Improving land-atmosphere data assimilation coupling in NWP

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NOAA's global NWP GFSv17 upgrade

- NOAA is undertaking a major upgrade to our global NWP system • GFSv17 is currently scheduled for launch in 2025, and will include (among
- many other upgrades):
 - Updating the Noah land surface model to Noah-MP Phased introduction of the Joint Effort for Data assimilation Integration (JEDI)
 - software as the DA platform
 - A major update to our land analysis
- The new land analysis will include:
 - JEDI, from station snow depth observations and (IMS) satellite snow cover
 - Upgrade of the current snow analysis, to an OI-based snow depth analysis in Introduction of a soil moisture and soil temperature analysis











A new soil moisture/soil temperature analysis for NOAA's global NWP

- The new soil analysis:
 - update, rather than implementing a separate land-only DA scheme
 - Initially based on assimilation of screen-level T and RH Performed by expanding the atmospheric DA system to also do the soil
- This presentation:
 - Demonstrate that we can update the soil moisture and temperature by expanding our atmospheric DA system
 - Test different options for coupling the land and atmosphere updates
 - Atmospheric DA uses the GSI Hybrid 4D-EnVar
 - For now, use only EnKF (LETKF) rather than the full hybrid DA to establish best coupling arrangement / use of screen-level observations







Land ensemble spread in NWP systems

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Ensemble Spread

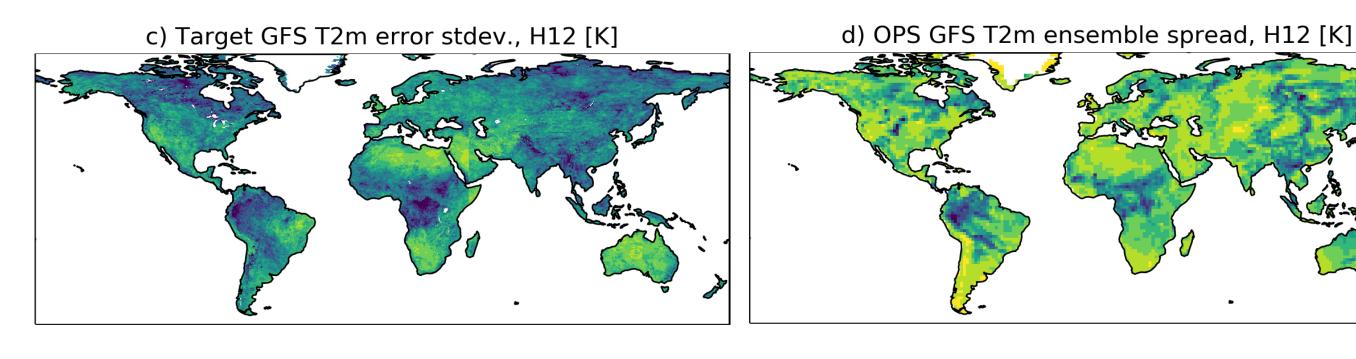
- NWP ensembles are underdispersed at the land surface
 - Expected, since ensembles are not explicitly perturbed to account for land model uncertainty
- Previous work: Tested different approaches to adding a scheme to represent forecast uncertainty at/ near land in NOAA's NWP ensemble system
- See: Draper, C., 2021, J. Hydromet



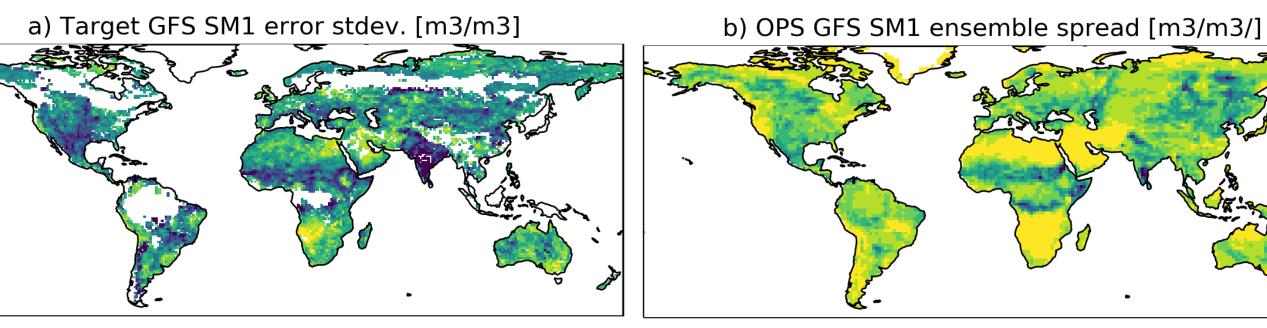




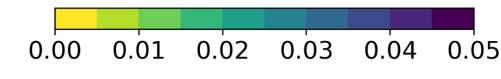




Boreal summer forecast soil moisture, layer 1 (SM1) error standard deviation [m3/m3]







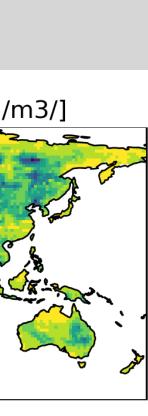
Boreal summer daytime model T_{SL} error standard deviation.

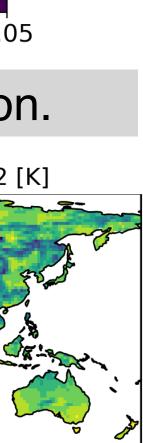


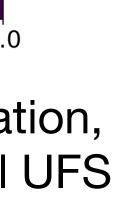
Target estimates, calculated using triple colocation (SM1), and from archived operational UFS comparison to ERA-5 anal. (T_{SL})

1.2 1.8 0.6 2.4 0.0

Ensemble standard deviation, output







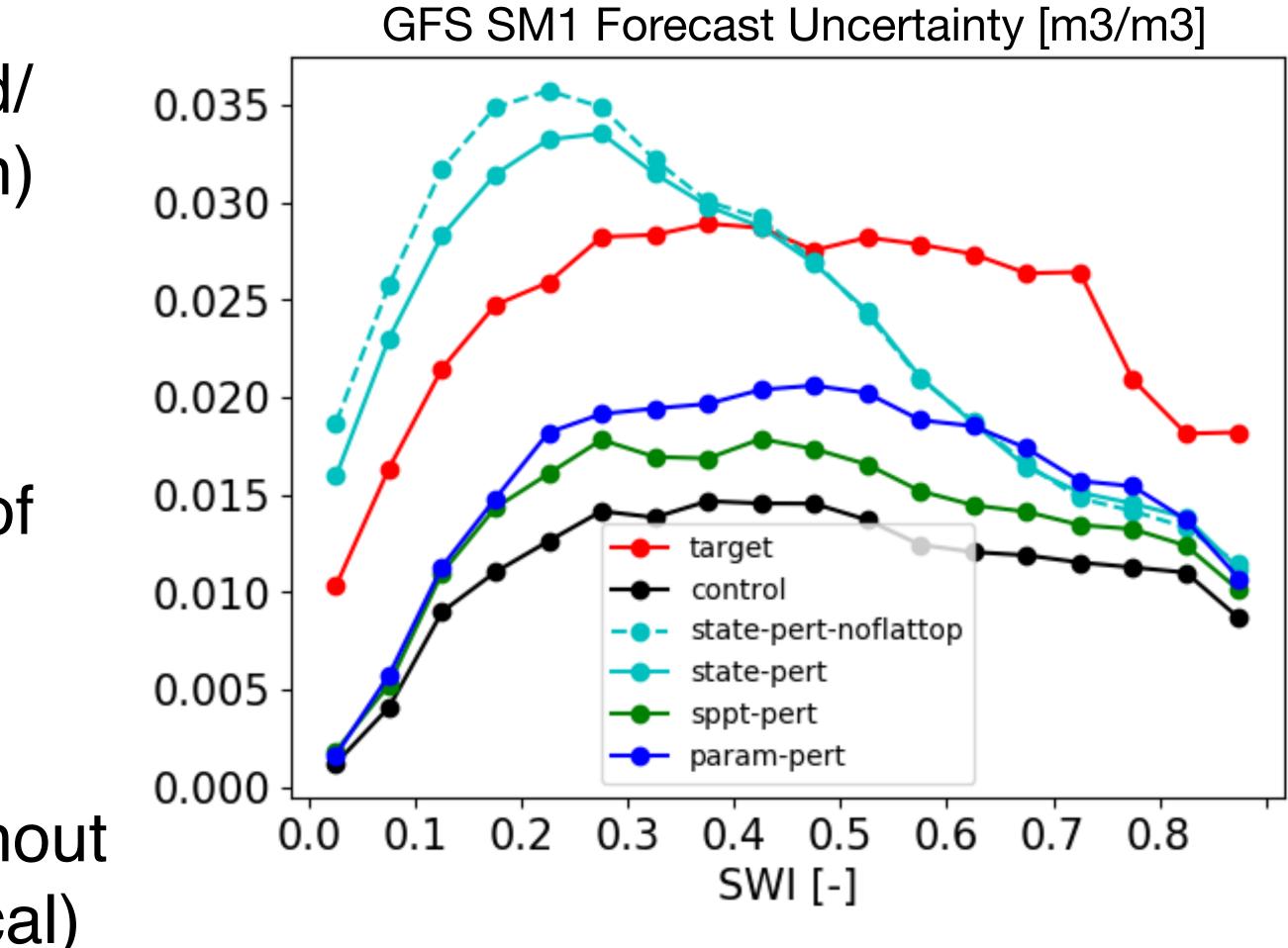




Ensemble Spread

- Recommended method is to perturb key model inputs controlling the land/ atmosphere fluxes (e.g. veg. fraction)
 - Generates reasonable spatial patterns in spread
 - Generates ensemble crosscovariances more representative of coupled land/atmosphere errors
- However, land is highly non-linear; difficult to obtain desired spread without changing ensemble mean (impractical)









Land/Atmosphere DA experiments



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Land/Atmosphere DA Experiments

	Update atmosphere*		Update soil moisture+temperature		
	from standard atmos obs	from screen-level obs	from standard atmos obs	from screen-level obs	
Control	X				
Screen	X	X			
SfcUpd	X		X		
Screen+SfcUpd	X	X	X	X	
SfcUpd-Weak	X			X	

* All experiments include bug-fixes/updates to the assimilation of conventional q obs.

- DA: GSI EnKF (LETKF)
- Model: GFSv17 (HR1 tag)
- Includes Noah-MP (new land model being introduced for GFSv17) Land model perturbation scheme not activated (still adapting to Noah-MP)
- Resolution: C192 (50 km), 127 atmos levels & 4 soil levels Period: 5-20 June, 2022 (eval last 10 days)
- Evaluation: assess impact on conventional (sondes, station observations) O-F for q, T

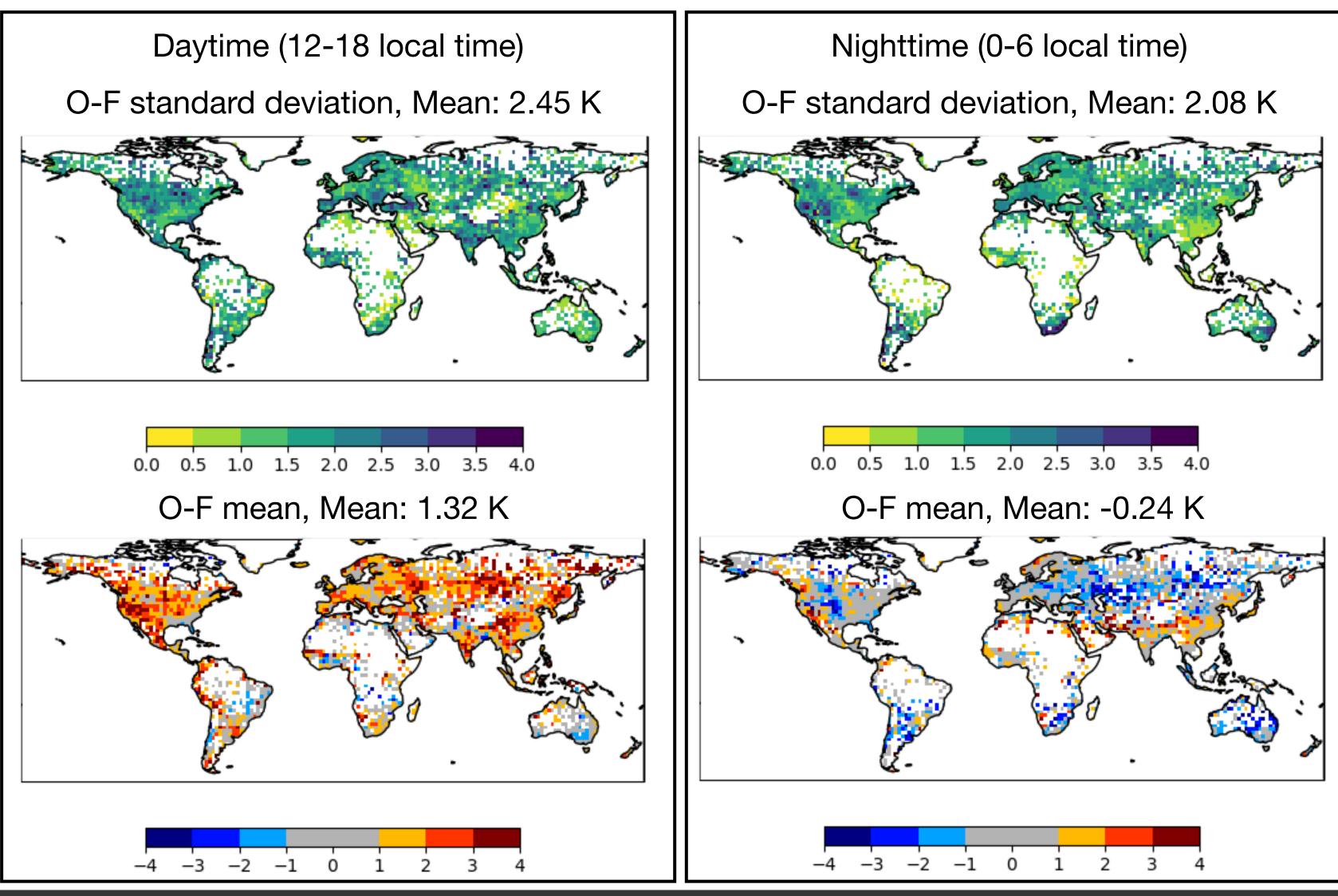








Control O-F for Screen-Level Temperature (Tsl)





