

# Lake model in OpenIFS

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and Souhail Boussetta



Caversham Lake

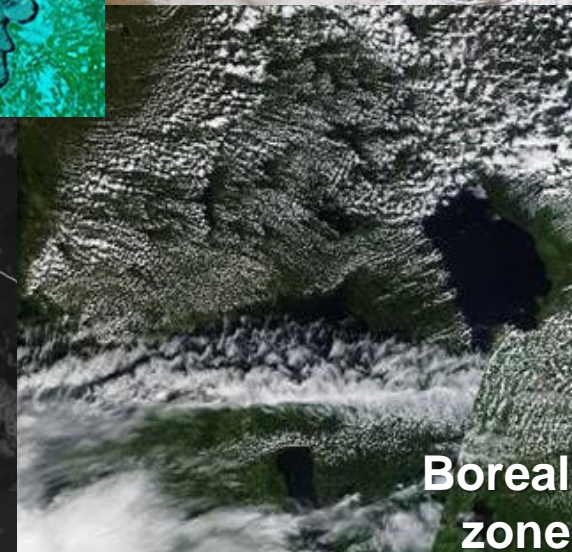
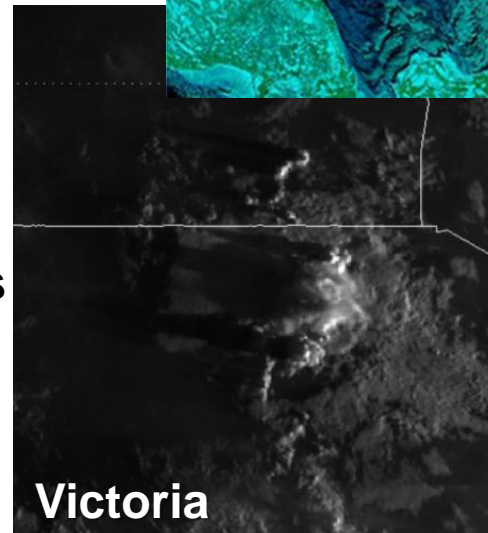
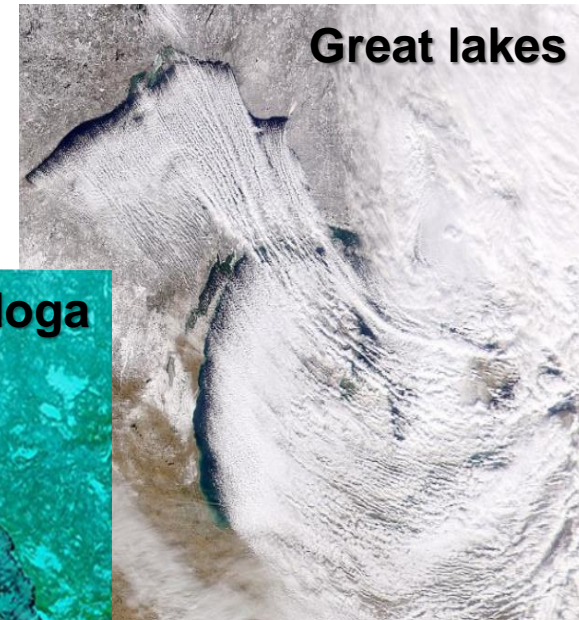
Photo from <https://en.wikipedia.org>

# Introduction: lakes & their numbers

- **Lake** – a significant **volume of water**, which occupies a depression in the land and has **no direct connection with the sea**.
- Inland water bodies are often referred to as **lakes** when the **lateral movement** of the water is **negligible**, and as **rivers** when there is a **sizeable lateral transport**. Here, term “lakes” is used in the broad sense of **any inland water body** which lateral movement of water is neglected (i.e. lakes, reservoirs, rivers and coastal waters), although a clear separation is often complex and varies in time.
- **Globally lakes** occupy about **3.7 %** of the land surface (Borre, 2014; Verpoorter et al., 2014).
- According to the latest calculations the **total number** of lakes with a water surface area not less than  $0.002 \text{ km}^2$ , is **117 million** (excluding Greenland and Antarctica), and their **combined area** is about **5 million  $\text{km}^2$**  (excluding the Caspian Sea) (Borre, 2014; Verpoorter et al., 2014).
- **Lakes are distributed very unevenly**. **Most** lakes are situated in **Boreal and Arctic climate zones  $45\text{-}75^\circ\text{N}$**  (Borre, 2014), namely in Canada, the Scandinavian Peninsula, Finland and Northern Russia and Siberia.

# Lakes influence local weather climate

- *During freezing/melting*: lake surface radiative and conductive properties and latent and sensible heat released to the atmosphere changes dramatically → **different surface energy balance**.
- *Great lakes (USA)*: **intensify** winter **snowstorms**.
- *Lake Ladoga (Russia)*: can **generate low clouds** → **increase in  $T_{2m}$  up to 10 °C** in neighbouring **Finland**.
- *Boreal zone (50-70 °N)*: usually cause a **decrease of summer precipitation**.
- *Lake Victoria (Africa)*: **generates** night convection with intensive **thunderstorms** → **death of thousands of fisherman every year**.





# Lakes can have global influence

- **Lakes** can also influence **global climate** by affecting the **carbon cycle** (Tranvik et al., 2009) and **methane CH<sub>4</sub> emissions** (Stepanenko et al., 2016).
- **Small shallow thermokarst lakes** located at Boreal and Arctic latitudes in the permafrost thaw area are rich in nutrients, which affect the **carbon dioxide CO<sub>2</sub> budget** (Walter et al., 2006; Walter et al., 2007; Stepanenko et al., 2012).
- **Small shallow** type of **lake** is most common (representing **~77 %** of the lakes **globally**), in general has a small surface area (0.002-0.01 km<sup>2</sup>) and a **big surface-to-volume ratio** →
- important as carbon dioxide **CO<sub>2</sub>** and methane **CH<sub>4</sub> degassing** takes place **through** the lake's **surface** (Borre, 2014; Verpoorter et al., 2014).



# Lake parameterization in NWP

- The **effect of lakes** is handled in Numerical Weather Prediction (NWP) and climate models through **parameterization**, which **needs** information on the **locations** of the lakes and their **morphological characteristics**.
- In **global models** lakes are necessary to correctly represent **surface boundary conditions** in NWP systems (forecast + data assimilation):
  - albedo (e.g. freezing lakes);
  - thermal capacity (linked to water mass);
  - evaporation (linked to water temperature, wind, surface roughness);
  - soil moisture realism.
- The **lake schemes** for NWP need to be **simple & fast**. Constraints:
  - **reduced set of equations** that can represent lake daily/seasonal/inter-annual variability;
  - **simplified hydrology**, as lake full hydrology is too sophisticated to be fully represented in current global climate models at 9-200 km resolution;
  - **reduced set of parameters** and variables that can be initialized also on global domain (e.g. by satellite observations);
  - **numerically stable** and **physically sound** at all latitudes.
- The **main target** of lake parameterization in NWP is to **reproduce** lake **surface water temperatures** (and therefore evaporation rates).

# ECMWF surface model status (ERA5 setup) and its evolution since ERA-Interim surface scheme

Implementation dates					
2007/11	2009/03-09	2010/11	2011/11	2012/06	2015/05
<p><b>Hydrology-TESEL</b></p> <p>Balsamo et al. (2009) Van den Hurk and Viterbo (2003)</p> <ul style="list-style-type: none"> <li>➤ Global Soil Texture (FAO)</li> <li>➤ New hydraulic properties</li> <li>➤ Variable Infiltration capacity &amp; surface runoff revision</li> </ul>	<p><b>NEW SNOW</b></p> <p>Dutra et al. (2010)</p> <ul style="list-style-type: none"> <li>➤ Revised snow density</li> <li>➤ Liquid water reservoir</li> <li>➤ Revision of Albedo and sub-grid snow cover</li> </ul>	<p><b>NEW LAI</b></p> <p>Boussetta et al. (2013)</p> <ul style="list-style-type: none"> <li>➤ New satellite-based Leaf-Area-Index</li> </ul>	<p><b>SOIL Evaporation</b></p> <p>Balsamo et al. (2011), Albergel et al. (2012)</p>	<p><b>H<sub>2</sub>O / E / CO<sub>2</sub></b></p> <p>Boussetta et al. 2013 Agusti-Panareda et al. 2015</p> <ul style="list-style-type: none"> <li>➤ Integration of Carbon/ Energy/ Water</li> </ul>	<p><b>FLake</b></p> <p>Mironov et al (2010), Dutra et al. (2010), Balsamo et al. (2012, 2010)</p> <ul style="list-style-type: none"> <li>➤ Extra tile (9) to for sub-grid lakes and ice</li> <li>➤ LW tiling (Dutra)</li> </ul>

