

The ECMWF radiation scheme

Mark Fielding and Robin Hogan

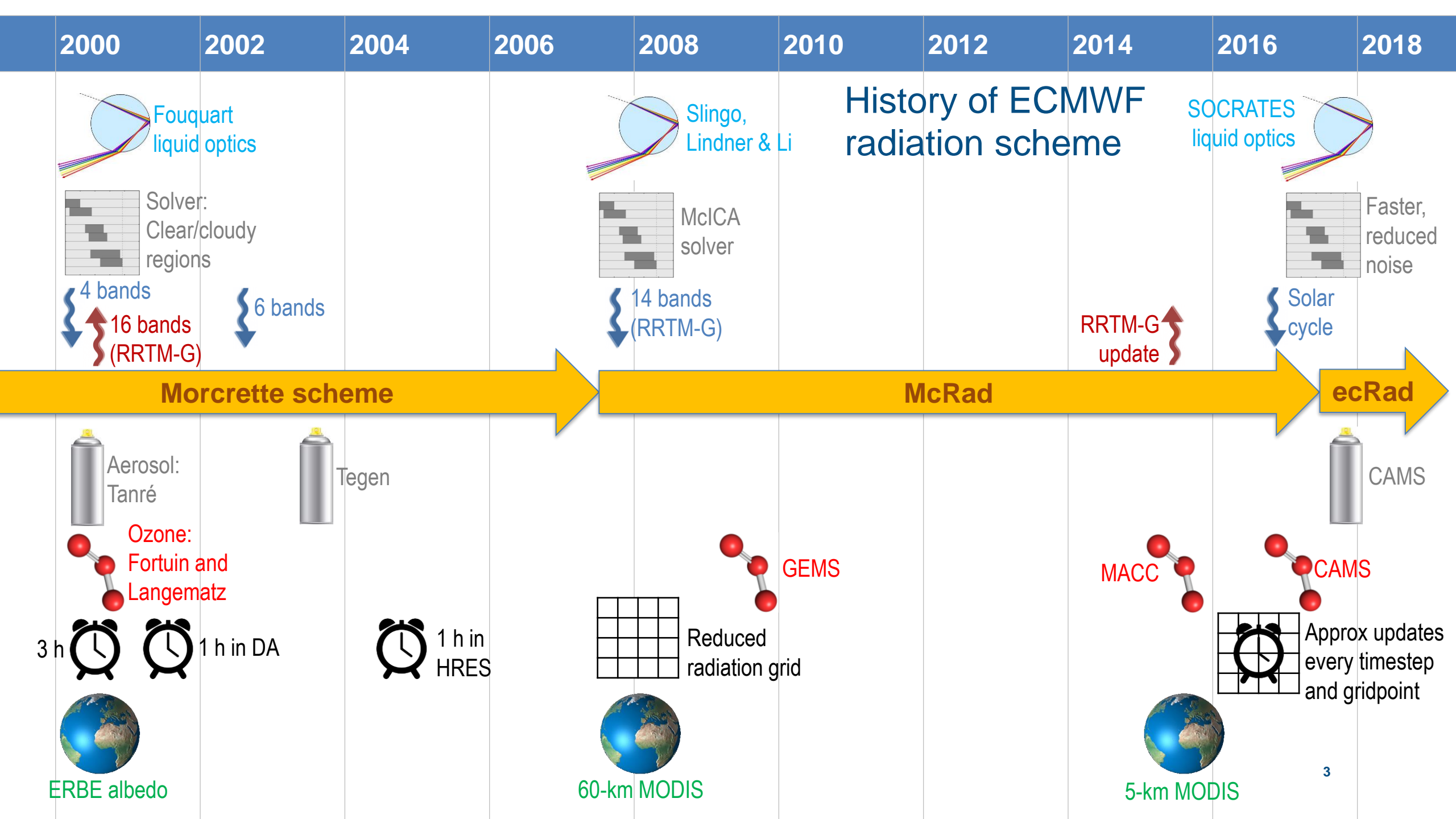
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With contributions from Alessio Bozzo



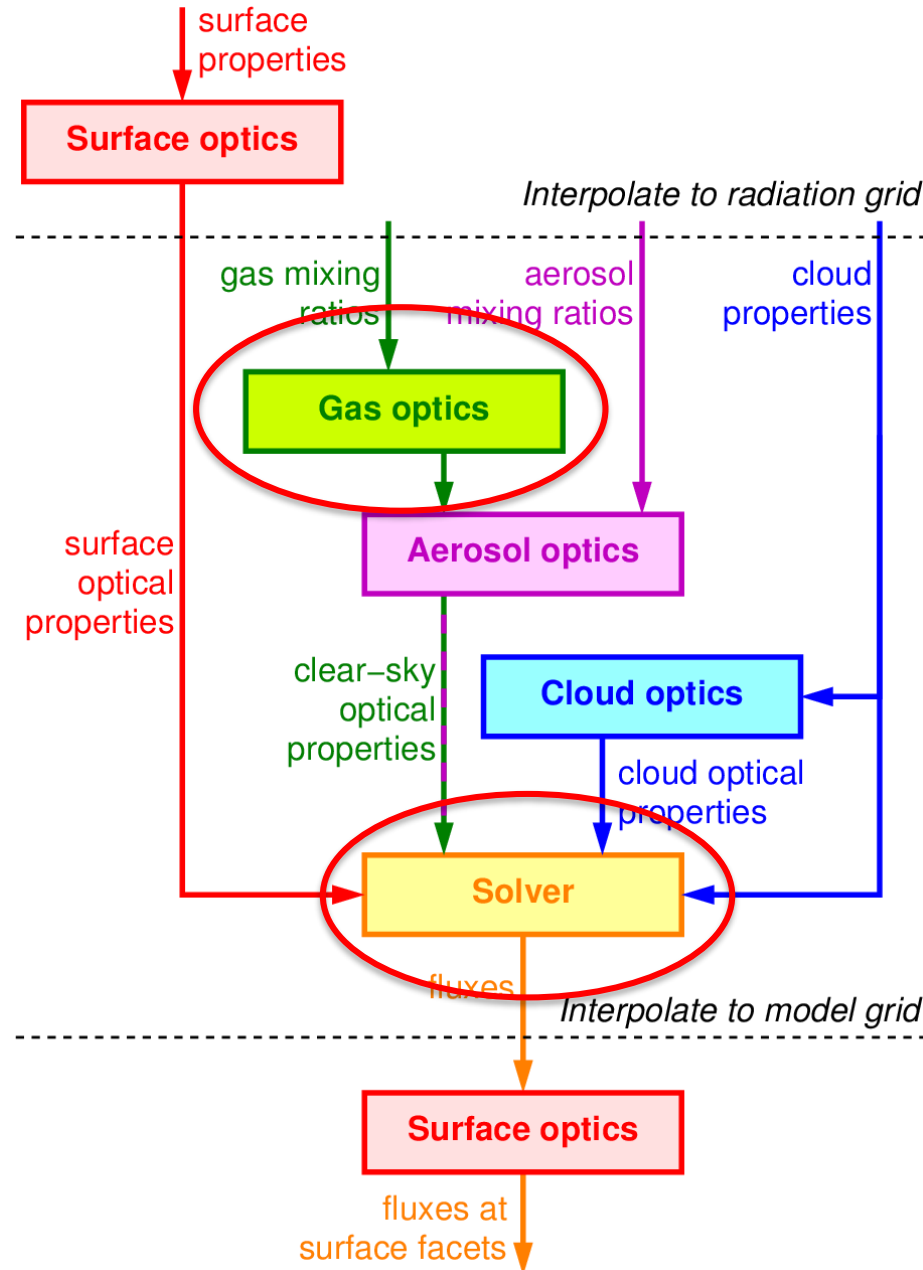
Outline

- History of the ECMWF radiation scheme
- Overview of the current scheme: ecRad
- Treatment of gases (Correlated-k) and clouds (McICA)
- Optimization considerations (spectral vs spatial vs temporal)
- Testing the radiative impact of changes to parameterizations
- Do aerosols matter for NWP?
- Remaining challenges



Modular design of ecRad

- Gas optics
 - RRTM-G (as before)
 - Plan to develop new scheme with far fewer spectral intervals
- Aerosol optics
 - Number of species set at run time and optical properties configured by NetCDF file
 - Supports Tegen and CAMS (prognostic & diagnostic)
- Cloud optics
 - Liquid clouds: more accurate SOCRATES scheme
 - Ice clouds: Fu by default, Baran and Yi available

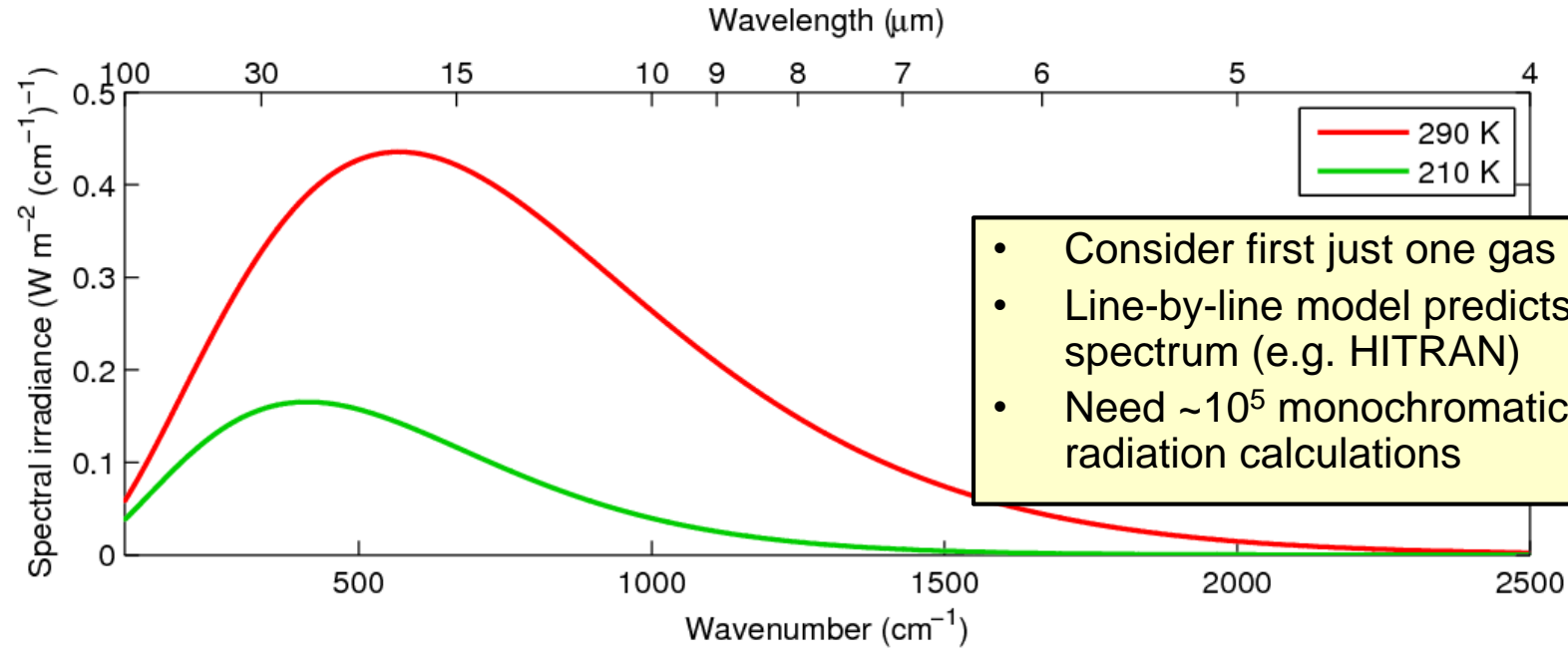


• Solver

- McICA, Tripleclouds or SPARTACUS solvers
 - SPARTACUS makes the IFS the only global model that can do 3D radiative effects
 - Better solution to longwave equations improves tropopause & stratopause
 - Longwave scattering optional
 - Can configure cloud overlap, width and shape of PDF
- Surface (*under development*)
- *Rigorous and consistent treatment of radiative transfer in urban and forest canopies*
- Offline version available for non-commercial use under OpenIFS license

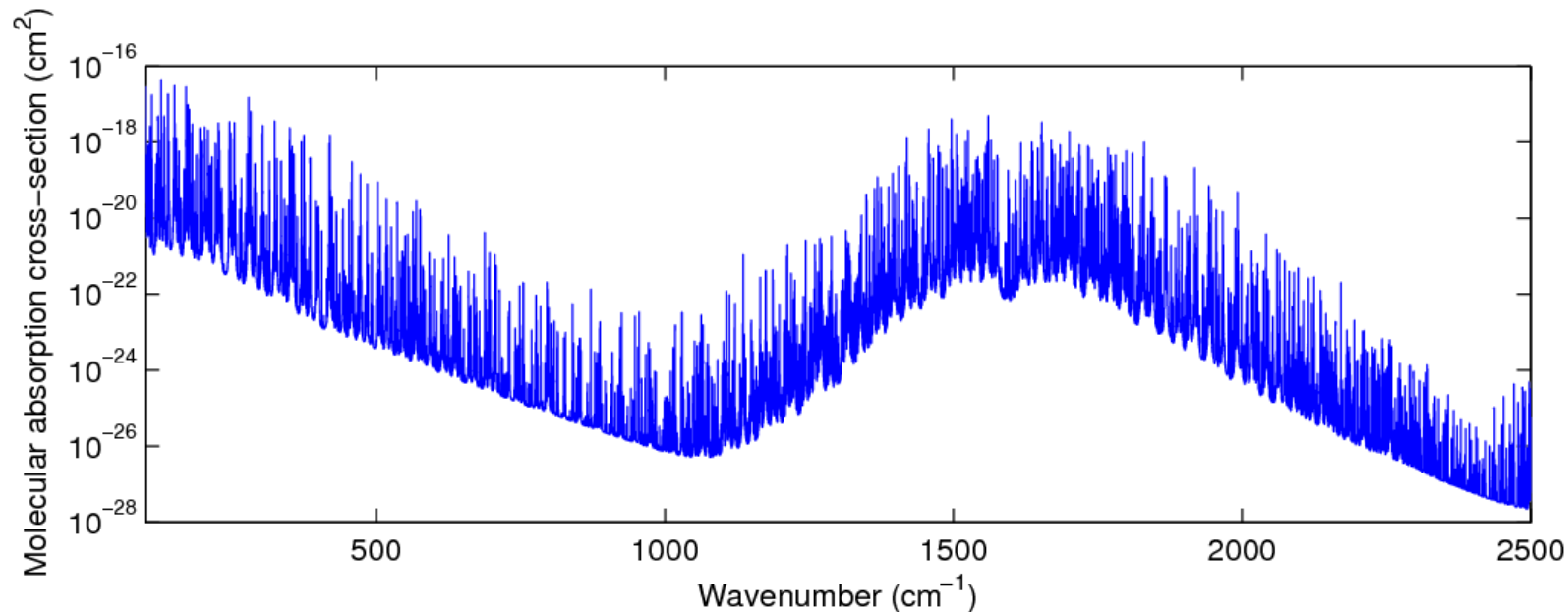
GAS OPTICS: How do we integrate across the spectrum?