- Substantial day-time cool model bias, lesser night-time warm bias
 - Sondes show similar bias, reduces rapidly away from surface
 - Noah-MP still being tuned; currently testing a potential solution to the diurnal T bias
- The T_{SL} daytime bias will • results in sub-optimal DA
 - Vertical T correlations much weaker during the day -> daytime T_{SL} obs expected to have lesser impact







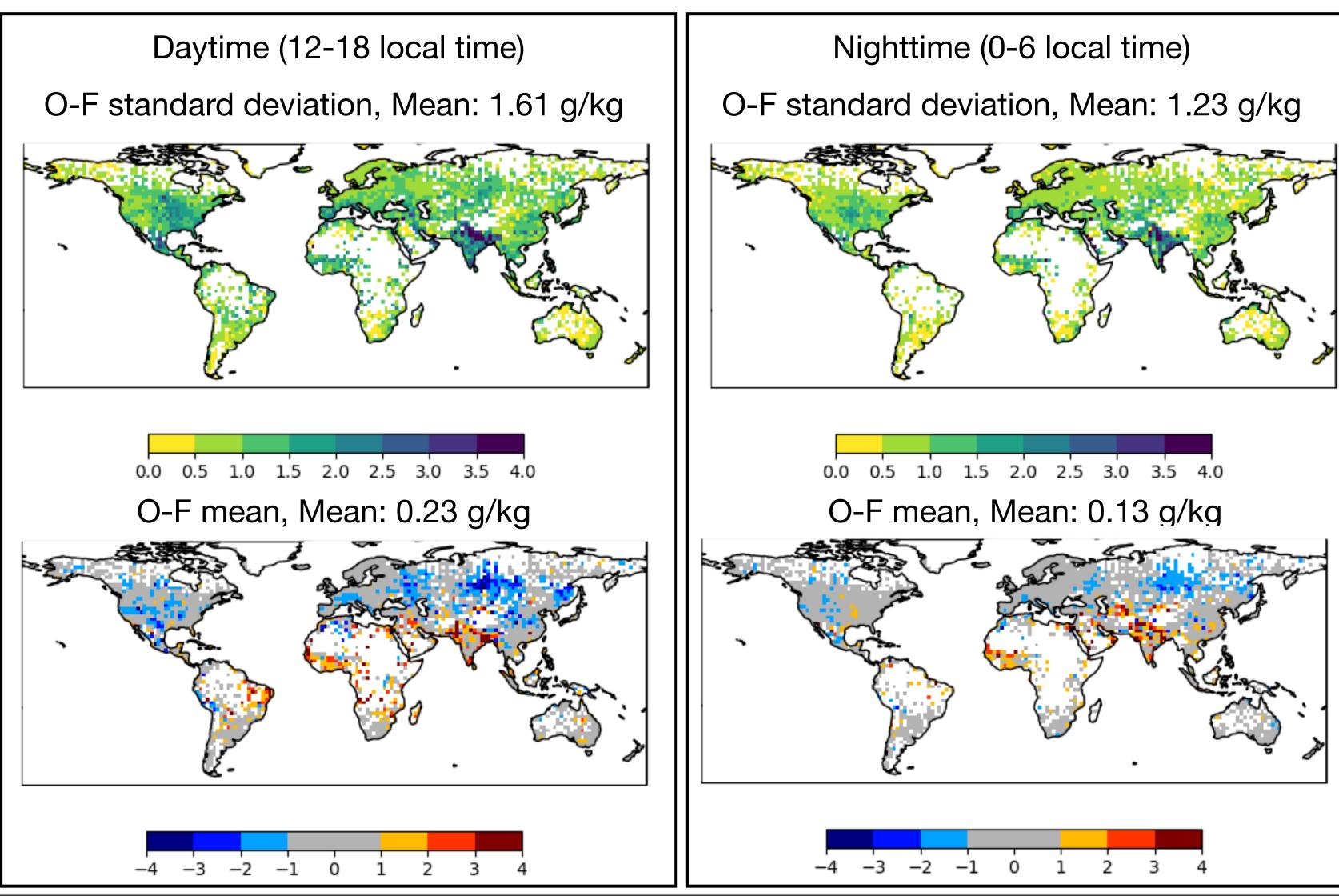








Control O-F for Screen-Level Humidity (qsl)

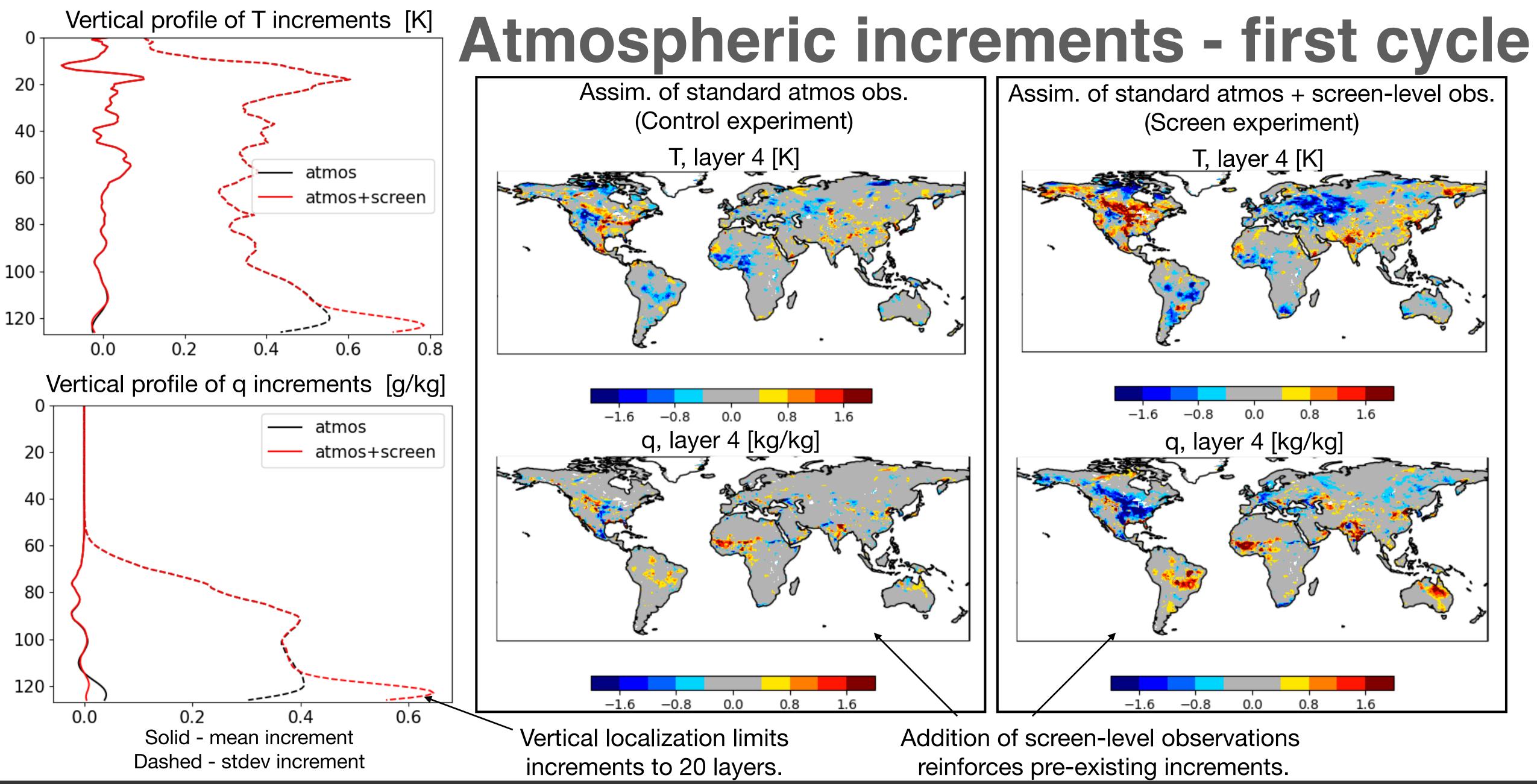


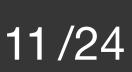
- Small wet bias in some regions, has minimal diurnal cycle

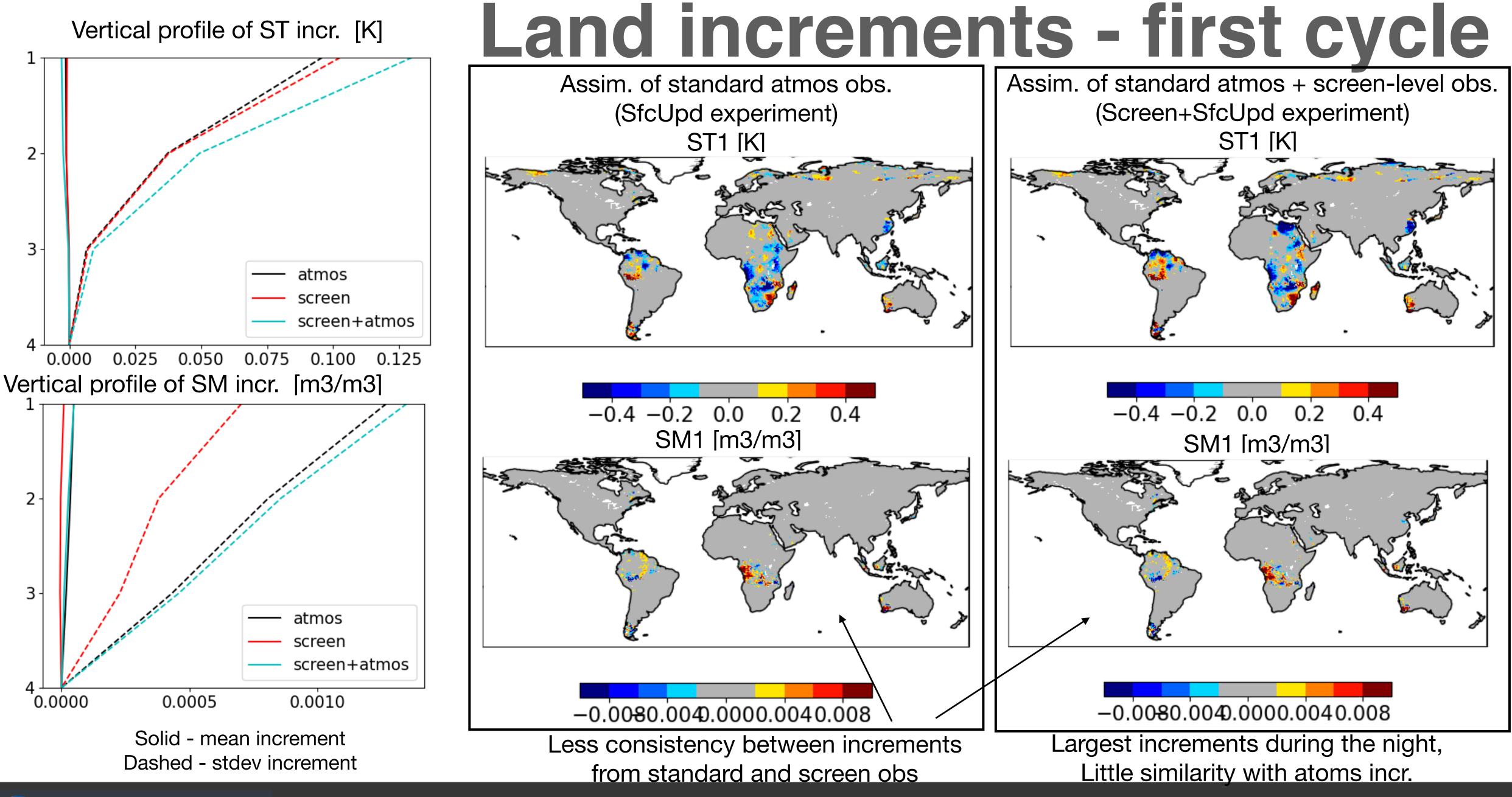




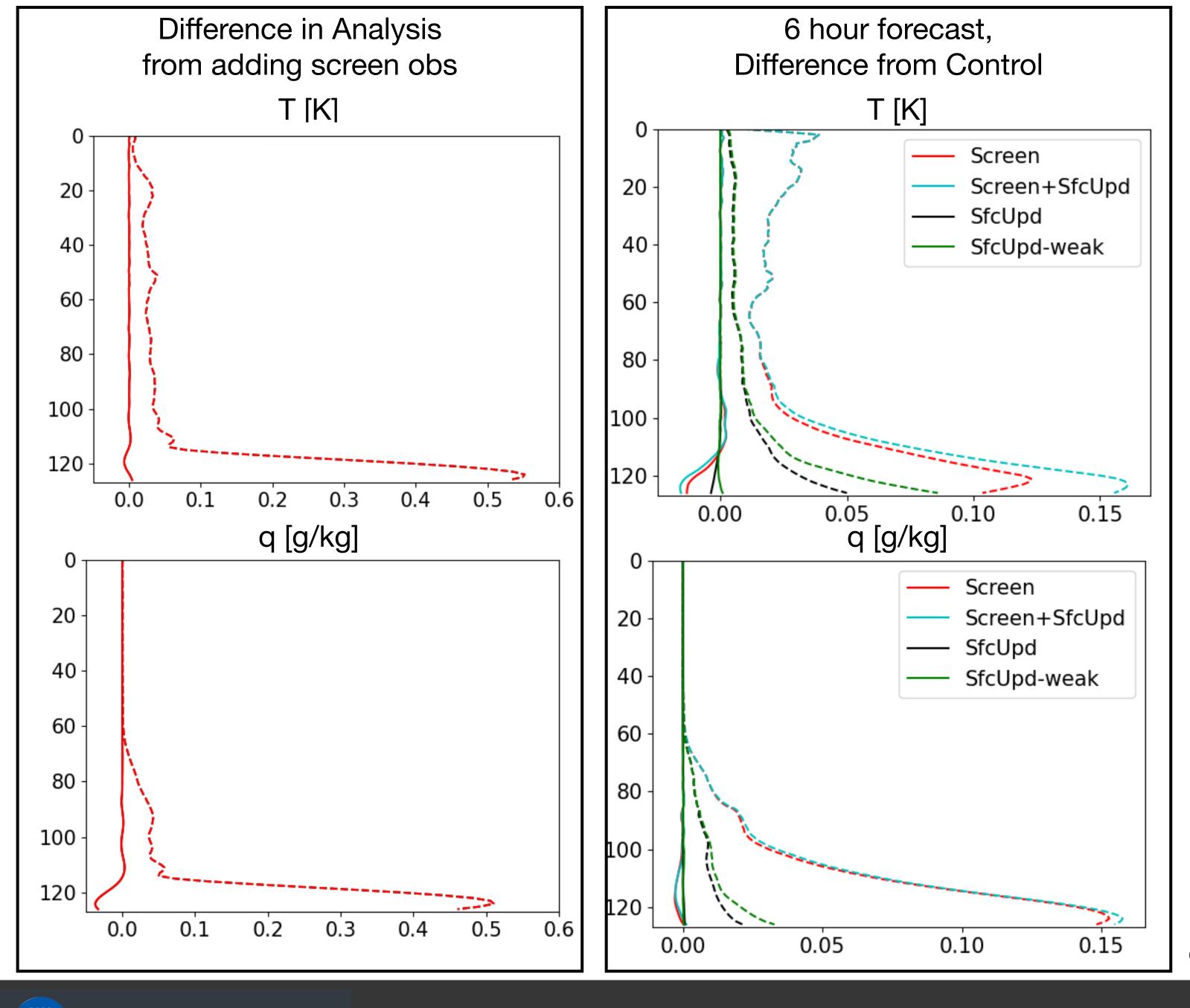












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Impact retained in subsequent forecast

