Sources of predictability beyond the deterministic limit

Sarah Keeley giving slides by Franco Molteni European Centre for Medium-Range Weather Forecasts

Outline

Persistent anomalies in the tropics and extra-tropics: examples from the last two decades

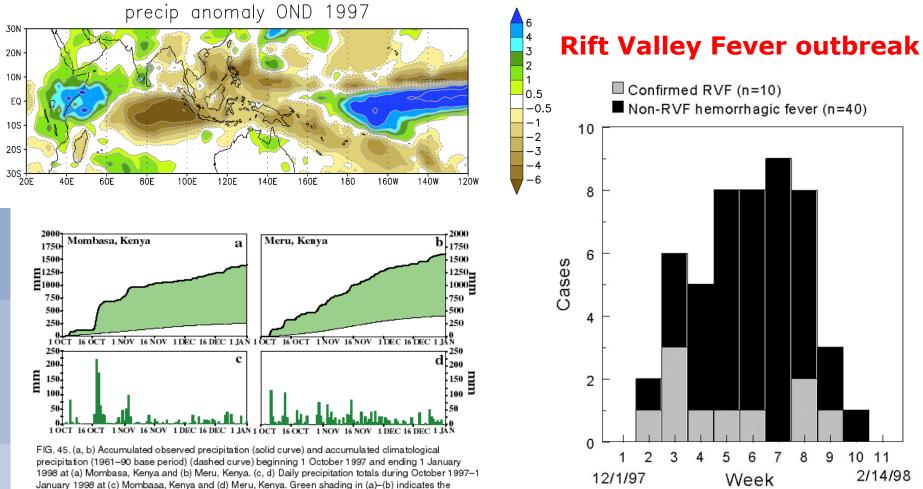
Beyond deterministic predictability in non-linear, chaotic systems: the role of variability in surface conditions and energy/water fluxes

Coupled ocean-atmosphere variability in the tropics and its teleconnections with the extra-tropical flow

Ensemble predictions in the "extended" range: estimates of predictability and actual skill

Predictability on the weekly/monthly time scale arising from sub-seasonal tropical variability and teleconnections

Oct-Dec 1997: floods in East Africa

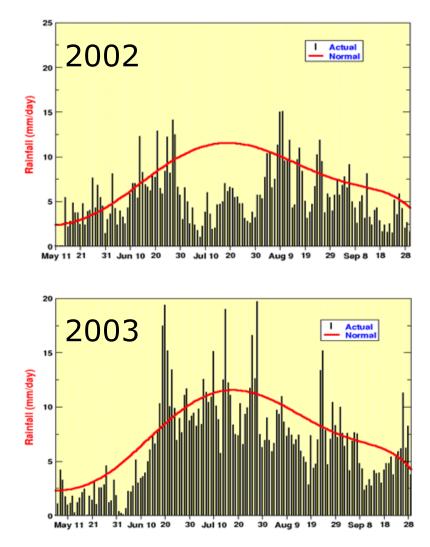


difference between the observed and normal accumulated rainfall.

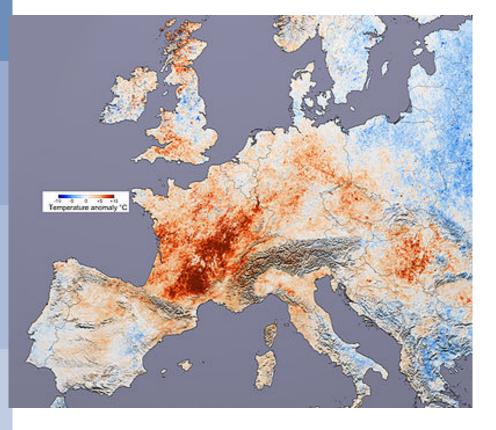
July 2002: drought in India



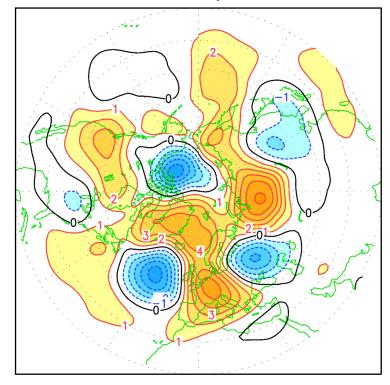
All-India Rainfall time series May - October