Planck function



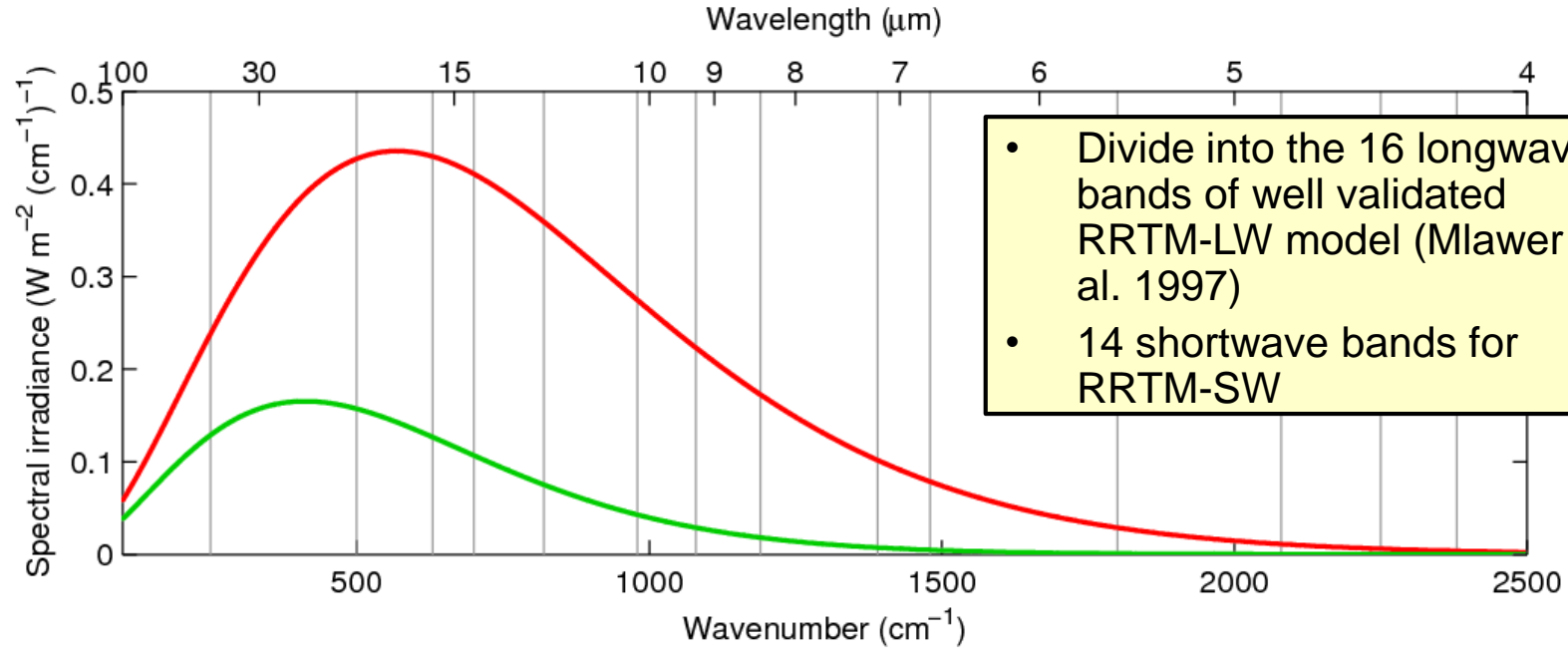
- Consider first just one gas
- Line-by-line model predicts spectrum (e.g. HITRAN)
- Need $\sim 10^5$ monochromatic radiation calculations