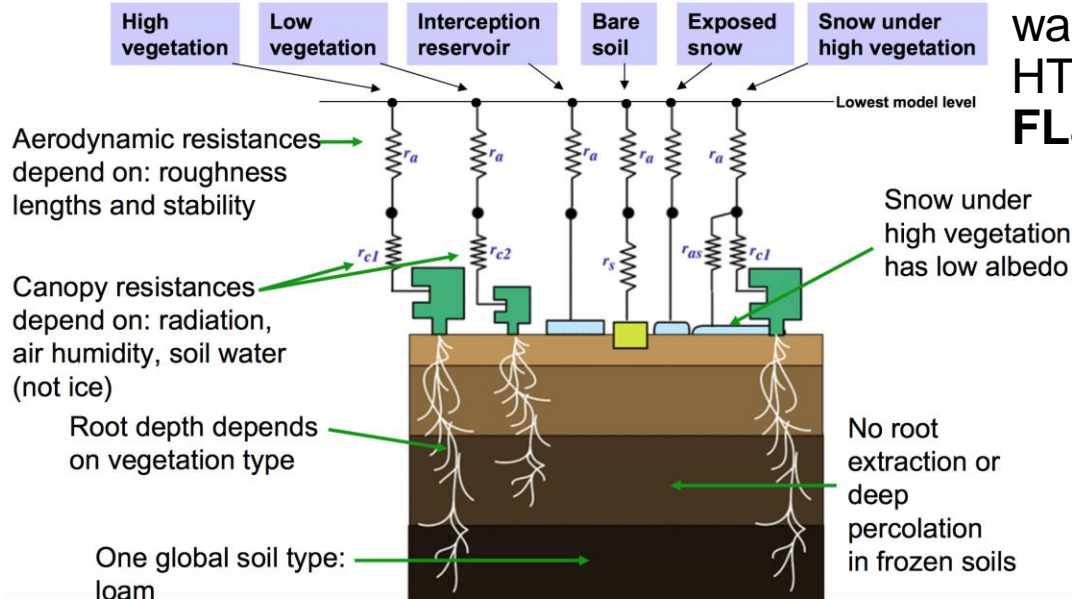


# Tiling scheme

- To represent **surface heterogeneity**, the Tiled ECMWF Scheme for Surface Exchanges over Land incorporating land surface hydrology (**HTESSEL**) is used (Balsamo et al., 2012; IFS Documentation, 2017).
- HTESSEL computes **surface turbulent fluxes** (of heat, moisture and momentum) and **skin temperature** over different tiles (vegetation, bare soil, snow, interception and water) and then calculates an **area-weighted average** for the grid-box to **couple with the atmosphere**.

Climatological land use data fields derived from 2'30" GLCC:

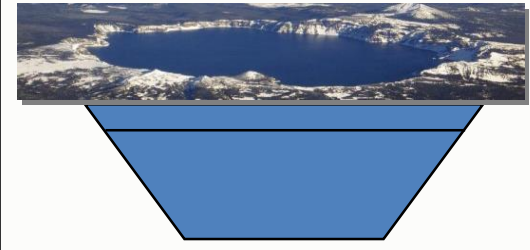
Low vegetation cover  
Low vegetation type      High vegetation cover  
High vegetation type



- A **new tile**, representing lakes, reservoirs, rivers and coastal waters, was introduced (Dutra et al., 2010) in HTESSEL in **2015** by including the **FLake model** (Mironov et al., 2006).

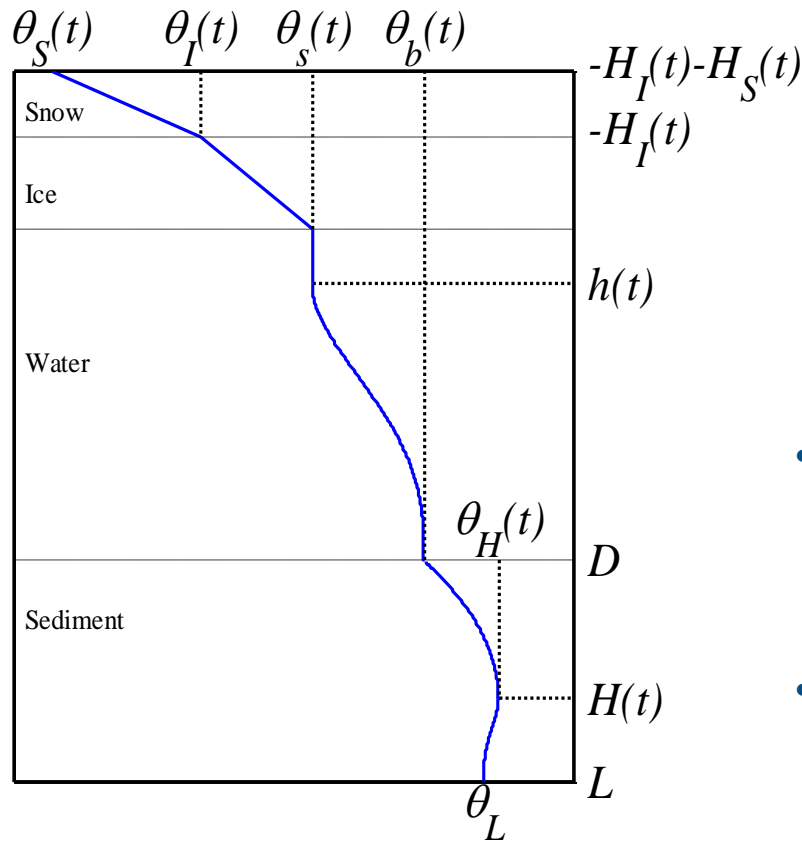
Balsamo et al. (2010, 2012, 2013)  
Dutra et al. (2010)  
Mironov et al (2010)

**Lake tile** (9)  
accounts sub-grid lakes



# FLake model: description

- The Fresh-water Lake model **FLake** (Mironov, 2008; Mironov et al., 2012) only accurately represents **fresh-water lakes**.



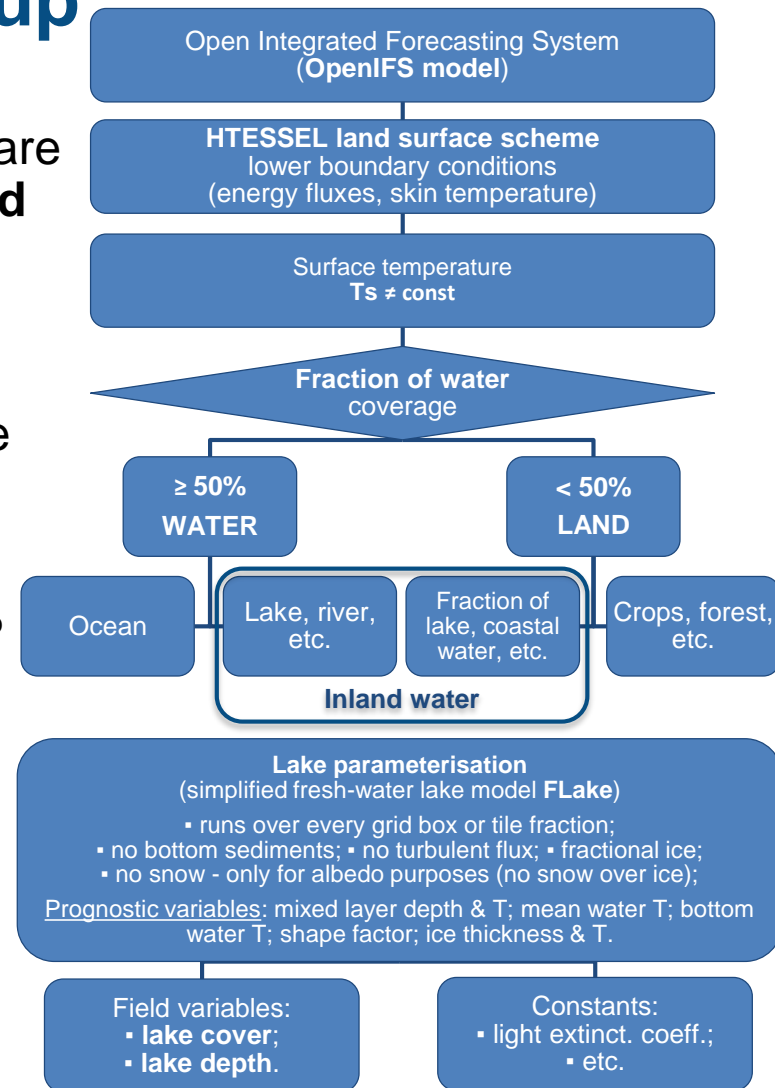
- FLake is a **one-dimensional** model, which uses an **assumed shape for the lake temperature profile** including the **mixed layer** (uniform distribution of temperature) and the **thermocline** (its upper boundary located at the mixed layer bottom, and the lower boundary at the lake bottom). The model also contains an *ice module*, a *snow module* and a *bottom sediments module*.
- FLake **requires the lake location and depth**, and lake *initial conditions*. **Depth is the most important** external parameter that FLake uses.
- FLake is a **computationally efficient shallow-water** model (it solves a number of ordinary differential equations) that incorporates most of the essential physics.

