

Seasonal Forecasting at ECMWF

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Sources of seasonal predictability

– KNOWN TO BE IMPORTANT:

- El Nino variability - biggest single signal
- Other tropical ocean SST - important, but multifarious
- Climate change - impact is substantial, especially in temperature forecasts, and must be accounted for
- Local land surface conditions - e.g. soil moisture in spring: dry soil warms up more quickly and is more prone to drought

– OTHER FACTORS:

- Volcanic eruptions - definitely important for large events, gives global cooling plus sometimes a winter warming in parts of the northern hemisphere
- Mid-latitude ocean temperatures - Complicated story, could be important in some regions
- Remote soil moisture/ snow cover - Unclear how large the effects might be
- Sea ice anomalies - definitely local effects, possibly weaker remote impacts
- Dynamic memory of atmosphere - most likely for first 1-2 months
- Stratospheric influences - polar vortex, solar cycle, QBO, ozone, ...

– Unknown or Unexpected - ???

Methods of seasonal forecasting

- Empirical forecasting
 - Use past observational record and statistical methods
 - Works with reality instead of error-prone numerical models 😊
 - Limited number of past cases means that it works best when observed variability is dominated by a single source of predictability 😞
 - A non-stationary climate is problematic 😞
- Two-tier forecast systems
 - First predict SST anomalies (ENSO or global; dynamical or statistical)
 - Use ensemble of atmosphere GCMs to predict global response
 - Some people still use regression of a predicted El Nino index on a local variable of interest
- Single-tier GCM forecasts
 - Include comprehensive range of sources of predictability 😊
 - Predict joint evolution of SST and atmosphere flow 😊
 - Includes indeterminacy of future SST, important for prob. Forecasts 😊
 - Model errors are an issue! 😞

2. How to make a numerical seasonal forecast

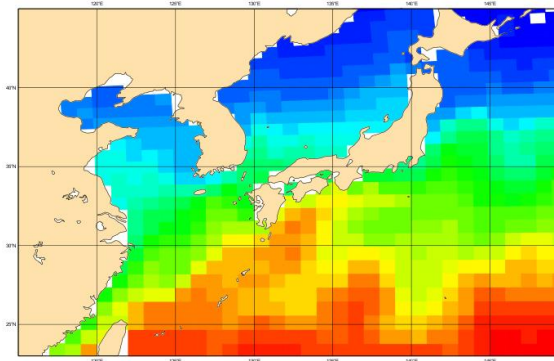
- Step 1: Build a comprehensive model
- Step 2: Create initial conditions
- Step 3: Run an ensemble forecast
- Step 4: Calibrate the output

Step 1: Build a comprehensive model

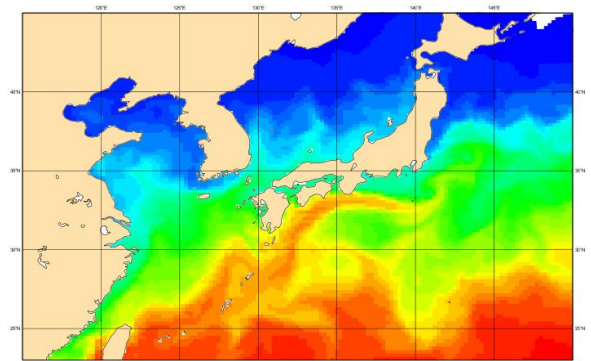
- Aim to have the best model possible, given the amount of computing resource available.
- Include relevant sources of predictability
- ECMWF forecasting model SEAS5, introduced in Nov 2017, includes:
 - IFS (atmosphere and land surface)
 - T_{CO}319L91 Cy43r1, 36km grid for physics, full stratosphere
 - All of the physical and dynamic processes of a world-class NWP model
 - Land surface model, multiple soil layers, different soil types, different vegetation types, snow, glaciers
 - Lake model, variable depths, variable mixed layer, surface and bottom temperatures, lake ice
 - Time varying tropospheric sulphate aerosol and stratospheric aerosol from volcanoes
 - Wave model
 - Ocean surface waves modify the interaction between ocean and atmosphere. Runs at 0.5 deg resolution.
 - NEMO (ocean)
 - Global ocean model, 0.25 deg resolution (eddy permitting), 75 vertical levels
 - LIM (sea-ice)
 - Single category ice, solved on same grid as ocean model

SEAS5 – ocean component

- Ocean model resolution upgraded from previous 1x1 deg to 0.25x0.25 deg
- Ocean vertical resolution improved from previous 42 levels to 75 levels
- High ocean resolution is needed to represent ocean eddies, and to better resolve the boundary currents that are important in the ocean, such as the Kuroshio in the Pacific (shown here) and the Gulf Stream in the Atlantic.