- Plots show difference in first forecast, from the control experiment, then in subsequent 6 hour forecast
- Impact of increments is not well retained in the subsequent forecast
 - Model error
- Adding updates to the surface states increases impact on T forecasts

solid lines - means dashed lines - stdevs













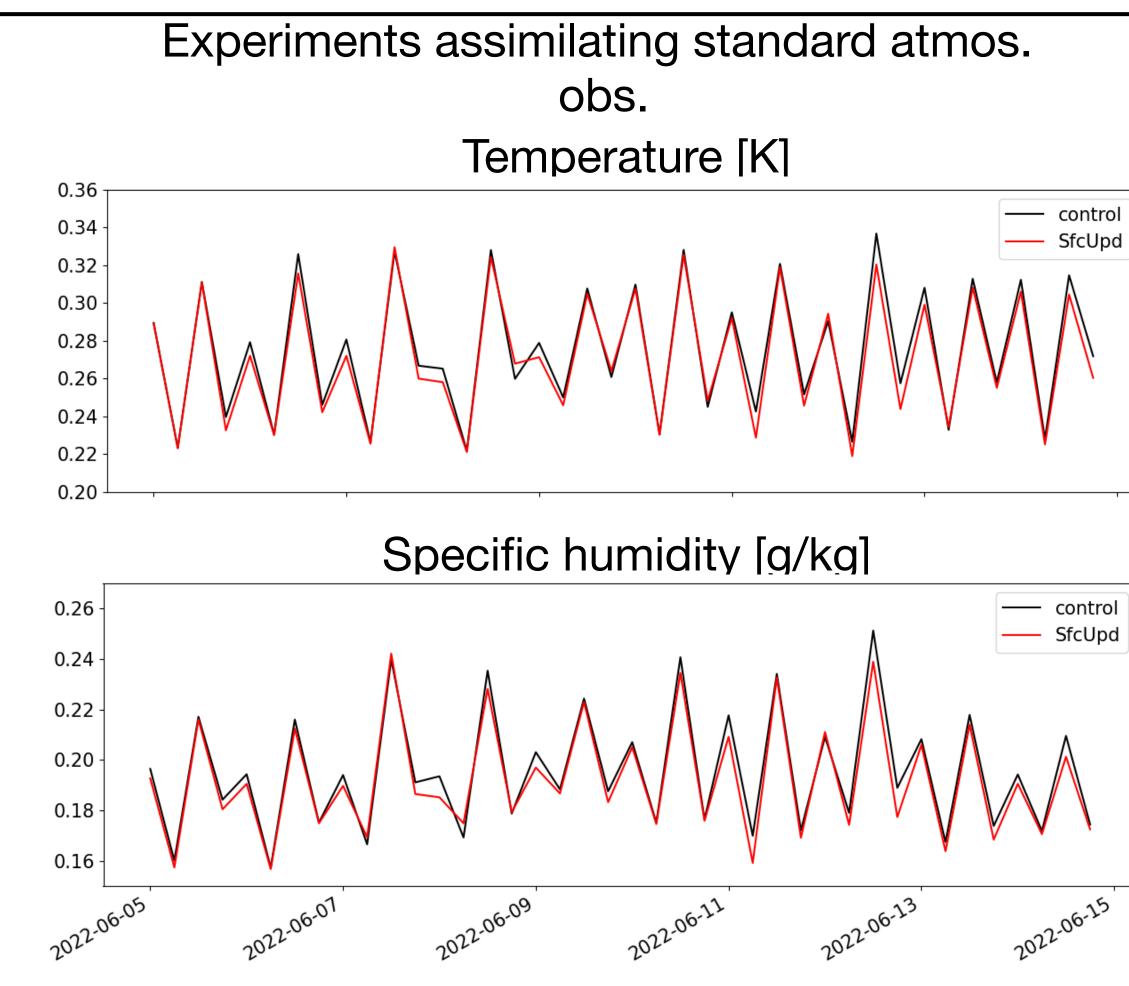




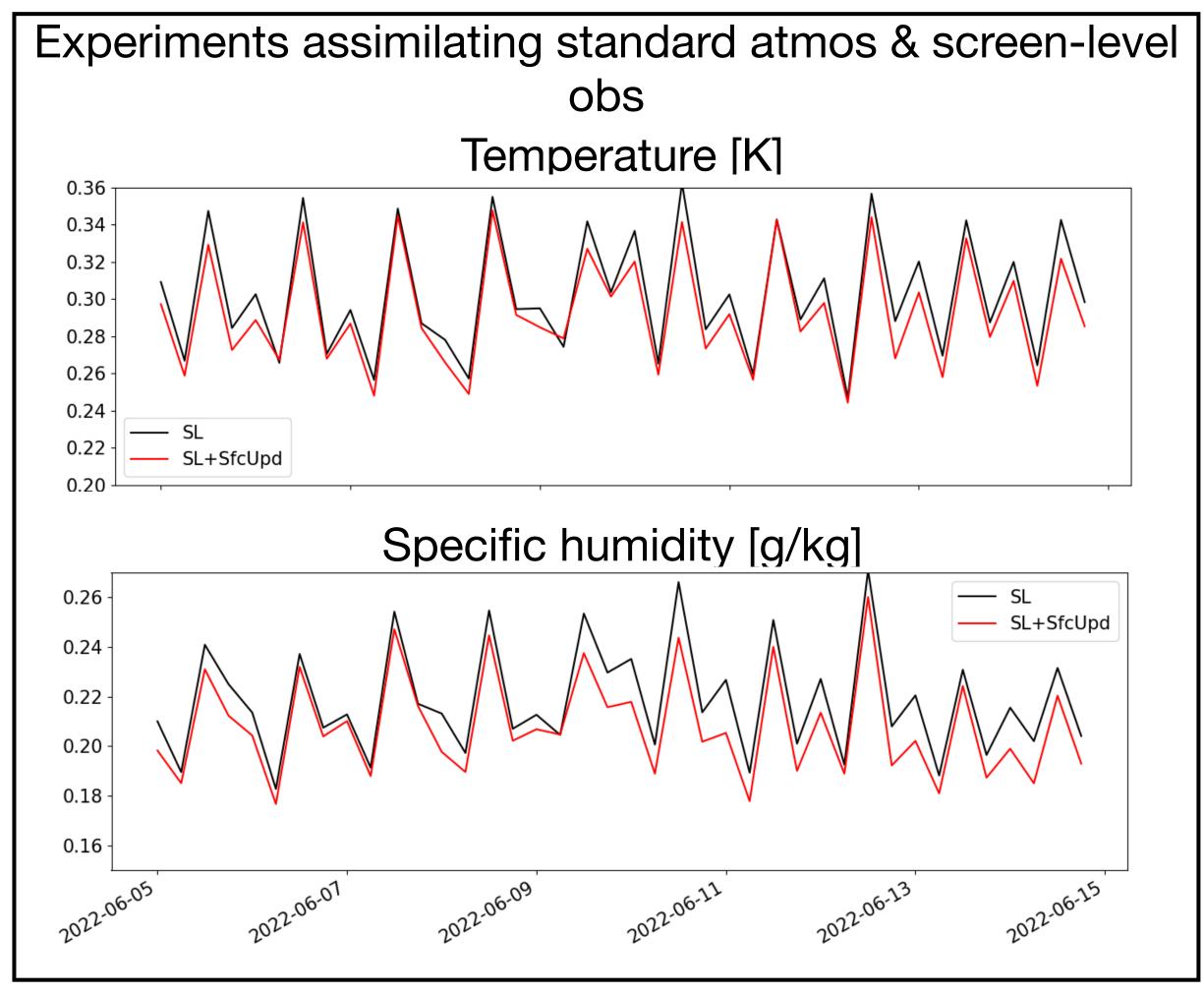
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Atmospheric Increment Timeseries

Time-series of sqrt(RMS increments in lowest 20 layers)



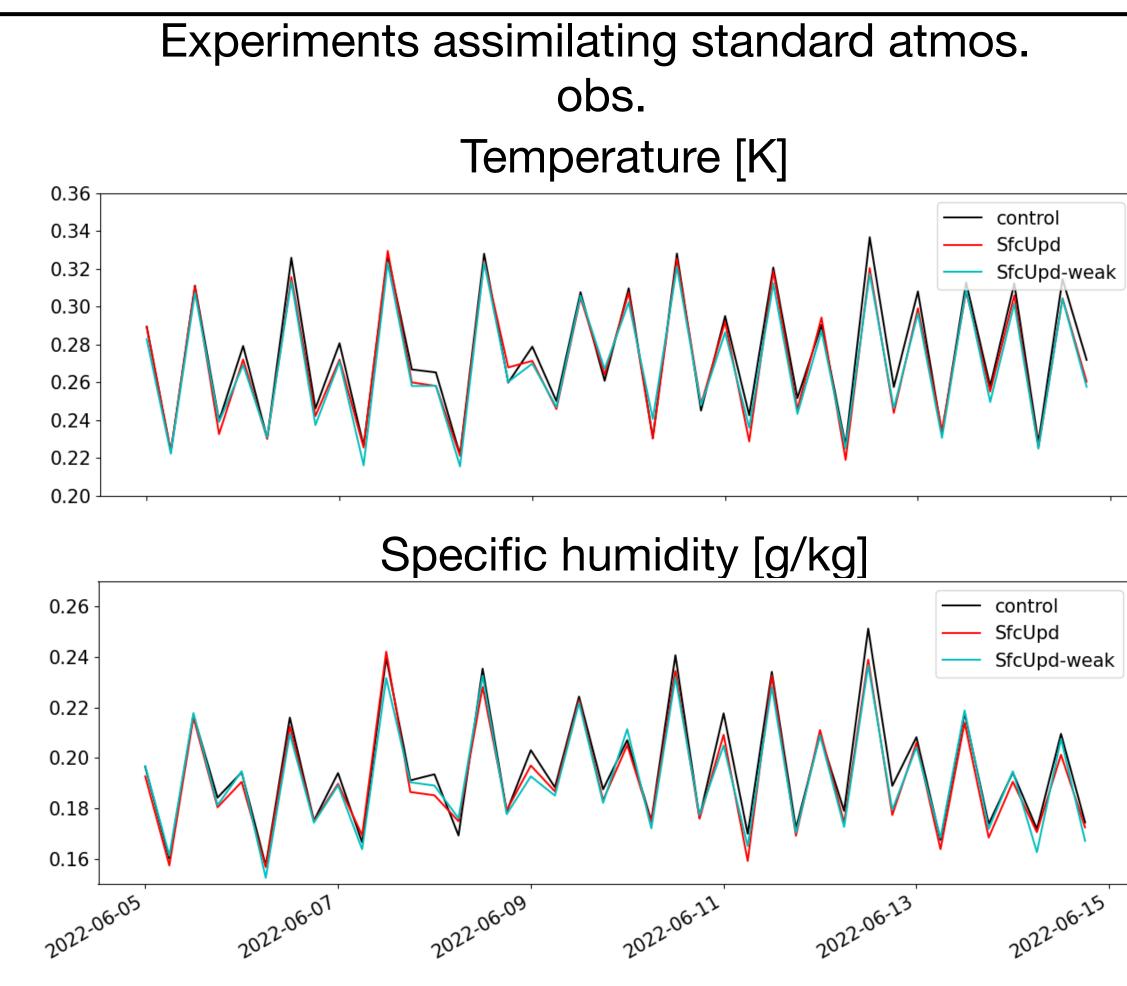




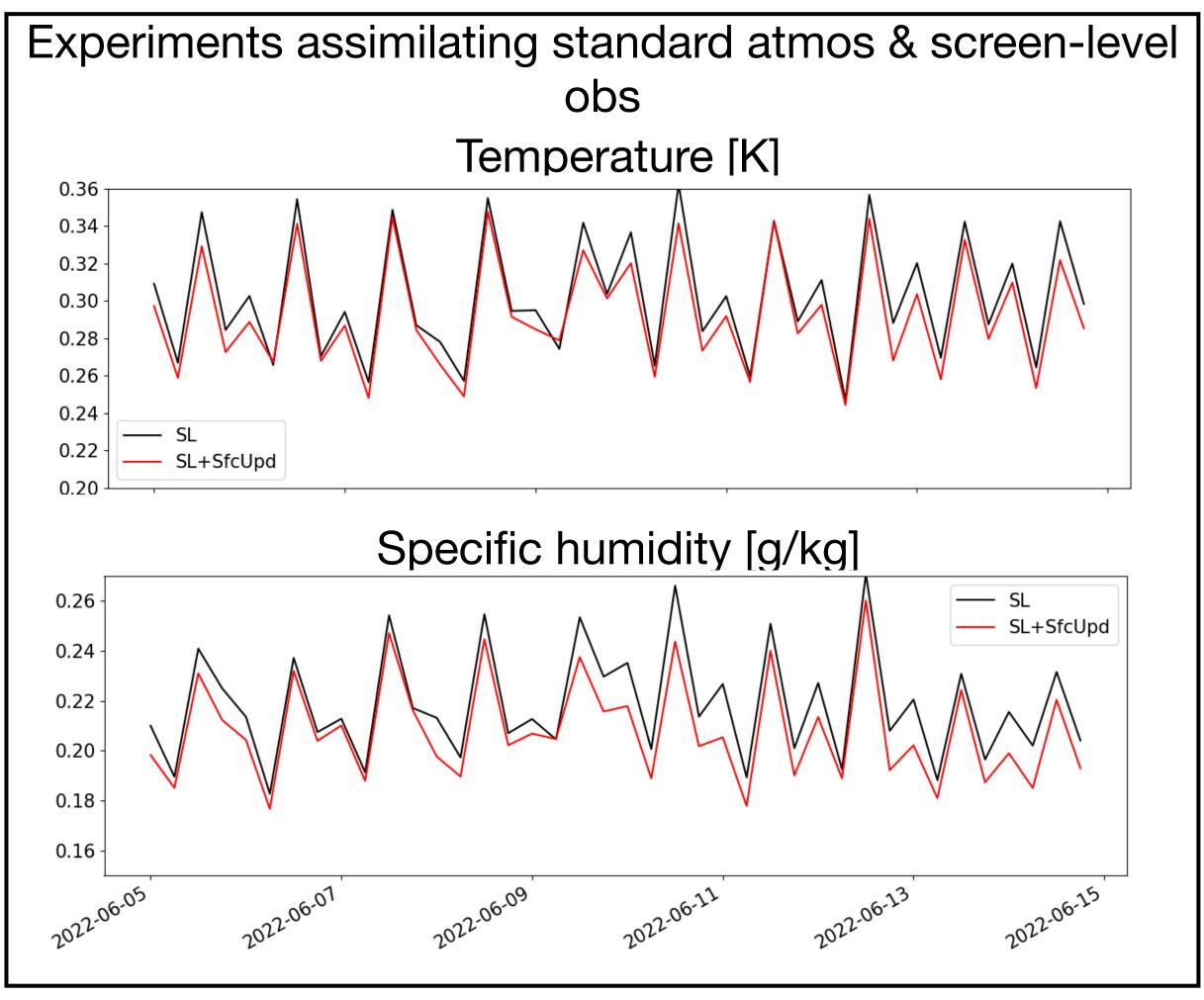


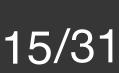
Atmospheric Increment Timeseries

Time-series of sqrt(RMS increments in lowest 20 layers)