- FLake web page includes an **Online FLake** version at <http://lakemodel.net>.



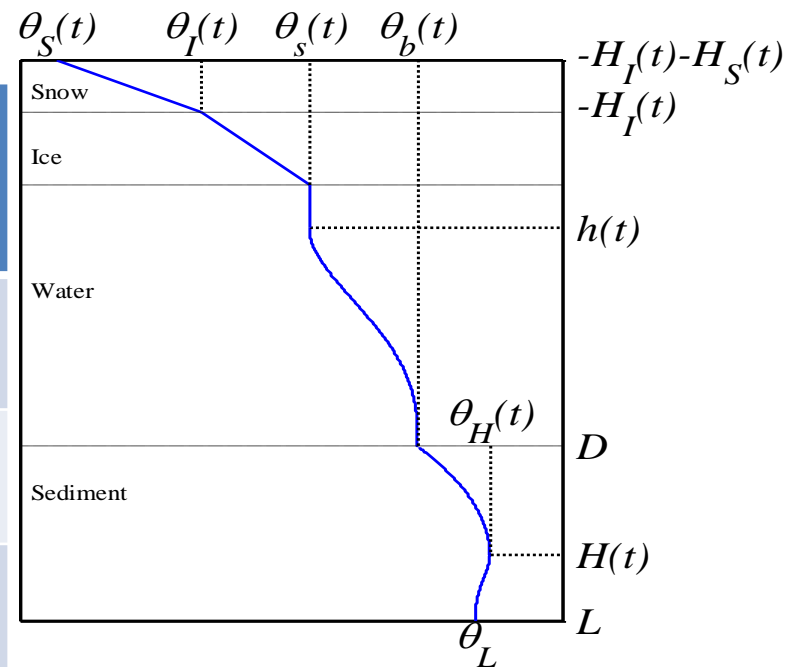
# FLake model: OpenIFS setup

- **FLake runs on each surface grid-point**, whether the simulation results in this point are used later or not (resulting **fields** are **stored** in the **MARS** archive).
- FLake runs with **no bottom sediment** and **snow modules** (snow accumulation over ice is not allowed and snow parameters are used only for albedo purposes).
- Lake **ice** can be **fractional** within a grid-box with inland water: **10 cm** of ice = 100% of a grid-box or tile is **covered with ice**; **0 cm** of ice = 100% of the grid-box is covered by **water**; linear interpolation in between) (Manrique-Sunen et al., 2013).
- The **water balance equation** is **not included for lakes** and the lake **depth** and surface **area** are kept **constant** in time (IFS Documentation, 2017).
- FLake also requires **fractional lake cover** and lake **depth** (preferably bathymetry; should be accurate and up-to-date, **global & continuous**).

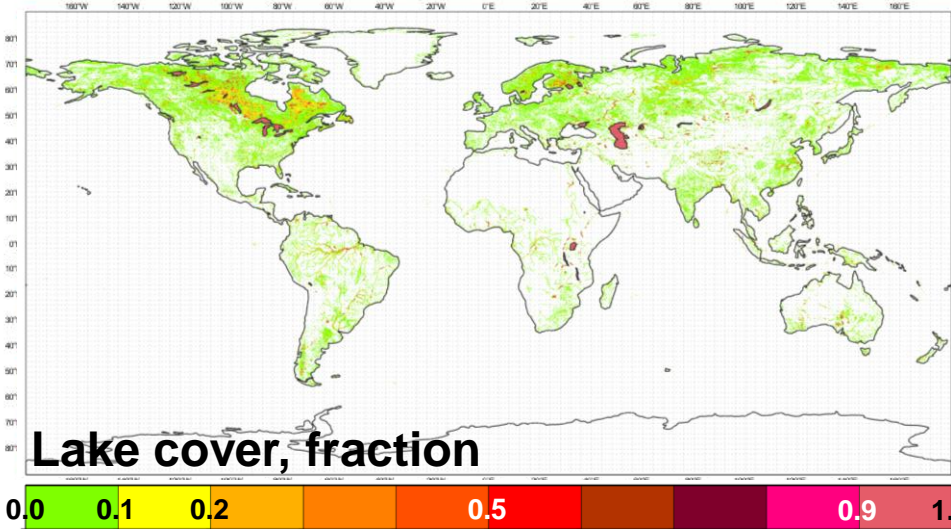


# FLake model: prognostic variables in MARS archive

GRIB short name	GRIB code	Full name
LMLT	8.228	Lake <b>mixed-layer temperature</b> (K)
LMLD	9.228	Lake <b>mixed-layer thickness</b> (m)
LBLT	10.228	<b>Temperature at the water-bottom sediment interface</b> (K)
LTLT	11.228	Lake <b>water column mean temperature</b> (K)
LSHF	12.228	Lake <b>shape factor</b> with respect to the temperature profile in the thermocline (dimensionless)
LICT	13.228	<b>Temperature at the lake snow-ice or air-ice interface</b> (K) (single level field)
LICD	14.228	Lake <b>ice thickness</b> (m)



# FLake model: global input fields

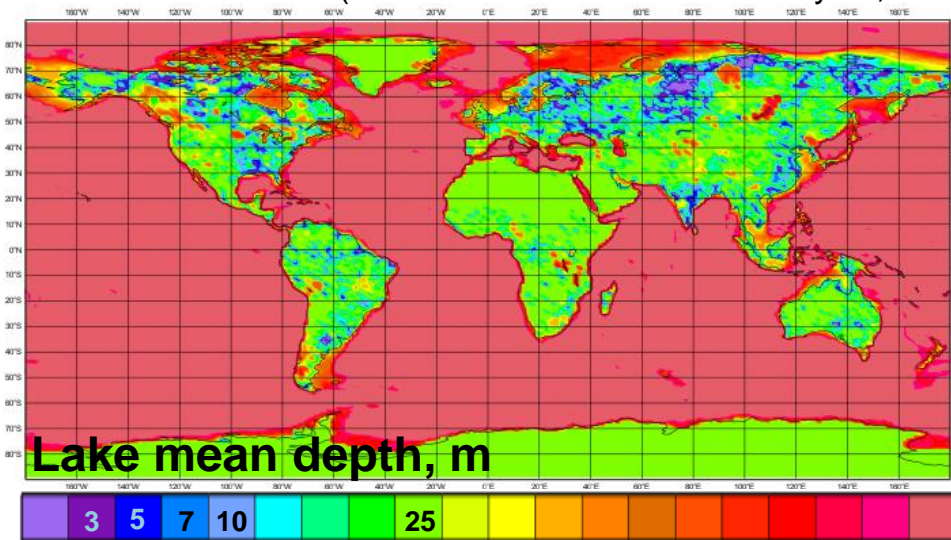


**Lake cover** (fraction) [CL / 26] - static field with **fraction** of lakes in every grid box. Values between **0 - no lake** and **1 - fully covered with lake**. CL is consistent with land-sea mask. Global field consists of:

- 60°S to 85°N - **GlobCover 2006** [300m]
- *Antarctica* - Radarsat Antarctic Mapping Project Digital Elevation Model Version 2 (**RAMP2**) (Liu et al., 2015) [1 km]
- *Arctic* - north of 85 °N **no land** is assumed

“Depth of a lake is the main parameter to which model is sensitive...”  
(Kourzeneva E. and Braskavsky D., 2006)

Lake depth, m	1-7	Annual cycle amplitude changes for 1 K when depth changes for	1 m
	7-16		2 m
	16-40		3 m



**Lake depth** (m) [DL / 7.228] - static field with **mean depth** value in each grid box in order to **ease the interoperability of lake output at diverse spatial resolutions** where the ratio of resolved/unresolved lakes and coastal water varies. Global field consists of:

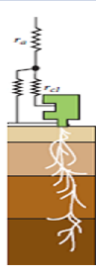

- *Over land*
  - ✓ lake mean depth – **GLDBv1** (in-situ ~13'000) [~1km]
  - ✓ bathymetry – **ETOPO1** [~2km] (Great lakes, Azov sea), **Calaveri** [~4km] (Caspian sea)
  - ✓ default = **25 m**
- *Over ocean* – **ETOPO1** [~2km]

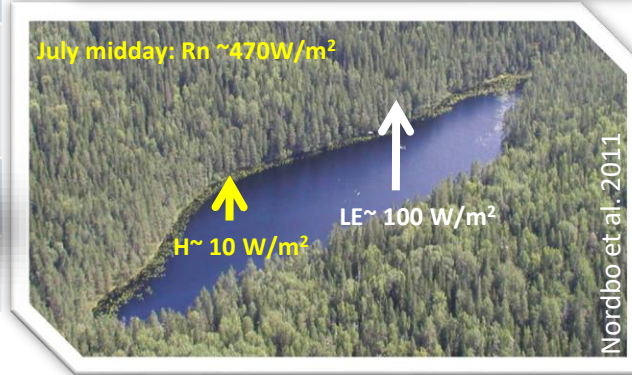
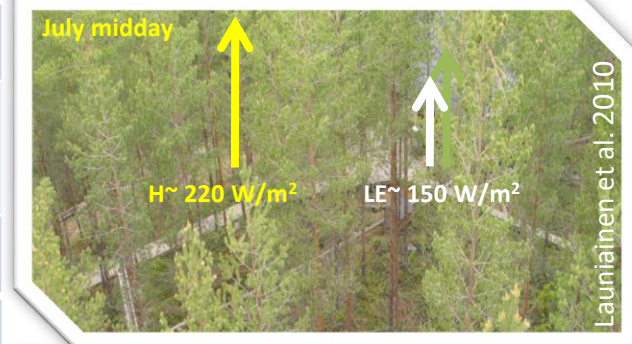
# Introducing FLake under the tiling approach

(Manrique-Suñén et al., 2013)

Southern Finland,  
stations are ~80 km apart

## In-situ data

<b>Hyytiala forest</b>	Location: 61°51'N, 24°17'E	
	Scots pine forest with lingonberry, blueberry and mosses underneath	
	Canopy height in 2006: 16.5 m	
<b>Lake Valkea-Kotinen</b>	Location: 61°51'N, 24°17'E	
	Small boreal lake surrounded by tall forest – strong channelling of flow along the lake	
	Area: 0.041 km <sup>2</sup> , elongated shape	
	Depth: mean = 2.5 m, max = 6.5 m	



## Model data

**Surface off-line** experiments single column (no feedback from surface to atmosphere)

**ERA-Interim** atmospheric forcing (~80 km), lakes not resolved

**IFS** model CY43R1, site characteristics from the nearest model grid point

**~9 km horizontal resolution** (Tco1279)

Length: January-December 2006

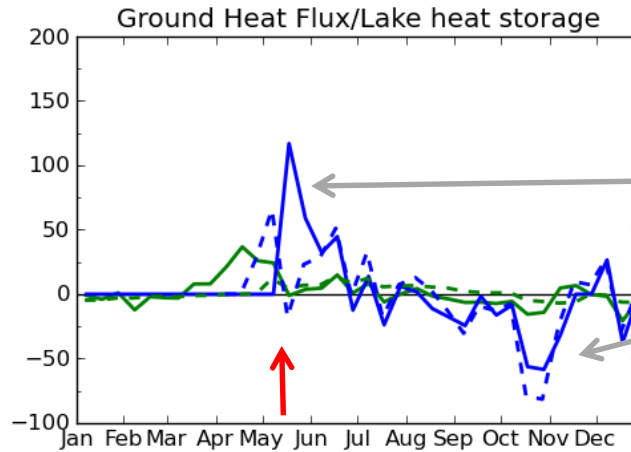
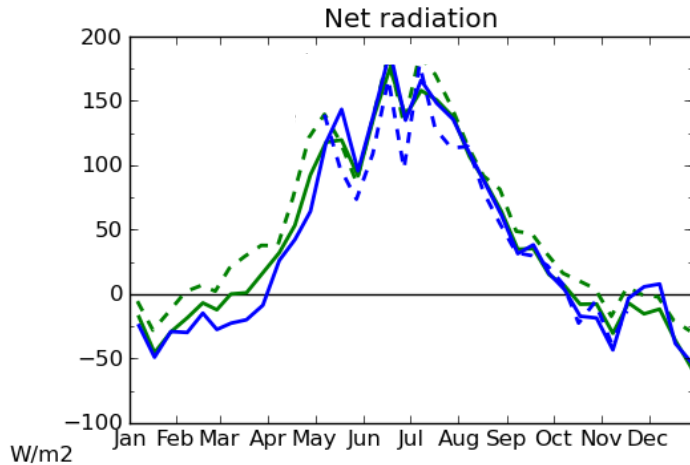
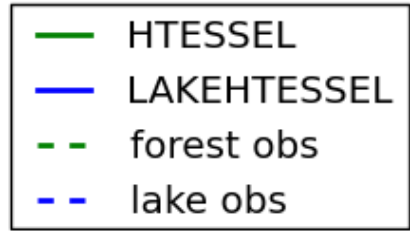
Prior cyclic yearly 3 time **spin up**



# Energy fluxes: Seasonal cycles

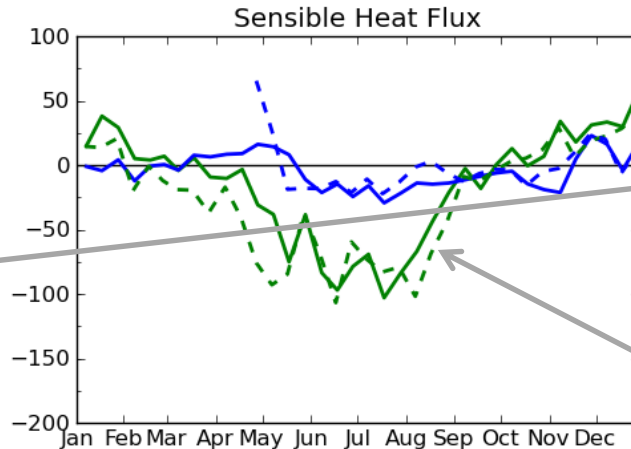
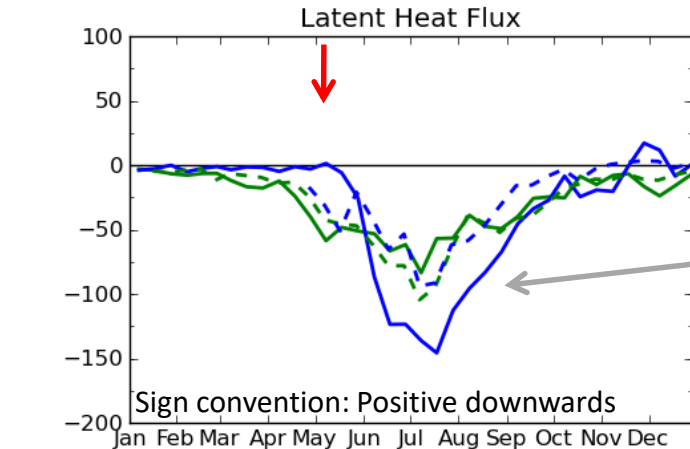
(Manrique-Suñén et al., 2013)