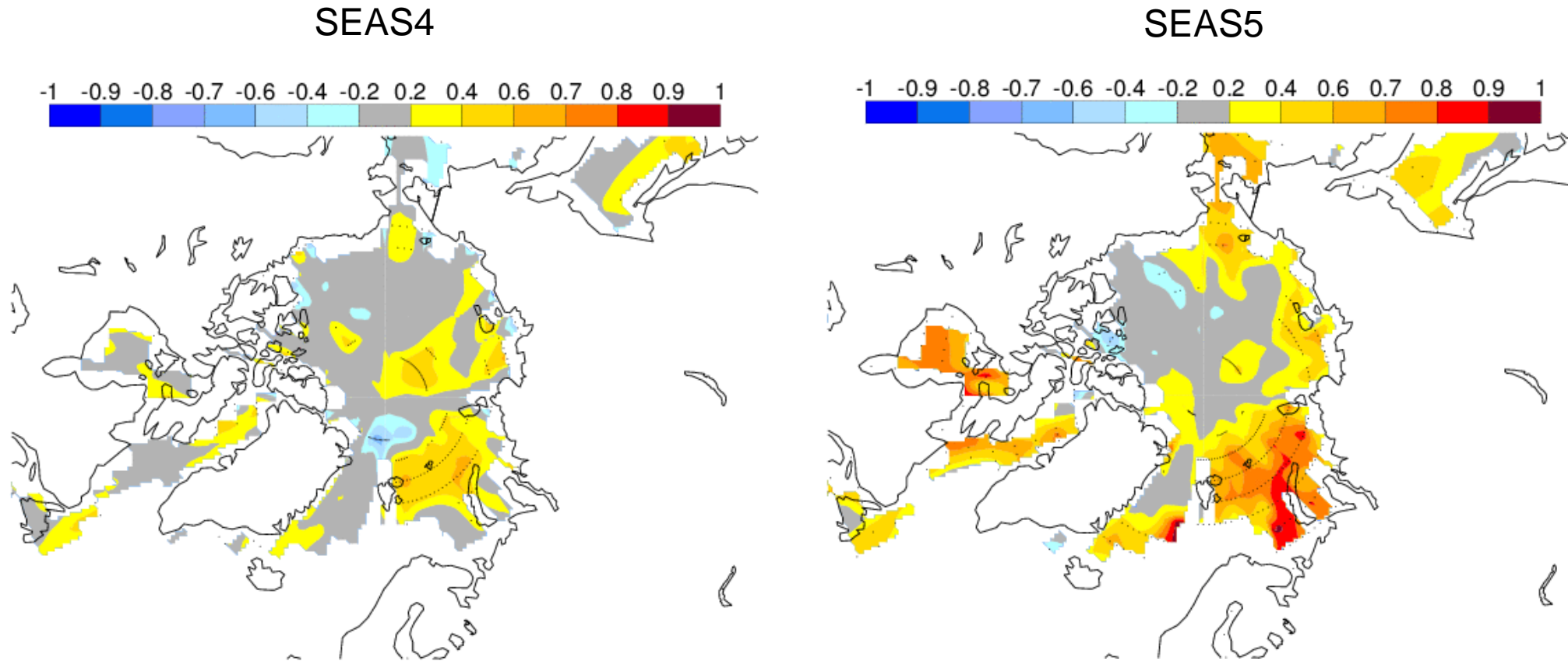


NEMO ORCA1 Z42



NEMO ORCA025 Z75

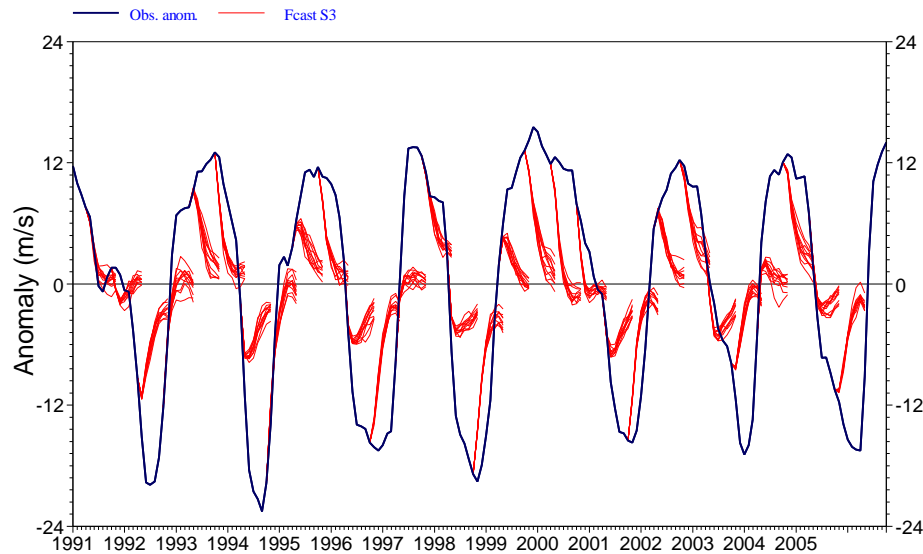
SEAS5 - Sea ice model



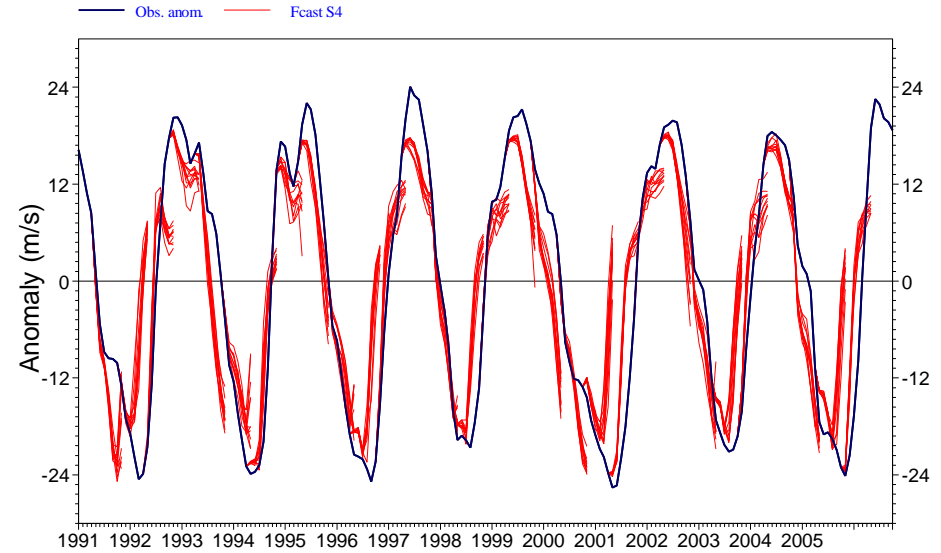
Sea ice cover predictability is improved when we include the interactive sea ice model, illustrated here with correlation scores for predictions of DJF sea ice cover.

