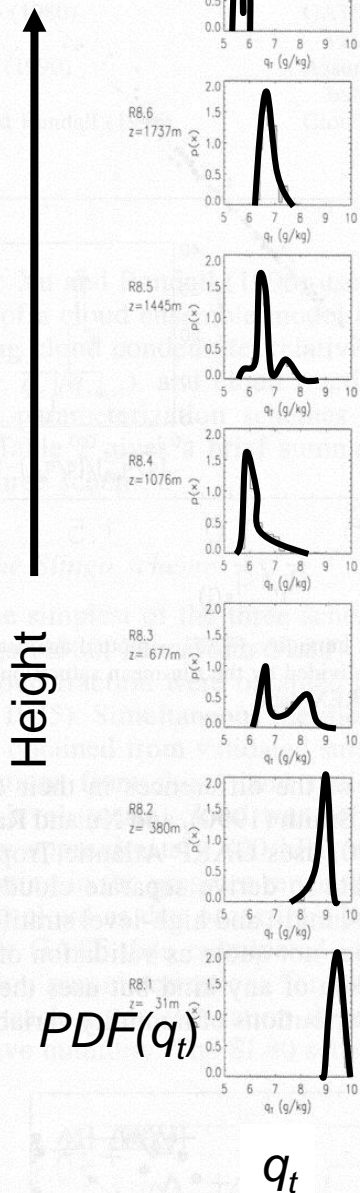


What do we observe?

- Limited observations to determine q_t PDF
 - Aircraft data
 - limited coverage
 - Tethered balloon
 - boundary layer only
 - Satellite
 - difficulties resolving in vertical
 - no q_t observations
 - poor horizontal resolution
 - Ground-based radar/Raman Lidar
 - one location
- Cloud Resolving models have also been used
 - realism of microphysical parametrization?

Modis image from NASA website





Wood and Field
JAS 2000
Aircraft
observations low
clouds < 2km

**Aircraft
Observed
PDFs**

Heymsfield and
McFarquhar
JAS 96
Aircraft IWC obs
during CEPEX

FIG. 2. Distributions of total water from penetrative cumulus during A

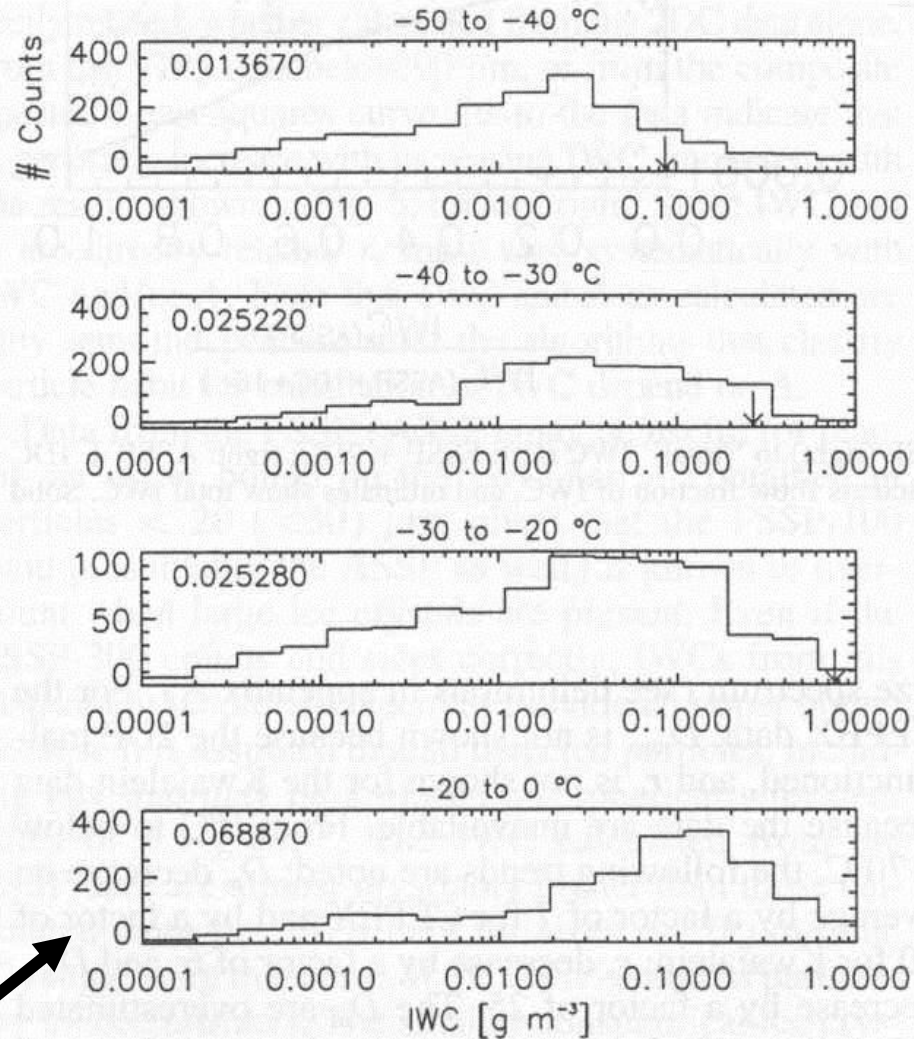


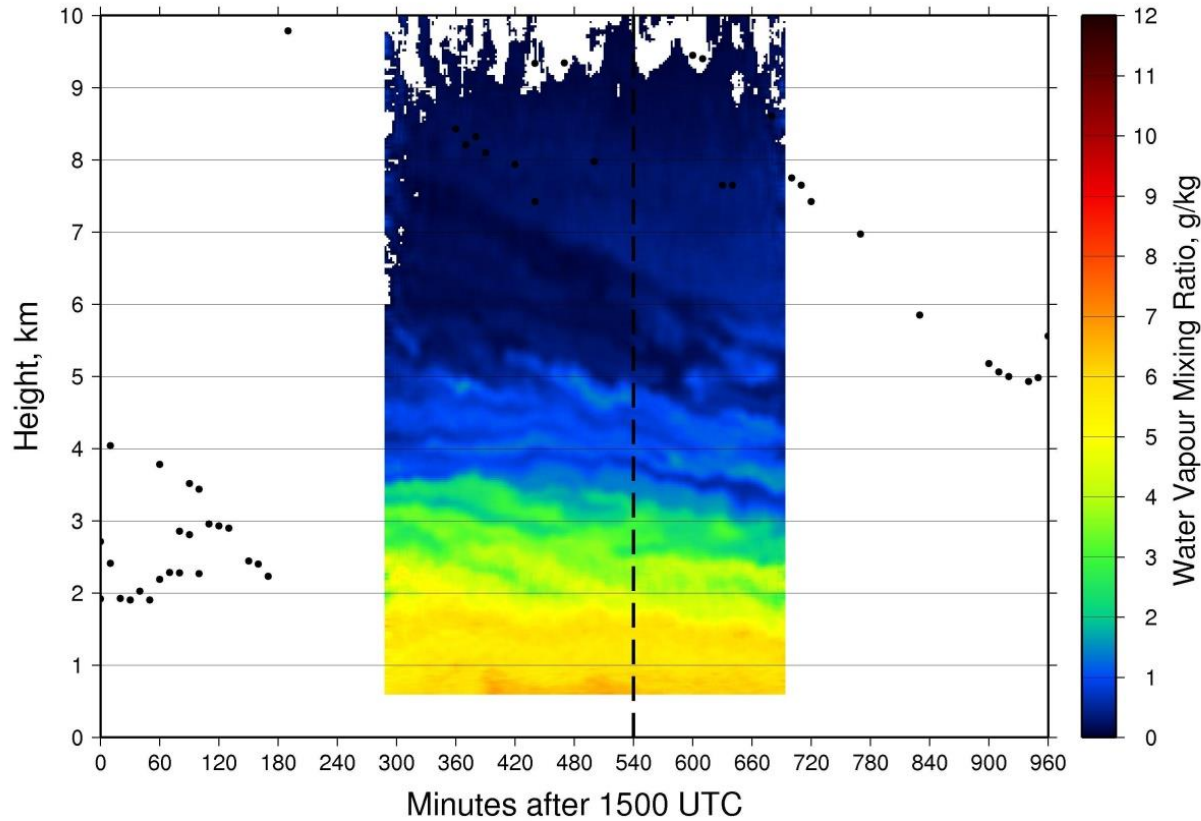
FIG. 9. Frequency distributions of IWC sorted by temperature from 2DC data during CEPEX. Each count (ordinate) represents 10-s average. Median IWC (g m^{-3}) in upper left corner of each panel. Arrows give saturation vapor density with respect to ice for midpoint of each temperature interval.