## Seasonal cycle of 10 day averages of energy fluxes



**Lake:** Energy is stored from May to August and released in autumn

Main difference between two sites is in energy partitioning into SH and G



**Lake:** Model overestimates the evaporation in summer by  $\sim 50 \text{ W m}^{-2}$

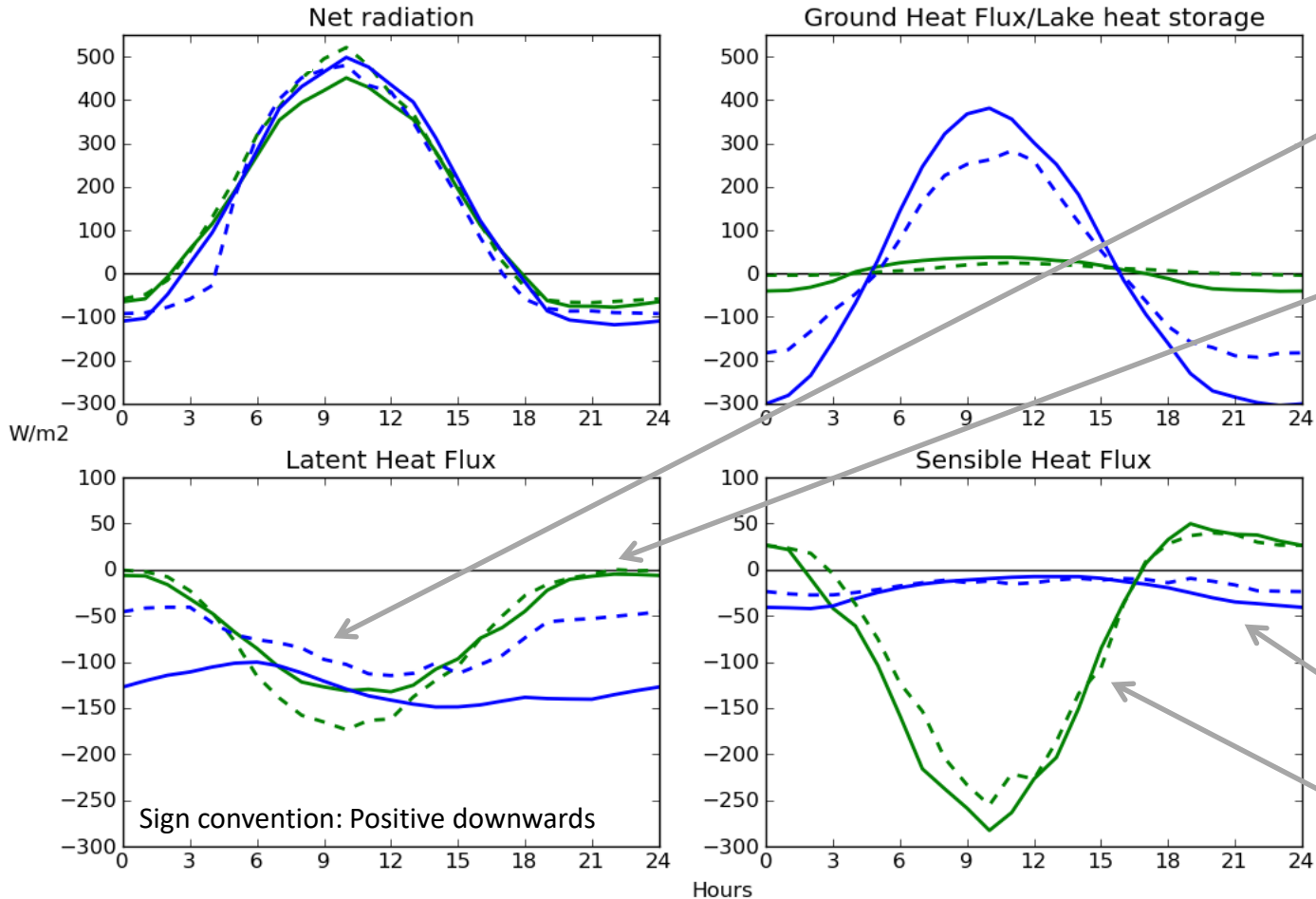
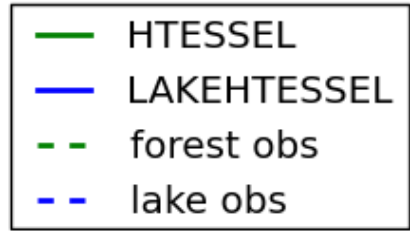
**Forest:** Upward SH flux in summer

**Timing of the lake's energy cycles is influenced by the ice cover break up, in the model it is delayed by 14 days → ice-initial condition will benefit from EO data constraint!**

# Energy fluxes: Diurnal cycles

(Manrique-Suñén et al., 2013)

## Monthly diurnal cycle of energy fluxes for July



**Lake LH diurnal cycle: overestimation in evaporation**

**Forest: evaporation is driven by vegetation → it is zero at night**

Main difference between two sites is in energy partitioning into SH and G

**Lake: SH maximum is at night**

**Forest: SH maximum is at midday**

Very good representation by the model of diurnal cycles and particularities of each surface!

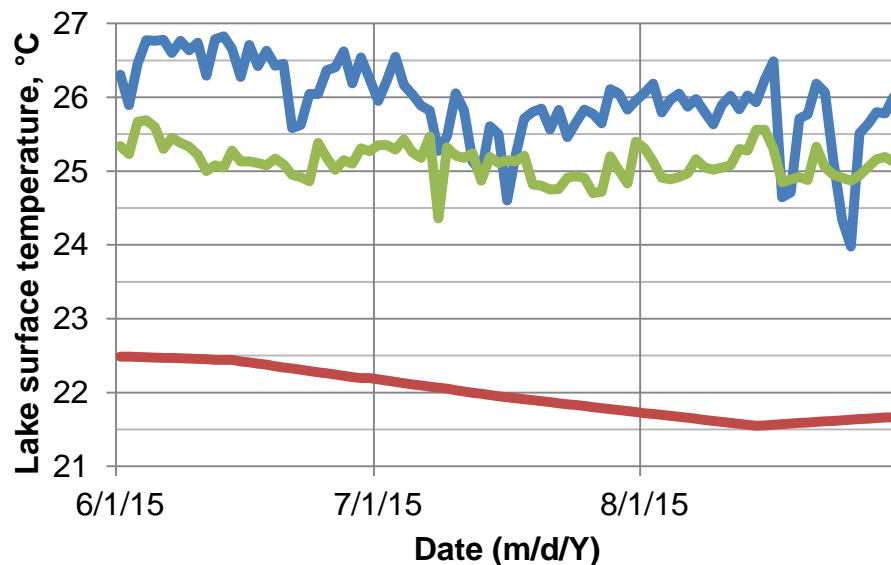
# Lake surface temperature: Diurnal cycles



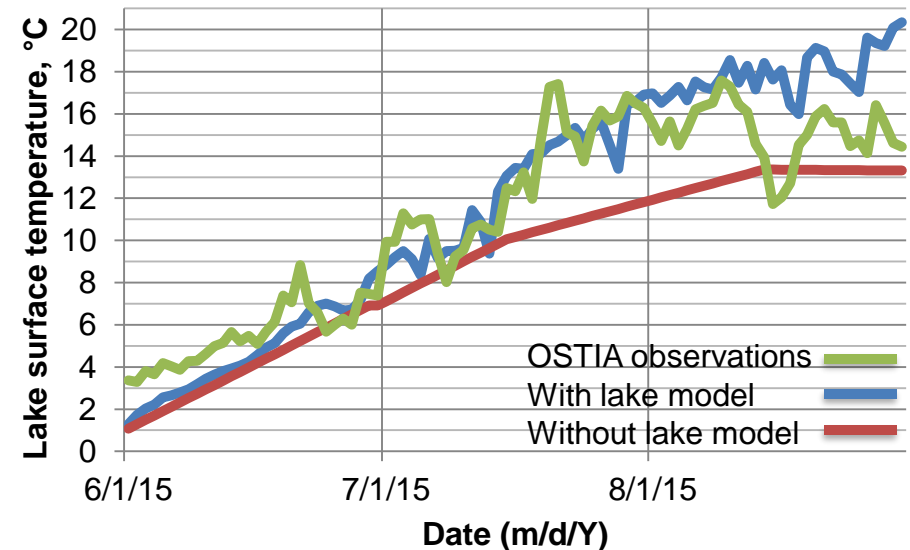
*Large improvements:*

- **reducing** the lake temperature **bias** for **big** and **deep** lakes (e.g. Lake Victoria, Great Bear, Titicaca);
- generally more **realistic** temperature **diurnal variability**.