Stratosphere - the QBO

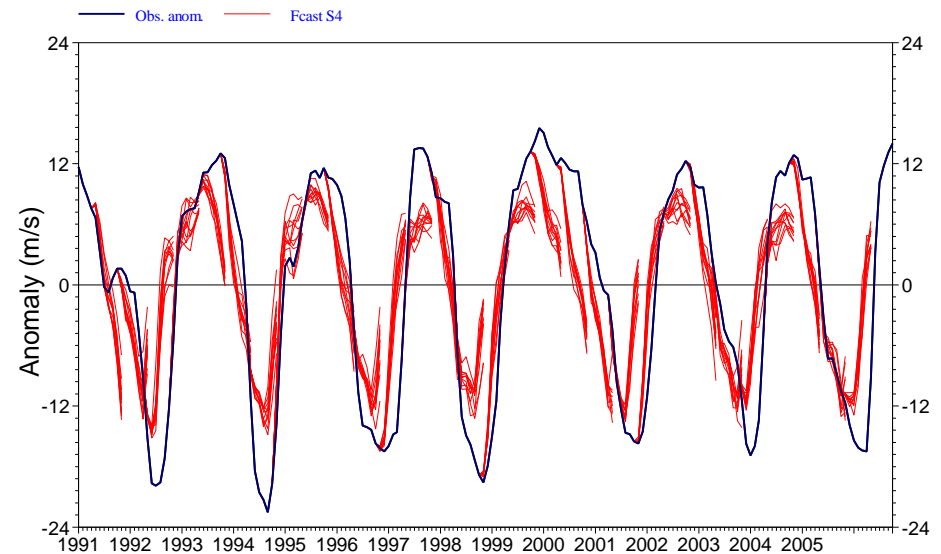
System 3



System 4



30hPa

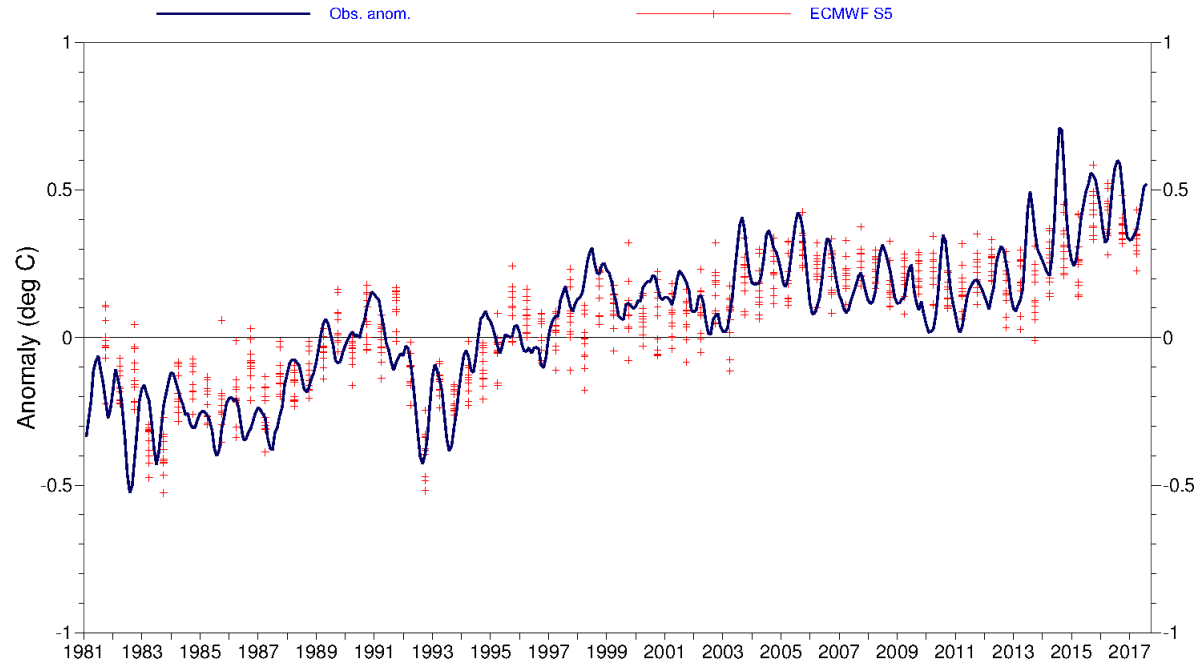


50hPa

Trends

NEXTR SST forecast anomalies

Bias corrected forecasts, mean for months 5-7, plotted at centre of verification period
Ensemble size is 11 SST obs: hd1_o12