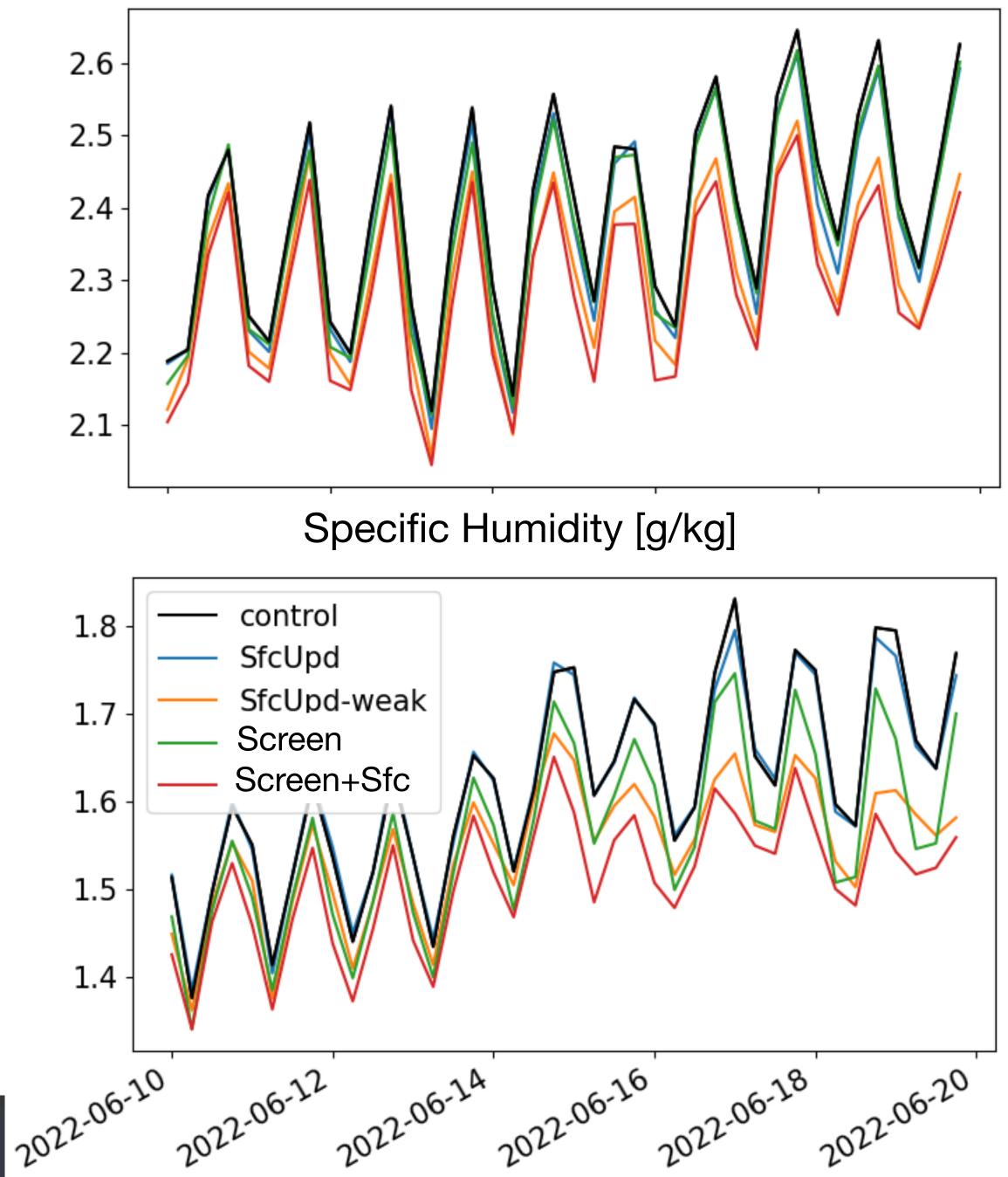








Temperature [K]



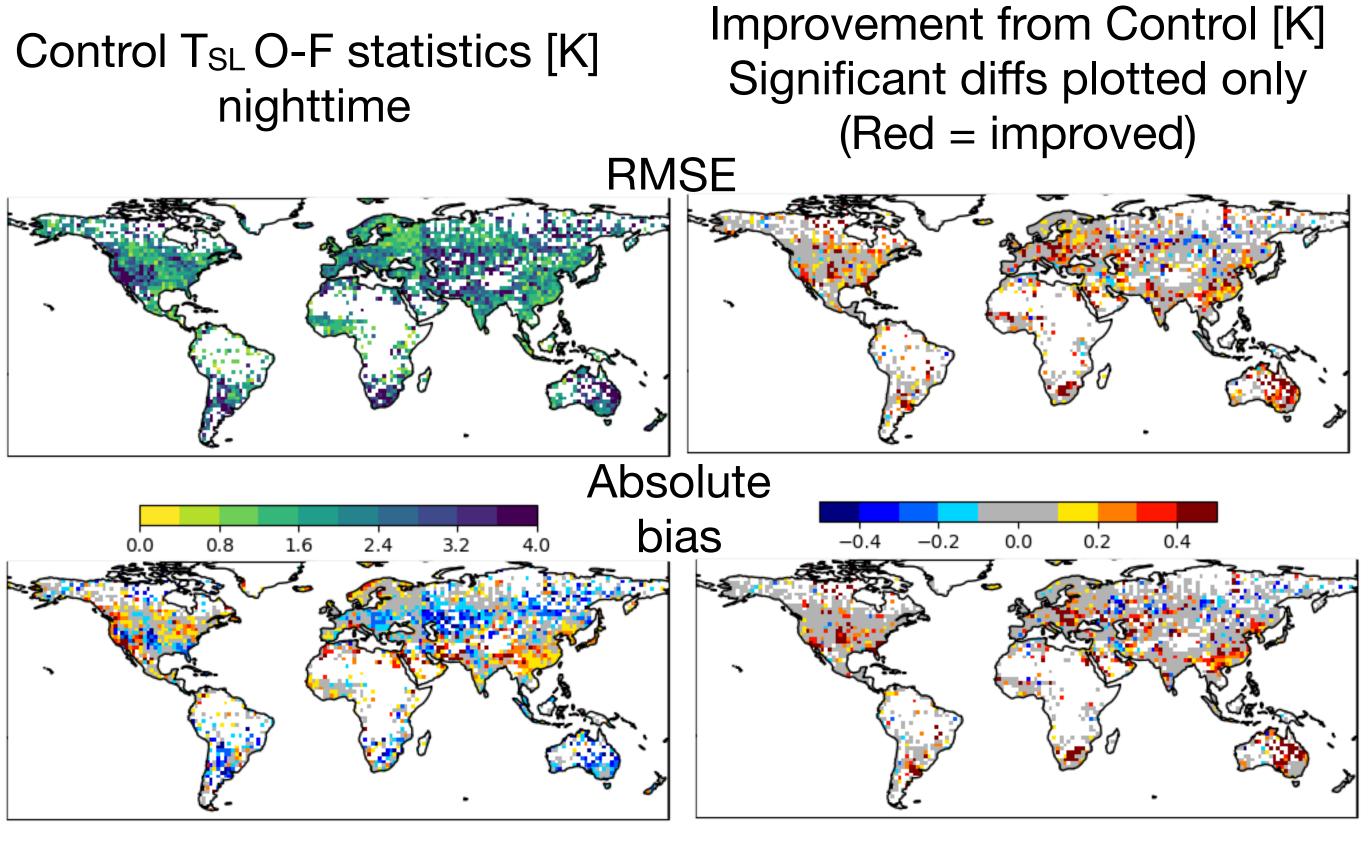
Screen-level sqrt(RMS O-F)

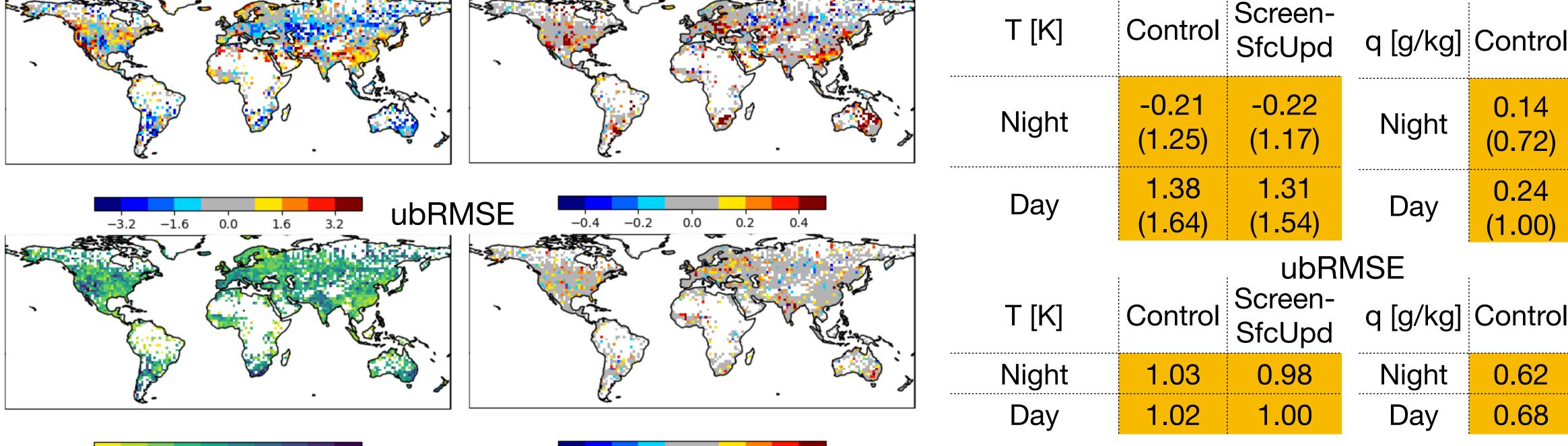
Mean RMSE			
	T [K]	q [g/kg]	
Control	2.39	1.62	
Screen	2.37	1.56	
SfcUpd	2.37	1.61	
Screen+SfcUpd	2.28	1.51	
SfcUpd - weak*	2.31	1.55	
shaded = sig. difference from Control			

- All experiments improve the O-F
- Best results from Screen+SfcUpd (3-4% reduction)
- Followed by SfcUpd-weak (screenlevel forecasts constrained more by updated surface than atmosphere)



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T [K]

Night

Day

Diurnal Screen-Level O-F statistics Screen-SfcUpd experiment

Screen-

SfcUpd

1.64

1.95

Control

1.75

2.05

RMSE

Bias (absolute)

q [g/kg] Control

Night

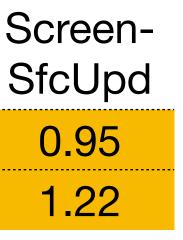
Day

1.02

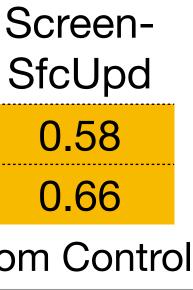
1.29

shaded = sig. difference from Control









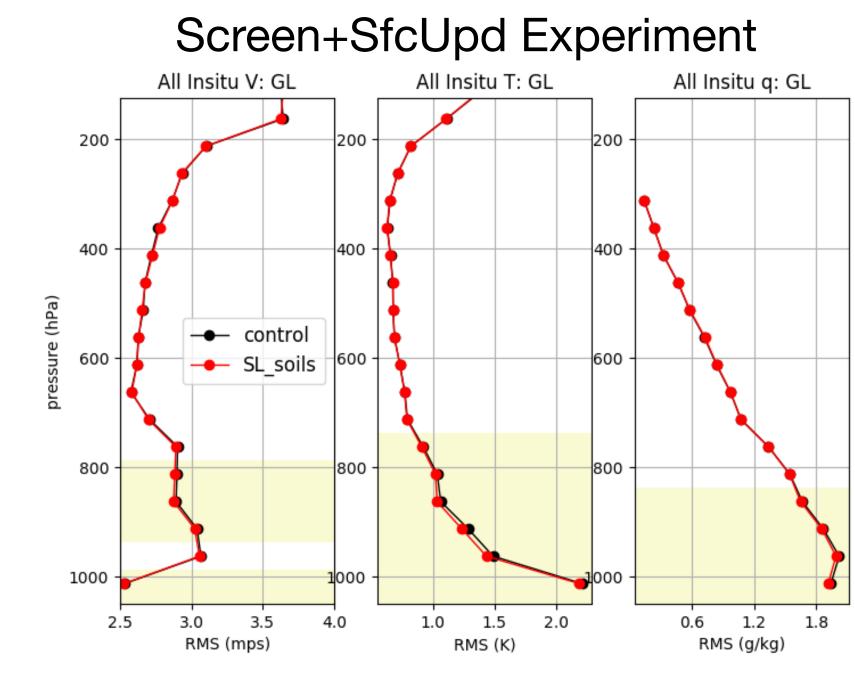


Sonde sqrt(RMS 0-F)