Building a statistical cloud scheme

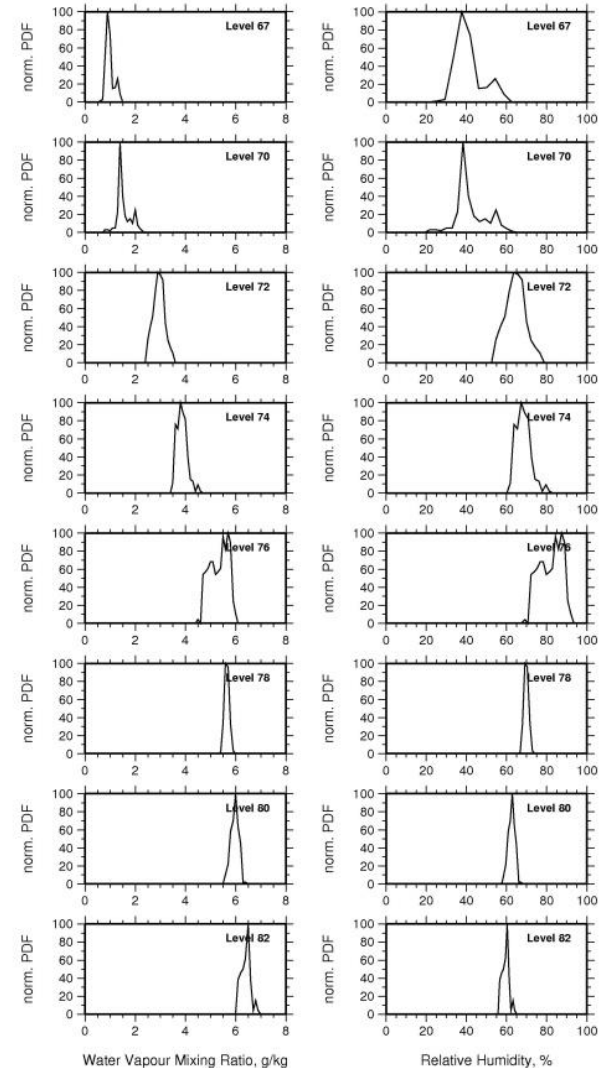


Observed PDF of water vapour/RH Raman Lidar

Raman-Lidar RAMSES – 15./16. May 2008



16. May 2008 00:00 UTC



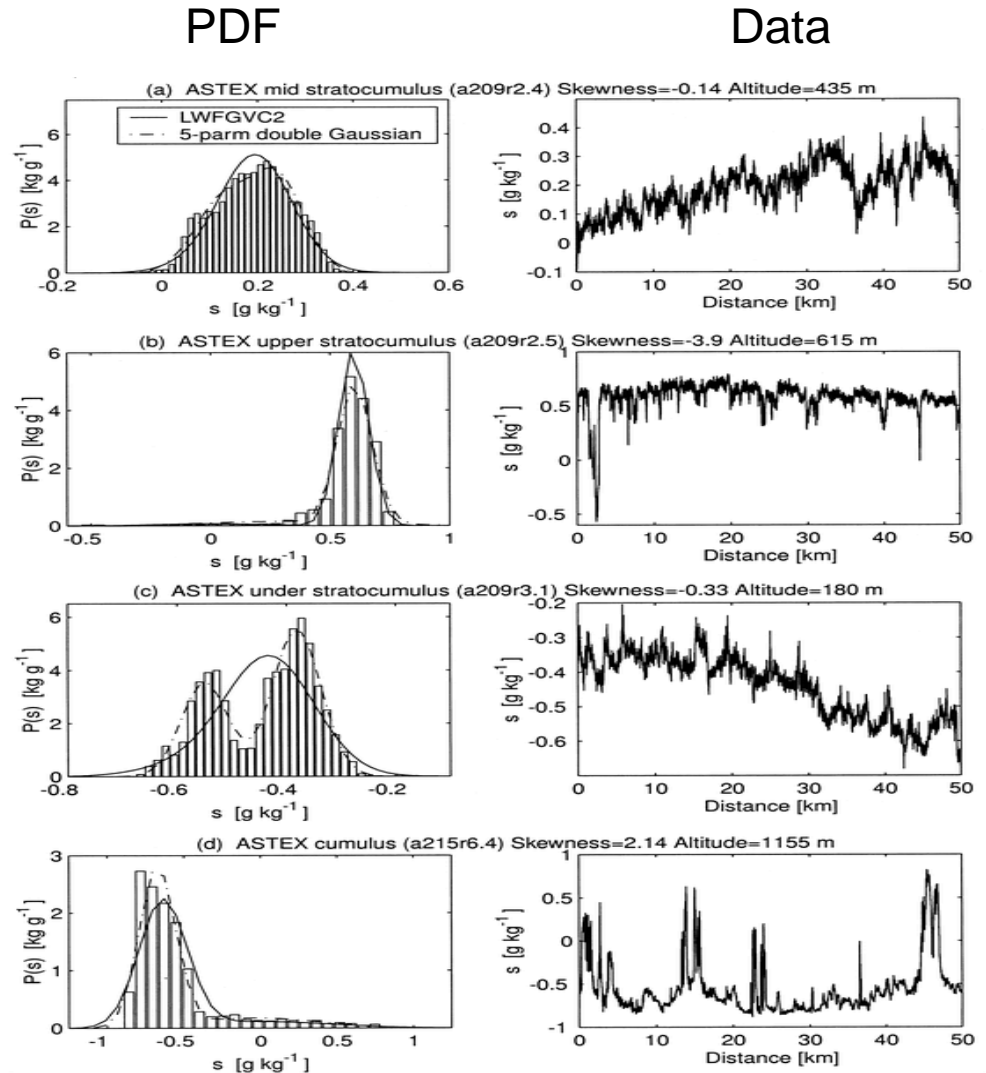
Building a statistical cloud scheme



Observed PDF example from aircraft

Example, aircraft data from Larson et al. 01/02

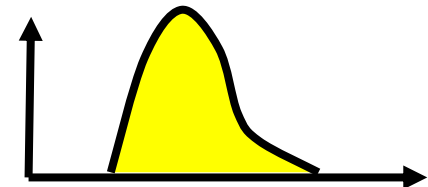
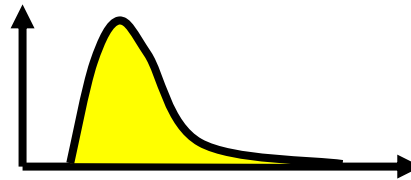
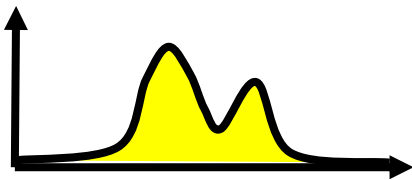
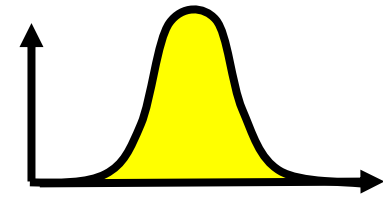
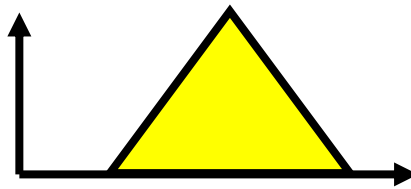
PDFs are mostly approximated by uni or bi-modal distributions, describable by a few parameters



Building a statistical cloud scheme



- Need to represent with a functional form, specify the:
 - (1) **PDF shape** (unimodal, bimodal, symmetrical, bounded?)
 - (2) **PDF moments** (mean, variance, skewness?)
 - (3) **Diagnostic or prognostic** (how many degrees of freedom?)



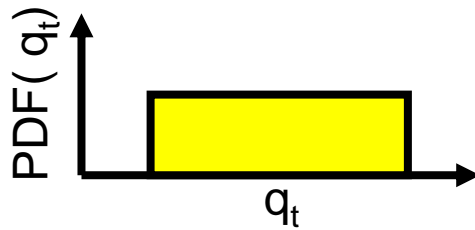
Building a statistical cloud scheme



(1) Specification of PDF shape

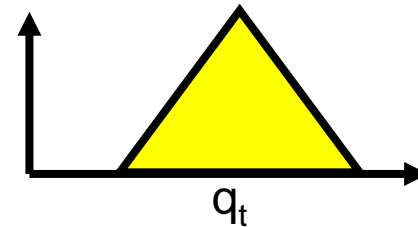
Many function forms have been used

symmetrical distributions:



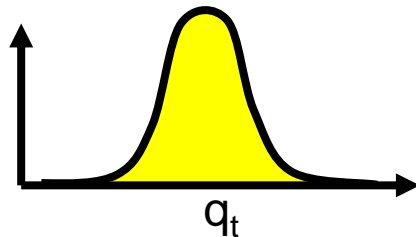
Uniform:

Letreut and Li (91)



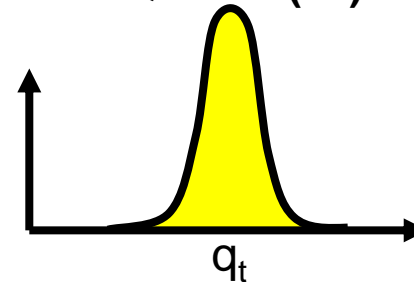
Triangular:

Smith QJRMS (90)



Gaussian:

Mellor JAS (77)



s^4 polynomial:

Lohmann et al. J. Clim (99)

Bounded

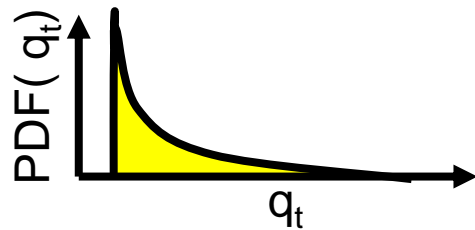
Unbounded:
Can clip, but need
additional
parameters

Building a statistical cloud scheme



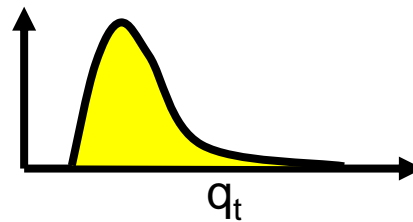
(1) Specification of PDF shape

skewed distributions:



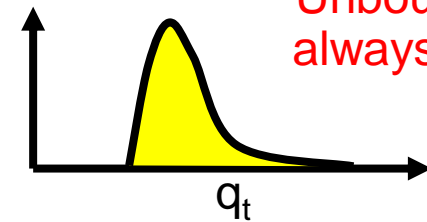
Exponential:

Sommeria and Deardorff
JAS (77)



Lognormal:

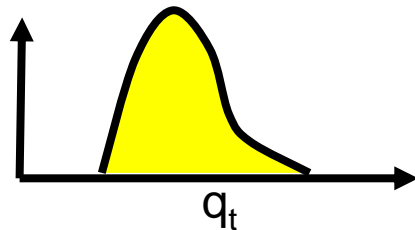
Bony & Emanuel
JAS (01)



Gamma:

Barker et al. JAS (96)

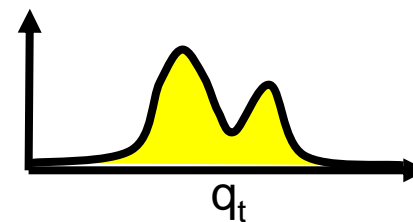
Unbounded,
always skewed



Beta:

Tompkins JAS (02)

Bounded, symmetrical or skewed



Double Normal/Gaussian:

Lewellen and Yoh JAS (93), Golaz et al.
JAS 2002

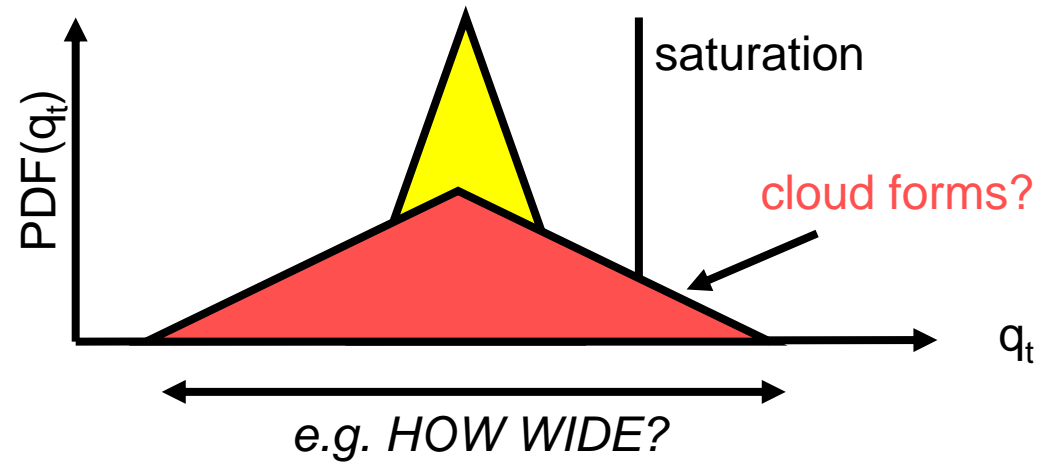
Building a statistical cloud scheme



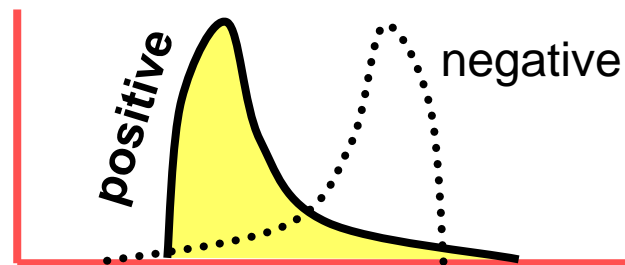
(2) Specification of PDF moments

Need also to determine the moments of the distribution:

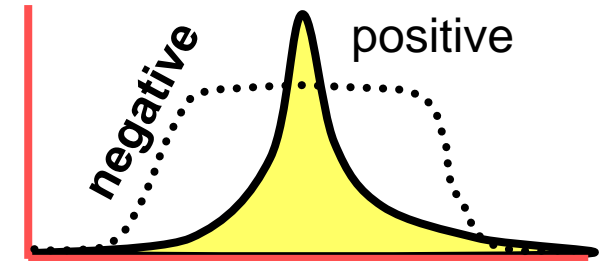
- Variance (Symmetrical PDFs)
- Skewness (Higher order PDFs)
- Kurtosis (4-parameter PDFs)



Skewness



Kurtosis



Moment 1=MEAN
Moment 2=VARIANCE
Moment 3=SKEWNESS
Moment 4=KURTOSIS

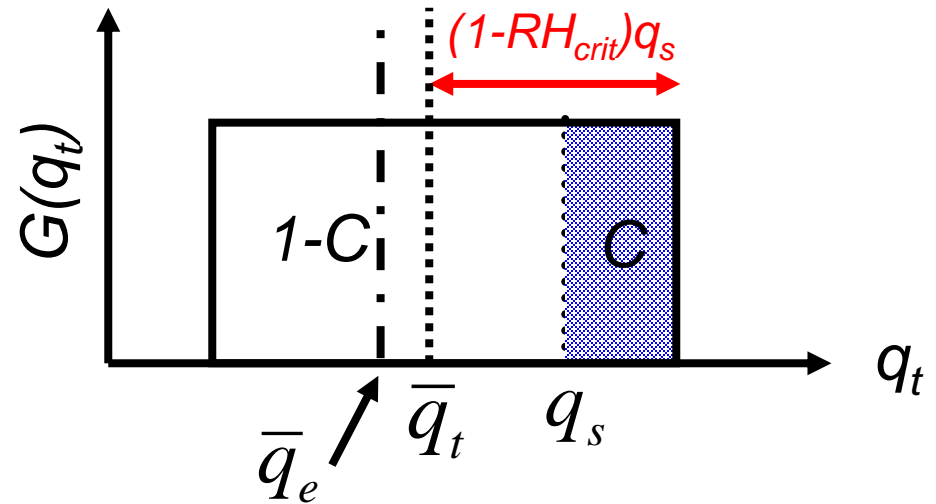
Functional form – needs to fit data but be sufficiently simple

Building a statistical cloud scheme



(3) Diagnostic or prognostic PDF moments

- Some schemes fix the moments (diagnostic e.g. Smith 1990) based on critical RH at which clouds assumed to form.
- Some schemes predict the moments (prognostic, e.g. Tompkins 2002). Need to specify sources and sinks.
- If moments (variance, skewness) are fixed, then statistical schemes are identically equivalent to a RH formulation
- e.g. uniform q_t distribution = Sundqvist formulation



$$\bar{q}_e = q_s (1 - (1 - RH_{crit})(1 - C))$$

$$\bar{q}_v = Cq_s + (1 - C)\bar{q}_e$$

$$RH = \frac{\bar{q}_v}{q_s} = 1 - (1 - RH_{crit})(1 - C)^2$$

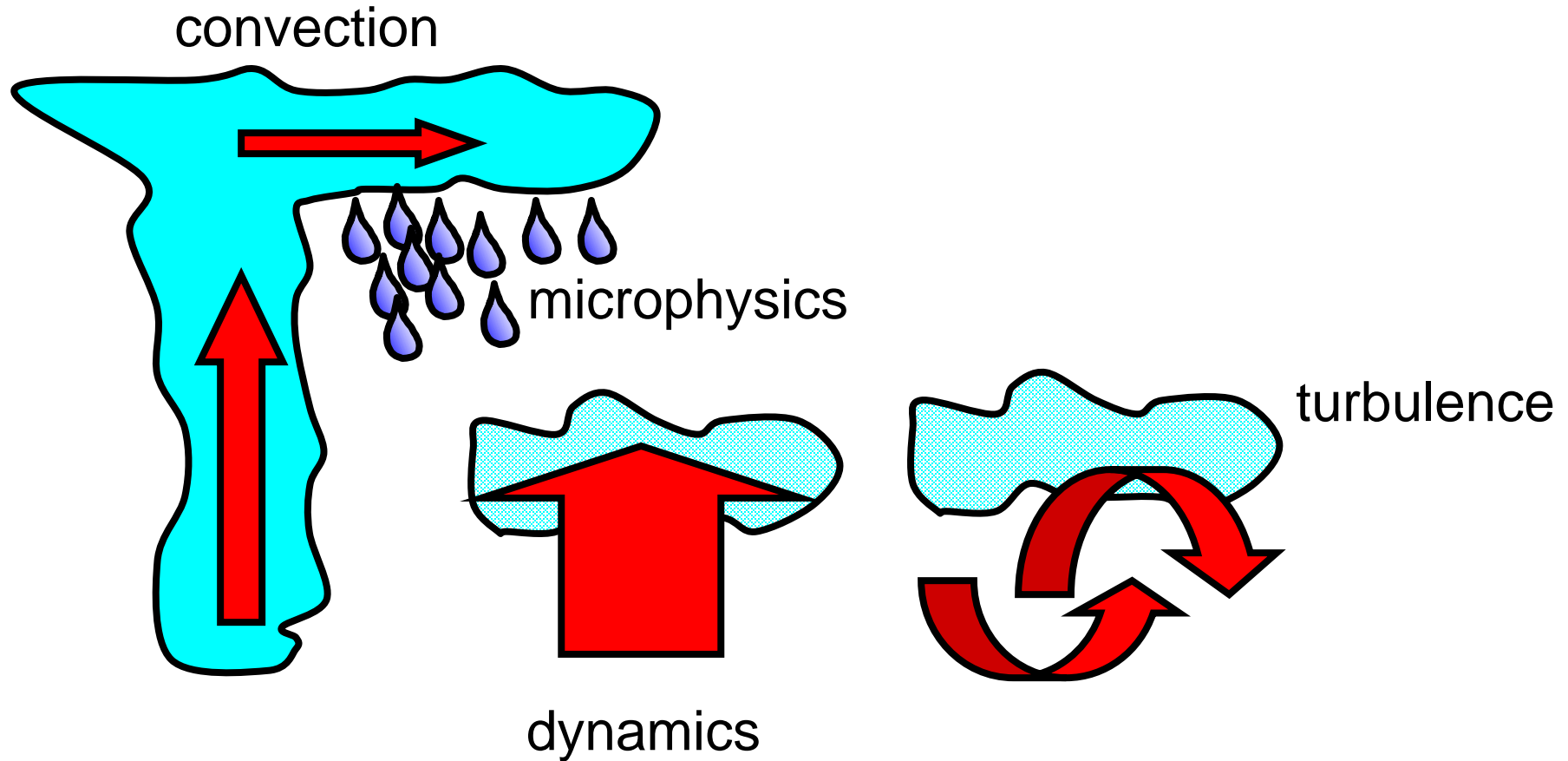
Sundqvist formulation!!!