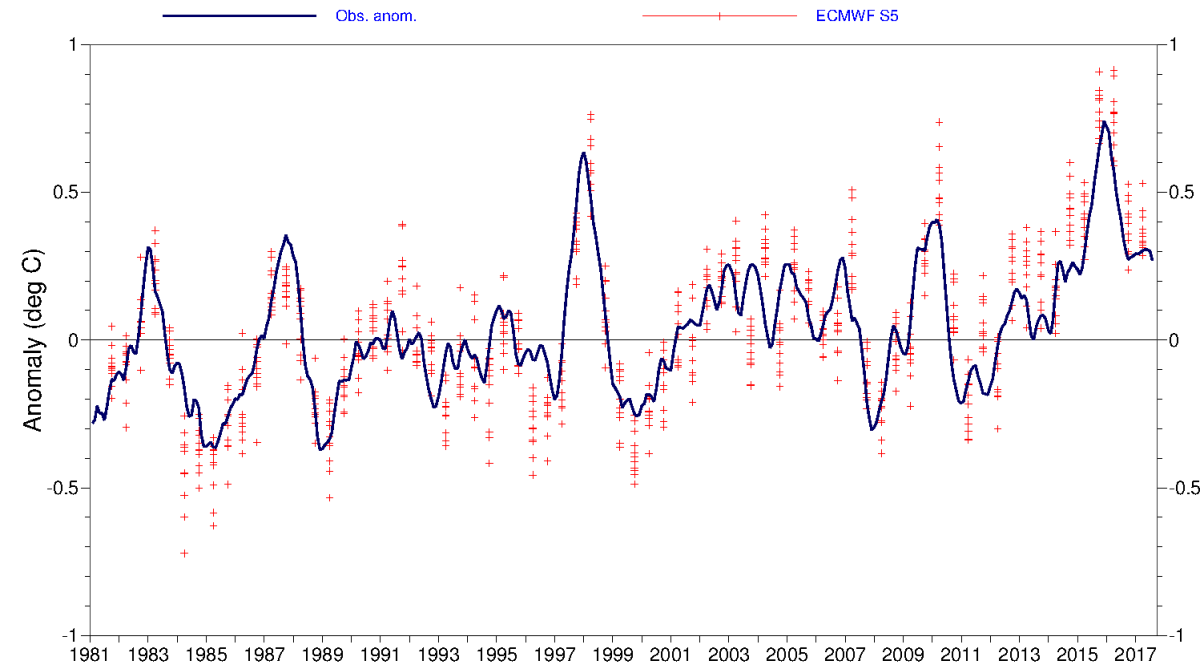


Capturing trends is important. Time-varying CO2 and other factors are important in this.

There is a strong link between seasonal prediction and decadal/ multi-decadal climate prediction.

TROPICS SST forecast anomalies

Bias corrected forecasts, mean for months 5-7, plotted at centre of verification period
Ensemble size is 11 SST obs: hd1_o12