Water vapour spectrum

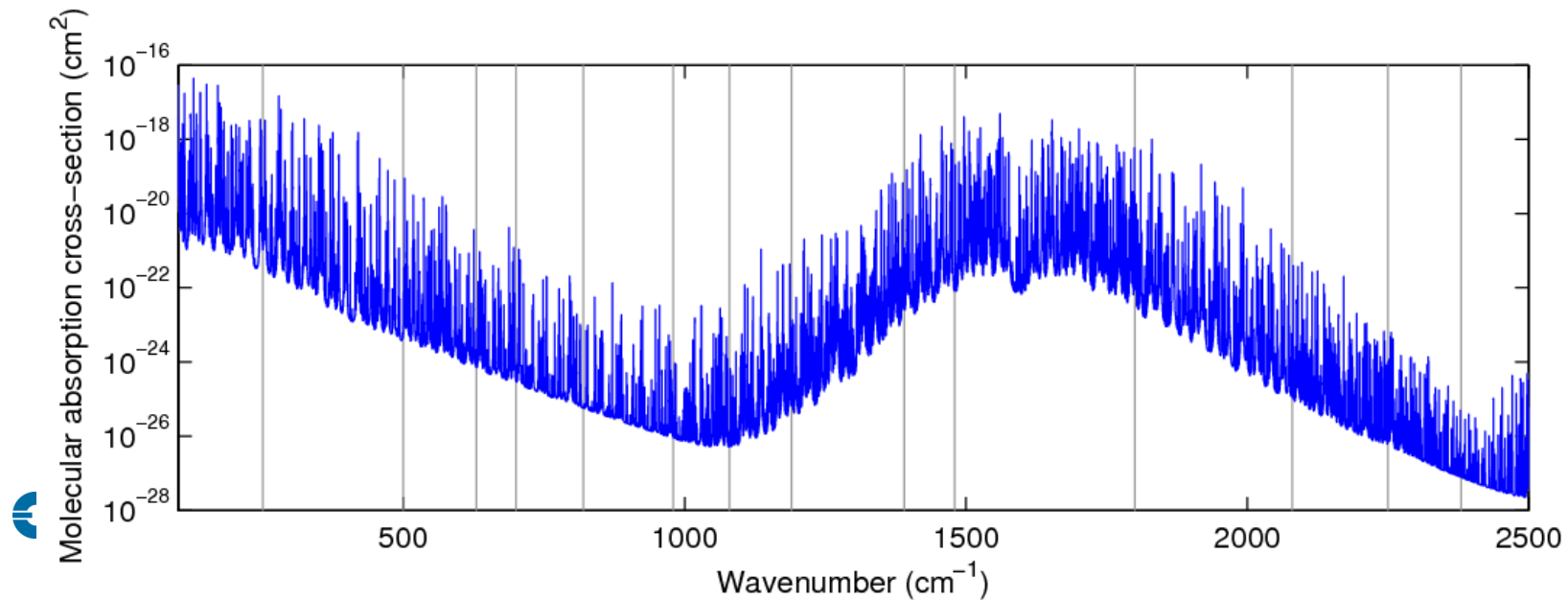


Divide into bands

Planck function

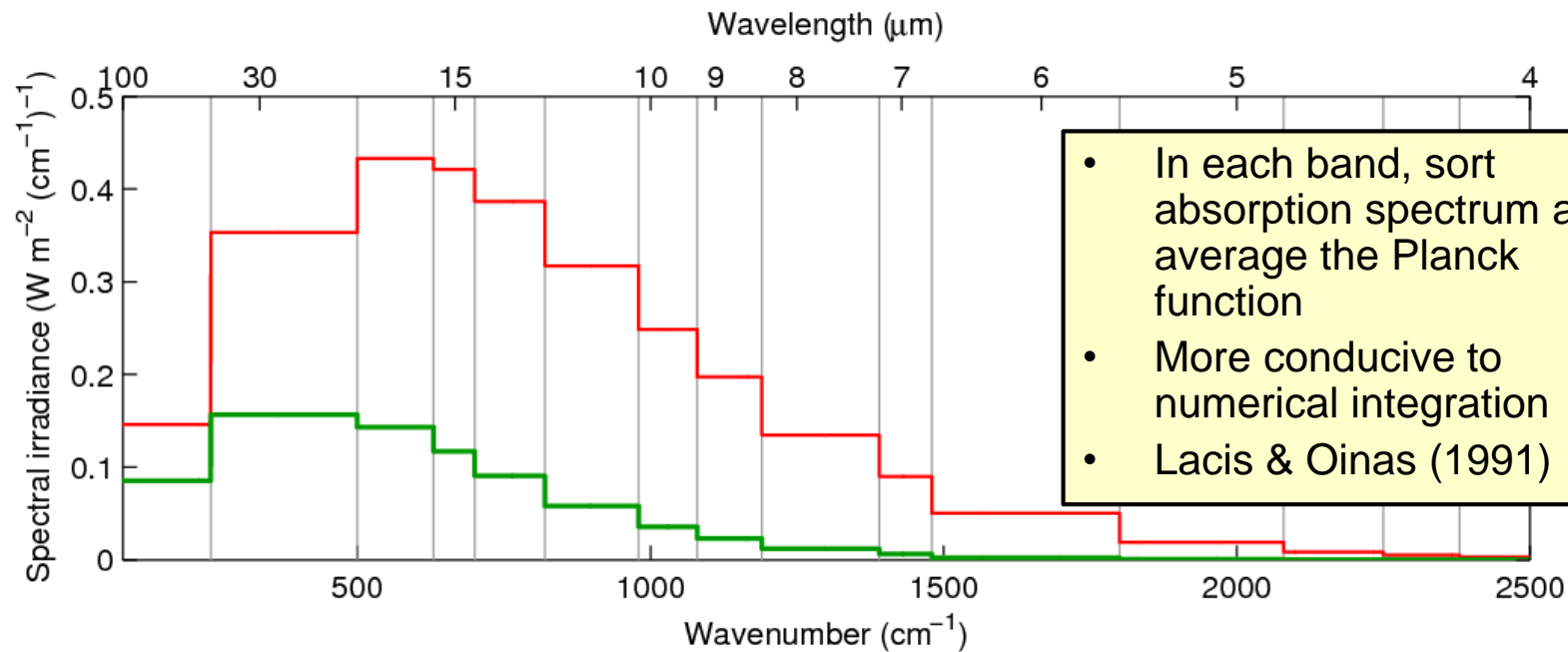


Water vapour spectrum

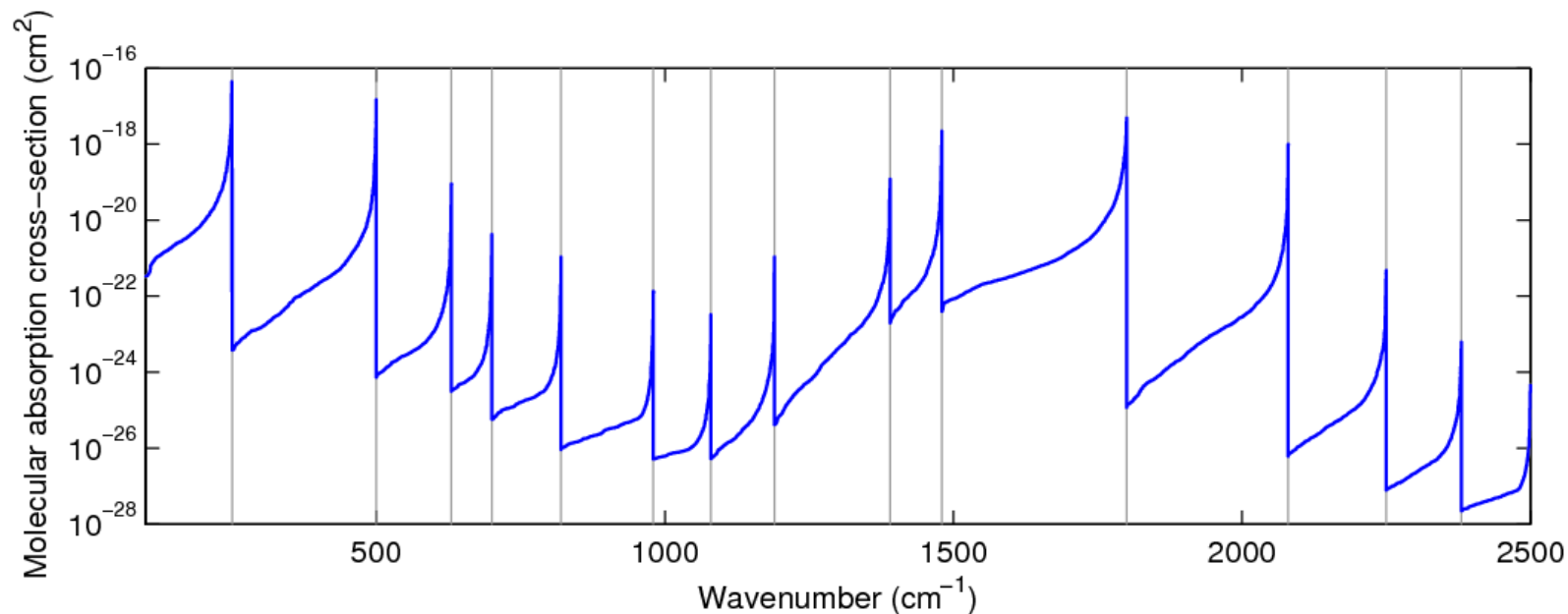


The k-distribution method

Planck function

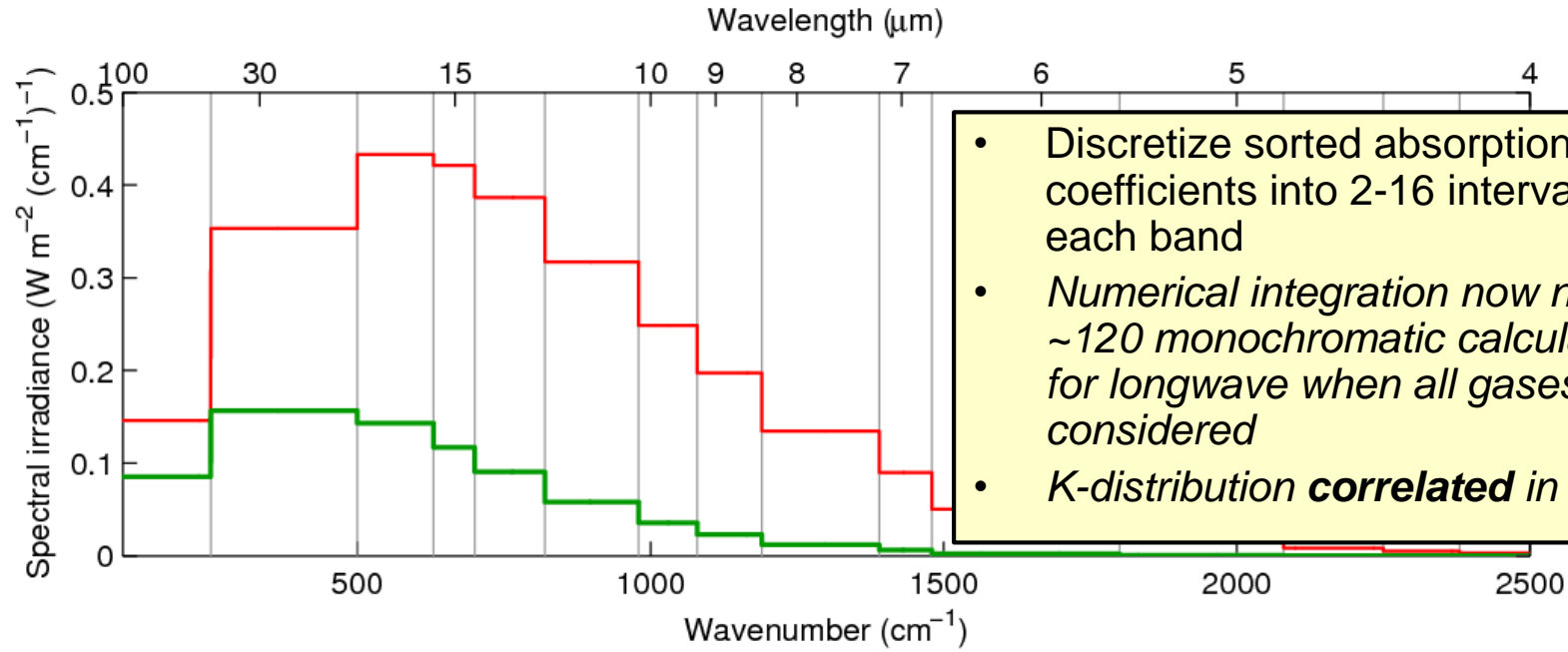


Water vapour spectrum

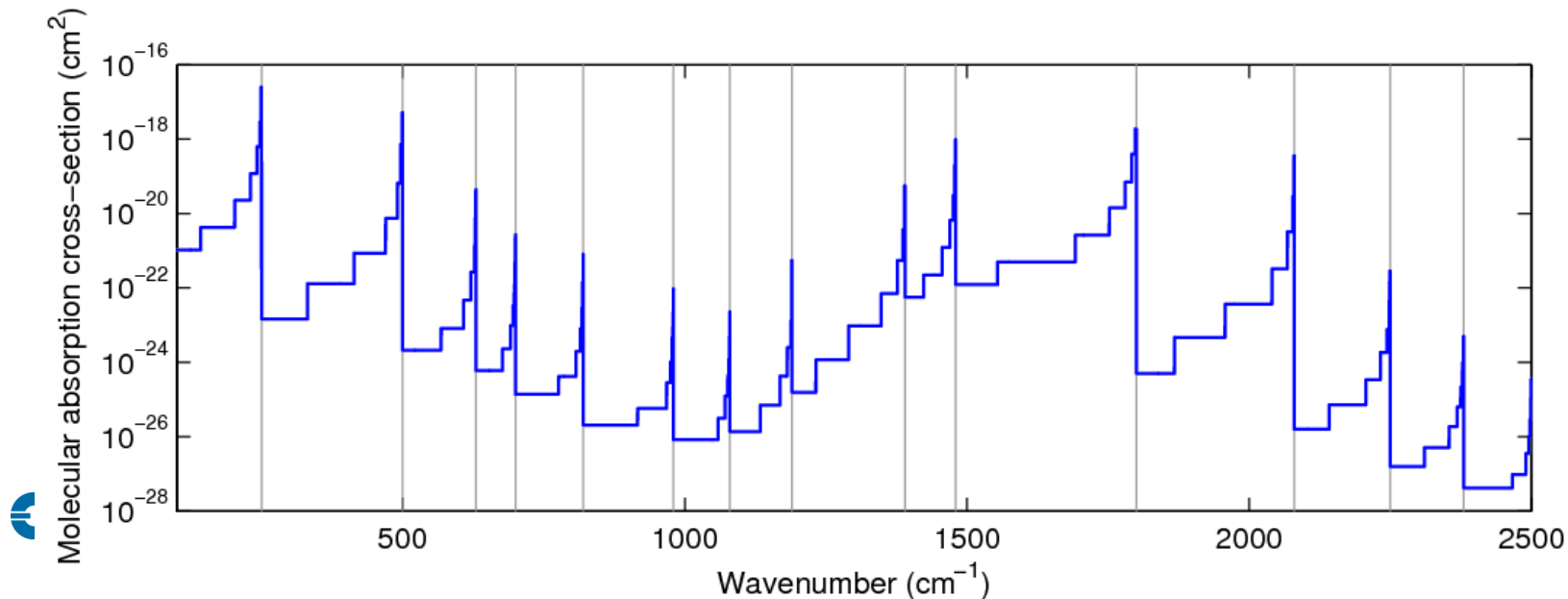


The correlated k-distribution (CKD) method

Planck function



Water vapour spectrum



RRTMG-LW configuration

	Spectral intervals cm^{-1}	Number of g-points	Gases included	
			Troposphere	Stratosphere
16 B A N D S	10–250	~100 microns	H ₂ O	H ₂ O
	250–500		H ₂ O	H ₂ O
	500–630		H ₂ O, CO ₂	H ₂ O, CO ₂
	630–700		H ₂ O, CO ₂	O ₃ , CO ₂
	700–820		H ₂ O, CO ₂ , CCl ₄	O ₃ , CO ₂ , CCl ₄
	820–980		H ₂ O, CFC11, CFC12	CFC11, CFC12
	980–1080		H ₂ O, O ₃	O ₃
	1080–1180		H ₂ O, CFC12, CFC22	O ₃ , CFC12, CFC22
	1180–1390		H ₂ O, CH ₄	CH ₄
	1390–1480		H ₂ O	H ₂ O
	1480–1800		H ₂ O	H ₂ O
	1800–2080		H ₂ O	
	2080–2250		H ₂ O, N ₂ O	
	2250–2380		CO ₂	CO ₂
	2380–2600		N ₂ O, CO ₂	
	2600–3000	~3 microns		H ₂ O, CH ₄

140 spectral points (too many? Perhaps...)

RRTMG-SW configuration

14
B
A
N
D
S

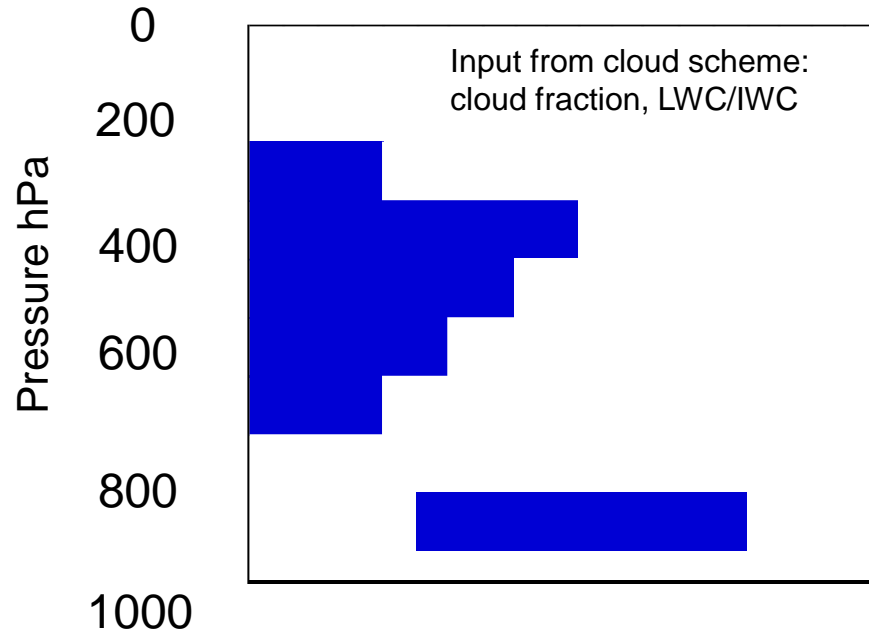
Spectral intervals cm^{-1}	Number of g-points	Gases included		
		Troposphere	Stratosphere	
800–2600	~3 microns	12	H ₂ O	CO ₂
2600–3250		6	H ₂ O, CH ₄	
3250–4000		12	H ₂ O, CO ₂	H ₂ O, CO ₂
4000–4650		8	H ₂ O, CH ₄	CH ₄
4650–5150		8	H ₂ O, CO ₂	CO ₂
5150–6150		10	H ₂ O, CH ₄	H ₂ O, CH ₄
6150–7700		10	H ₂ O, CO ₂	H ₂ O, CO ₂
7700–8050		2	H ₂ O, O ₂	O ₂
8050–12850		10	H ₂ O	
12850–16000		8	H ₂ O, O ₂	O ₂
16000–22650		6	H ₂ O	
22650–29000		6		
29000–38000		8	O ₃	O ₃
38000–50000	0.2 microns	2	O ₃ , O ₂	O ₃ , O ₂

112 spectral points (too many? Perhaps....)

Improving the representation of cloud radiative effects (1)

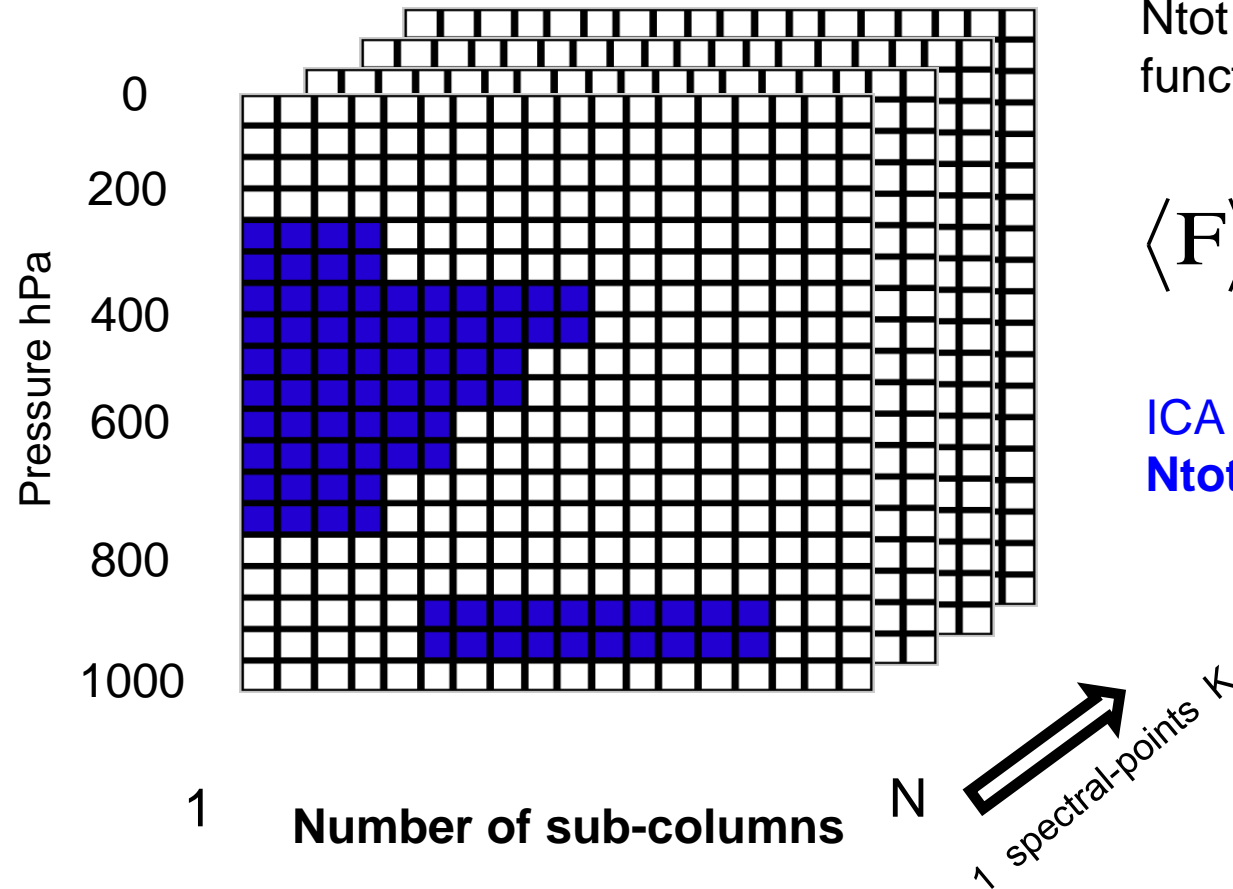
most models until ~2000

Easy way to tackle the problem: compute the clear and cloud part of the grid box (according to cloud fraction and overlap at each level) and merge fluxes



Improving cloud radiative effects (2)

independent column approximation ICA (if we had infinite computing power)



K = number of spectral intervals (g-points)

$\langle F \rangle$ average flux in the grid box

N = number of independent sub-columns

N_{tot} = total number of transmission function computations

$$\langle F \rangle = \frac{1}{N} \sum_{n=1}^N F_n$$

ICA RT scheme:

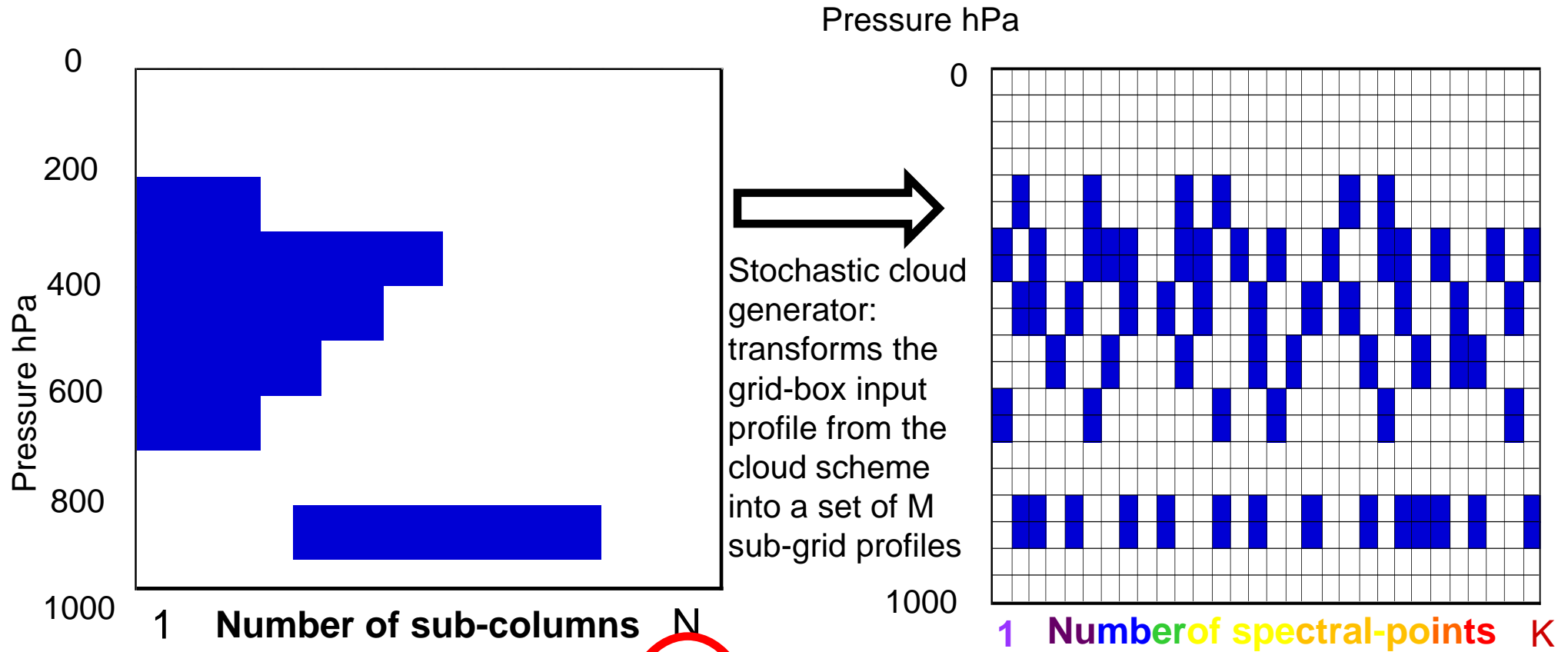
$$N_{tot} = N * K \sim O(10^3)$$

Improving cloud radiative effects (3)

Monte Carlo Independent Column Approximation McICA

Barker et al. (2003),
Pincus et al. (2003)

Cloud generator: Raisanen et al.
(2004)



McICA: approximates $\langle F \rangle = \frac{1}{N} \sum_{n=1}^N \sum_{k=1}^K c_k F_{n,k}$ into $\langle F \rangle \sim \sum_{k=1}^K c_k F_{n_k,k}$

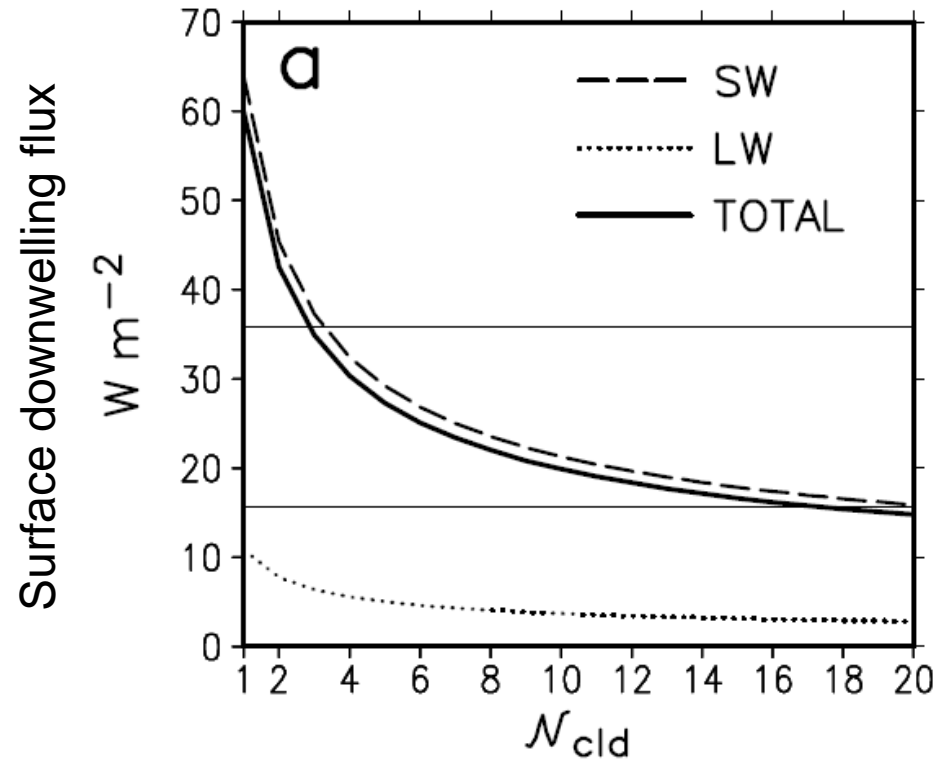
randomly assigning a different cloud profile for each spectral-point from the distribution of M profiles created by a cloud generator

Advantages of McICA

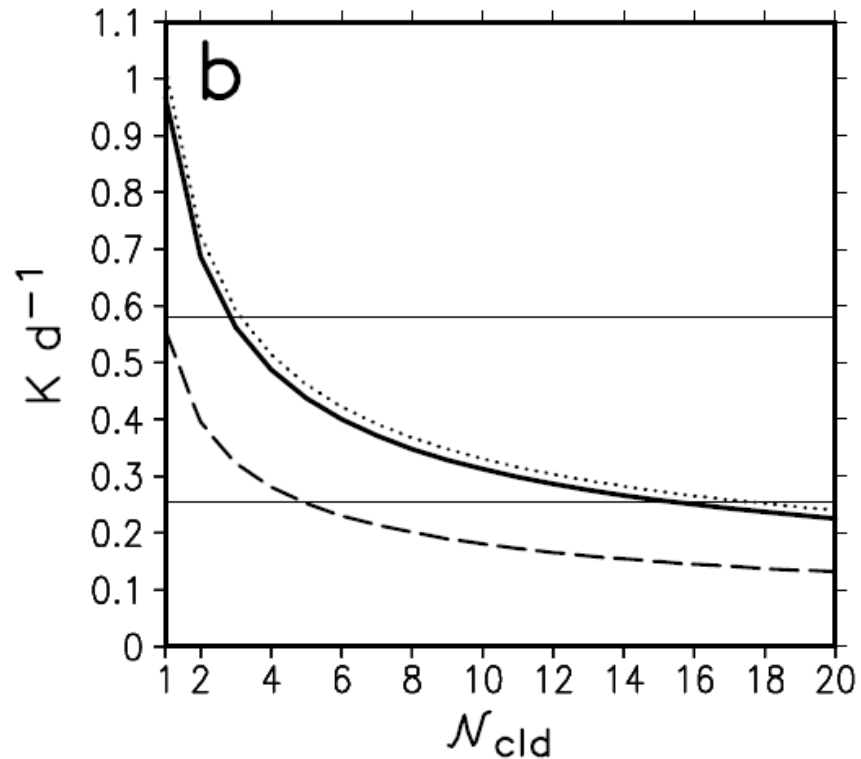
- Approximates a full 'ICA' calculation in an intuitive way
- Each sub-column is fast to compute: cloud fraction is either 1 or 0
- Easy to implement different overlap schemes
- Easy to implement subgrid-cloud inhomogeneity scheme
- Efficient when optimized

Disadvantages of McICA (1)

- McICA is inherently noisy, particularly for LW heating rates



Number of sub-columns per g-point



Räsänen and Barker (2004)

Disadvantages of McICA (2)

- Two optimisations have been suggested:

Spatial sampling:

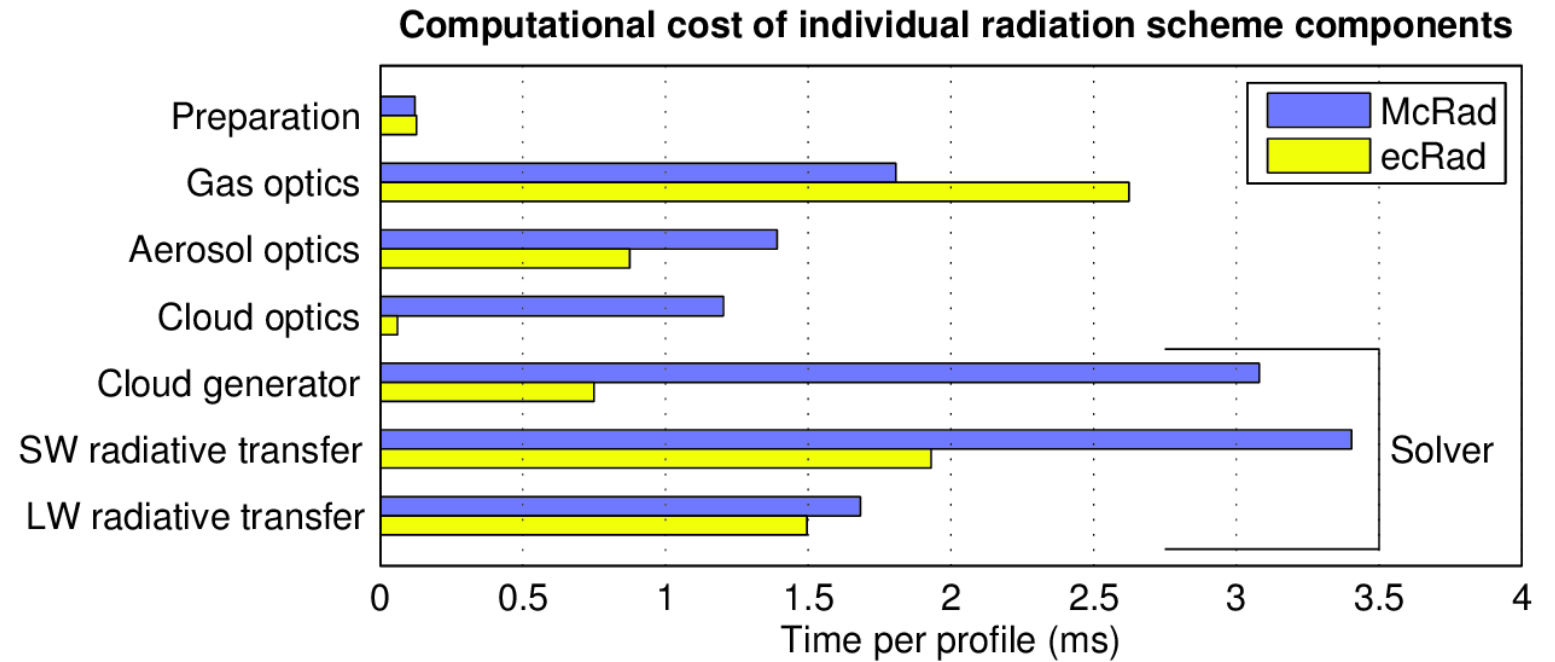
- Given that clear-sky fluxes are computed at every g-point for diagnostic purposes, only need to sample cloudy profiles
- Compute the all-sky fluxes according to the diagnosed total integrated cloud fraction

Spectral sampling:

- Some g-points have a greater contribution to cloud radiative effect than others
- Use additional sub-columns for g-points that contribute the most
- Requires prior knowledge of which g-points are most sensitive. Could be different depending on situation and what metric you want to minimize (e.g., surface or TOA?)

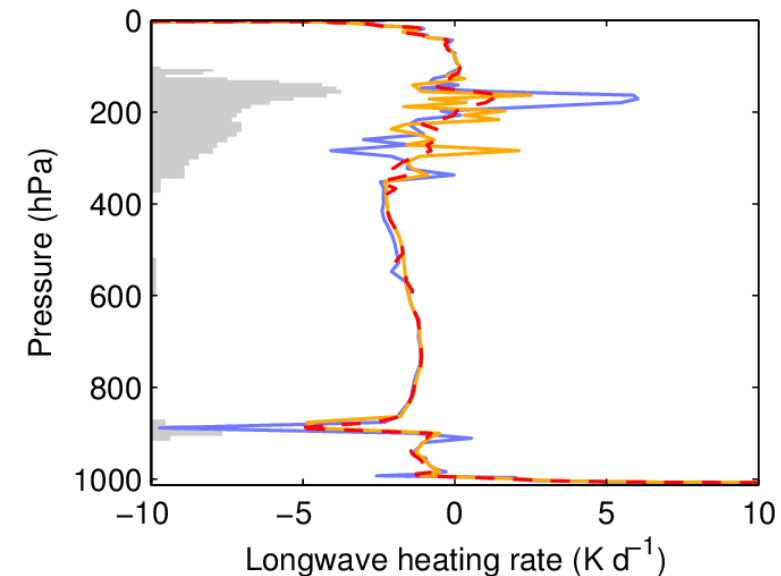
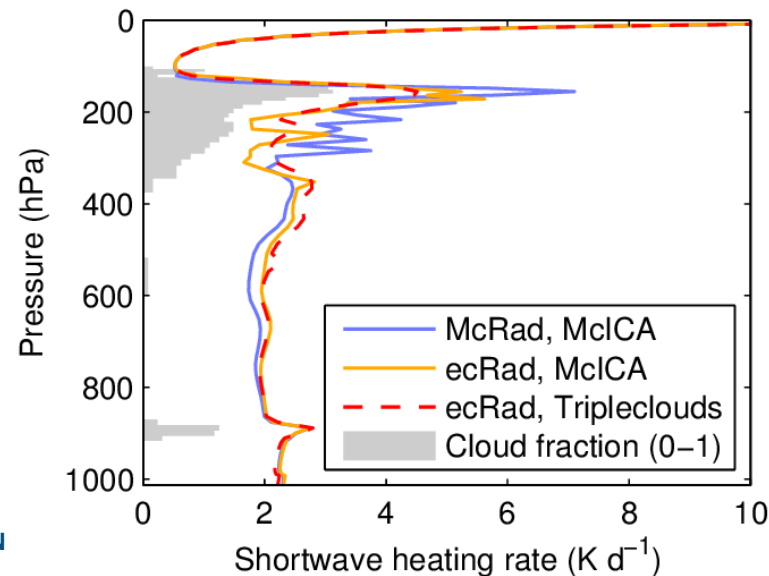
ecRad: Improved efficiency