$$\therefore C = 1 - \sqrt{\frac{1 - RH}{1 - RH_{crit}}}$$

Building a statistical cloud scheme

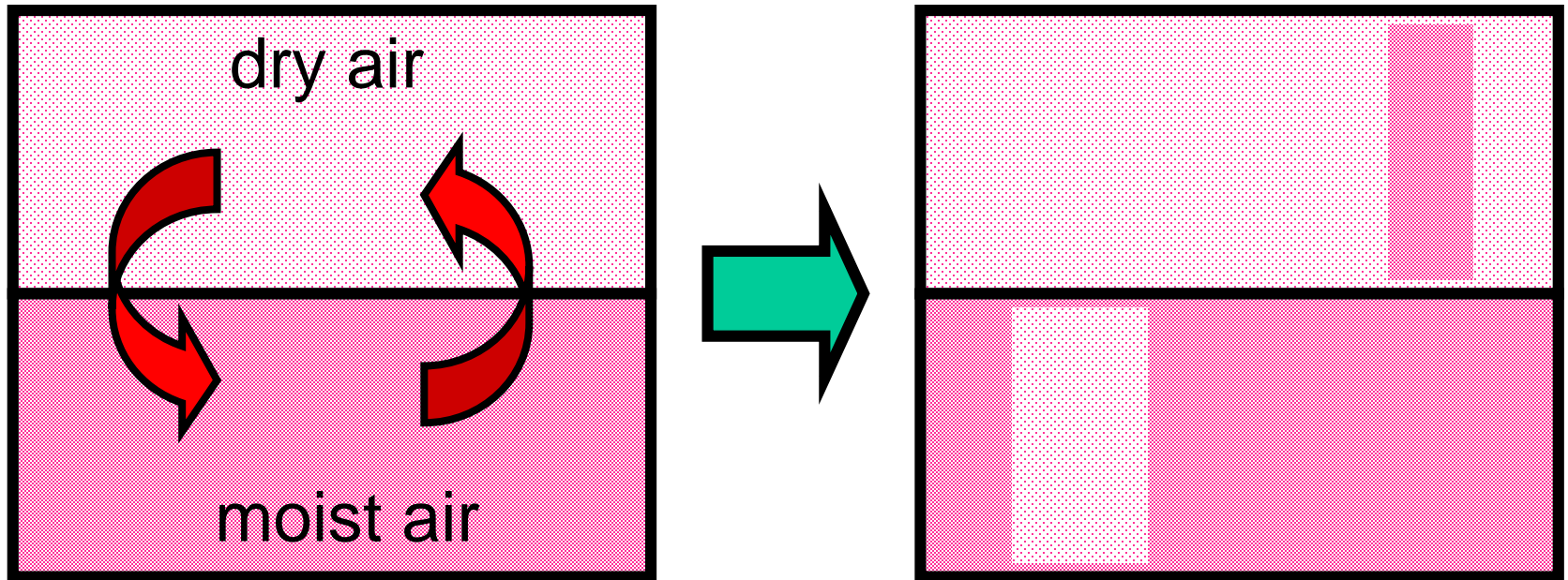


Processes that can affect PDF moments



Example: Turbulence

In presence of vertical gradient of total water, turbulent mixing can increase horizontal variability



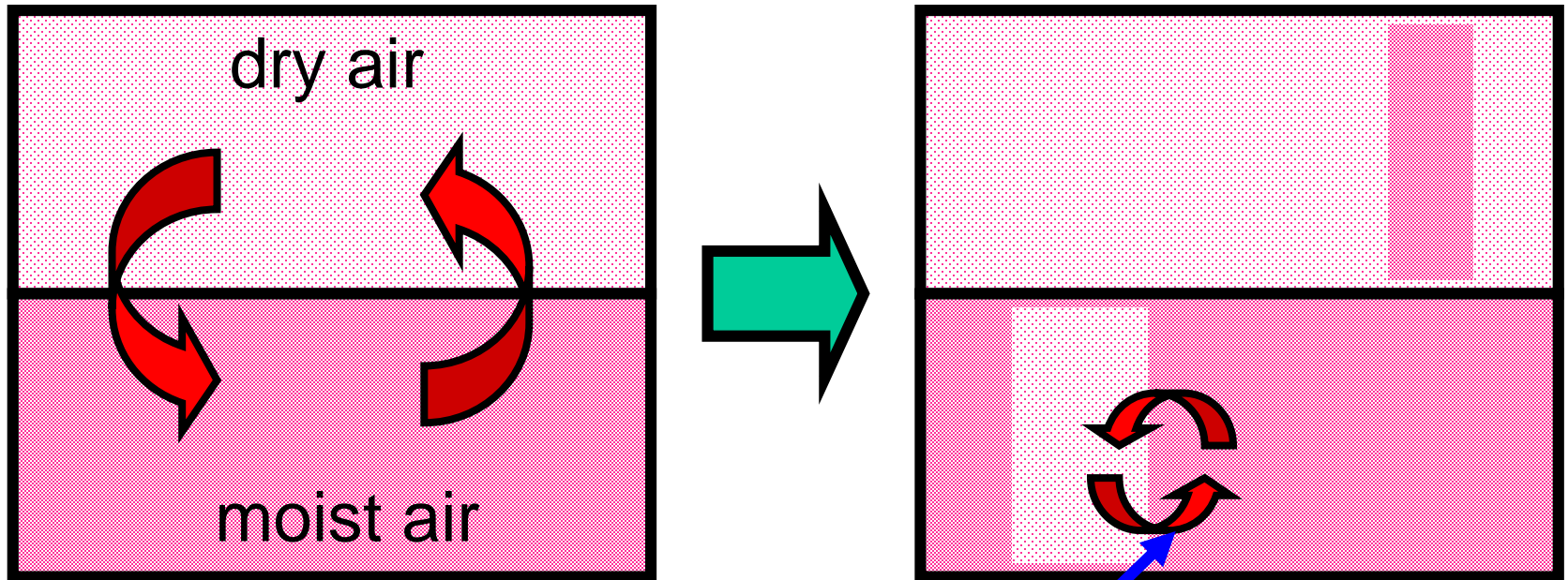
Rate of change
of total water
variance

$$\frac{d \overline{q_t'^2}}{dt} = -2 \overline{w' q_t'} \frac{d \overline{q_t}}{dz}$$



Example: Turbulence

In presence of **vertical gradient** of total water, turbulent mixing can **increase horizontal variability**



while **subgrid mixing** in the **horizontal plane** naturally **reduces** the **horizontal variability**

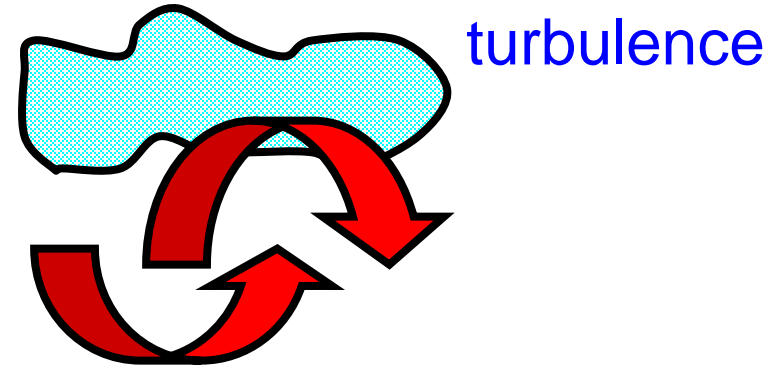
$$\frac{d \overline{q_t'^2}}{dt} = - \frac{q_t'^2}{\tau}$$