Mean RMSE at ~900 hPA			
	T [K]	q [g/kg]	
Control	1.29	1.67	
Screen			
SfcUpd			
Screen+SfcUpd	1.24	1.65	
SfcUpd - weak*			

 Screen+SfcUpd: Small, but consistent improvement (1-3%)



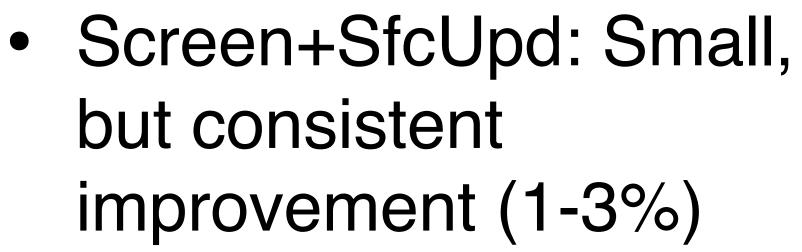


shaded = significant difference from Control

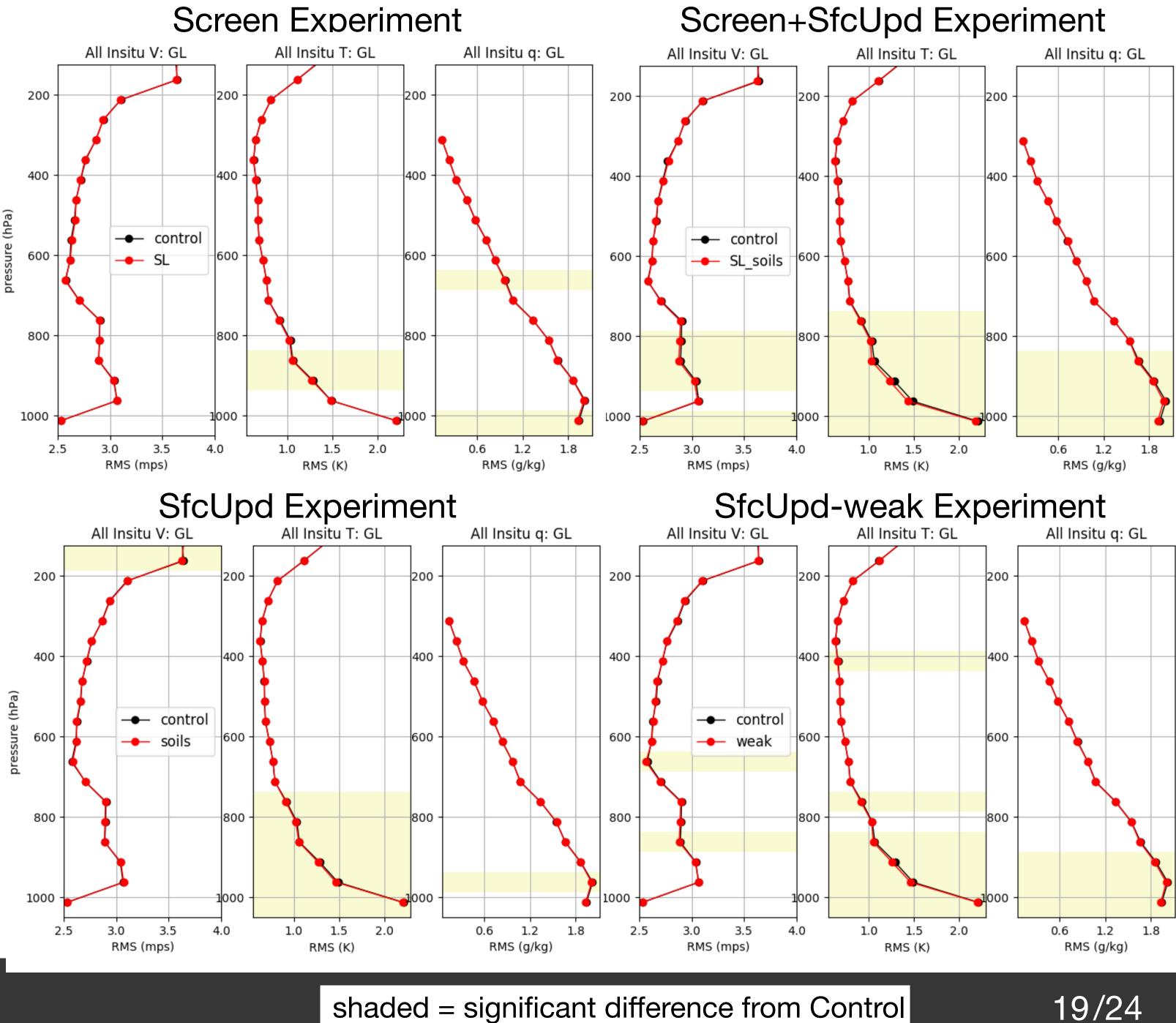


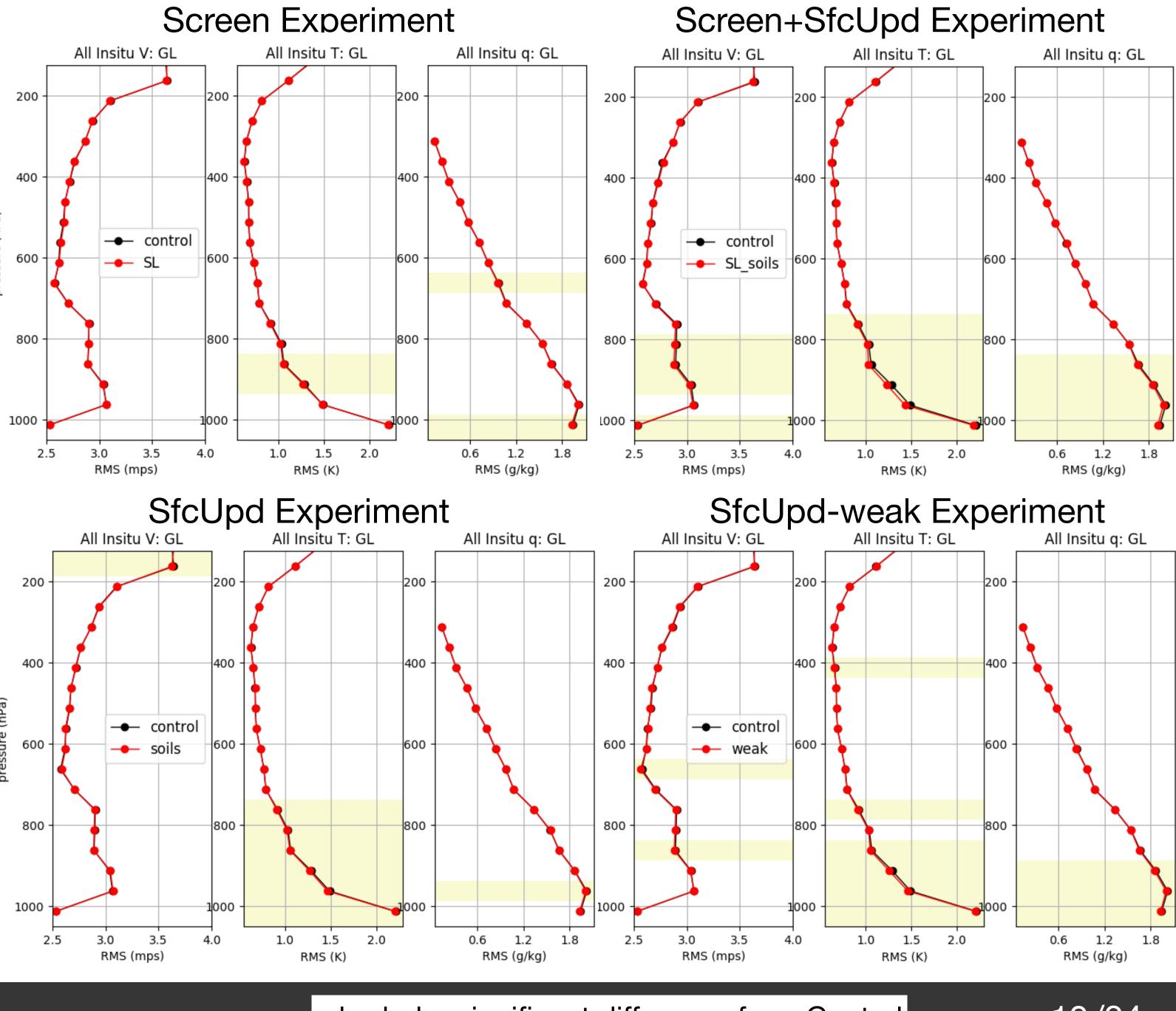
Sonde sqrt(RMS O-F) Mean RMSE at ~900 hPA

		q
	[K]	[g/kg]
Control	1.29	1.67
Screen	1.28	1.66
SfcUpd	1.27	1.65
Screen+SfcUpd	1.24	1.65
SfcUpd - weak*	1.25	1.66



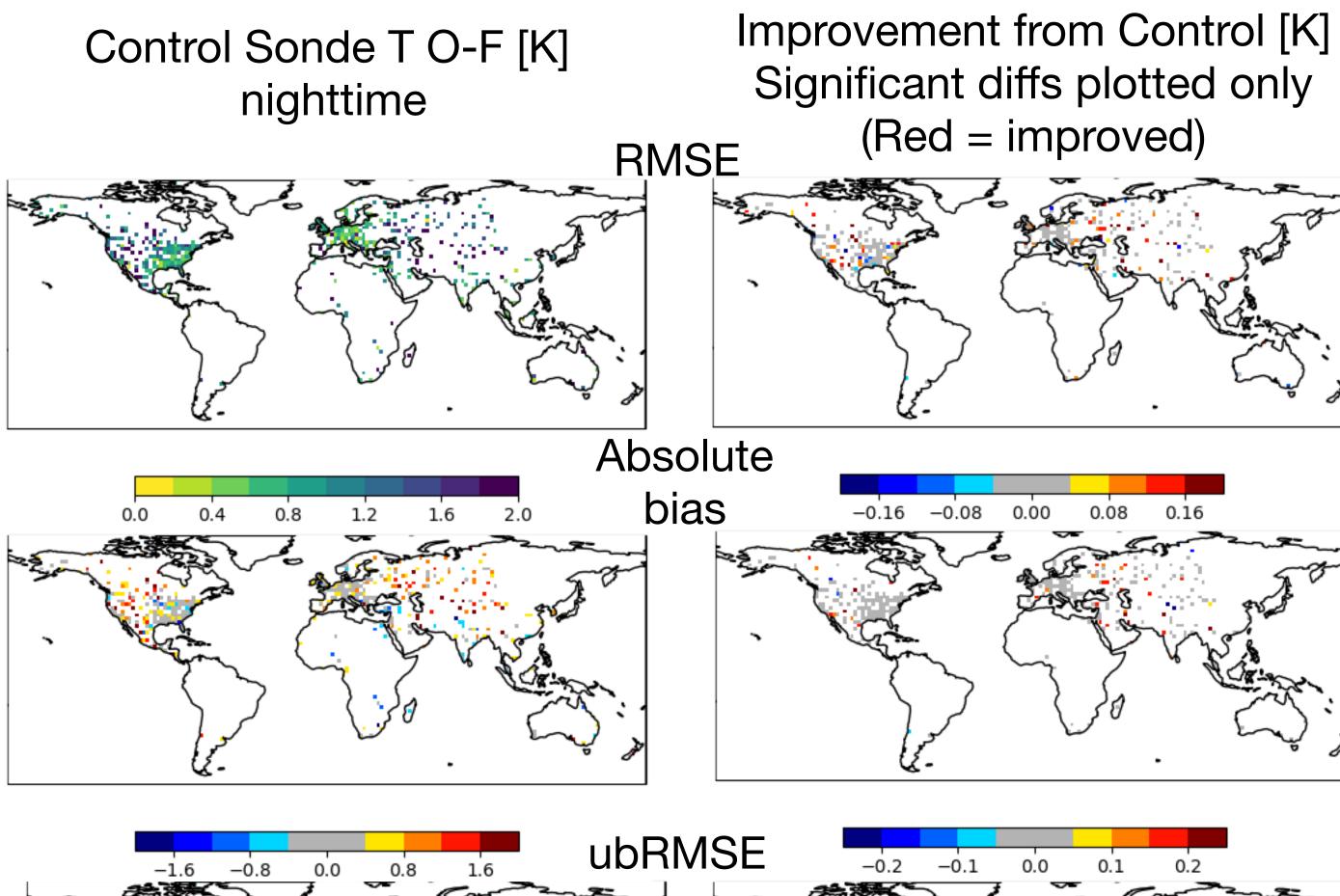
Coming from the surface update (with screen-obs)

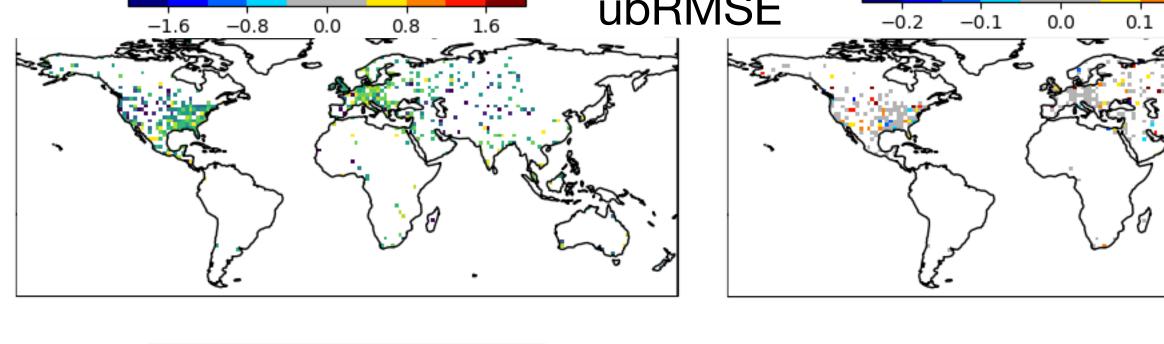






shaded = significant difference from Control







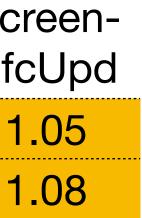


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Diurnal Sonde (1100-800 hPA) O-F statistics Screen-SfcUpd experiment

son and	RMSE					
Cherry Cherry	T [K]	Control	Screen- SfcUpd	q [g/kg]	Control	Screen- SfcUpd
	Night	0.97	0.94	Night	1.08	1.05
• •	Day	1.11	1.06	Day	1.10	1.08
0.16	Bias (absolute)					
VWS &	T [K]	Control	Screen- SfcUpd	q [g/kg]	Control	Screen- SfcUpd
	Night	0.37 (0.64)	0.35 (0.61)	Night	0.09 (0.59)	0.13 (0.71)
0.2	Day	0.49 (0.79)	0.44 (0.74)	Day	0.12 (0.57)	0.16 (0.69)
ST ST	ubRMSE					
V V S S S S S S S S S S S S S S S S S S	T [K]	Control	Screen- SfcUpd	q [g/kg]	Control	Screen- SfcUpd
	Night	0.59	0.58	Night	0.78	0.76
-	Day	0.62	0.61	Day	0.71	0.70
^{0.16} shaded = sig. difference from Control						