Lake Victoria (Africa)

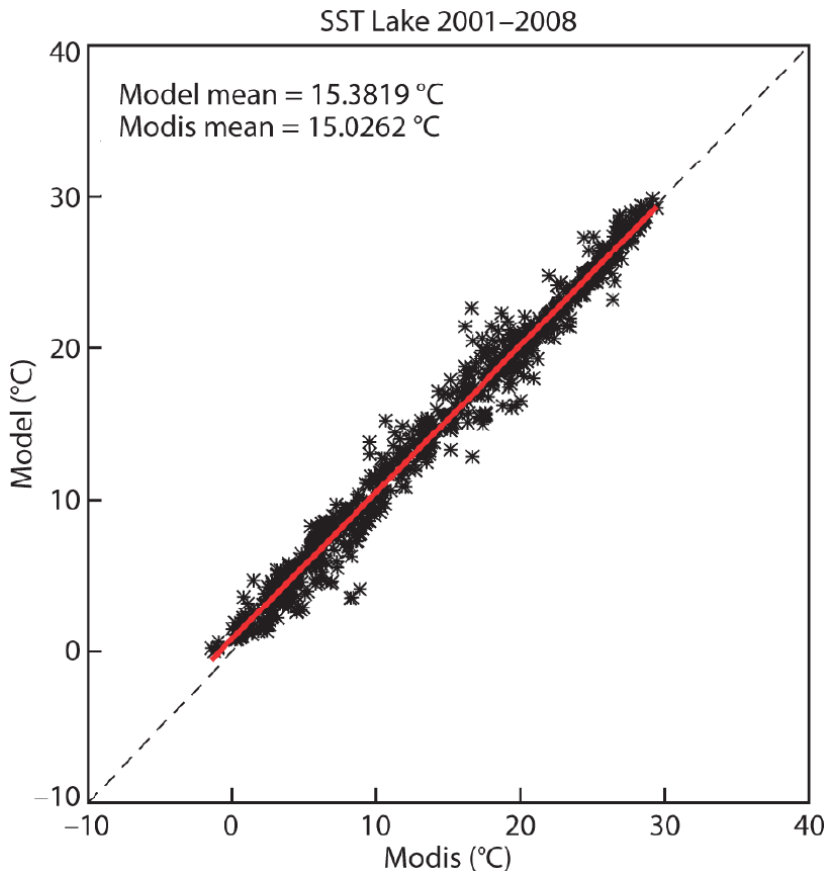


Lake Baikal (Russia)



# Lakes surface temperature: global validation

(Balsamo et al., 2012)

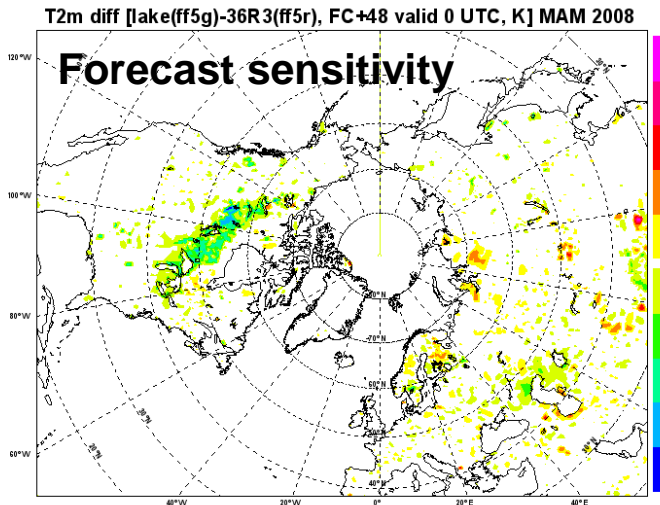


- **In-situ:** *MODIS* Terra/Aqua satellite global composite based on the Level 3 Mapped Thermal IR sea surface temperature product, which senses the *sea/lake water temperature*, resolution  $\sim 4$  km.
- **Model:** *FLake* model driven by *ERA-Interim* 3-hourly atmospheric forcing, resolution  $\sim 80$  km.
- **Period:** 2001.01.01-2008.12.31.
- **Comparison** in terms of annual mean surface water temperature values:
  - ✓ largely *unbiased* simulation over grid points where the model lake fraction is  $\geq 10$  %:
    - ❖ **good correlation** between **modelled** and **observed** annual mean  $R = 0.98$ ;
    - ❖ **BIAS** (modelled – observed) is **reduced**, is  $< 0.3$  K;
  - ✓ largest differences are found over Caspian sea and southern regions of the North-American Great Lakes (positive BIAS) and over Norwegian lakes (negative BIAS) → consistent with model intrinsic **limitations over deep waters**.

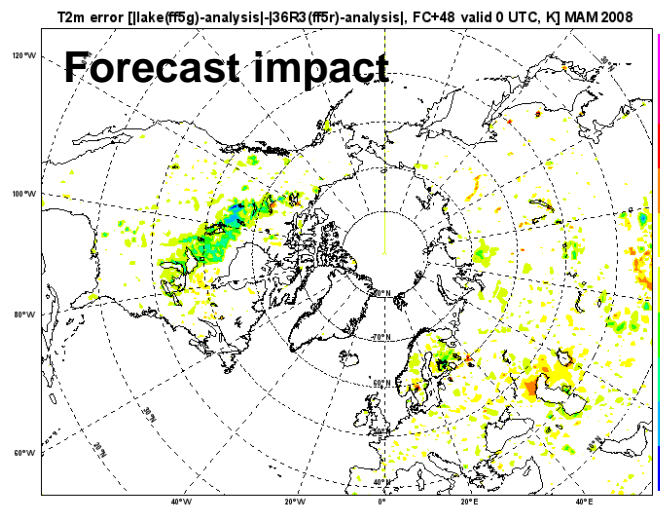


# Impact of lakes in NWP forecasts

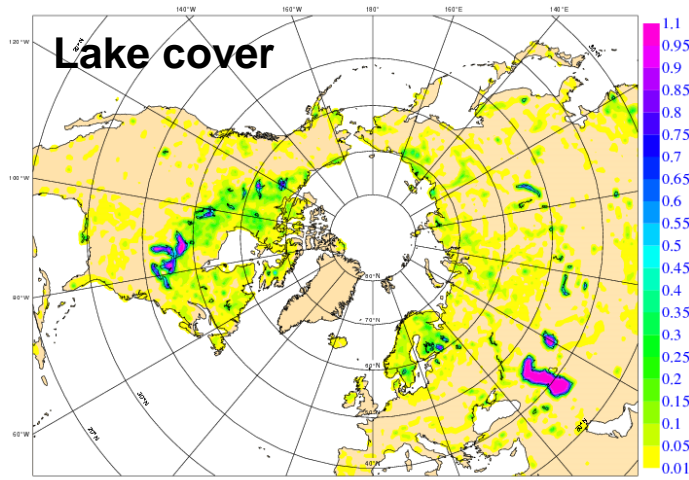
(Balsamo et al., 2012)



Cooling 2m temperature  
Warming 2m temperature



Improves 2m temperature  
Degrades 2m temperature



Forecasts sensitivity and impact show **spring-cooling over lake** areas with benefit to the 2-meter temperatures forecasts (day-2, 48-hour forecast).

ERA-Interim forced runs of FLake model are used for **lake model climatology** generation, which serves as **initial conditions in forecasts experiments**.

Here is shown **spring sensitivity** and **error impact** on **temperature** when **lake** model is **activated**.