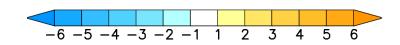


Summer 2003: European heat-wave

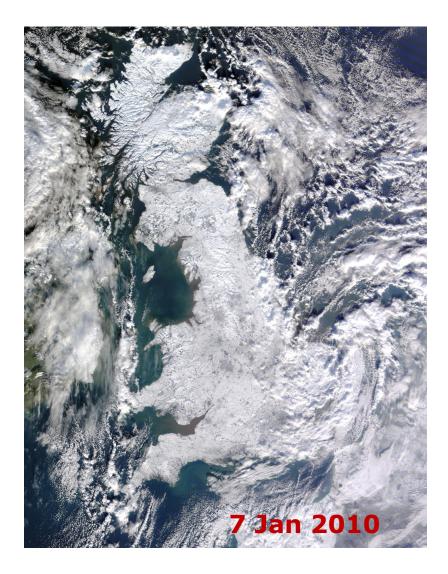


Z 500 anomaly JJA 2003

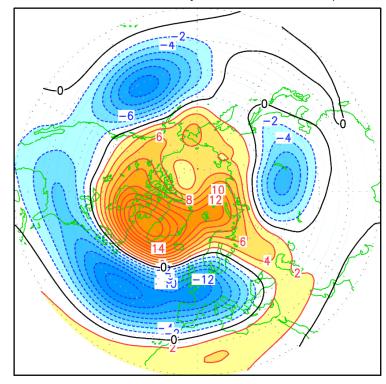


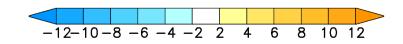


Winter 2009-2010: cold anomaly over N. Europe

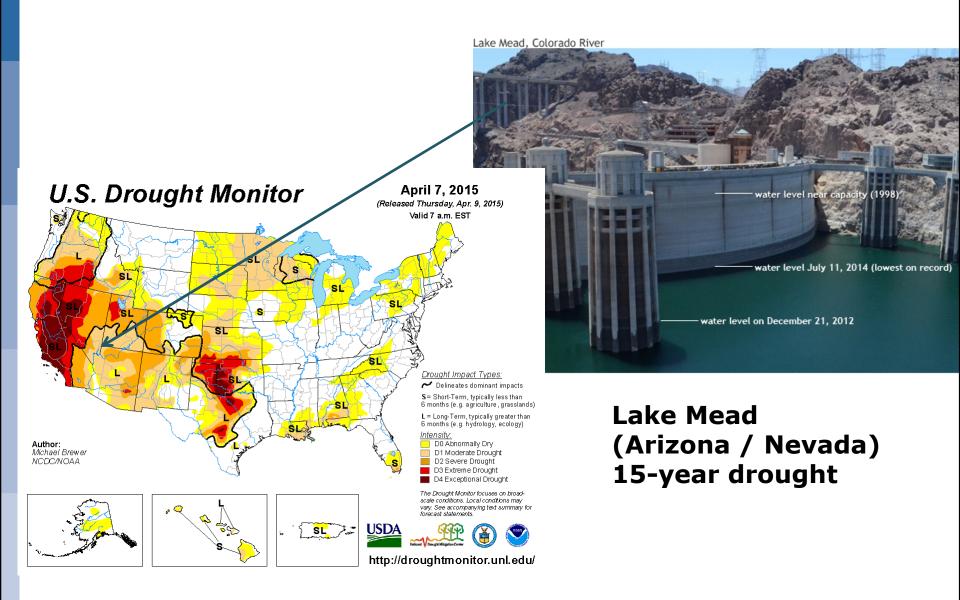


Z 500 anomaly DJF 2009/10





Drought in South-western USA

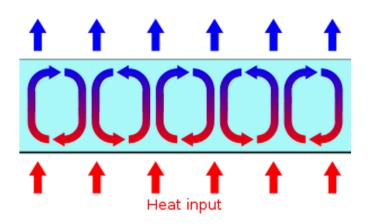


Chaotic behaviour in non-linear systems 3-variable model of Rayleigh-Benard convection (Lorenz 1963) $dX/dt = \sigma (Y - X)$