Building a statistical cloud scheme



Predicting change of q_t variance due to turbulence

If a process is fast compared to a GCM timestep, an equilibrium can be assumed, e.g. turbulence



$$\frac{d \overline{q_t'^2}}{dt} = \underbrace{-2 \overline{w' q_t'}}_{\text{Source}} \frac{d \overline{q_t}}{dz} - \underbrace{\frac{q_t'^2}{\tau}}_{\text{Dissipation}} \xrightarrow{\text{Local equilibrium}} q_t'^2 = -\tau 2 \overline{w' q_t'} \frac{d \overline{q_t}}{dz}$$

Example: Ricard and Royer, Ann Geophys, (93), Lohmann et al. J. Clim (99)

- **Disadvantage:**

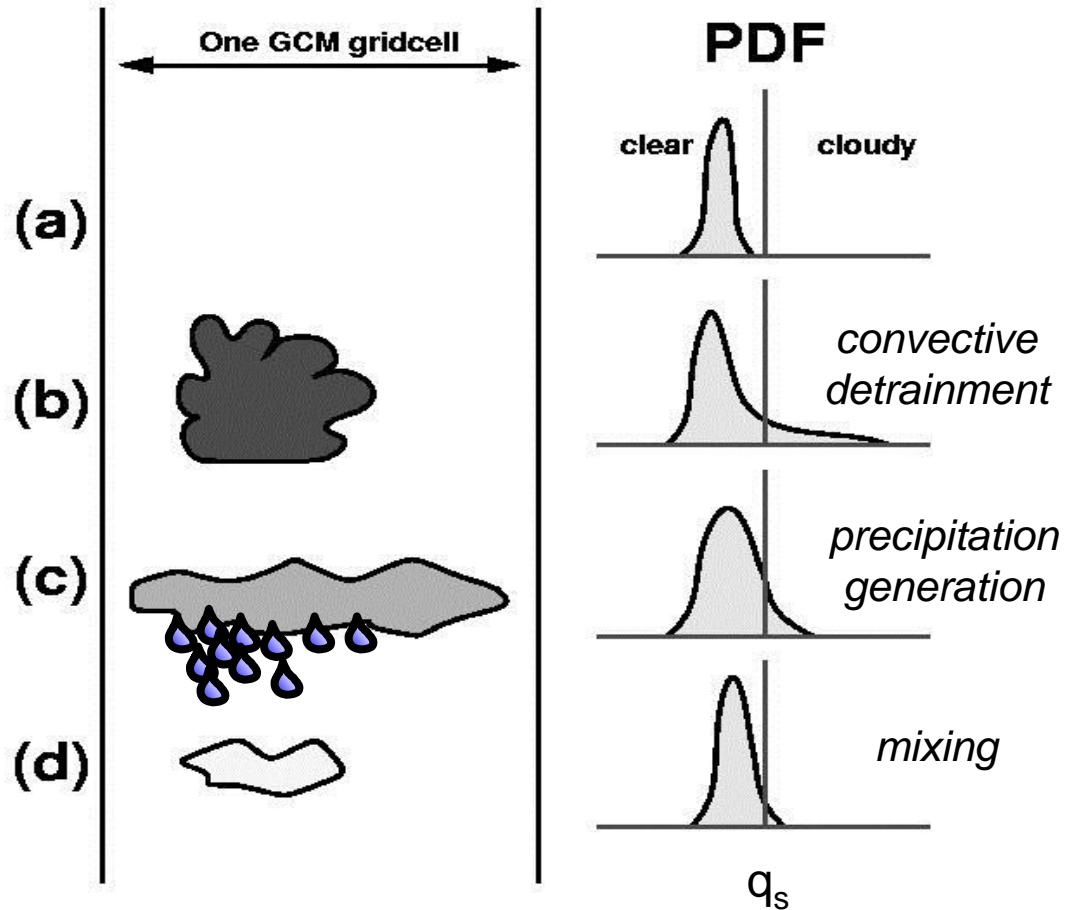
- Can give good estimate in boundary layer, but above, other processes will determine variability, that evolve on slower timescales

Building a statistical cloud scheme



Example: Tompkins (2002) prognostic PDF

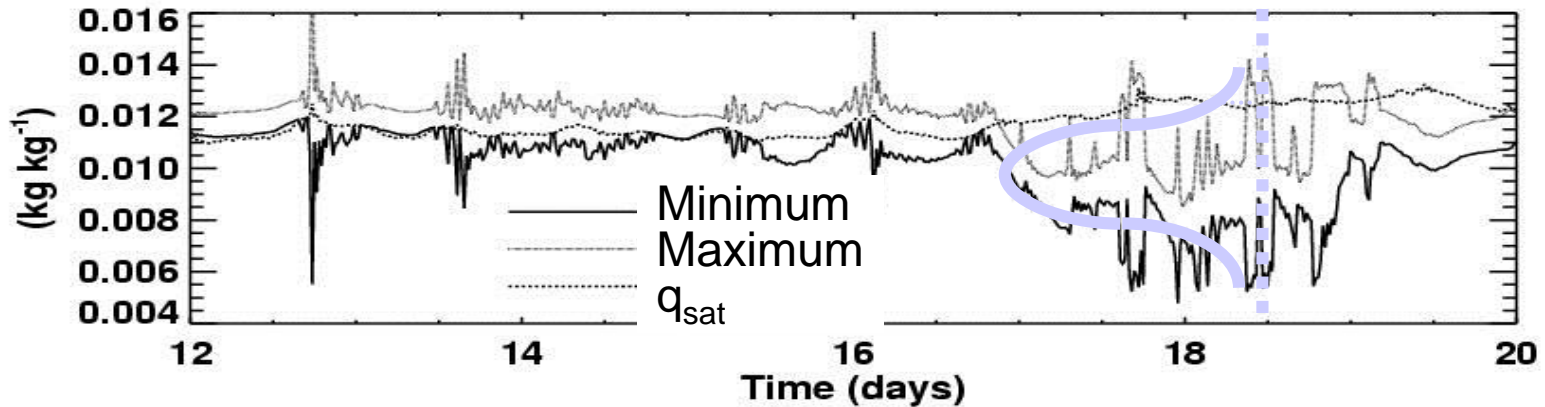
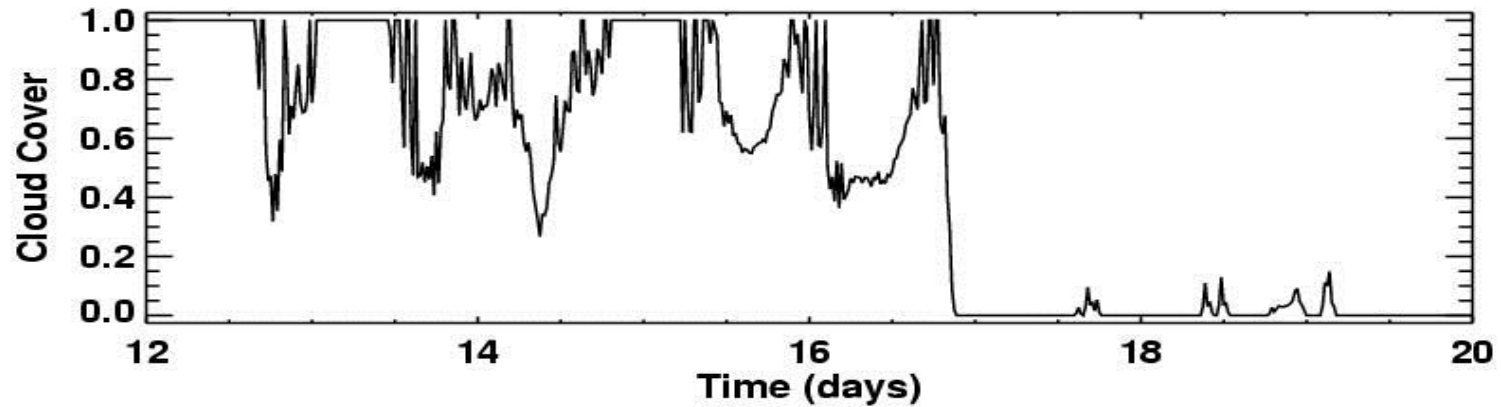
- Tompkins (2002) prognostic statistical scheme (implemented in ECHAM5 climate GCM).
- Prognostic equations are introduced for variables representing the **mean**, **variance** and **skewness** of the total water PDF.
- Some of the sources and sinks are rather ad-hoc in their derivation!