- 31-35% faster than McRad
- Much faster treatment of cloud optics and cloud generator
- with longwave scattering by clouds alone (LWscat=1) is only 3% slower



ecRad: Improved accuracy

- As well as being much faster, reformulation of McICA scheme generates less stochastic noise
- Better LW fluxes in the stratosphere



Radiation in ECMWF: temporal and spatial resolution

- Even if optimised and much faster than before ecRad is computationally expensive
- Operationally, the RT code runs on a spatial grid coarser (6-10 times fewer gridpoints) than the rest of the physics and with a longer time step
 - HiRes 10 day forecasts: model Tco1279 (~9 km at the equator) ecRad Tco511 (~20 km).
Model time step 10 min, ecRad 1h
 - Ensemble forecasts: model Tco639 (~16 km), ecRad Tco255 (~30 km). Model time step 20min, ecRad 3h
- LW flux constant between radiation calls, SW flux adjusted for correct solar zenith angle and path length. Minor impact on diurnal cycle, larger for the 3h time step

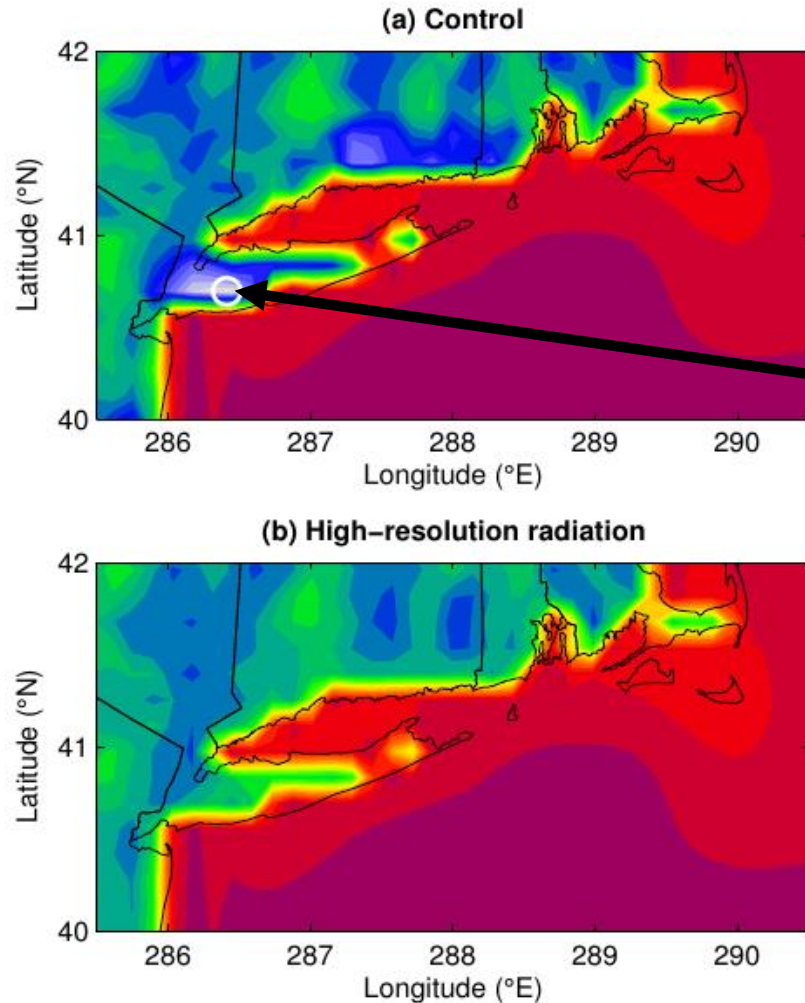
Are we using our computer time wisely?

- Temporal, spatial and spectral resolution in various global NWP models:

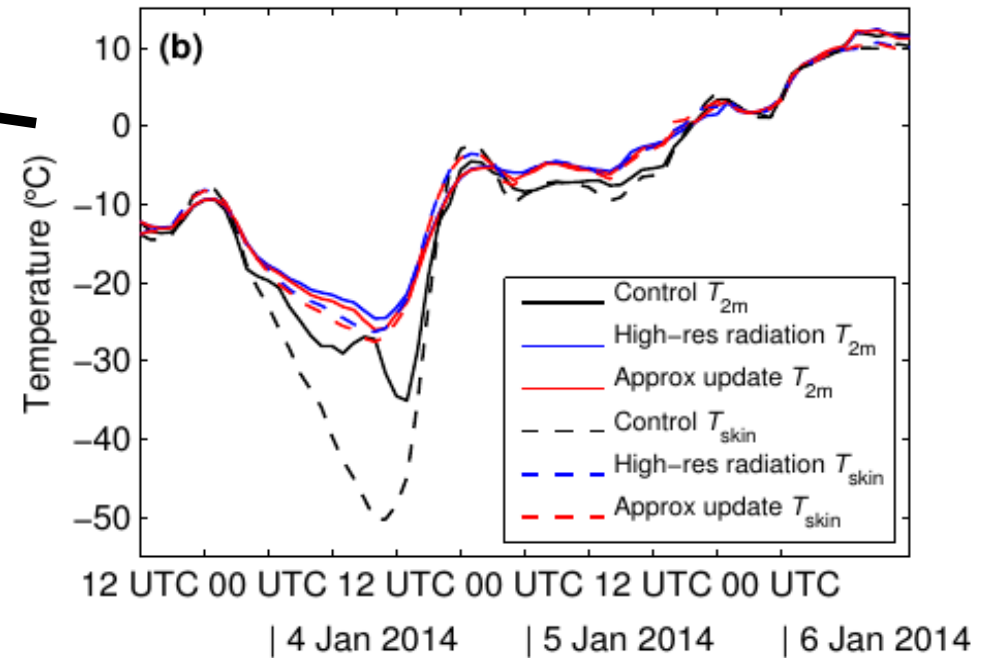
Centre	Radiation timestep (h)		Horiz. coarsening		Bands		Spectral intervals	
	HRES	ENS	HRES	ENS	SW	LW	SW	LW
ECMWF	1	3	10.24	6.25	14	16	112	140
NCEP	1	1	1	1	14	16	112	140
DWD	0.4	0.6	4	4	14	16	112	140
Météo France	1	1	1	1	6	16	–	140
Met Office	1	1	1	1	6	9	21	47
CMC	1	1	1	1	4	9	40	57
JMA	1	1 (SW), 3 (LW)	4	4	16	11	22	156
FSCK	–	–	–	–	2	1	~ 15	~ 32

- **ECMWF** has lowest spatial resolution for radiation
 - Experiments show this barely degrades forecasts (unlike 3-h radiation timestep)
- **Met Office** NWP model uses 3.7 times fewer g-points than RRTM-G
- **Full-spectrum correlated-k** estimates of coarsest possible spectral resolution

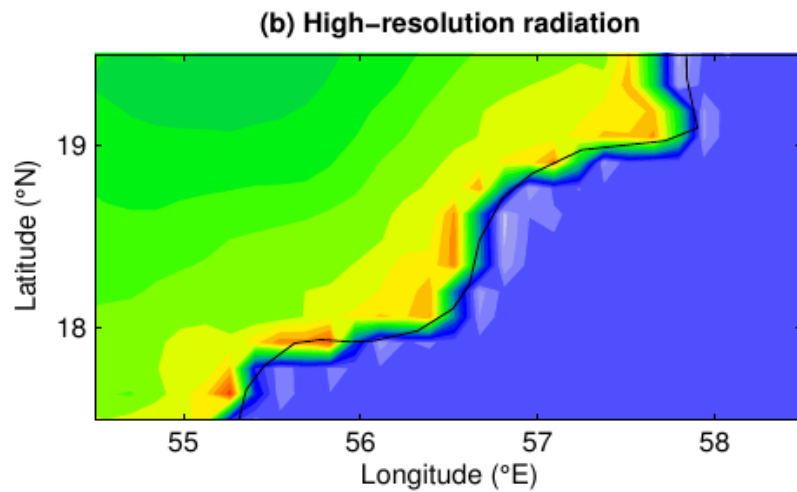
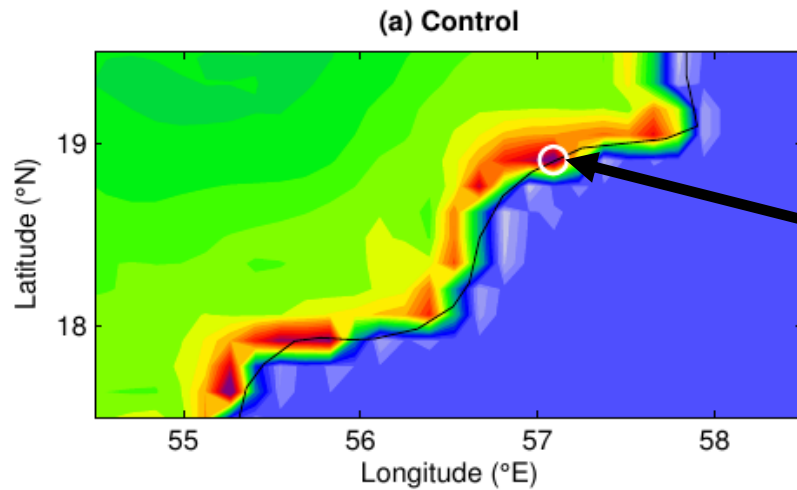
Mitigating extreme temperature errors at coastlines a cold case



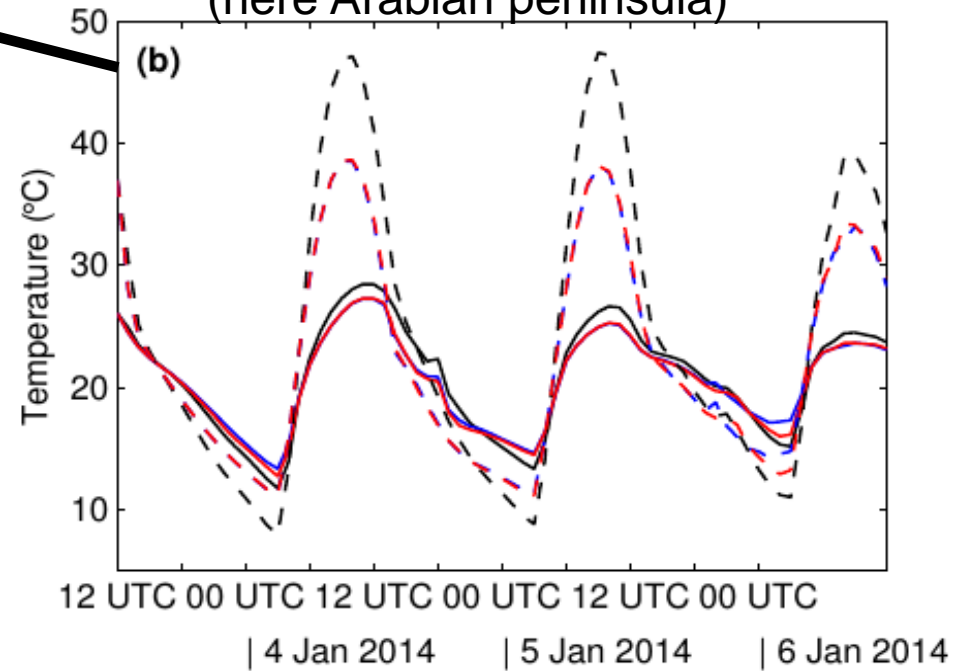
Approximate solution of the surface energy balance taking into account local albedo and changes in the skin temperature reduces the impact of the coarse radiation grid and time step



Mitigating extreme temperature errors at coastlines a warm case

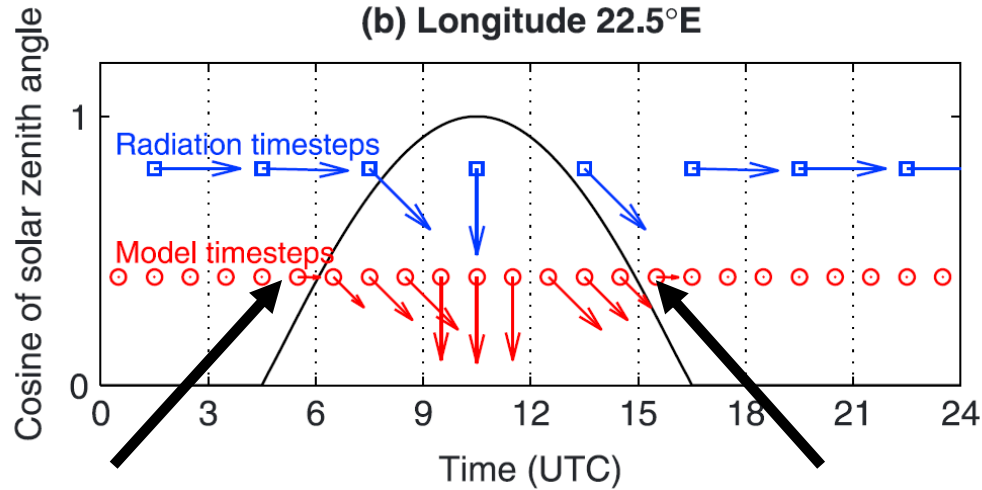


Extreme temperatures at the boundary between desert areas and ocean (here Arabian peninsula)



Hogan&Bozzo 2015

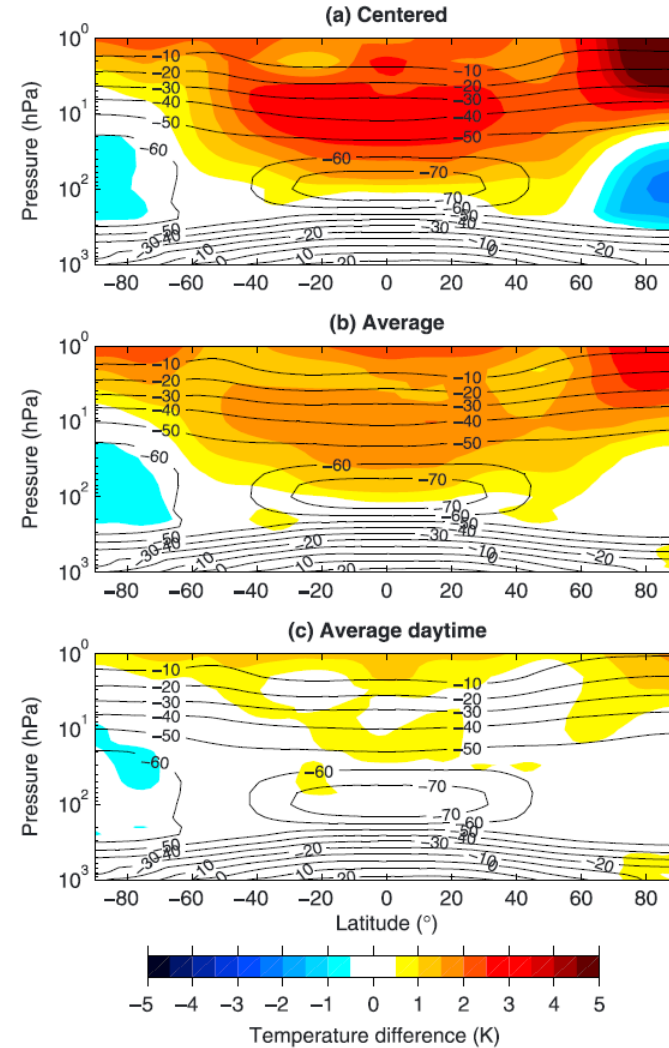
Accounting for solar zenith angle between time steps



SZA averaged between time steps: too shallow Sun's elevation at dusk and dawn -> too large path length -> too large O3 absorption in the stratosphere.