- dY/dt = -XZ + rX Y
- dZ/dt = X Y b Z

Unstable stationary states

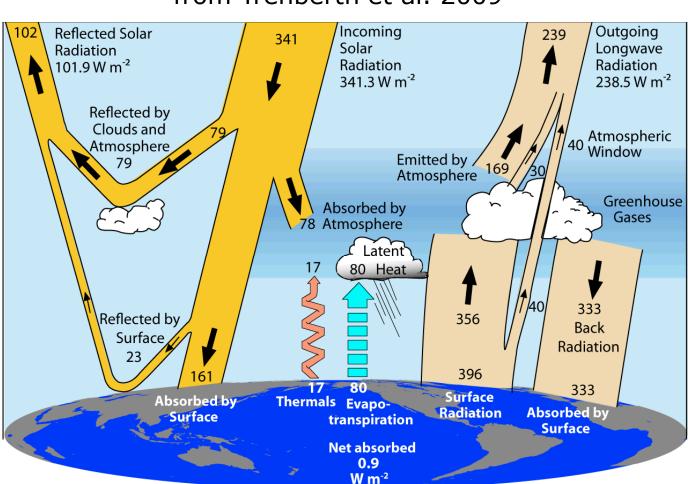
X = Y = Z = 0 $X = Y = \pm [b (r - 1)] \frac{1}{2}, Z = r - 1$







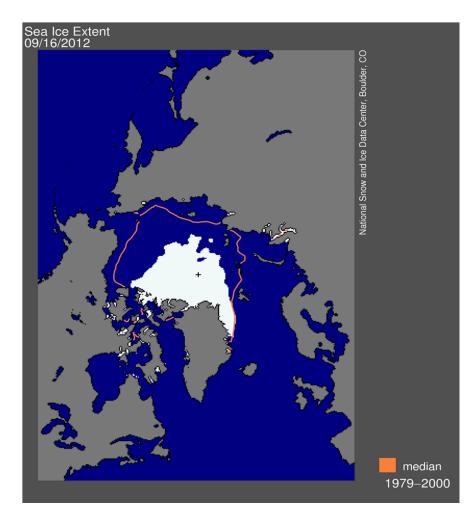
The global energy cycle (2000-2004)



from Trenberth et al. 2009

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Sea ice: Interaction of climate change and natural variability

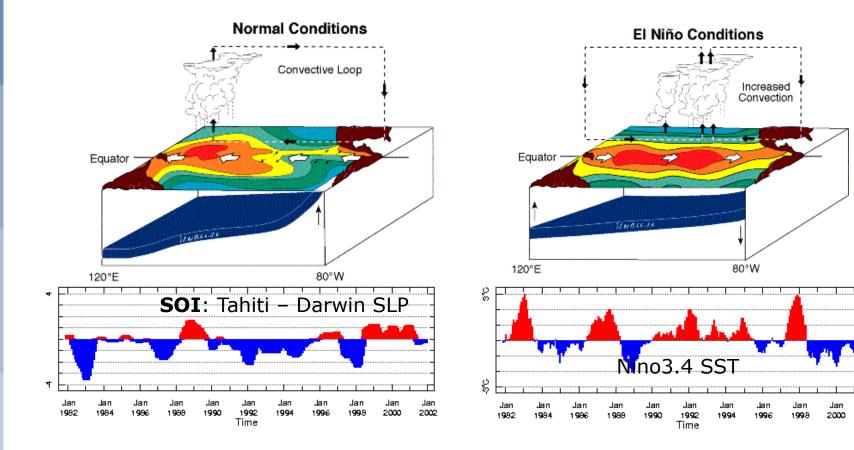


Record minimum in Arctic sea-ice extent: 16/9/2012 (from NSIDC)



El Niño and the Southern Oscillation

Walker and Bliss (1932); Bjerknes (1969)

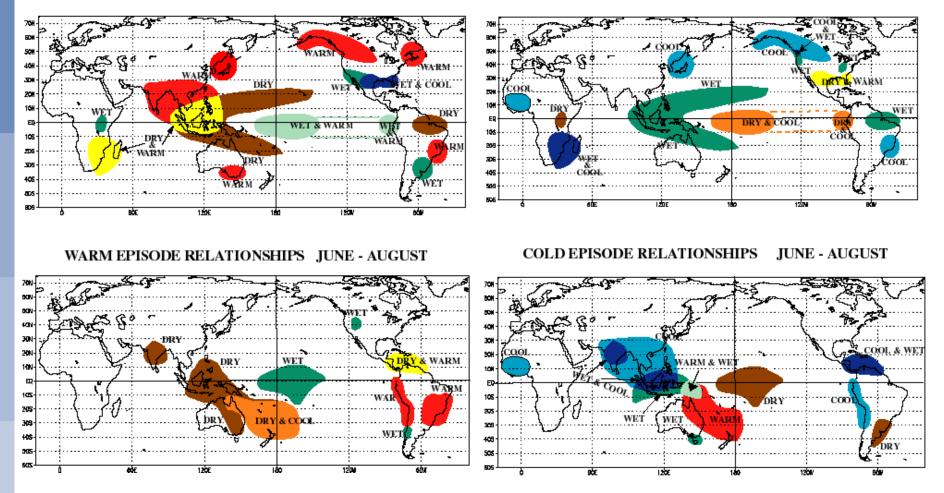


Jan

2002

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ENSO impacts: rainfall and temperature