Prognostic statistical scheme in action

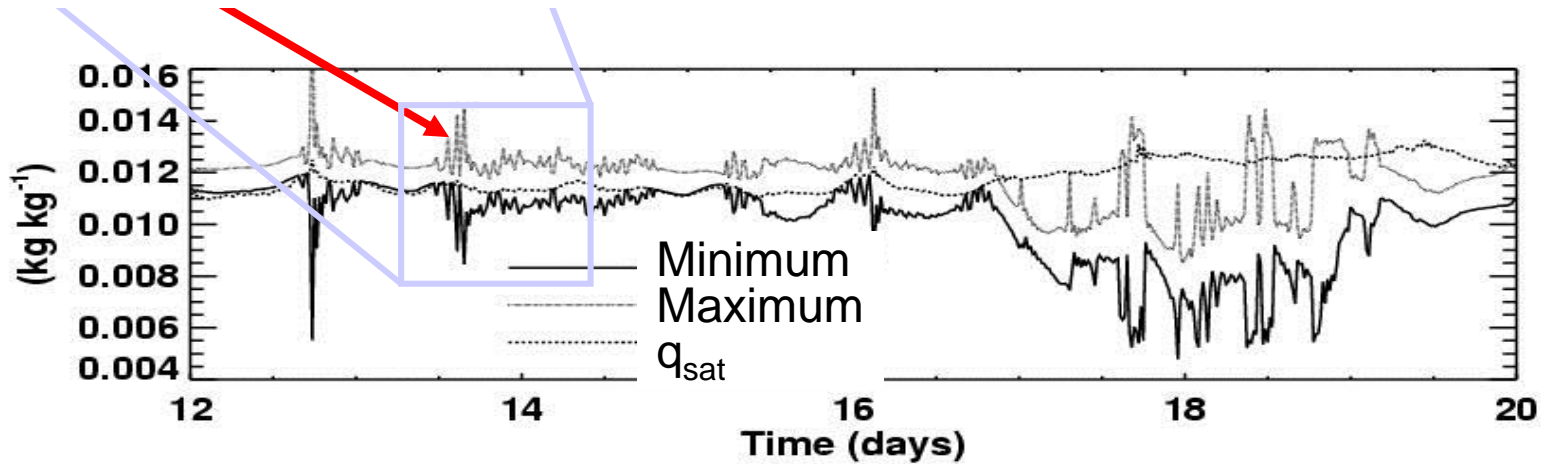
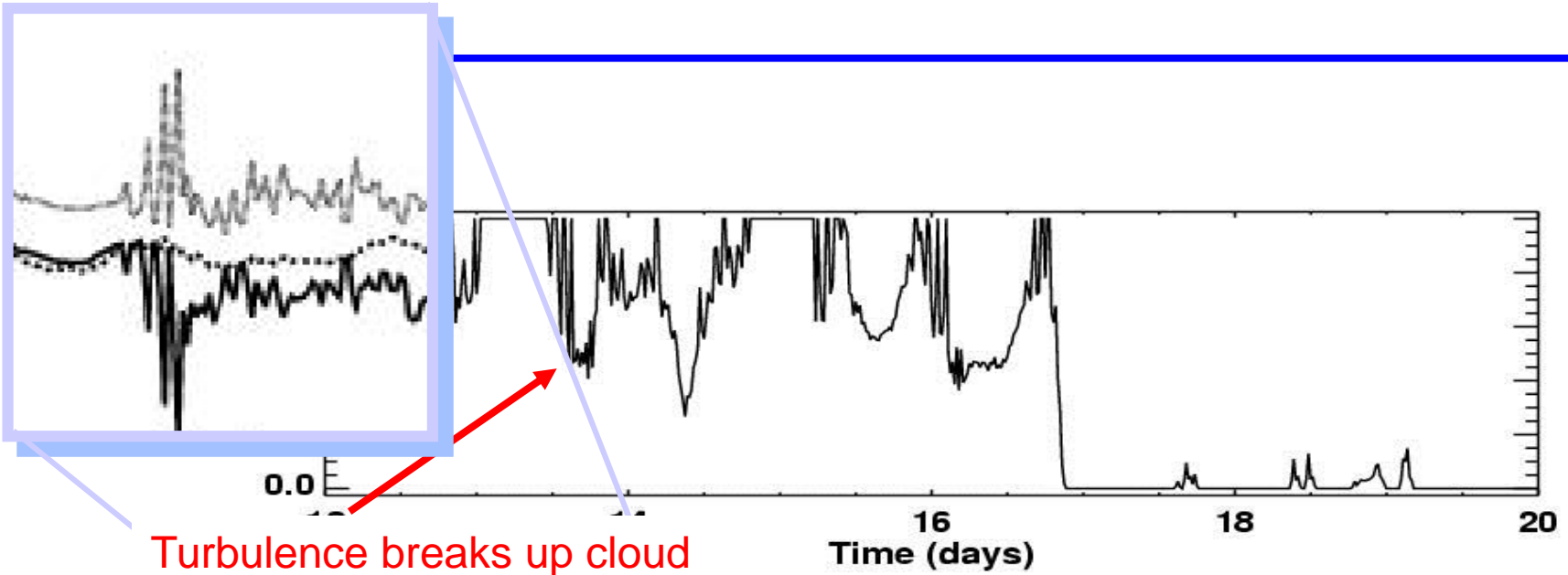


Evolution of stratocumulus cloud – Tompkins (2002)

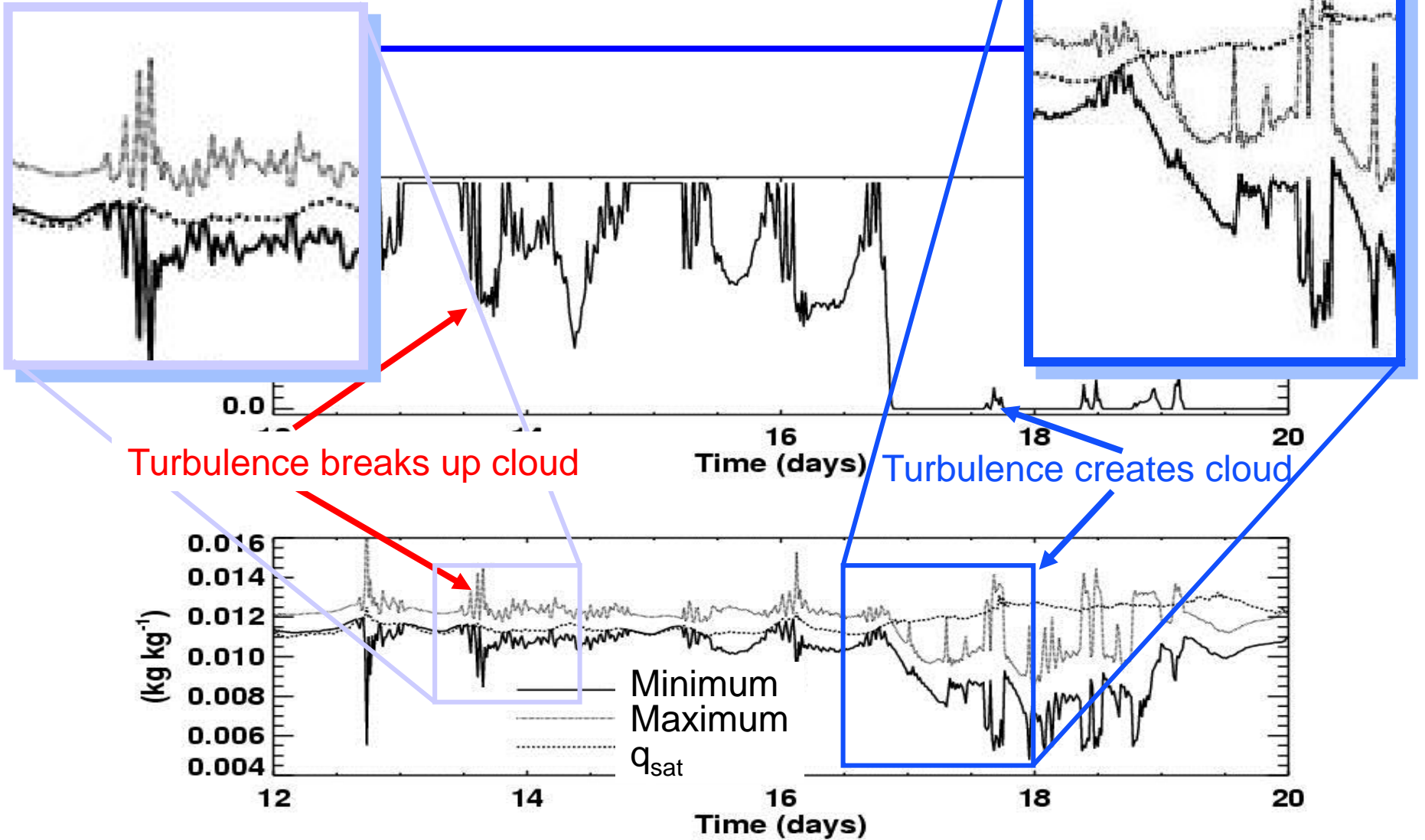




Prognostic Statistical Scheme in action



Prognostic Statistical Scheme in action



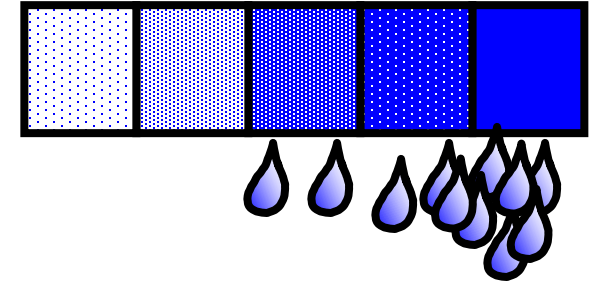
Building a statistical cloud scheme



Predicting change of q_t variance due to precipitation

- Change in variance due to precipitation

$$\frac{d \overline{q_t'^2}}{dt} = \overline{P' q_t'} = \int_{q_t = q_{sat}}^{q_{t_max}} P' q_t' G(q_t) dq_t$$



Where P is the precipitation generation rate, e.g:

$$P = K q_l \left(1 - e^{-\left(\frac{q_l}{q_{lcrut}}\right)^2} \right)$$

- However, the tractability depends on the PDF form for the subgrid fluctuations of q_t , given by G .