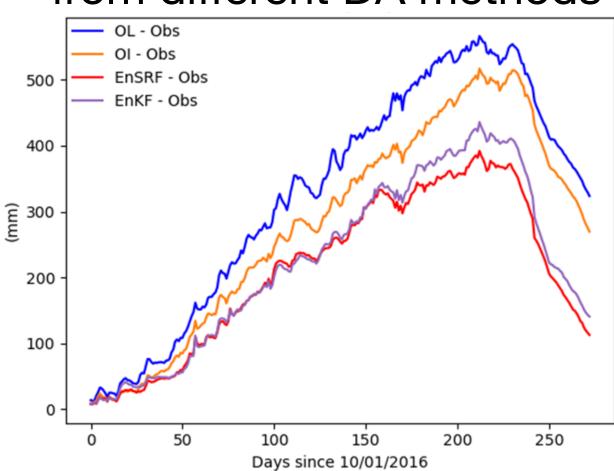


EnKF snow depth assimilation

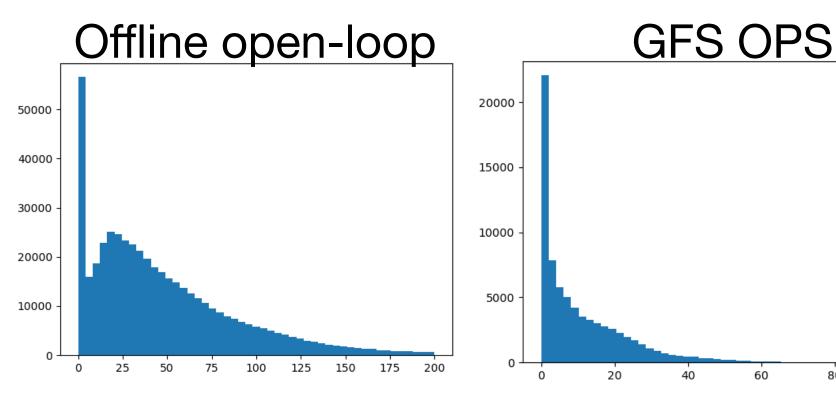
- NOAA is replacing our current snow depth assimilation with an OI-based scheme Offline (land-only) experiments show can
- get better performance, in terms of snow depth O-F, from EnKF than OI
- Also working towards unifying the snow depth DA with the atmos DA
 - Obtaining sufficient spread in the NWP ensemble may be difficult



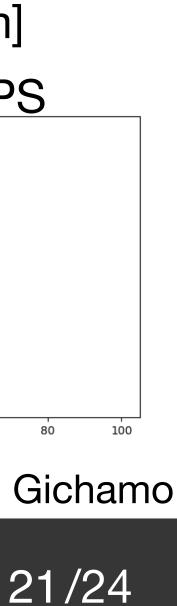
Snow depth sqrt(RMS O-F) [mm] from different DA methods



Ensemble stdev snow depth [mm]



With Tseganeh Gichamo



Conclusions (1/2)

- The GFSv17 soil analysis is being designed to use the same DA as for the atmosphere
 - Possible now, since atmospheric DA uses ensemble-based methods at/close to model resolution
- Tested different coupling options for assimilating screen-level obs and updating soil states (moisture, temperature)
 - than atmosphere
 - Also benefit to updating the land states (even without the screen-level obs)
 - obs)
 - Using this approach to develop the new soils analysis at NOAA
 - good

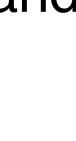


• Clear benefit to assimilating the screen-level obs, with more benefit from assimilation into land

• Greatest benefit from assimilating screen-level obs into both atmos and land using a single coupled update (reminder: not really assimilating land obs here; screen-level obs are interface

• Weakly coupled experiment (assimilate screen-level observations into surface only) nearly as

• Land and atmospheric dynamics are very different: weakly coupled approach more flexible











Conclusions (2/2)

- Next steps:
 - Check DA benefit holds with latest model version (reduced diurnal T bias)
 - Add land perturbation scheme
 - Test using full 4D-EnVar, rather than pure EnKF, for atmospheric update (and ultimately, the soil update too)
 - Hybrid may be more appropriate to land model problem
- Out-standing questions:
 - Do these results hold if assimilating true land obs (satellite soil moisture, snow depth, etc)?
 - Here, control experiment has no soil update. How would this approach compare to one of the established land DA methods?









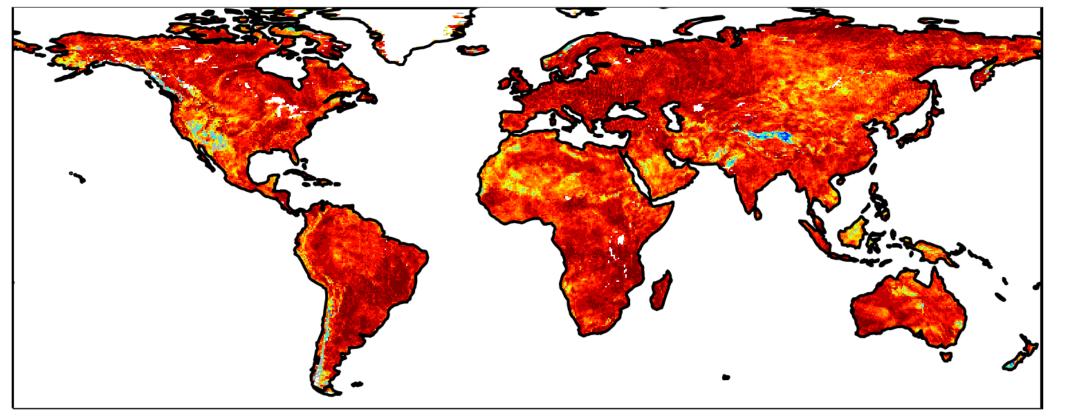
Thanks for Listening clara.draper@noaa.gov



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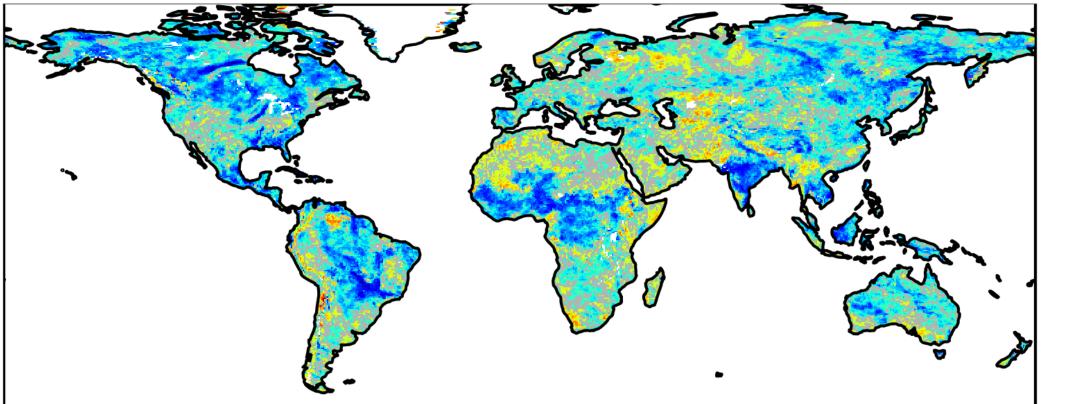
Vertical correlations for updating soil states

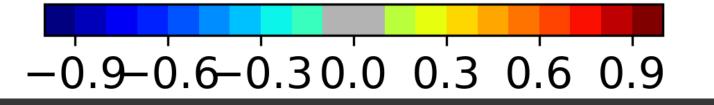
Ensemble correlation (T_{SL}, ST1)

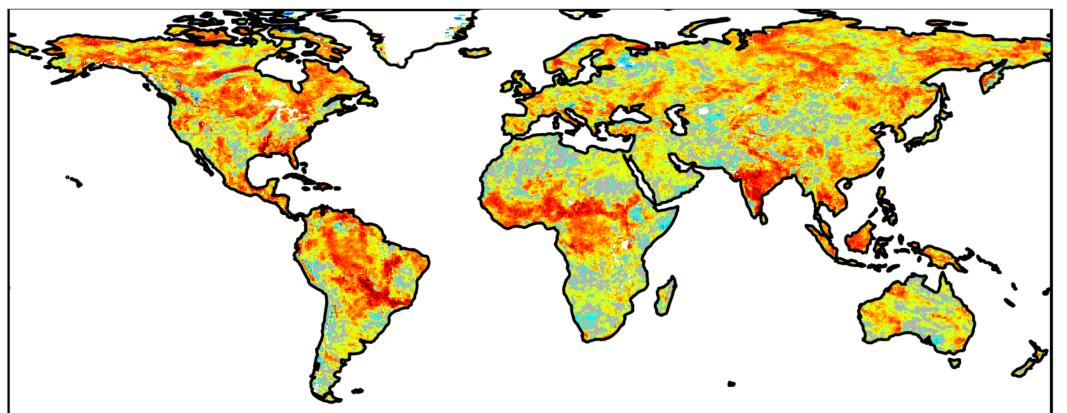


-0.9-0.6-0.30.0 0.3 0.6 0.9

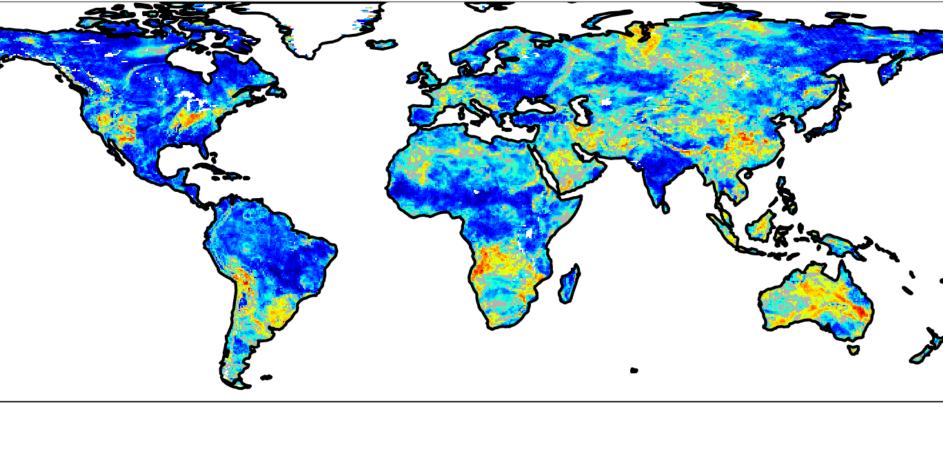
Ensemble correlation (T_{SL}, SM1)







Ensemble correlation (RH_{SL}, ST1)



-0.9-0.6-0.30.0 0.3 0.6 0.9

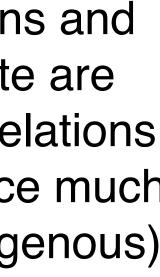
Ensemble correlation (RH_{SL}, SM1)



•Soil temperature: strong correlation with T_{SL}, often with RH_{SL}

- •Soil moisture: correlations strong in some regions; smaller / noisy in other regions
- •Note: GSI humidity observations and control state are RH (q correlations near surface much less homogenous)

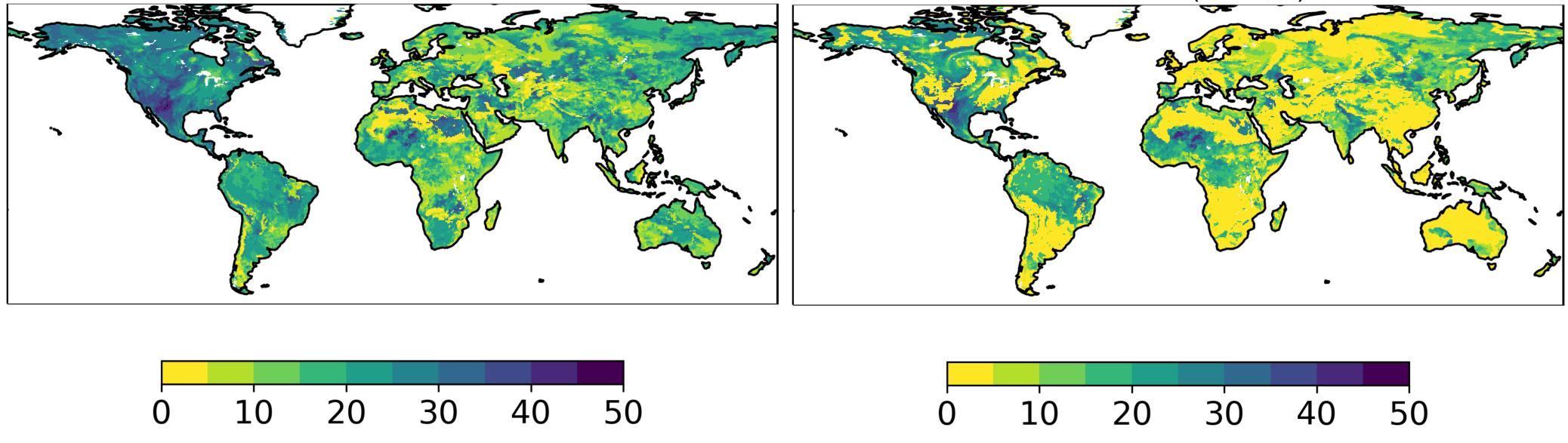


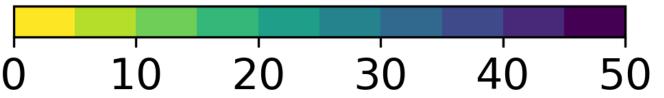




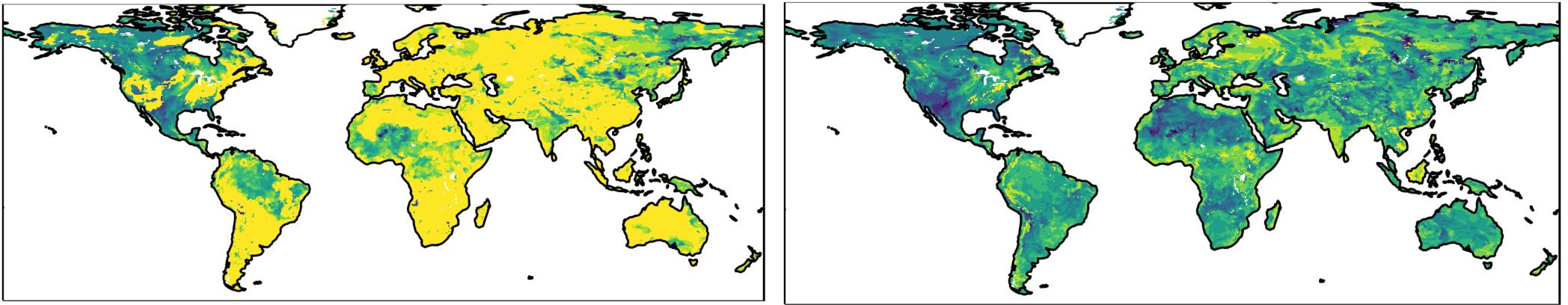
Vertical correlations for updating atmospheric states