# Impact of lakes in NWP analysis cycles (1)

## AN cycling and initialisation: temperature scores

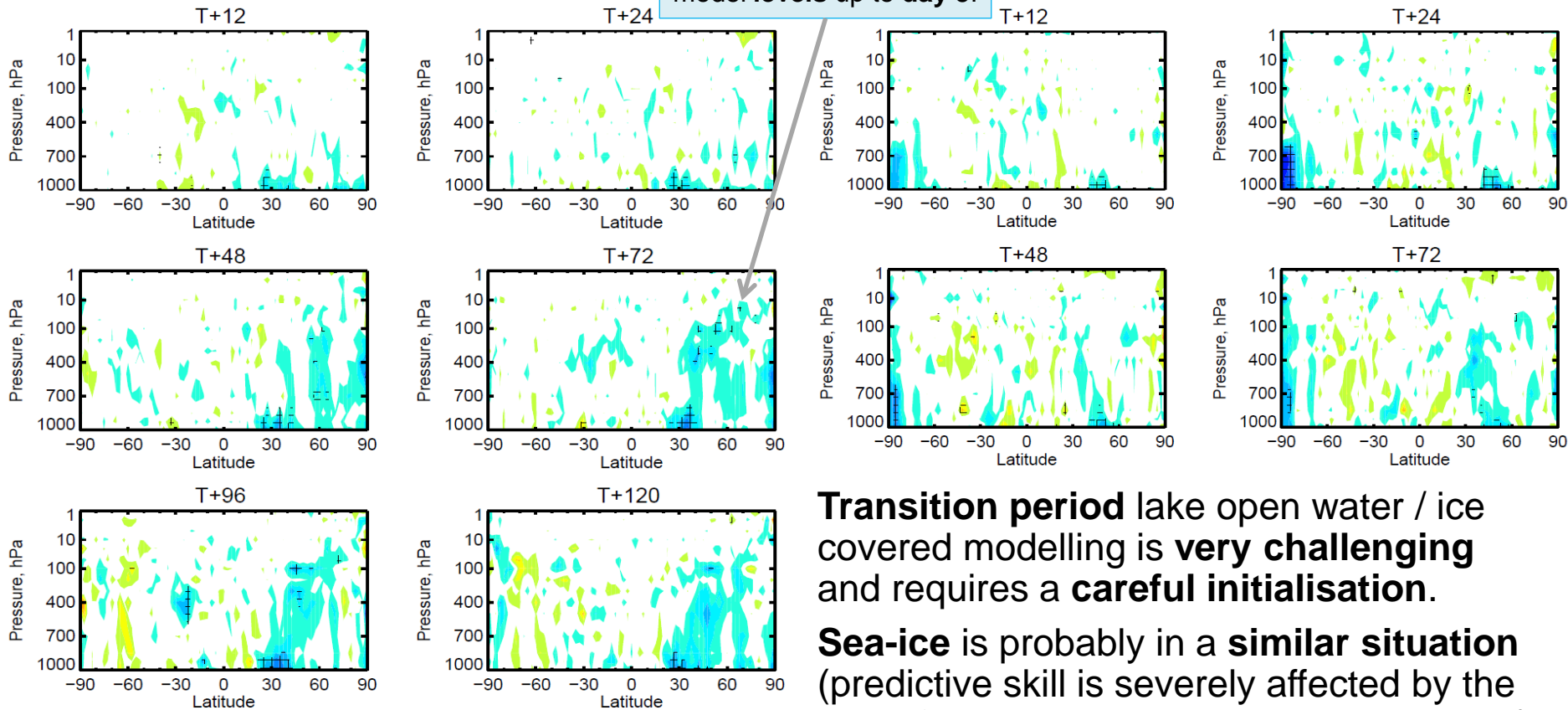
### Summer experiment

15-Jun-2013 to 5-Jul-2013

Propagation of the positive impact to higher model levels up to day 5!

### Winter experiment

1-Dec-2013 to 31-Dec-2013



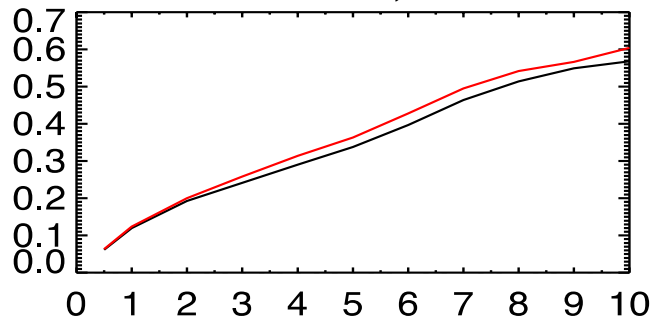
**Transition period lake open water / ice covered modelling is very challenging and requires a careful initialisation.**

**Sea-ice is probably in a similar situation (predictive skill is severely affected by the lack of atmospheric predictability in winter).**

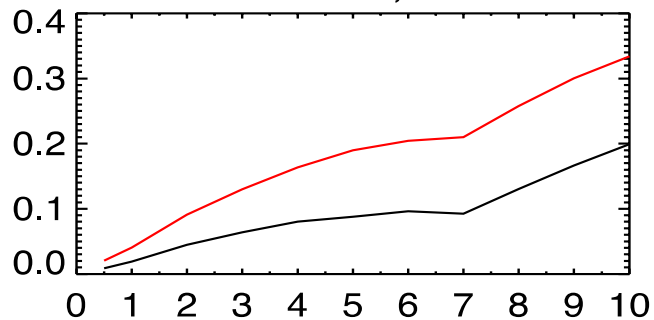


# Impact of lakes in NWP analysis cycles (2)

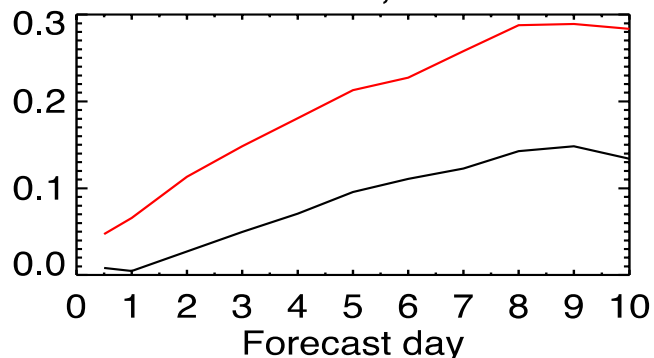
T: 20° to 90°, 500hPa



T: 20° to 90°, 850hPa

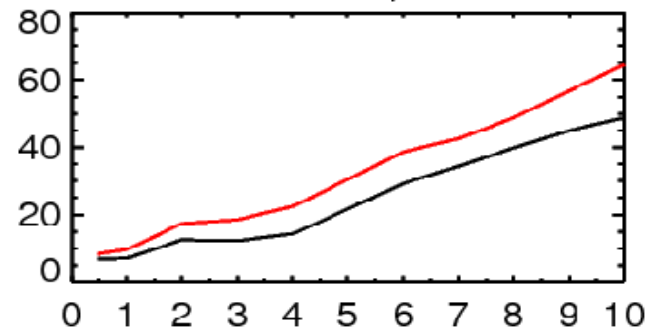




T: 20° to 90°, 1000hPa



- During **summer lake impact** is not confined to the surface layer but **propagates upwards reducing** the mean model **temperature error** over **Northern hemisphere** (e.g. at 850 hPa).
- Part of the **signal** is also **detected** in **Z500** (geopotential height at 500 hPa), which gets a lot **attention from meteorologists!**

Z: 20° to 90°, 500hPa



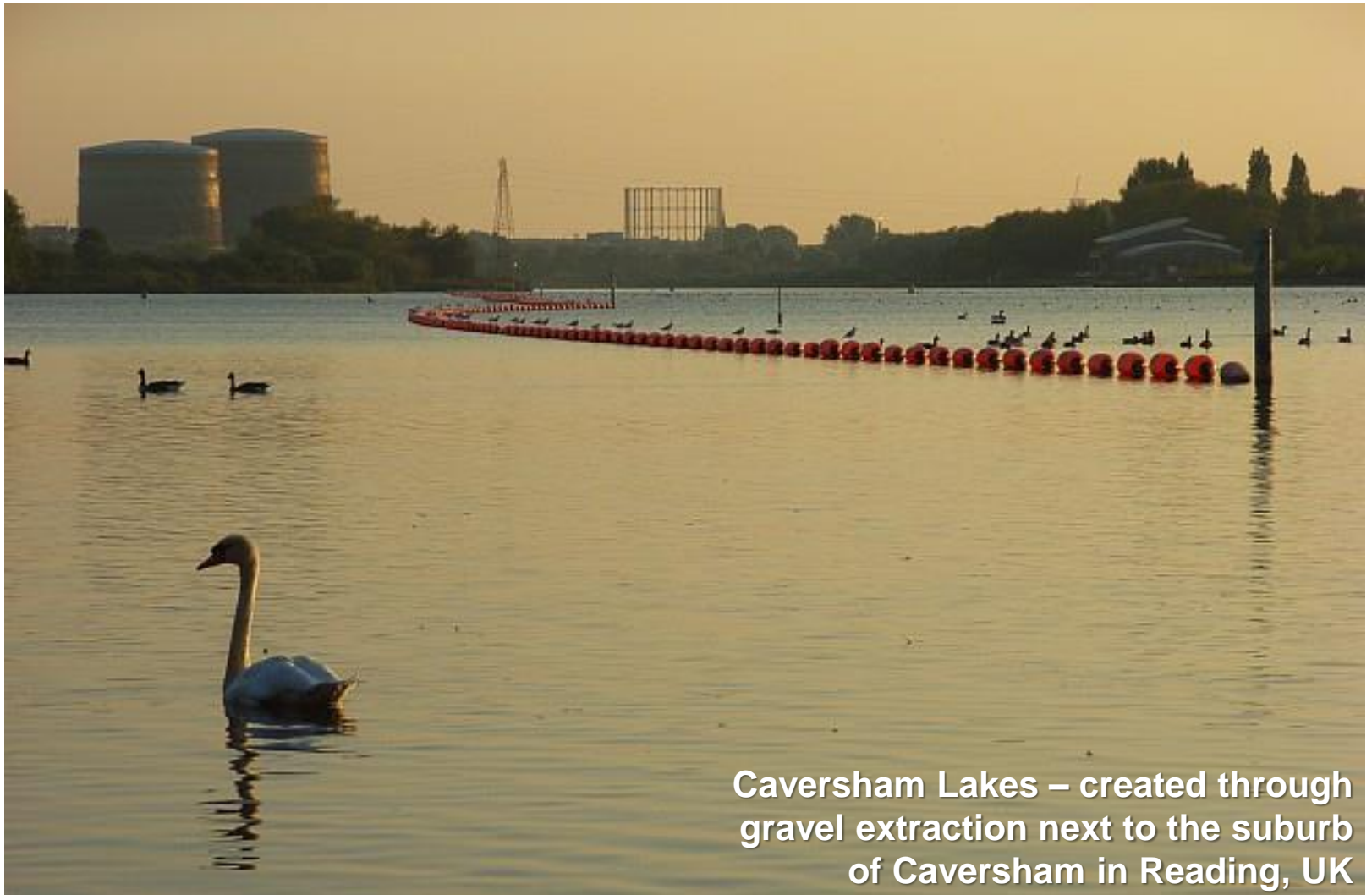
mean error of CY40R3 with lakes   
mean error of CY40R3 without lakes 