WARM EPISODE RELATIONSHIPS DECEMBER - FEBRUARY

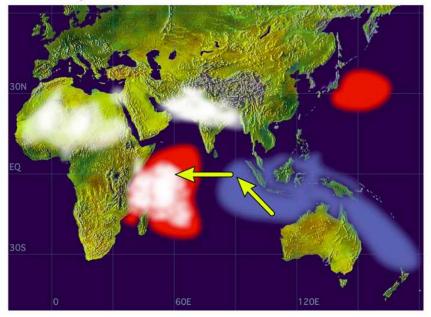
COLD EPISODE RELATIONSHIPS DECEMBER - FEBRUARY

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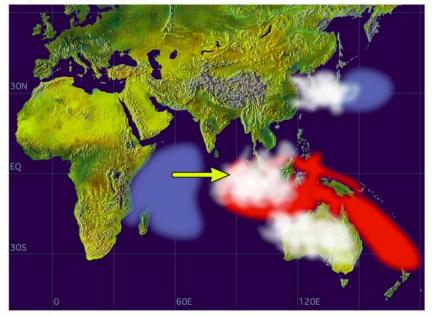
The Indian Ocean Dipole (or I.O. Zonal Mode)

Saji et al. (1999) Webster et al. (1999)

Positive Dipole Mode



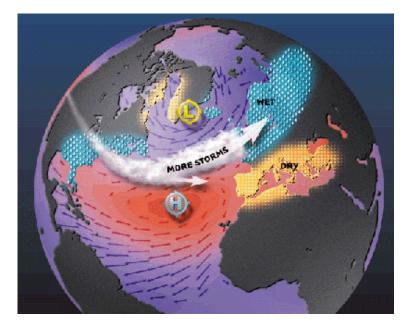
Negative Dipole Mode



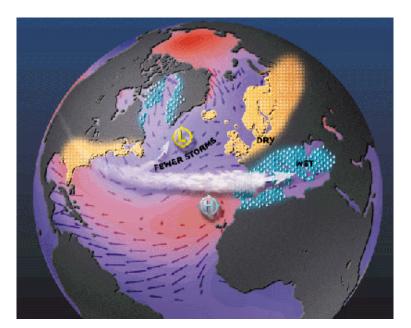


The North Atlantic Oscillation

Walker and Bliss (1932) Van Loon and Rogers (1978)



Positive NAO phase

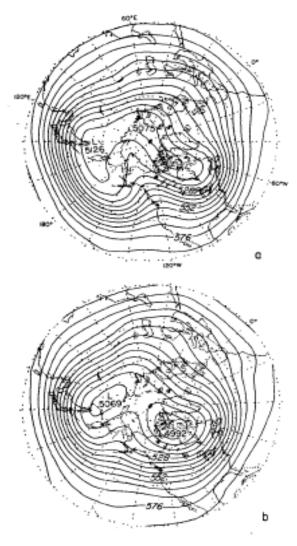


Negative NAO phase



The Pacific /North American (PNA) pattern

500-hPa height composites from Wallace and Gutzler 1981



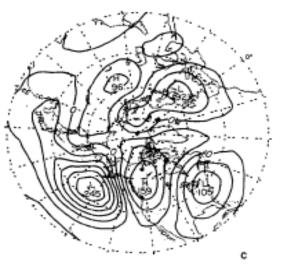
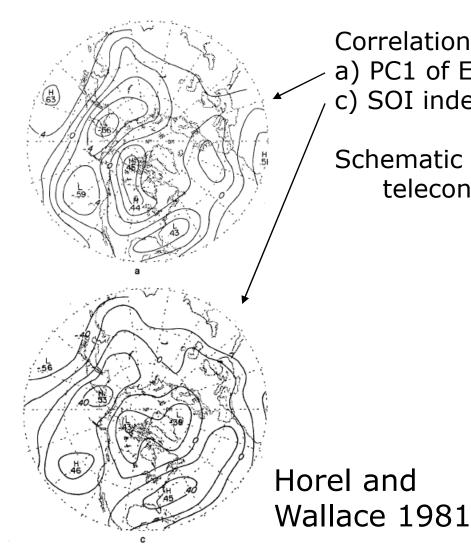


FIG. 17. As in Fig. 13 except for the Pacific/ North American pattern.