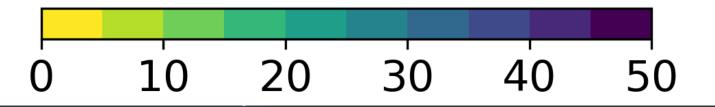
Level at which correlation (T_{SL}, T) falls below 0.5





Level at which correlation (T_{SL}, RH) falls below 0.5



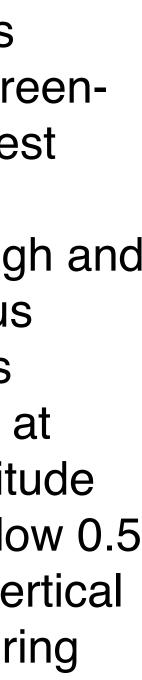


Level at which correlation (RH_{SL}, T) falls below 0.5

Level at which correlation (RH_{SL}, RH) falls below 0.5

30 10 20 40 50

- Correlations between screenlevel at lowest model level generally high and homogenous
- Plots shows model level at each magnitude reduces below 0.5
- Strongest vertical profile is during night

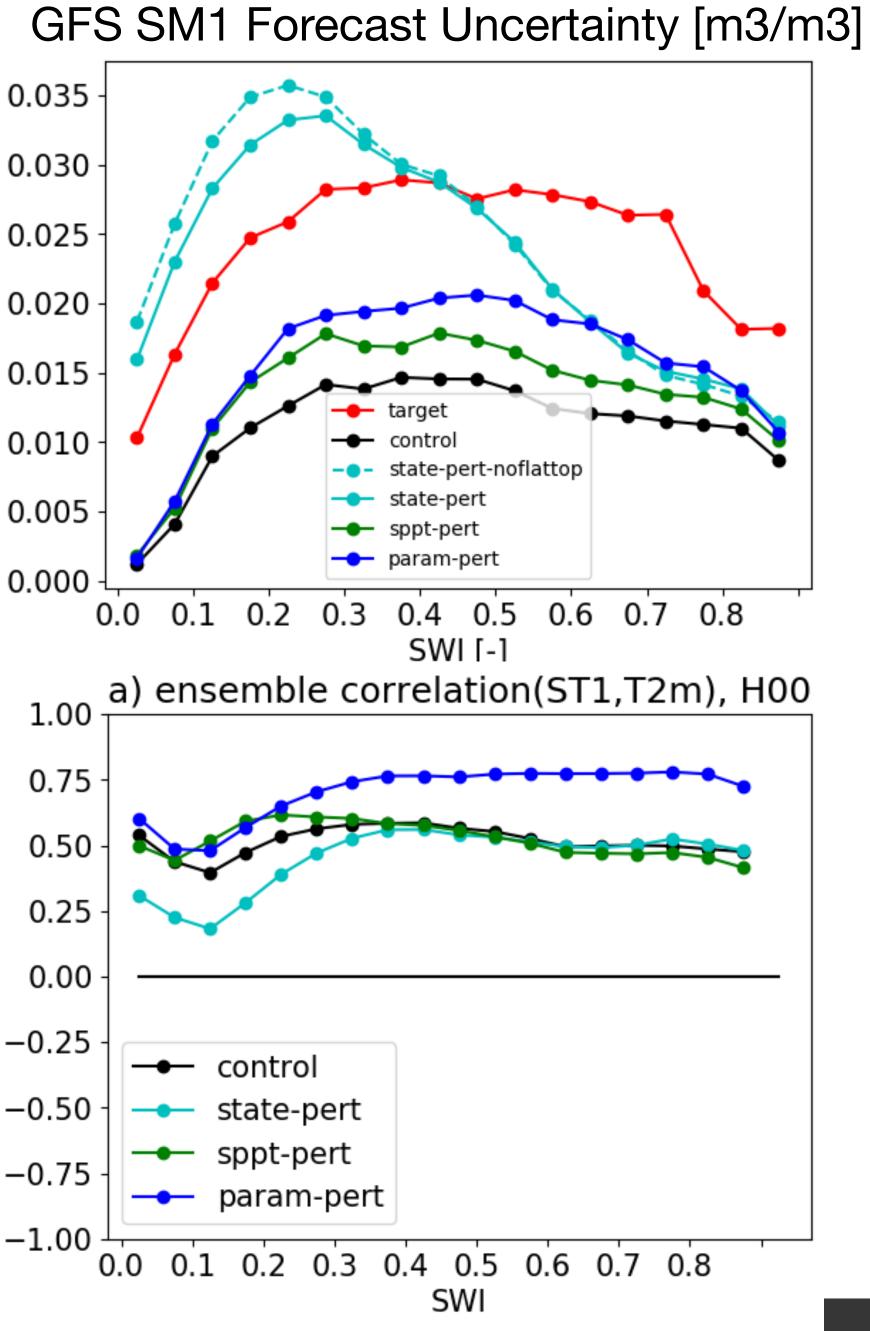




Accounting for land model error in **NWP ensembles**

- No information gained on model error growth / instability by adding perturbations to the soil moisture states
 - Resulting ensemble spread function of state perturbations added and local model persistence
- SPPT not well suited to soil moisture
- In a coupled data assimilation system applying perturbations to one component only will gives ensembles with higher cross-component covariances where that component is driving the coupling, and lower covariances where the other component is driving the coupling
- Recommended method to account for land model error in NWP ensembles is to perturb key parameters controlling the land/atmosphere fluxes (in these experiments, vegetation fraction)
 - Generates reasonable spatial patterns in ensemble spread
 - Generates ensemble cross-covariances more representative of errors in land/ atmosphere coupled model
- Caveat: Land is highly non-linear; difficult to obtain sufficient spread to represent forecast uncertainty without inducing large changes in ensemble mean (impractical)





Adding Land Model Uncertainty

- Test methods drawn from atmospheric and land ensemble DA communities:
 - State-pert: Stochastically perturb the soil moisture content (SMC) and soil temperature content (STC) at • each time step (standard approach used in land-only ensemble data assimilation systems)
 - <u>SPPT-pert</u>: Apply stochastically perturbed physics tendencies (SPPT) scheme to SMC and STC ulletMotivation: use model physics to provide relationship between SM and ST deltas
 - <u>Param-Pert</u>: Stochastically perturb key model parameters controlling the land /atmosphere fluxes (here: ulletvegetation fraction) Motivation: physically consistent perturbations in the land and atmosphere
- Tested each in a suite of data assimilation experiments:
 - 30 member ensemble at ~0.5 degrees (C192), run 30 days from July 10, 2019 •
 - Atmospheric data assimilation is cycled every 6 hours, using hybrid 3DEnVar DA •
 - Assimilating the standard atmospheric obs, using standard atmospheric stochastic physics

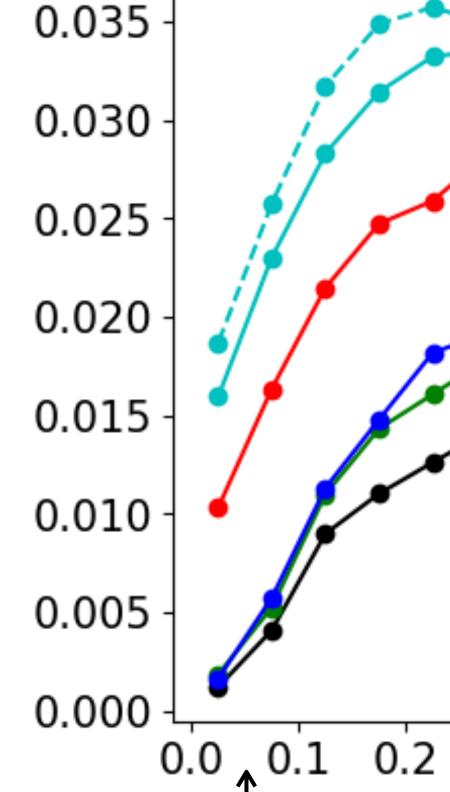






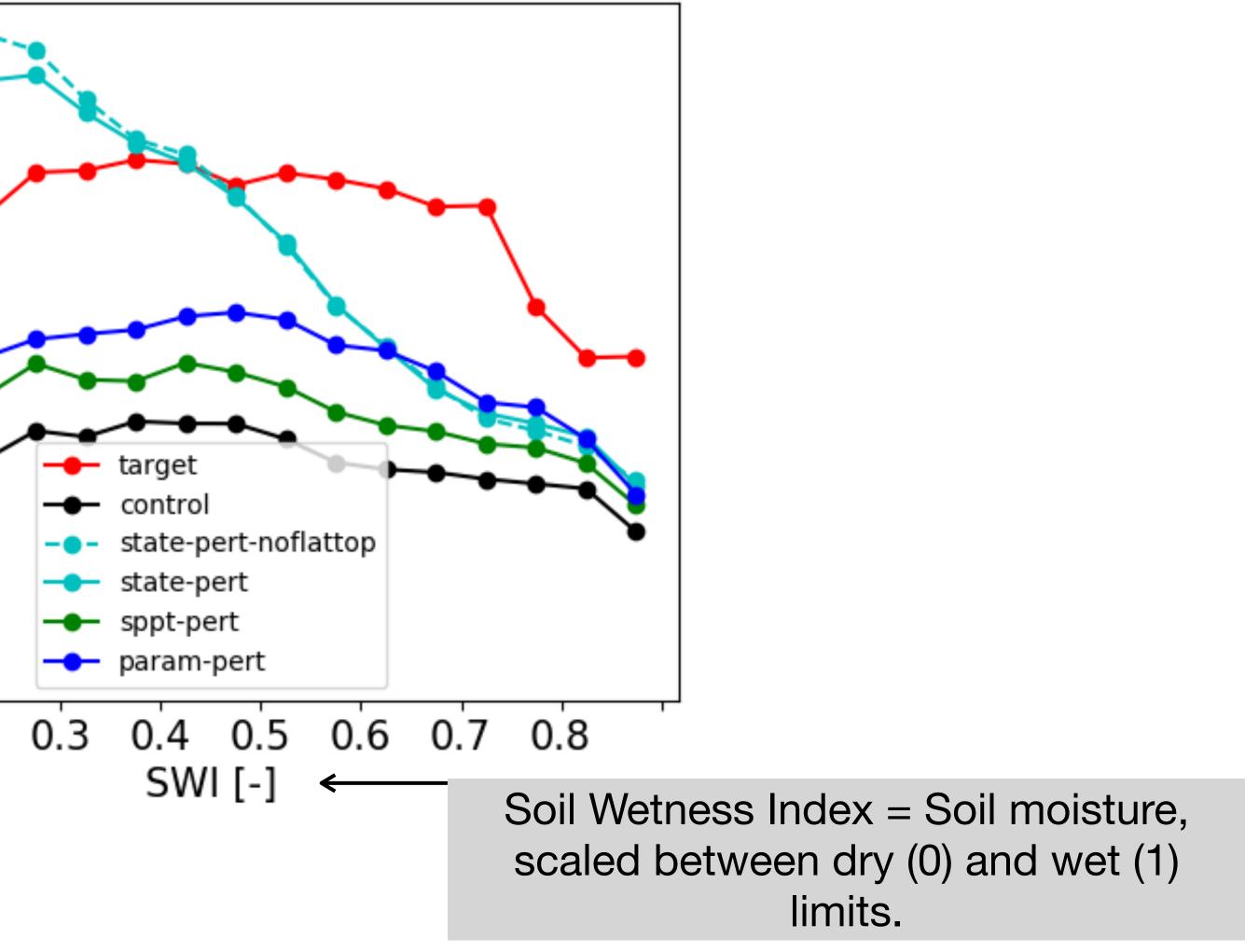
Ens. Spread in Soil Moisture Layer 1 (SMC1)





Target (red) is best estimate of forecast error standard deviation (c.f, independent obs). Others are ensemble-based estimates from each experiment.