Fix: computing the averaged SZA between time steps but only for the sun lit portion

Larger errors for radiation time step ≥ 3 h



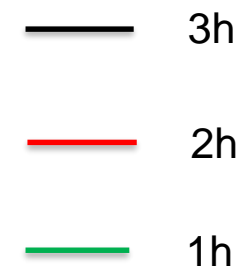
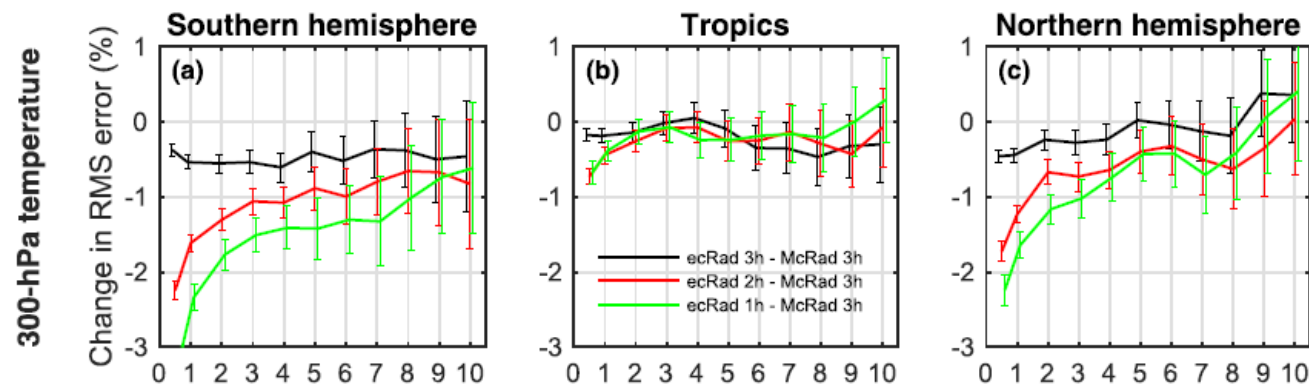
Hogan and Hirahara, 2015

Impacts of changing a radiation scheme

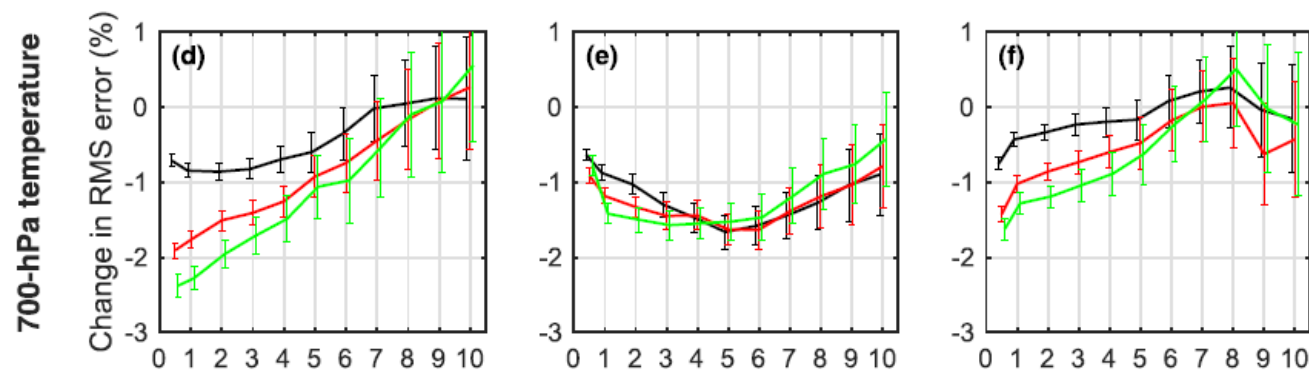
- Forecast skill
 - Set of 10-days forecasts generally over summer and winter seasons
 - Evaluation against analysis and observations
 - Impact on forecast skill scores at different lead times
- Mean model climate
 - Set of small ensembles of 1-year long forecast
 - Good to evaluate mean model biases
 - Coupled and un-coupled experiments

Impact of radiation timestep on forecast skill scores

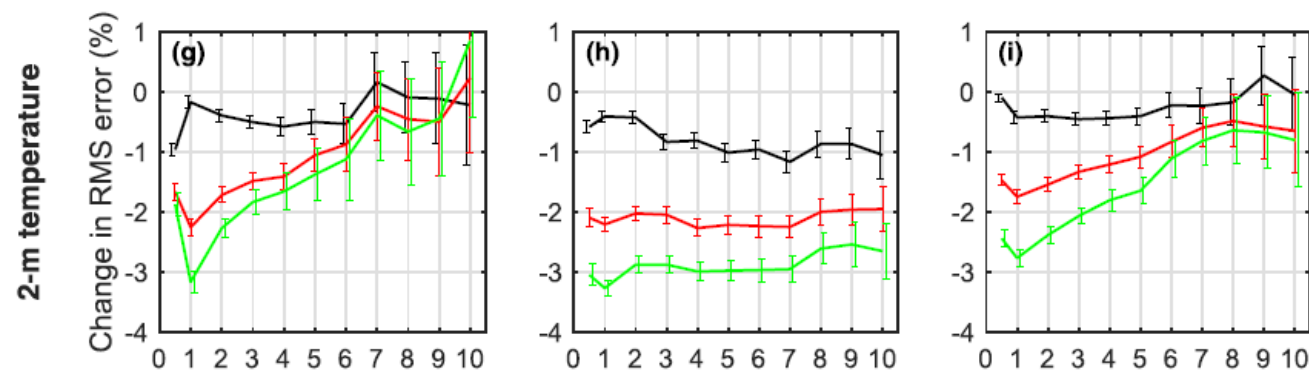
300 hPa temperature



700 hPa temperature



2-m temperature



Normalised change in RMSE using ecRad with respect to the old McRad scheme

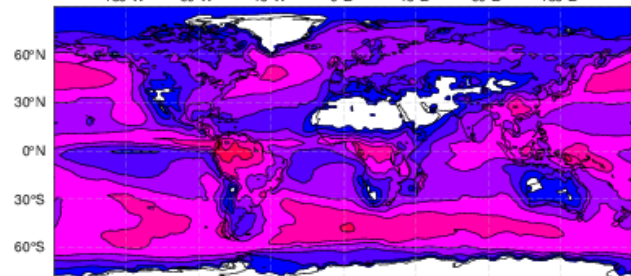
Values <0 mean that the IFS performs better using ecRad



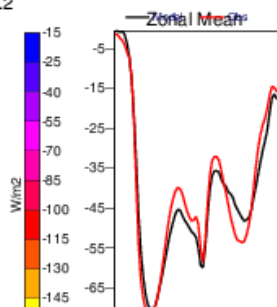
Current radiation biases in ECMWF model (46r1) “climate” runs

TOA SWCF Radiation

TOA swcf h3le Sep 2000-2004 nmon=12 nens=4 Mean: -47.1 50S-50N Mean: -49.2

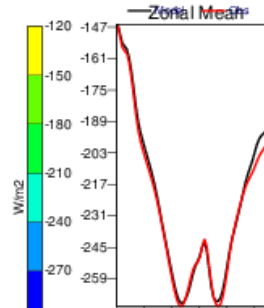
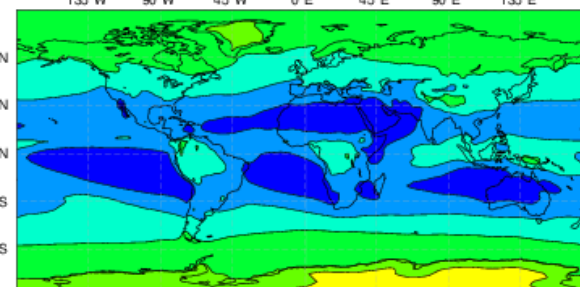


Model



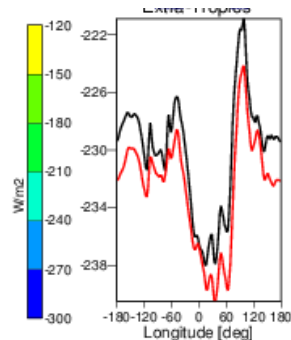
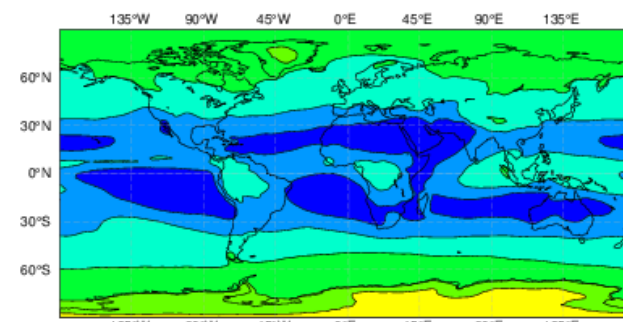
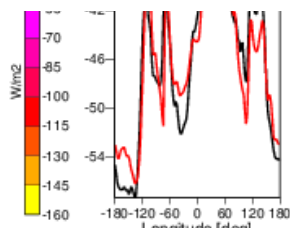
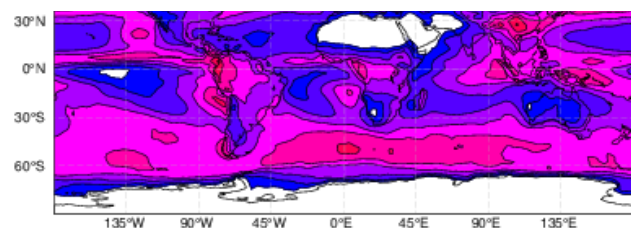
Outgoing LW Radiation

TOA lw h3le Sep 2000-2004 nmon=12 nens=4 Mean: -239 50S-50N Mean: -251



Too much cloudiness over tropical oceans
 Too much reflection at TOA, too little downward SW radiation at the ocean surface. Not enough reflection in stratocumulus regions and close to Antarctica

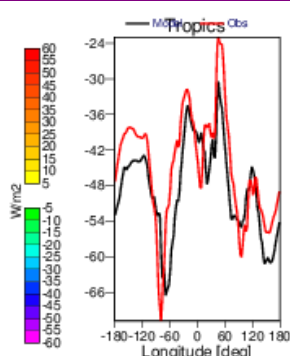
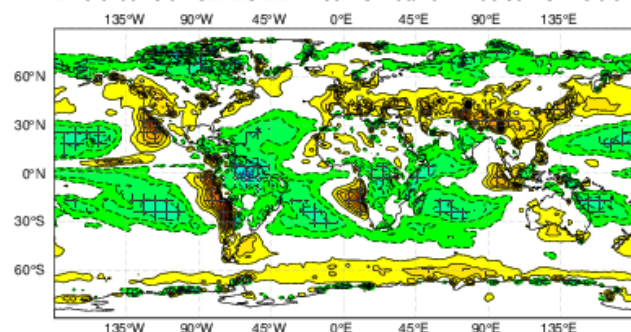
CERES



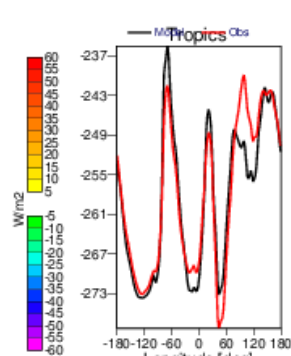
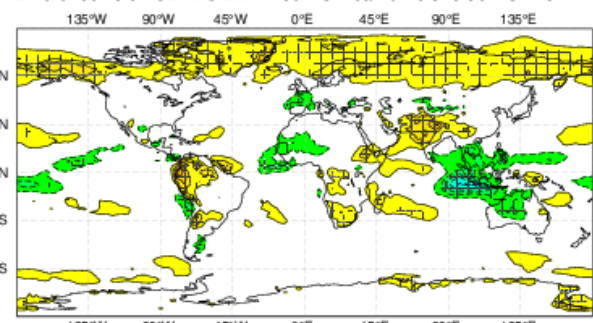
Cold bias in Arctic, corrected previous errors over continents (too much OLR)