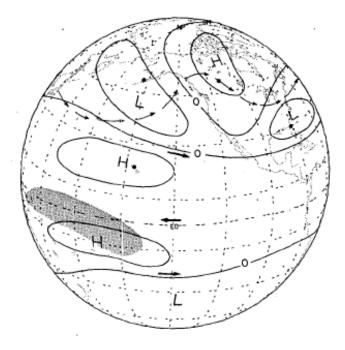


Teleconnections with ENSO



Correlation of 700hPa height with a) PC1 of Eq. Pacific SST c) SOI index

Schematic diagram of tropical-extratropical teleconnections during El Niño





Multiple flow regimes in non-linear models

Charney and DeVore 1979: multiple steady states of a loworder barotropic model with sinusoidal bottom topography

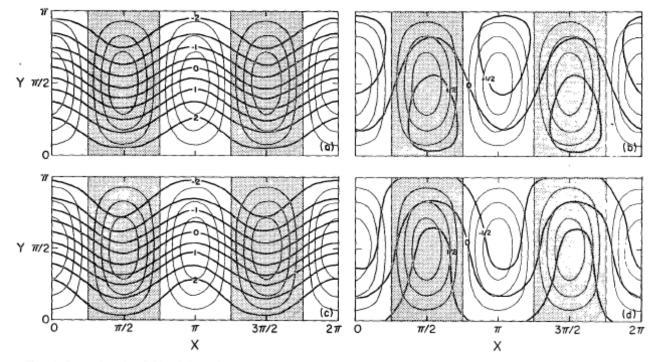
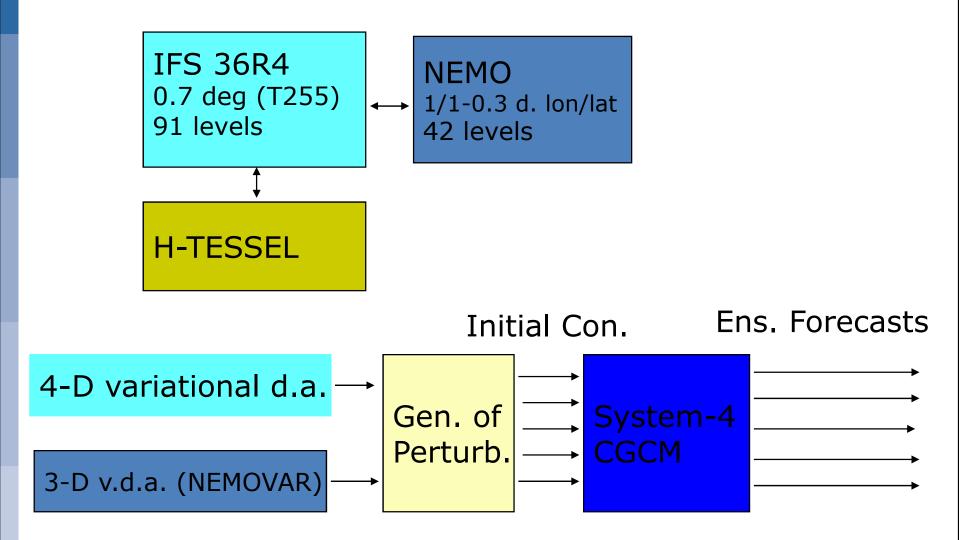


FIG. 4. Streamfunction fields of the stable first mode equilibria of a topographically forced flow for $k = 10^{-2}$, $L/a = \frac{1}{4}$, n = 2, $h_0/H = 0.2$ and $\psi_A^* = 0.2$; for the spectral model above resonance (a) and slightly below resonance (b); and for the grid-point model above resonance (c) and slightly below resonance (d). The nondimensional topographic heights are shown with light lines; the contour spacing is 0.05 units, with negative regions shaded.