GFS SM1 Forecast Uncertainty [m3/m3]



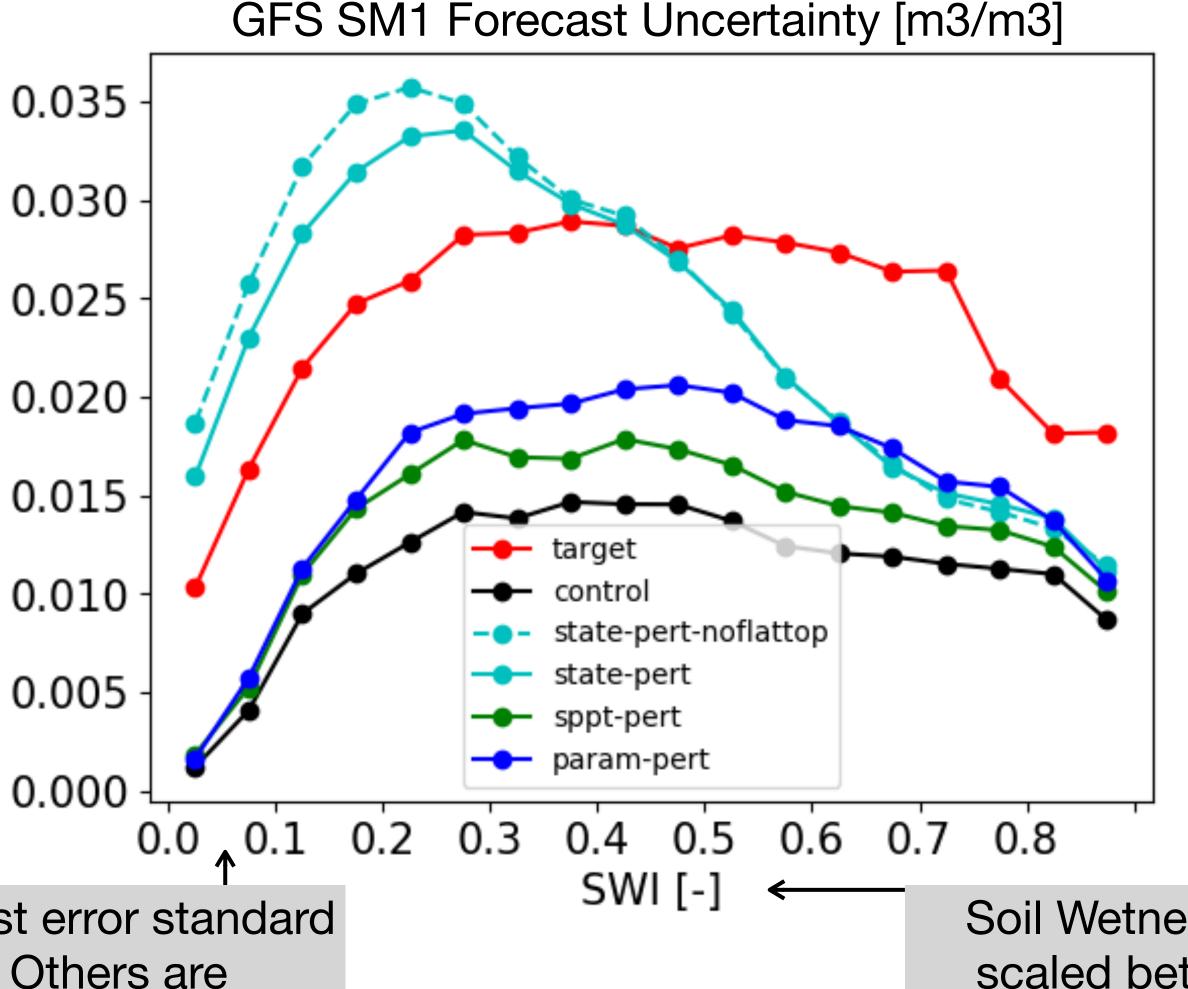






Ens. Spread in Soil Moisture Layer 1 (SMC1)

- State-pert induces too much spread in dry regions. Due to soil moisture memory being longer in dry conditions.
- SPPT-pert can induce only a small amount of spread. Inherent limitation of the method.



Target (red) is best estimate of forecast error standard deviation (c.f, independent obs). Others are ensemble-based estimates from each experiment.

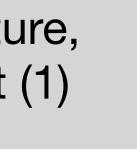
Param-pert looks reasonable. Spread could be inflated by perturbing additional variables.

Soil Wetness Index = Soil moisture, scaled between dry (0) and wet (1) limits.







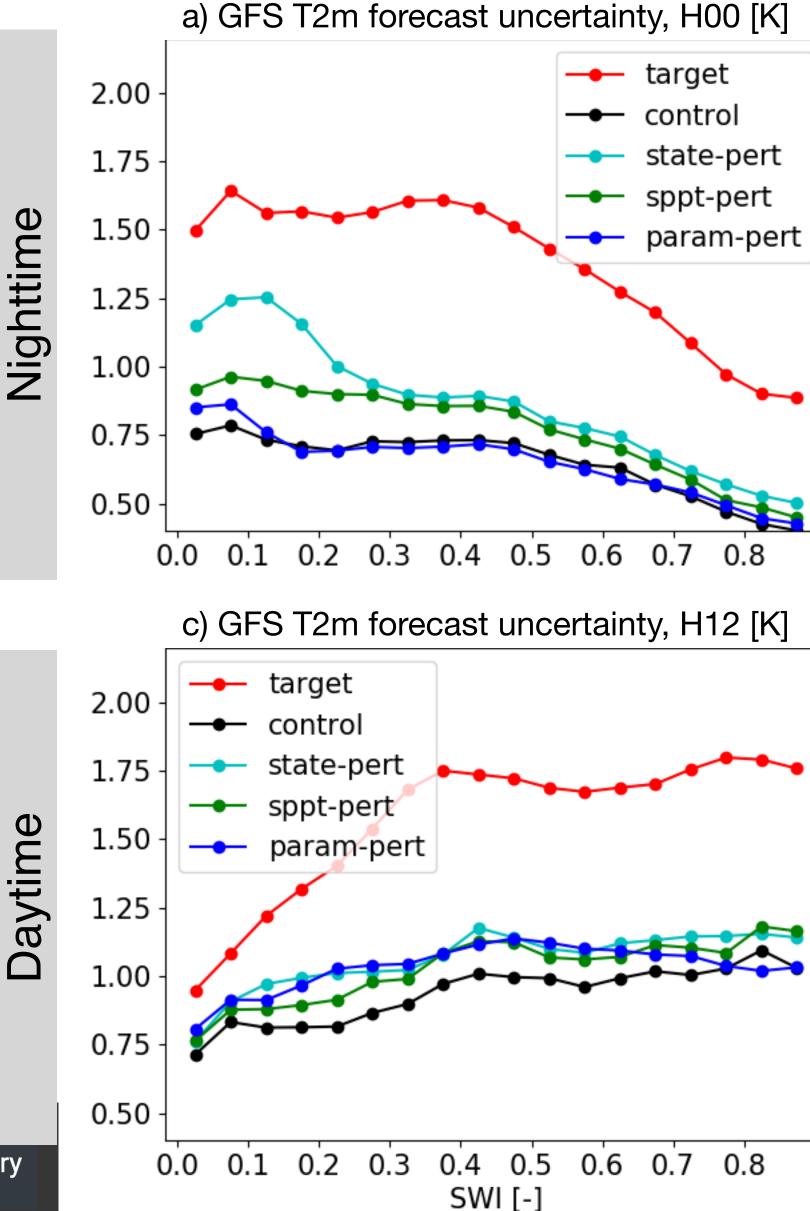




Ens. Spread in 2m Temperature and Specific Humidity

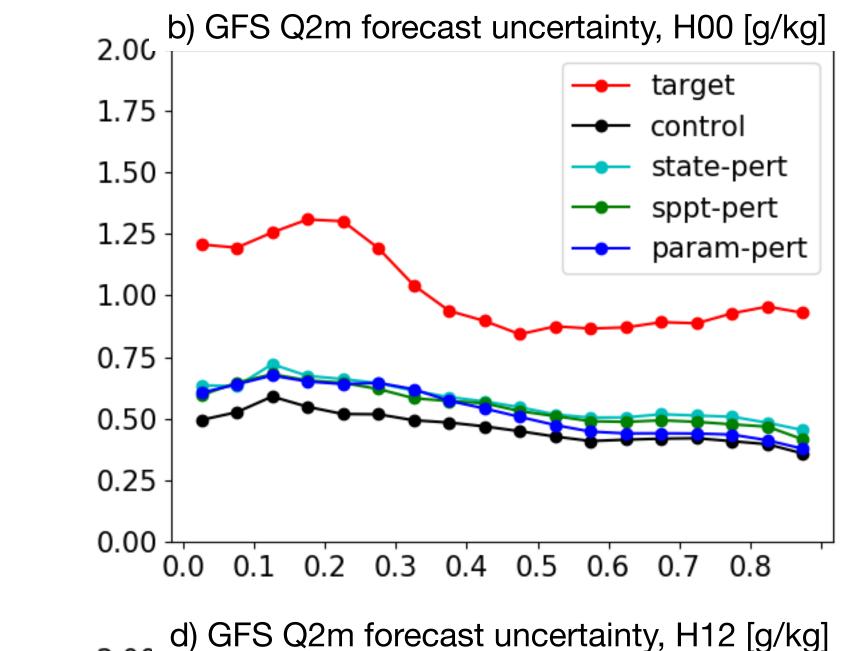
2m Temperature

Induced spread is generally limited in all experiments



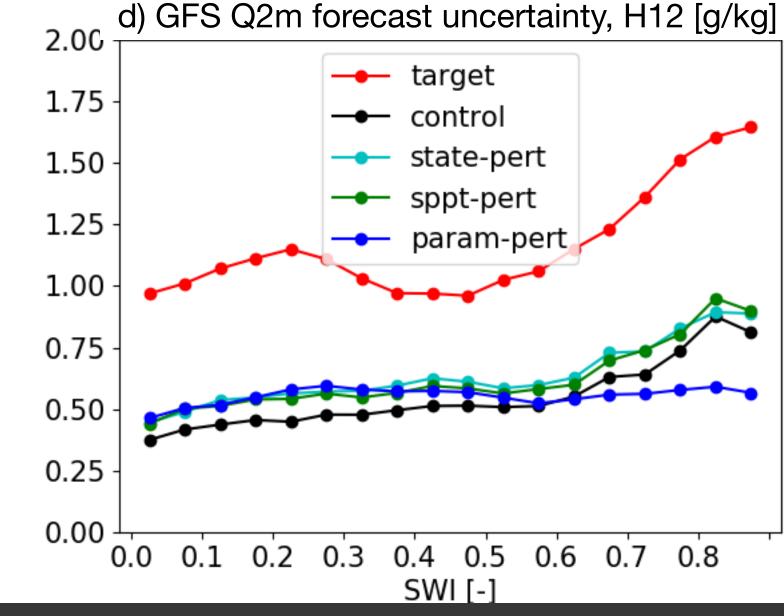
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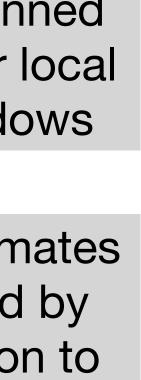


Results binned into 6 hour local time windows

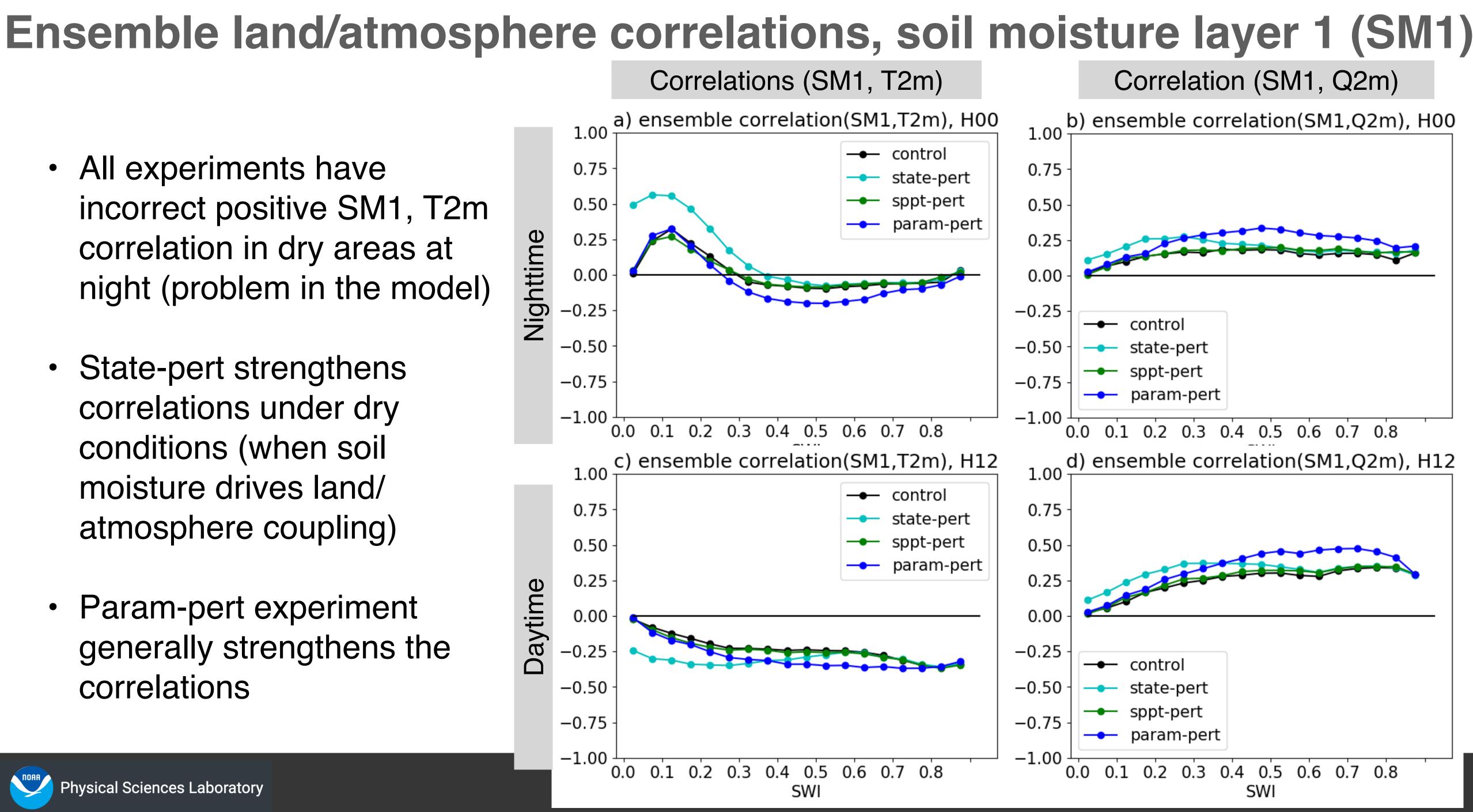
Target estimates calculated by comparison to ERA-5 analysis.







- All experiments have • incorrect positive SM1, T2m correlation in dry areas at night (problem in the model)
- State-pert strengthens correlations under dry conditions (when soil moisture drives land/ atmosphere coupling)
- Param-pert experiment generally strengthens the correlations





- State-pert weakens the ST1, T2m correlations (atmosphere is driving) the land/atmosphere coupling)
- Param-pert experiment again generally strengthens the correlations