Model-CERES

Difference h3le - CERES-EBAF 50N-S Mean err -1.69 50N-S rms 9.81

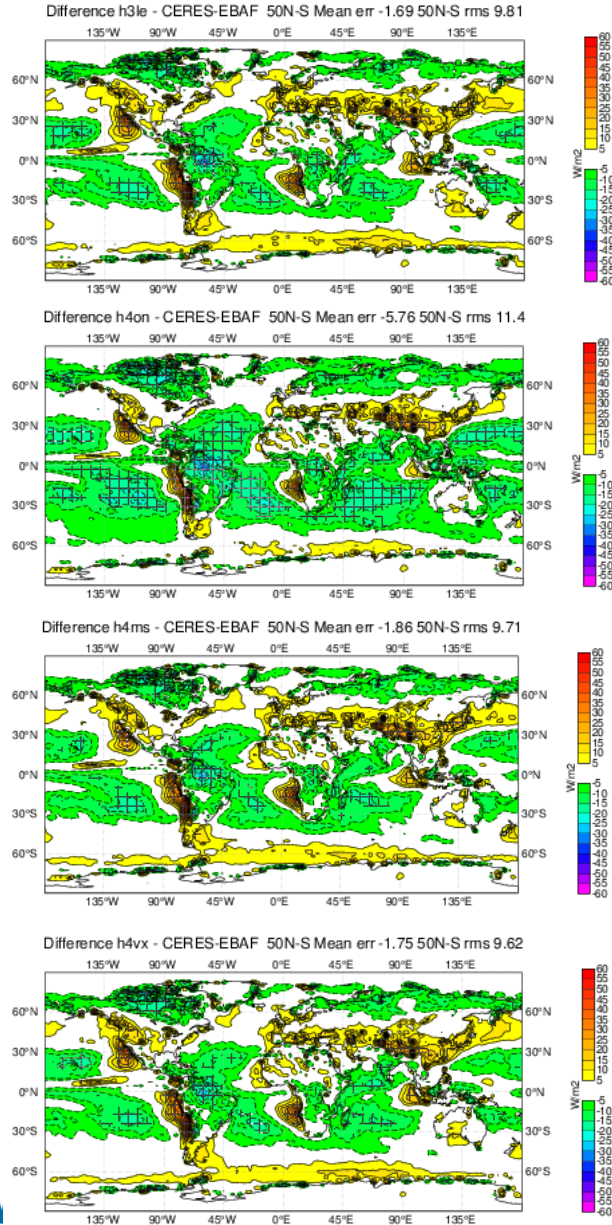


Difference h3le - CERES-EBAF 50N-S Mean err 0.543 50N-S rms 4.74

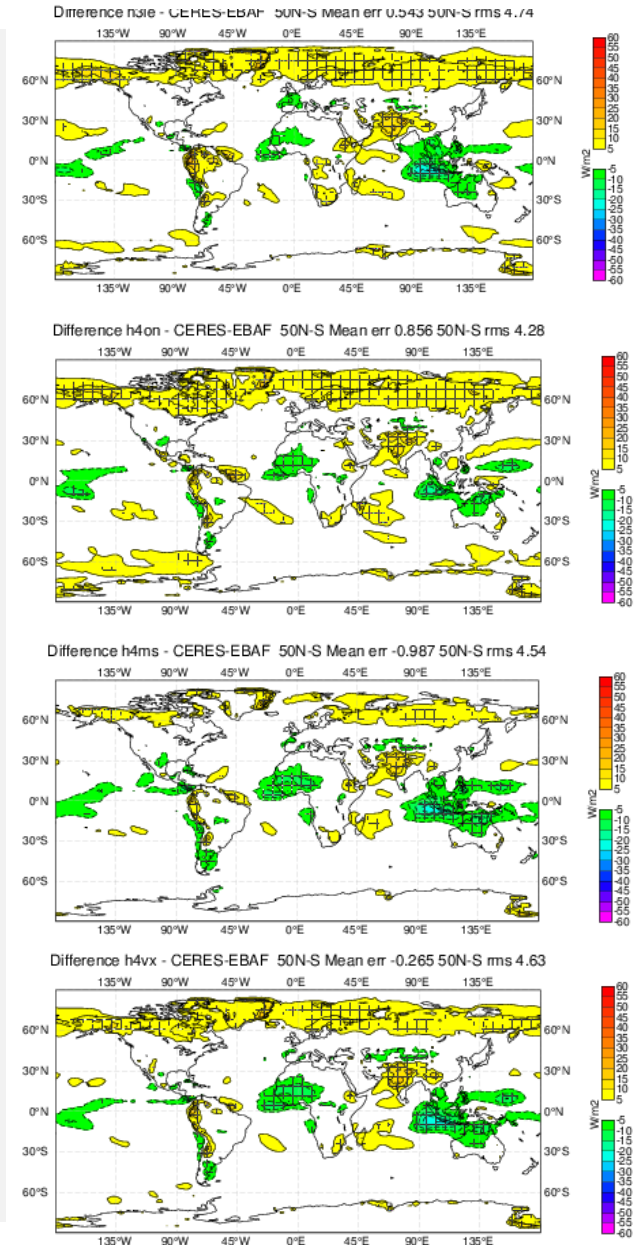


Higher or lower? Impacts of cloud radiative property changes

TOA net SW Radiation (down)



Outgoing LW Radiation



Control: CY46r1

Liquid cloud effective radius x 0.8

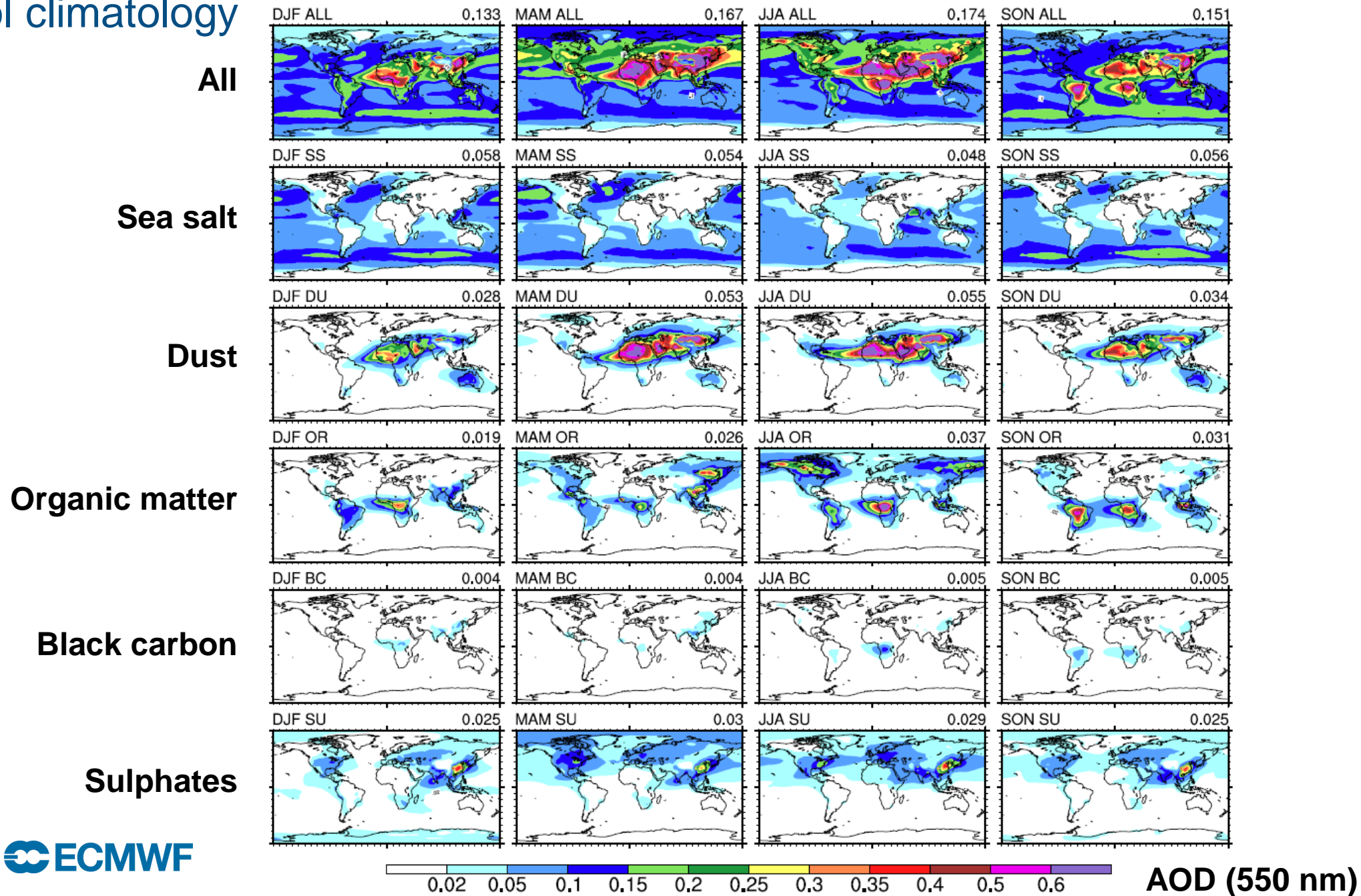
LW scattering OFF

Yi scattering

Aerosols in NWP

- While aerosols play an important role in determining climate, their day to day variability is probably of secondary importance for NWP (e.g., prediction of temperature, winds).
- However, the radiative effect of aerosols is important to include as it can be significant, particularly for absorbing aerosols.
- Within IFS, CAMS monthly climatologies are used to account for direct effects.
- Indirect effects of aerosols are partially accounted for in cloud scheme (e.g., parameterizations of N_d from wind speed or land/sea mask).

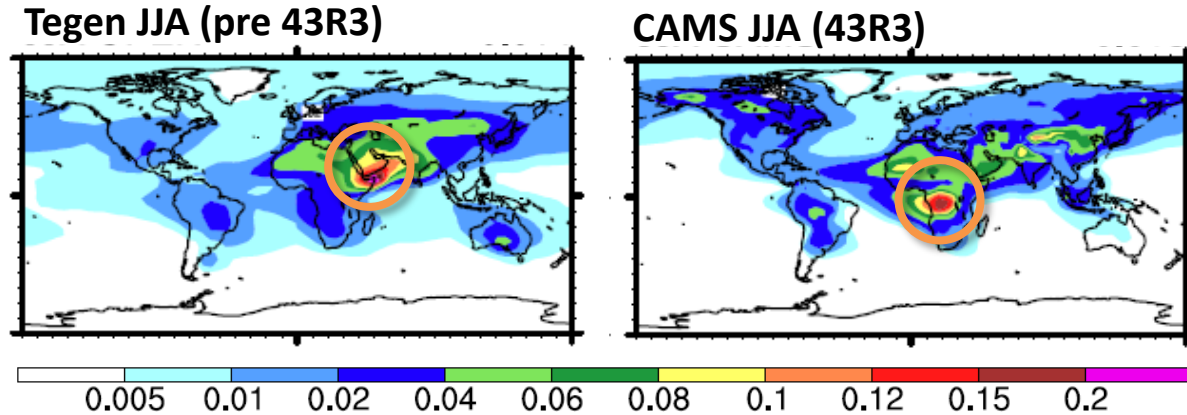
CAMS aerosol climatology



Aerosols

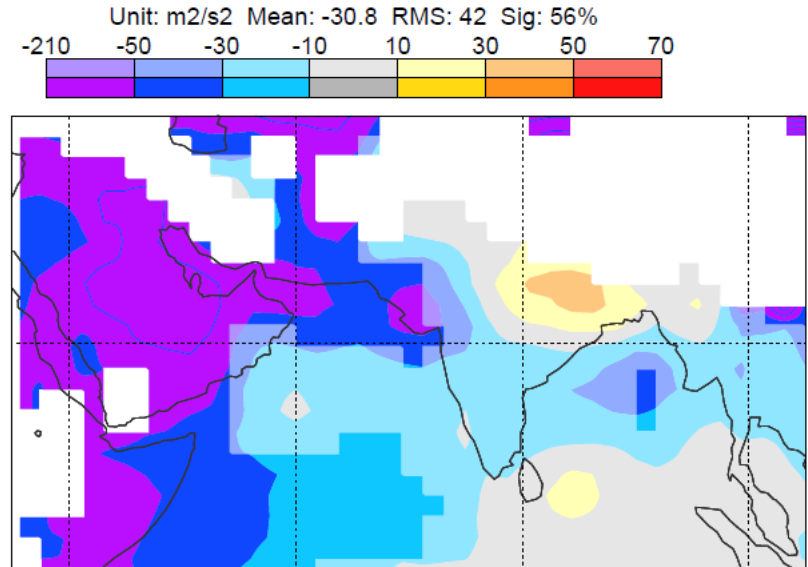
Bozzo et al. (2017)

- Atmospheric forcing depends on *absorption* optical depth:

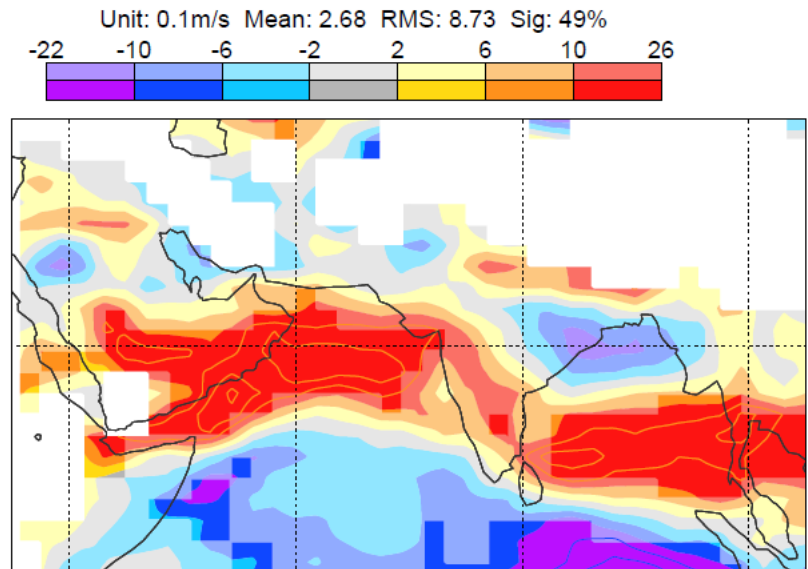


- Reduced absorption over Arabia in new CAMS climatology weakens the overactive Indian Summer Monsoon, halving the overestimate in monsoon rainfall
- Increased absorption over Africa degraded 850-hPa temperature, traced to excessive biomass burning in CAMS
- *We can measure the impact of aerosols on the tropical atmosphere more easily than the absorption optical depth itself! Use to provide information on aerosol errors?*

(a) Tegen climatology: geopotential *bias*



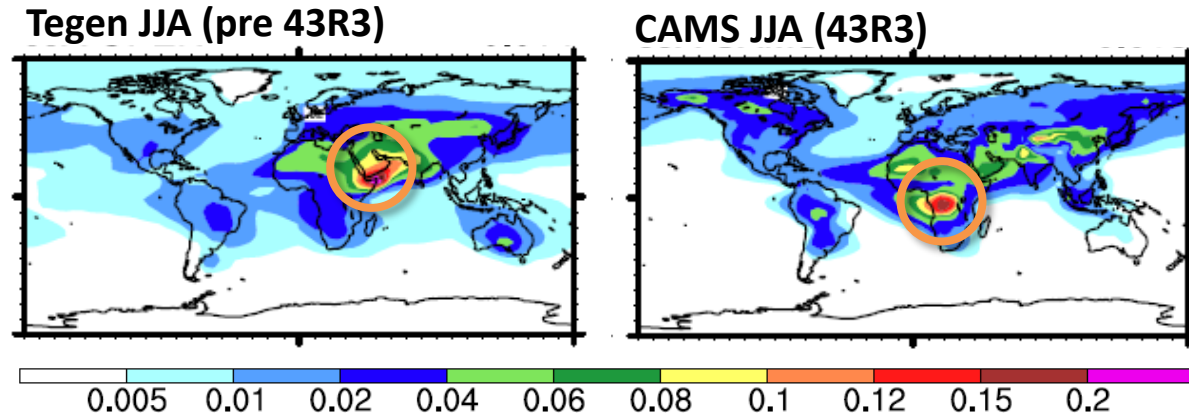
(c) Tegen climatology: zonal wind *bias*