The ECMWF Seasonal fc. system (Sys-4)



ECMWF System 4: main features

Operational forecasts

51-member ensemble from 1st day of the month
released on the 8th
7-month integration

Experimental ENSO outlook

13-month extension from 1st Feb/May/Aug/Nov15-member ensemble

Re-forecast set

30 years, start dates from 1 Jan 1981 to 1 Dec 2010



Variability in an ensemble of time-evolving fields

$$\boldsymbol{F}(t, m) = \boldsymbol{F}(t', j, m)$$

t' = time within yearj = 1, N (no. of years) m = 1, M (no. of ens. members)

Climatology:
$$F_{cl}(t') = \{F(t', j, m)\}_{j, m}$$

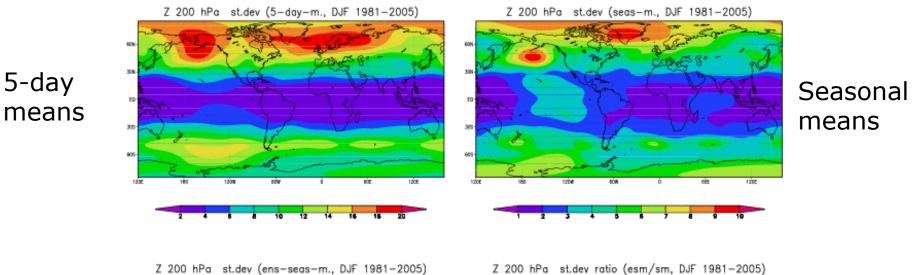
Anomaly: $A(t', j, m) = F(t', j, m) - F_{cl}(t')$

Seasonal mean anomaly: $\mathbf{A}_{s}(j, m) = {\mathbf{A}(t', j, m) }_{t'}$ Ens./seas. mean anomaly: $\mathbf{A}_{es}(j) = {\mathbf{A}(t', j, m) }_{t', m}$

- for an observation/analysis dataset, M =1 !
- in a "perfect" model environment, the average correlation between *A_{es}* and *A_s* is equal to the ratio of their standard deviations.

Variability of Z 200hPa in DJF from seasonal ensembles

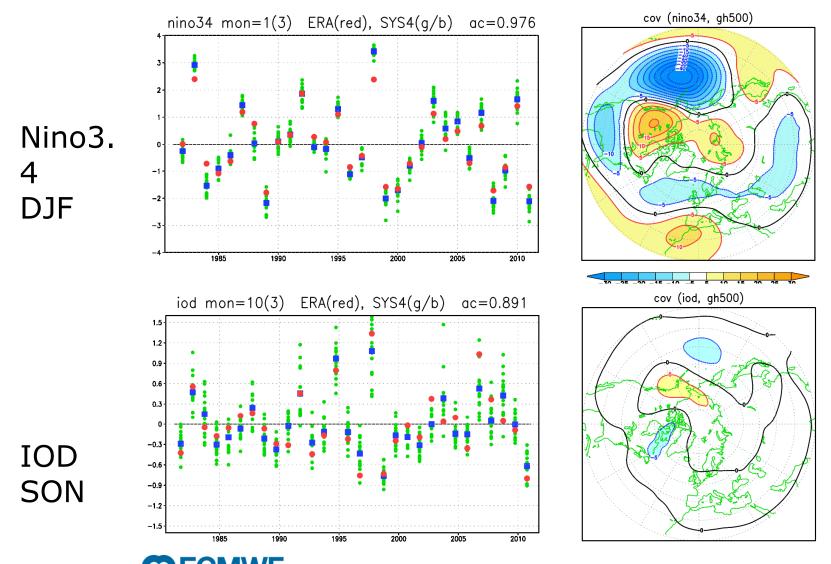
Standard deviation from 11-member ensembles, DJF 1981/2005



Seasonal - ensemble means - be and the second of the secon

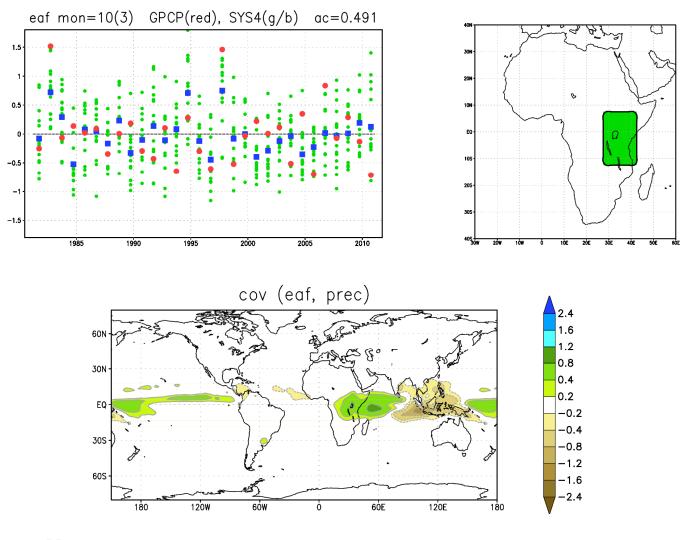
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Prediction of tropical SST anomalies in Sys4