Some further issues for GCMs



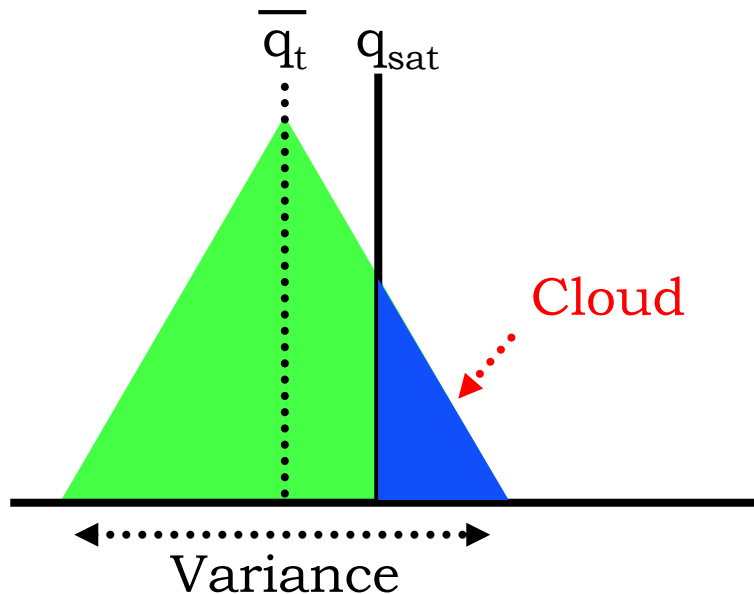
- **If we assume a 2-parameter PDF for total water, which prognostic variables should we use ?**
- How do we treat the ice phase when supersaturation is allowed ?
- How do we treat sedimentation ?

Prognostic statistical PDF scheme: Which prognostic variables/equations?

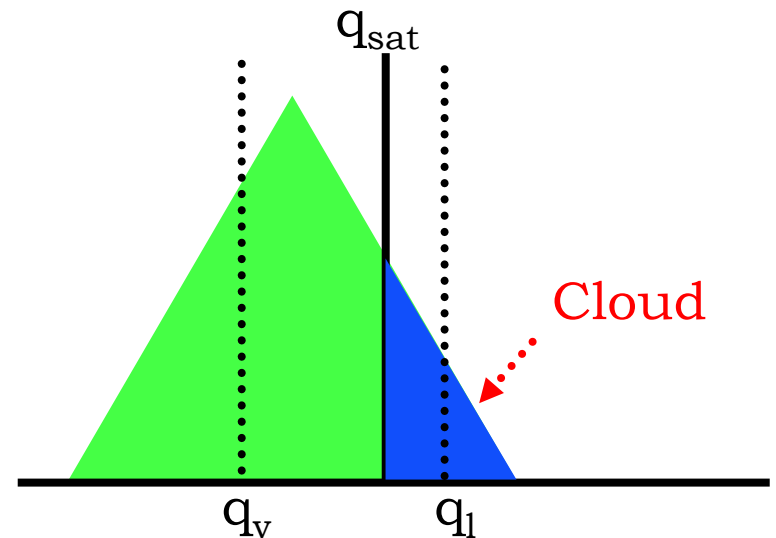


Take a 2 parameter distribution & partially cloudy conditions

- (1) Can specify distribution with
- (a) Mean
 - (b) Variance of total water



- (2) Can specify distribution with
- (a) Water vapour
 - (b) Cloud water mass mixing ratio

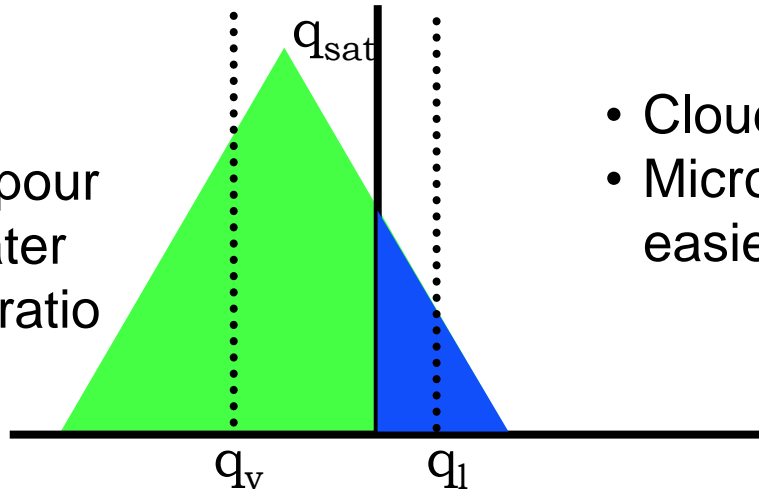


Prognostic statistical scheme:

(1) Water vapour and cloud water ?



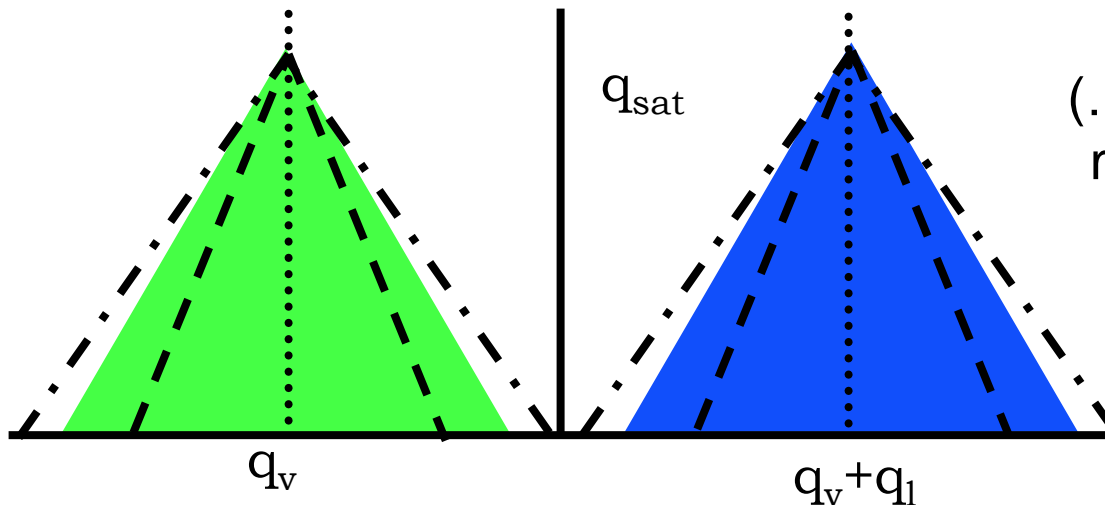
(a) Water vapour
(b) Cloud water
mass mixing ratio



- Cloud water budget conserved.
- Microphysical sources and sinks easier to parametrize.

But problems arise in...

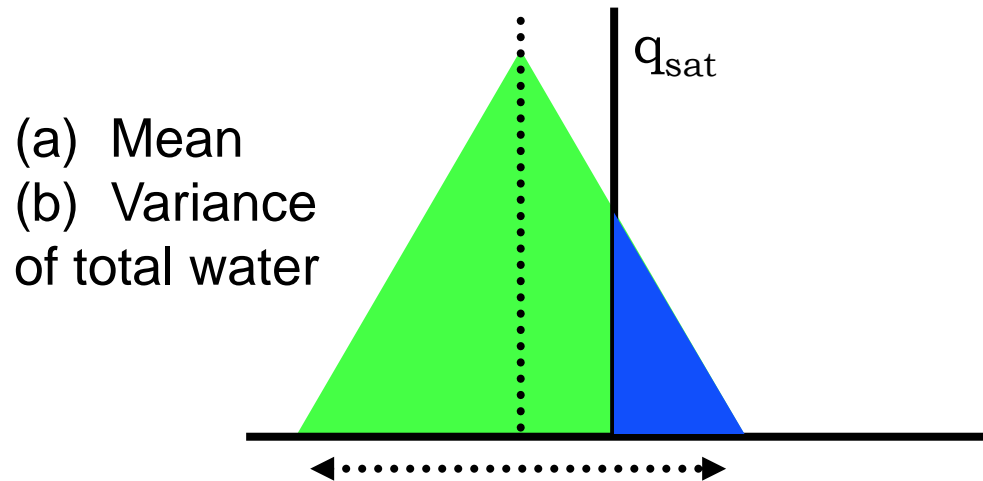
Clear sky
conditions
(turbulence)



Overcast
conditions
(...convection +
microphysics)

Prognostic statistical scheme:

(2) Total water mean and variance ?



- “Cleaner solution”.
- But conservation of liquid water may be difficult (eg. advection)
- Need to parametrize those tricky microphysics terms!



Some further issues for GCMs

- If we assume a 2-parameter PDF for total water, which prognostic variables should we use ?