Aerosols

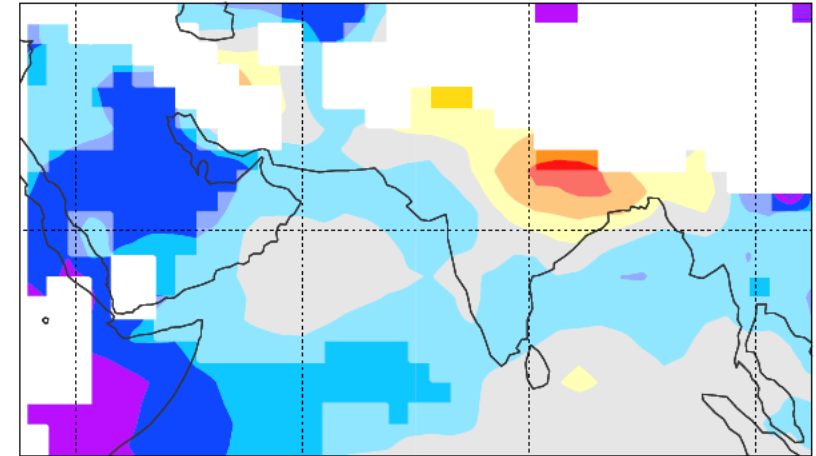
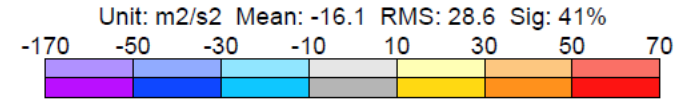
Bozzo et al. (2017)

- Atmospheric forcing depends on *absorption* optical depth:

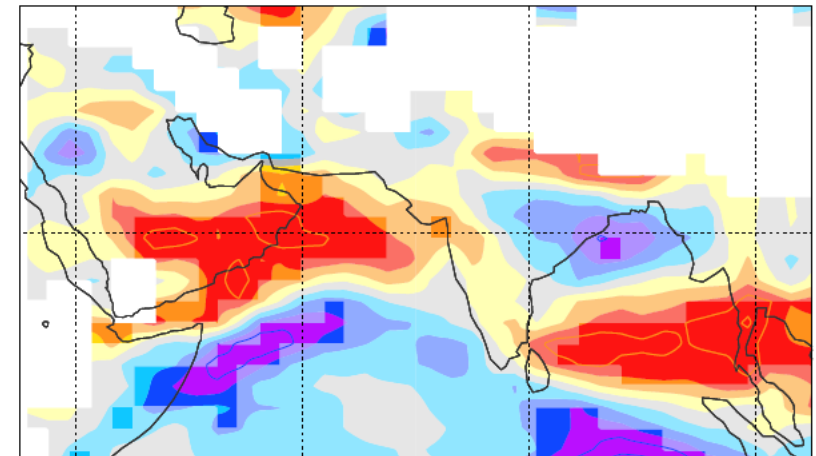
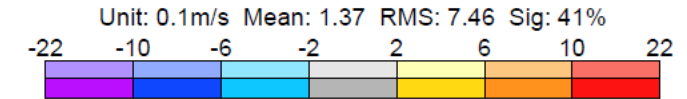


- Reduced absorption over Arabia in new CAM5 climatology weakens the overactive Indian Summer Monsoon, halving the overestimate in monsoon rainfall
- Increased absorption over Africa degraded 850-hPa temperature, traced to excessive biomass burning in CAM5
- *We can measure the impact of aerosols on the tropical atmosphere more easily than the absorption optical depth itself! Use to provide information on aerosol errors?*

(b) CAM5 climatology: geopotential *bias*



(d) CAM5 climatology: zonal wind *bias*



Five “Grand Challenges” for radiation in NWP models

Solar spectrum

Water vapour biases

Middle atmosphere

Ozone

Code optimization

GPUs

Efficiency

Spatial/temporal/spectral resolution

Clouds

Overlap

Sub-grid heterogeneity

3D effects

Particle size

Water vapour continuum

Clear-sky absorption

Longwave scattering

Optical properties

Aerosols

Sea emissivity

Snow albedo

Forests

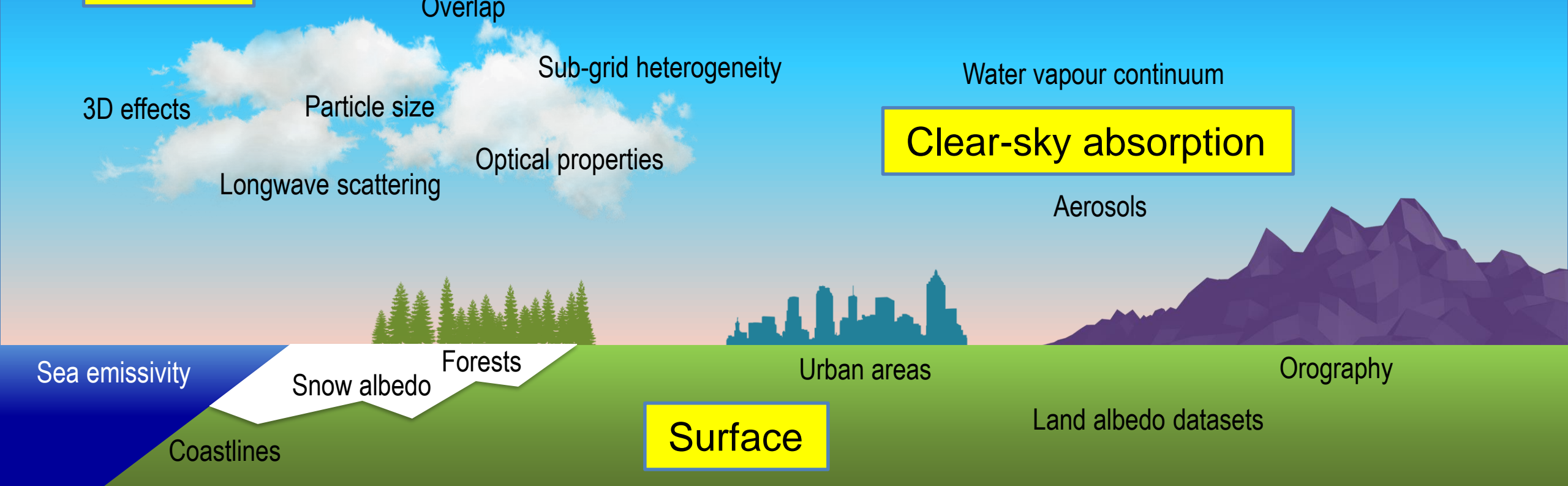
Urban areas

Orography

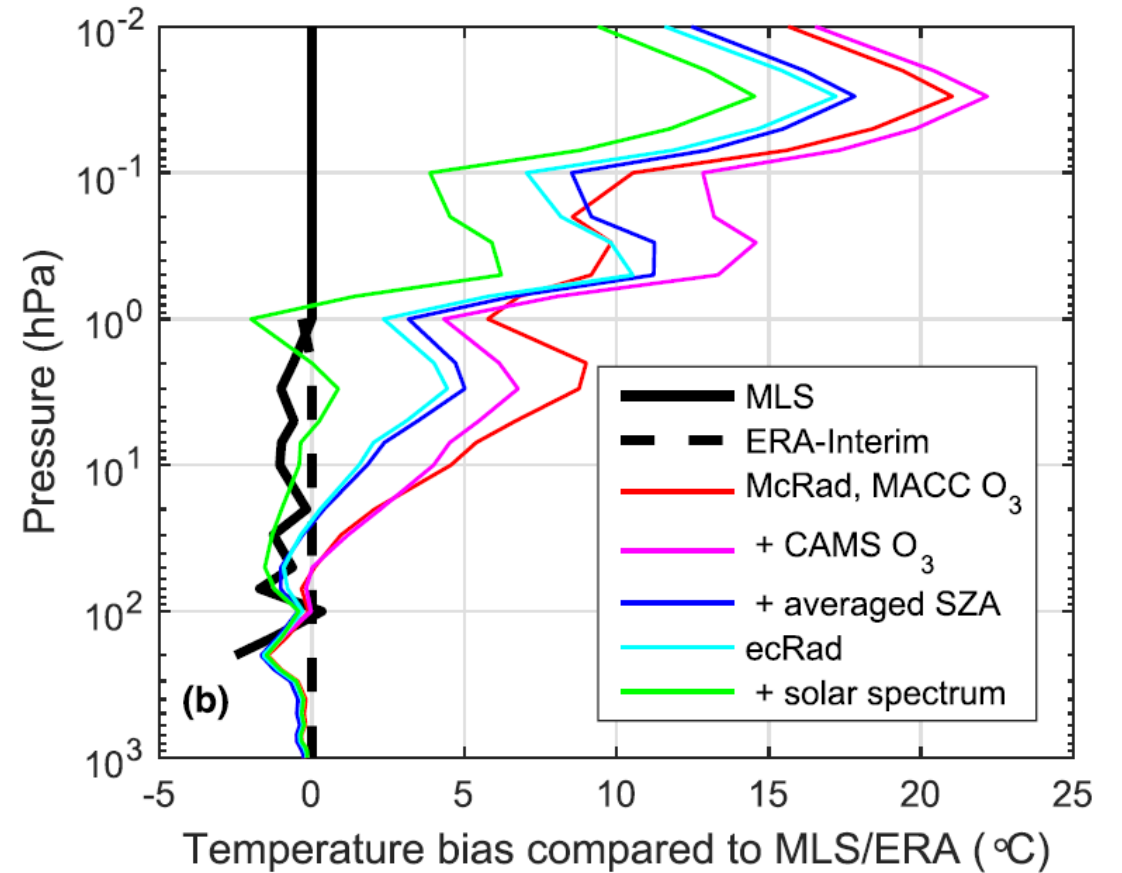
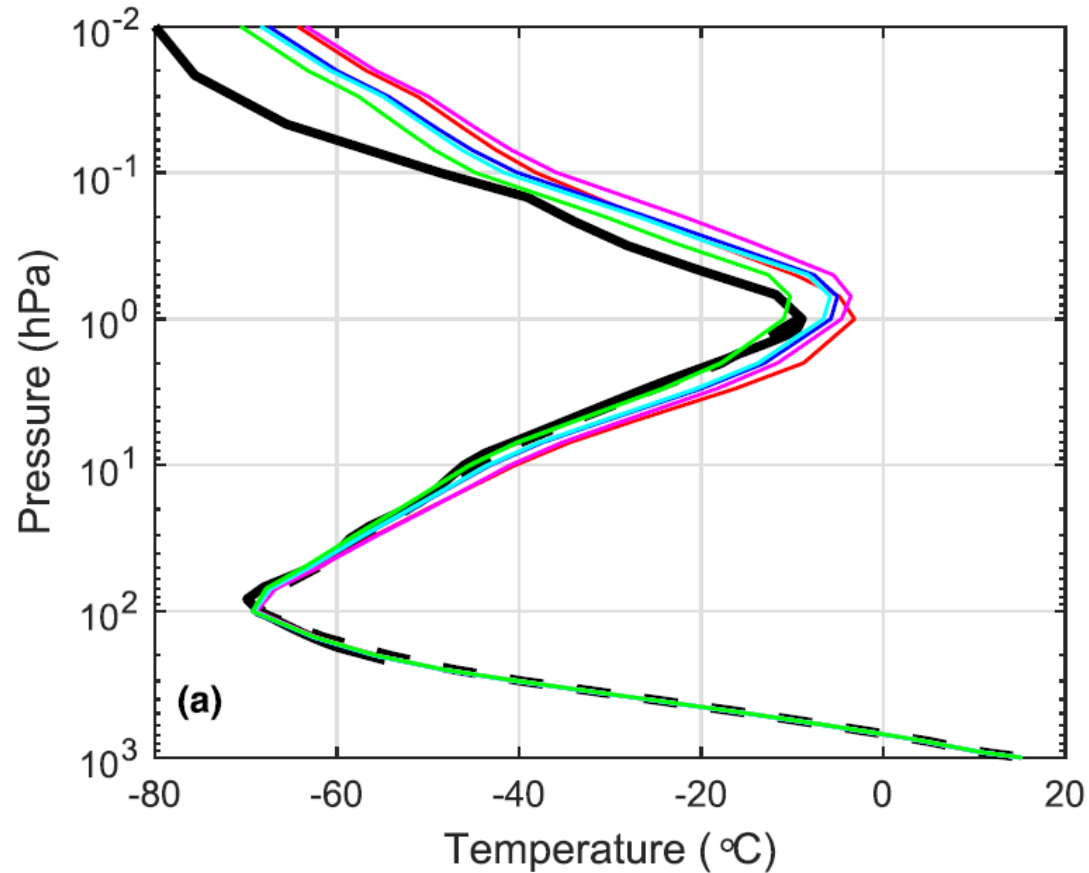
Coastlines

Surface

Land albedo datasets



Improving the middle atmosphere in the IFS



Hogan and Bozzo (2018)

Summary and outlook

- Global tropospheric climate of the IFS is excellent, but need concerted effort on many fronts to tackle much larger regional and stratospheric biases
- New ecRad scheme is good platform for future developments, but interaction and consistency between schemes is also very important
- Intriguing impacts of radiative heating on predictive skill: water vapour and stratosphere-troposphere coupling, and aerosols and monsoon systems
- Five main Grand Challenges in the coming years:
 1. Overhaul surface treatment, including 3D interactions with cities and forests
 2. Package of physically-based improvements to clouds
 3. Role of aerosols in predictability; upgrade water vapour continuum
 4. Remove middle-atmosphere temperature bias via new UV solar spectrum
 5. Much more efficient gas optics and spectral integration

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Radiation in numerical weather prediction

Robin J. Hogan, Maike Ahlgrimm, Gianpaolo Balsamo, Anton Beljaars, Paul Berrisford, Alessio Bozzo, Francesca Di Giuseppe, Richard M. Forbes, Thomas Haiden, Simon Lang, Michael Mayer, Inna Polichtchouk, Irina Sandu, Frederic Vitart and Nils Wedi

Research, Forecast and Copernicus Departments

Paper to the 46th Science Advisory Committee, 9–11 October 2017

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Europäisches Zentrum für mittelfristige Wettervorhersage
Centre européen pour les prévisions météorologiques à moyen terme

TECHNICAL MEMORANDUM

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ECRAD: A new radiation scheme for the IFS

Robin J. Hogan and Alessio Bozzo

Research Department

November 2016

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