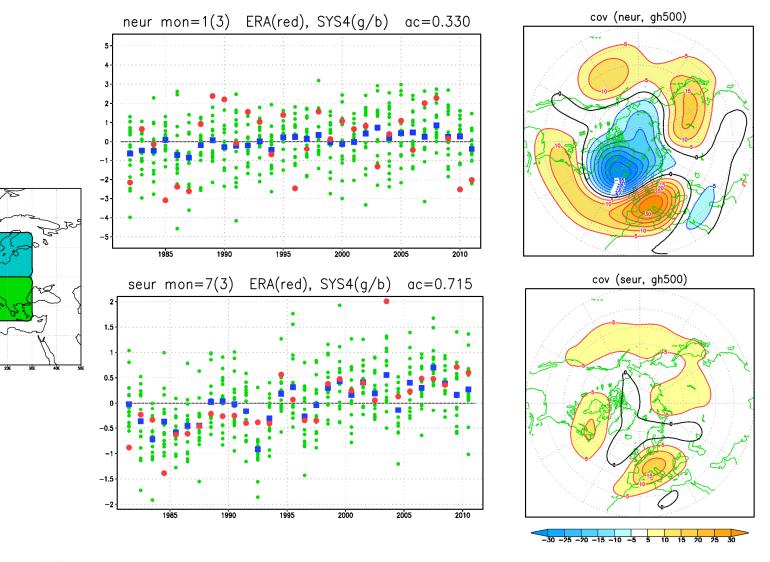
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Prediction of tropical rainfall in Sys4: East Africa (SON)



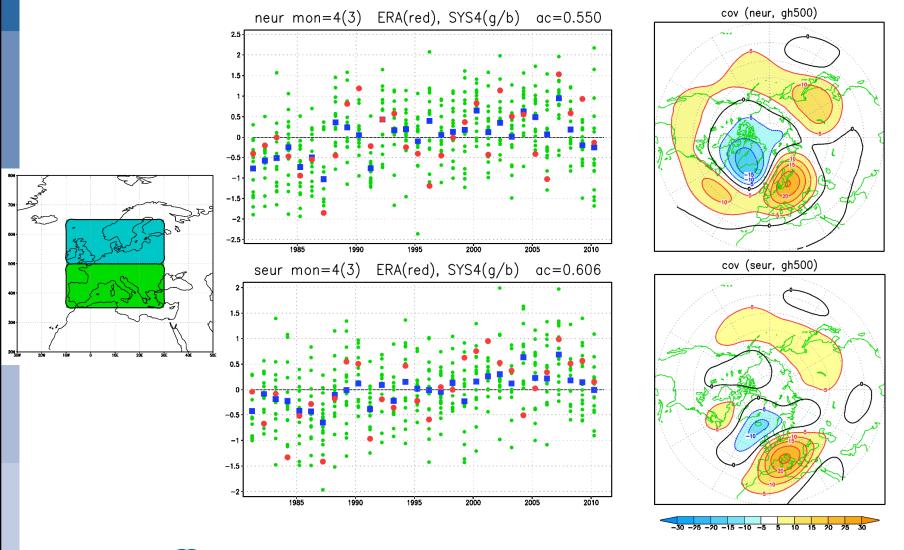


Prediction of 2-m temperature in Sys4: Europe (DJF, JJA)



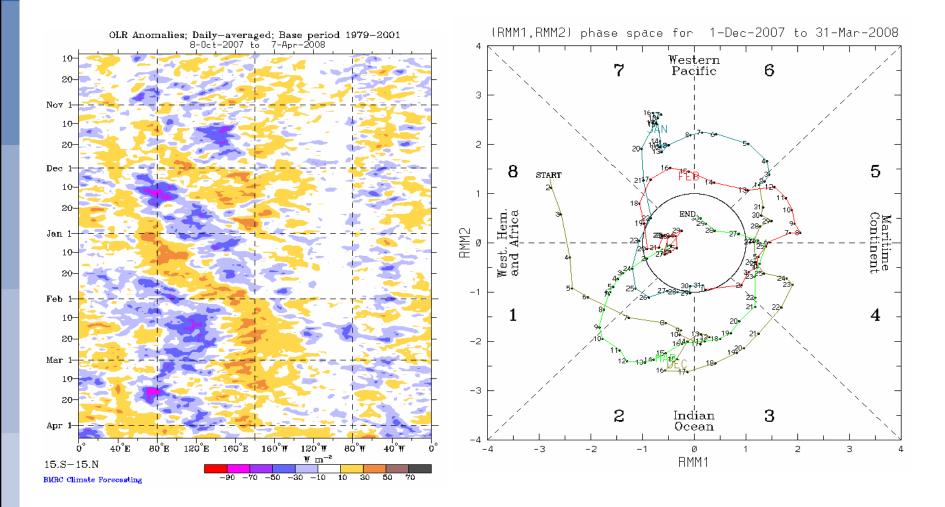
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Prediction of 2-m temperature in Sys4: Europe (MAM)