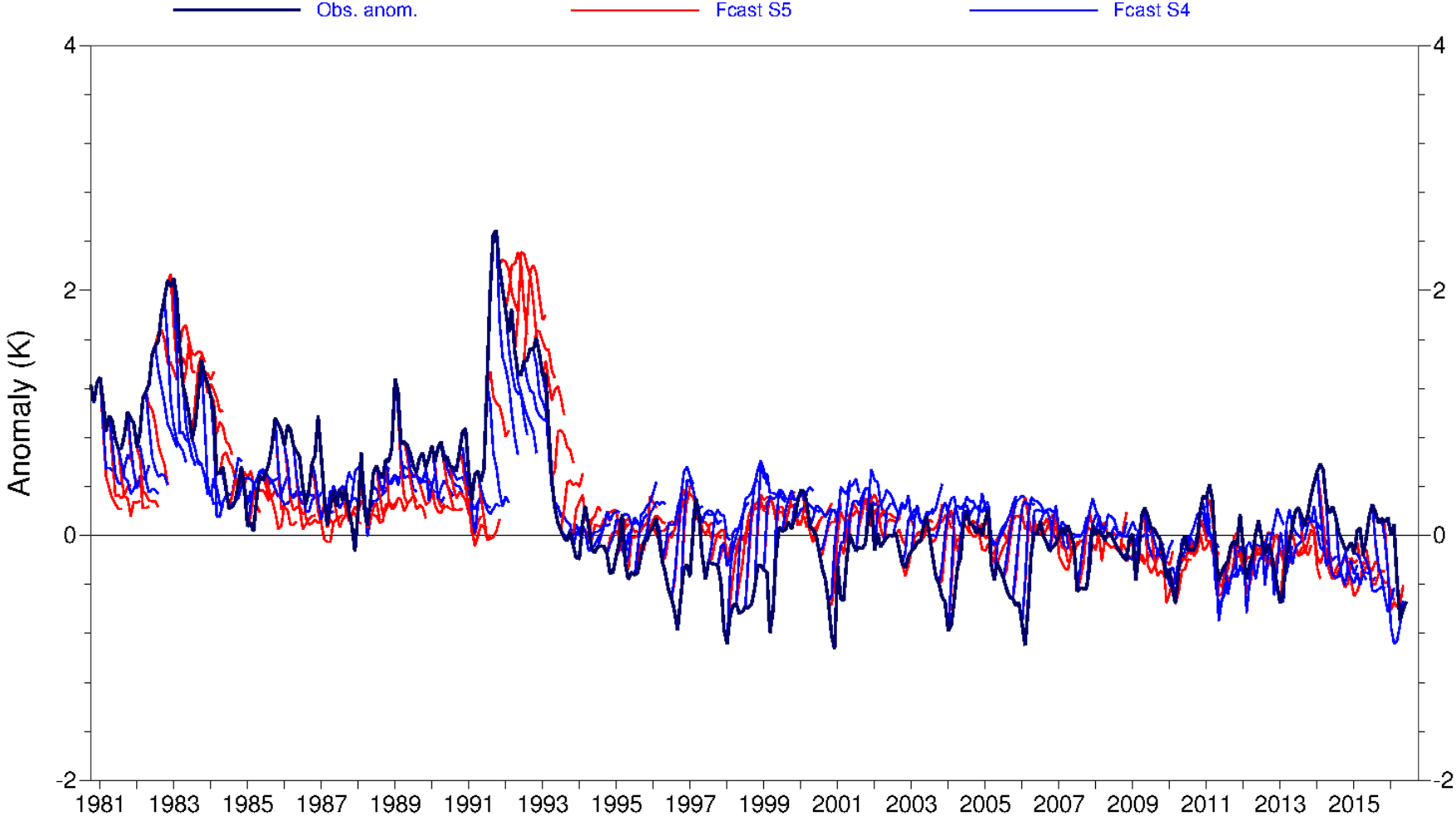


Stratosphere – volcanic aerosols

GLOBAL T50 forecast anomalies

Bias corrected forecasts at month 7

Ensemble sizes are 5 (0001) and 5 (0001) T50 obs: ec_erai