# Summary & Outlook

- **The ECMWF land surface scheme and its extension to lakes**
  - ✓ The introduction of subgrid lakes and coastal waters enhances the capacity of representing natural Earth surface heterogeneity.
- **Benefits of considering sub-grid lakes**
  - ✓ Each tile has its process description (no ad-hoc or effective parameters).
  - ✓ All inland water bodies considered independently from their size, shape and depth.
- **Atmospheric forecast impact**
  - ✓ The introduction of interactive lakes has beneficial impact on forecast accuracy.
  - ✓ Impact is significant and detected in Northern Hemisphere scores.
- **The verification phase**
  - ✓ Use of satellite based lake temperature and lake ice information has potential to allow a routine verification of lakes that can foster further improvements.



# Thank you for your attention!



Caversham Lakes – created through gravel extraction next to the suburb of Caversham in Reading, UK



Photo from [www.touropia.com](http://www.touropia.com)

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# Lake surface temperature: point validation

Lakes verification in the first three full months June-July-August 2015 (91-days AN vs OSTIA-lake) of operations → large improvements on the majority of lakes according to the OSTIA satellite products.

Lake AFRICA	RMSE	BIAS	CORRELATION	Mean Model	Mean Obs	St.Dev. Model	St.Dev. Obs
Victoria_IFS41R1	<b>0.957</b>	<b>0.826</b>	0.491	25.665	24.849	0.554415	0.230933
Victoria_IFS40R1	3.157	-3.14	0.328	21.743	24.849	0.322463	0.230933

Lake CANADA	RMSE	BIAS	CORRELATION	Mean Model	Mean Obs	St.Dev. Model	St.Dev. Obs
Great_Bear_IFS41R1	<b>2.875</b>	<b>1.877</b>	0.927	5.225	3.368	3.87317	1.96852
Great_Bear_IFS40R1	5.401	4.598	0.894	7.916	3.368	4.45394	1.96852

Lake S. AMERICA	RMSE	BIAS	CORRELATION	Mean Model	Mean Obs	St.Dev. Model	St.Dev. Obs
Titicaca_IFS41R1	<b>0.611</b>	<b>-0.425</b>	0.822	12.322	12.742	0.739826	0.482809
Titicaca_IFS40R1	3.804	-3.789	0.752	8.995	12.742	0.463688	0.482809

Lake EU	RMSE	BIAS	CORRELATION	Mean Model	Mean Obs	St.Dev. Model	St.Dev. Obs
Ladoga_IFS41R1	<b>2.45</b>	<b>2.051</b>	0.958	14.207	12.178	4.22985	4.60613
Ladoga_IFS40R1	1.443	-0.295	0.984	11.886	12.178	3.3881	4.60613

Lake sub-grid EU	RMSE	BIAS	CORRELATION	Mean Model	Mean Obs	St.Dev. Model	St.Dev. Obs
Haukivesi_IFS41R1	<b>1.706</b>	<b>-0.02</b>	0.807	15.188	15.207	2.24239	2.88615
Haukivesi_IFS40R1	2.915	-2.733	0.964	12.504	15.207	3.44774	2.88615

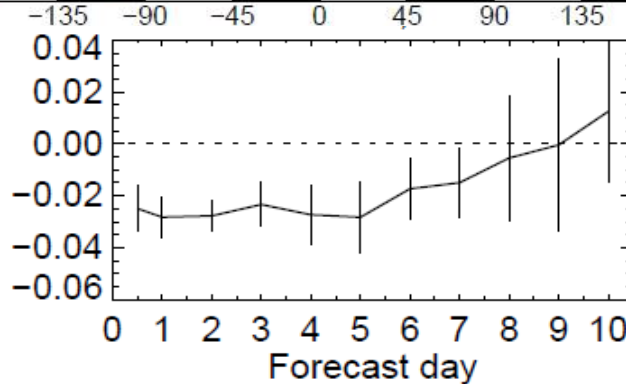
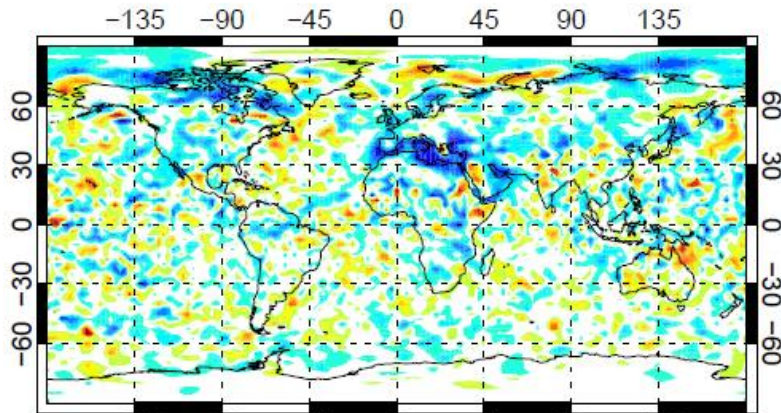
# Impact of lakes in NWP analysis cycles (3)

## AN cycling and initialisation: temperature scores

### Summer experiment

15-Jun-2013 to 5-Jul-2013

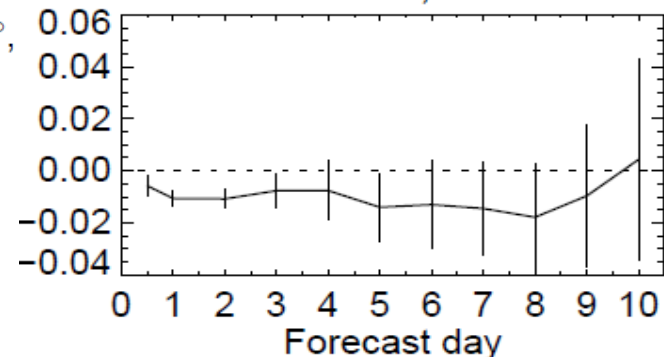
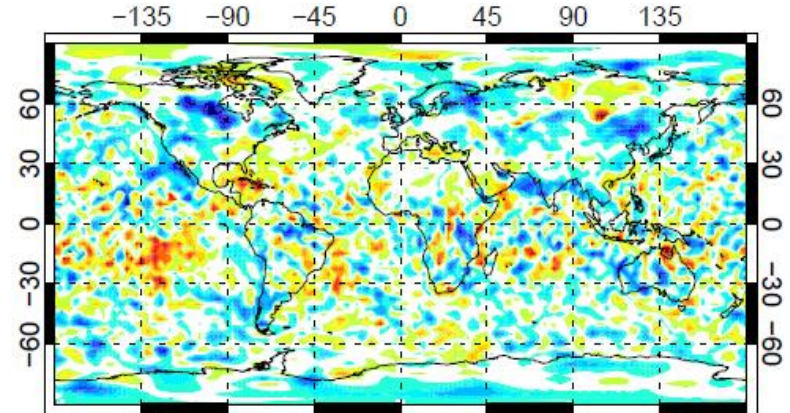
T+48; 1000hPa



T: 20° to 90°,  
1000hPa

### Winter experiment

1-Dec-2013 to 31-Dec-2013



- In lake and coastal water proximity 2-meter temperature forecast is improved.
- During summer there is 2-3 % of relative improvement in 1000 hPa temperature RMSE, with significance of up to 7 days.
- During winter RMSE impact is also positive and around 1 %.