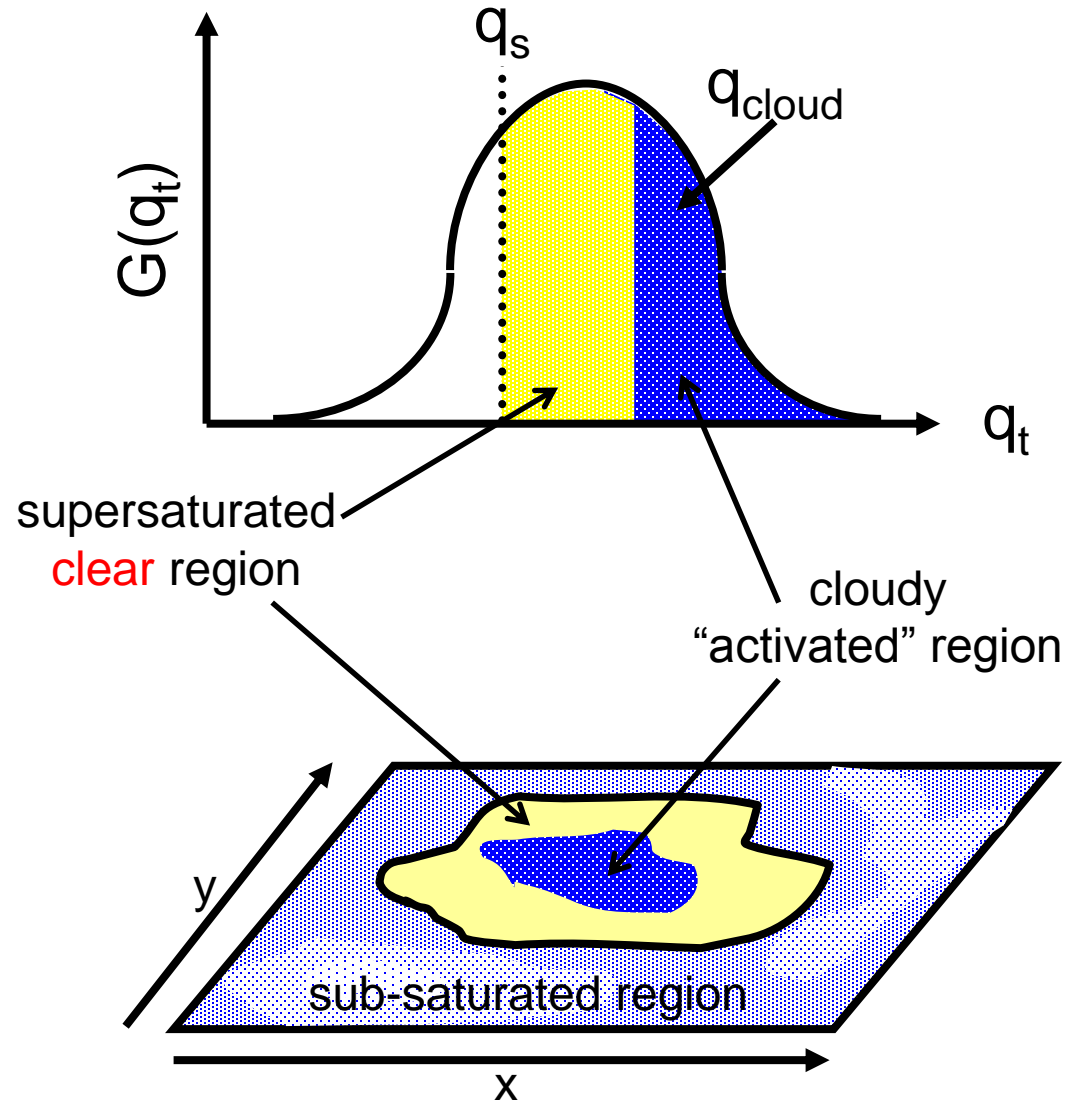


- **How do we treat the ice phase when supersaturation is allowed ?**
- How do we treat sedimentation ?

Prognostic statistical PDF scheme: How do we treat ice (and mixed-phase) cloud?



If supersaturation allowed, then the equation for cloud-ice no longer holds



$$q_i \neq \int_{q_s}^{\infty} (q_t - q_s) G(q_t) dq_t$$



Some further issues for GCMs

- If we assume a 2-parameter PDF for total water, which prognostic variables should we use ?
- How do we treat the ice phase when supersaturation is allowed ?



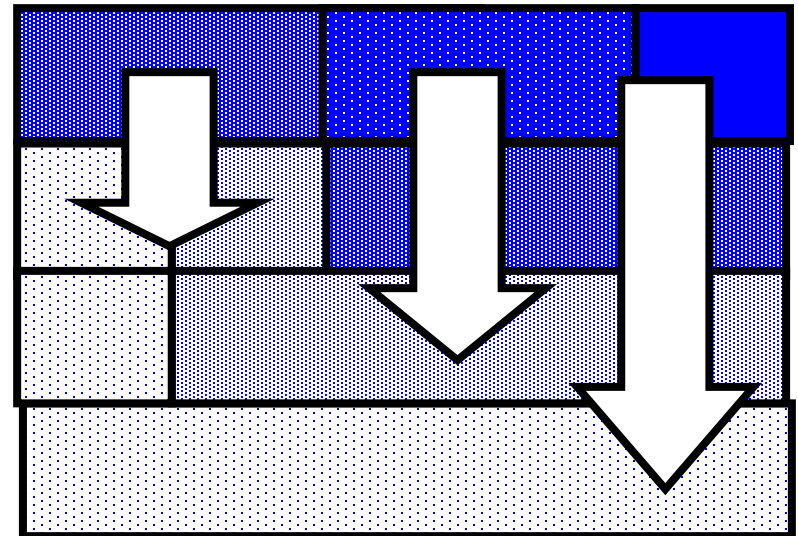
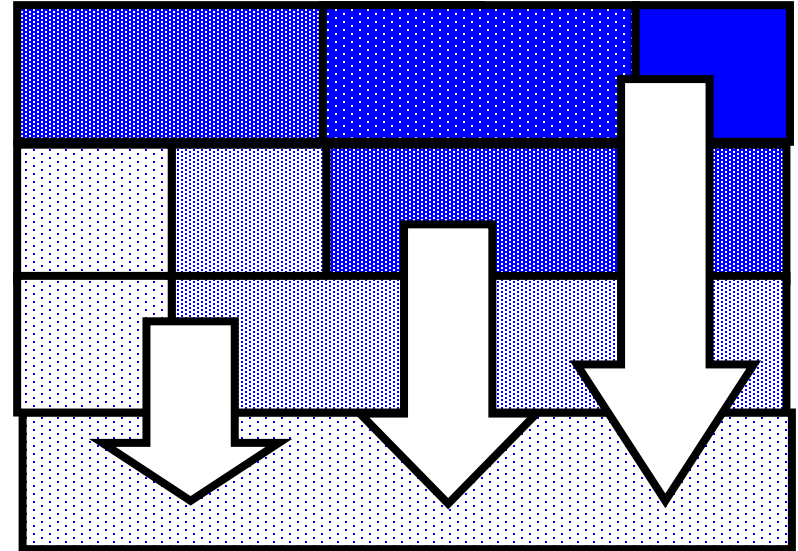
- **How do we treat sedimentation ?**

Prognostic statistical PDF scheme: How do we treat sedimentation ?



Can quickly get untractable !

- E.g: Semi-Lagrangian ice sedimentation
- Source of variance is far from simple, also depends on overlap assumptions
- Would really also like to retain the sub-flux variability too



Prognostic statistical PDF scheme: Knowing the PDF....



- **Advantages**

- Information concerning subgrid fluctuations of humidity and cloud condensate is available (for all parametrizations)
- Use of underlying PDF means cloud variables (condensate, cloud fraction) are always self-consistent.

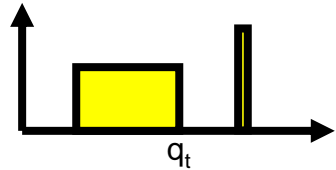
- **Challenges...**

- Deriving these sources and sinks rigorously is difficult, especially for higher order moments for more complex PDFs!
- Limited observations to define PDF
- If variance and skewness are used instead of cloud water and humidity, conservation of the latter is not ensured.
- Is a fixed PDF shape, even with variable moments, able to represent the wide range of variability in the atmosphere?
- How do we treat the ice phase, supersaturation, mixed-phase cloud, sedimentation? These are important questions!

Sub-grid cloud parametrization

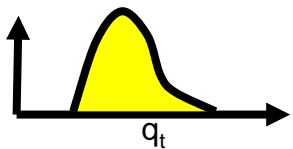


Current status in GCMs...?



Uniform-delta:
Tiedtke (1993)

- The ECMWF global NWP model has prognostic water vapour, cloud water and cloud fraction (for the warm phase). With a uniform function for heterogeneity in the clear air and a delta function (homogeneous) in-cloud ([more next time](#)....)
- The UK Met Office global NWP model (PC2 scheme) also has prognostic water vapour, cloud water and cloud fraction (for the warm phase).
- Many other operational global NWP/climate models have diagnostic sub-grid cloud schemes, e.g. NCEP GFS: Sundquist et al. (1989)
- Research is ongoing for statistical schemes with prognostic PDF moments (e.g. Tompkins scheme tested in ECHAM, CLUBB being tested in CAM).



Beta:
Tompkins (2002)

Summary



Representing subgrid scale heterogeneity

- Representing sub-gridscale heterogeneity in GCMs is important for cloud formation, microphysical processes, radiation etc.
- Many different approaches have been tried, with varying degrees of complexity to represent the variability observed in the atmosphere.
- More degrees of freedom allow greater flexibility to represent the real atmosphere, but we need to have enough knowledge/information to understand and constrain the problem (form of pdf/sources/sinks)!
- Cloud, convection and BL turbulence are all part of the subgrid heterogeneity – active research into unified schemes.
- Statistical prognostic PDF schemes have many advantages but challenges remain for clouds other than warm-phase boundary layer cloud!
- However, we should continue to strive for a **consistent representation of this heterogeneity** for all processes in the model.



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