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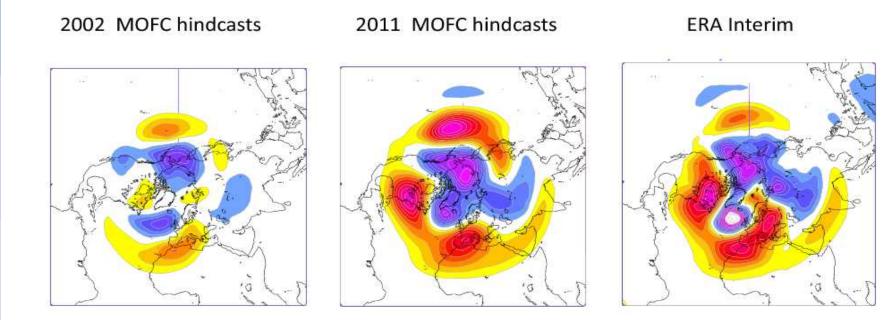
Sub-seasonal variability: the Madden-Julian Oscillation



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MJO teleconnections in October-March

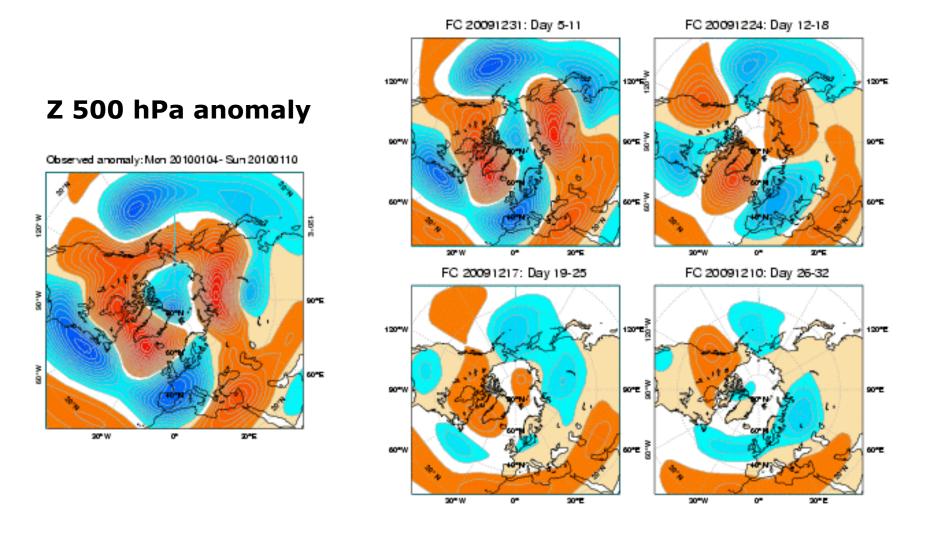
500 hPa height, MJO phase 3 + 10 days



from Vitart 2014



ECMWF monthly fc. for 4-7 Jan. 2010



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Conclusions

Regional anomalies in atmospheric flow and weather parameters may persist on time scales longer than the deterministic predictability limit, and have substantial societal impacts.

The possibility of performing probabilistic predictions of these events arises from the interaction of the atmospheric flow with slowly varying anomalies in surface conditions, which modify the energy and water sources for the atmosphere.

In the extratropics, persistent anomalies can be generated by (linear) teleconnections with tropical variability (eg ENSO) but also from the alternation of different (non-linear) flow regimes.

Ensemble prediction systems provide an <u>estimate</u> of long-range predictability based on the ratio of ensemble spread and ensemble-mean variability.

Predictability over Europe: limited by strong internal variability during winter (but with significant teleconnections on the *sub-seasonal scale*), higher in other seasons when internal variability is reduced.

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