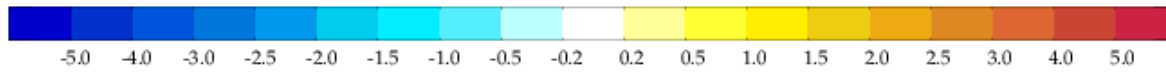
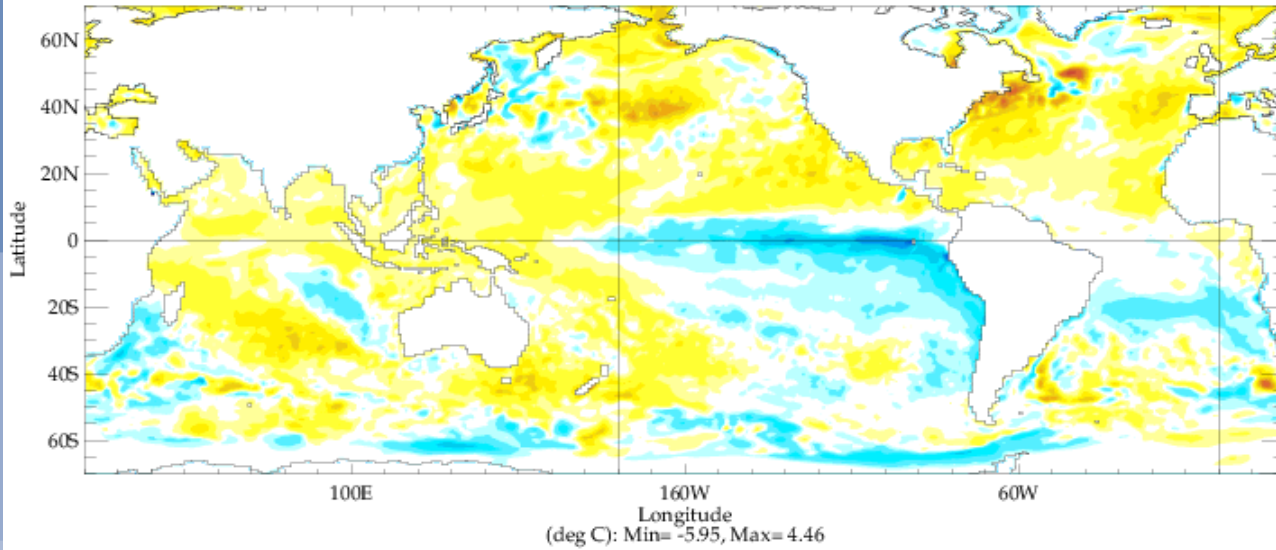


Step 2: Create initial conditions

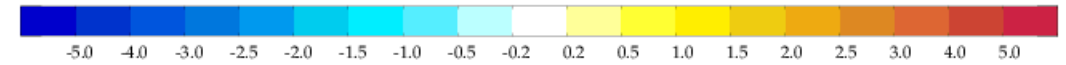
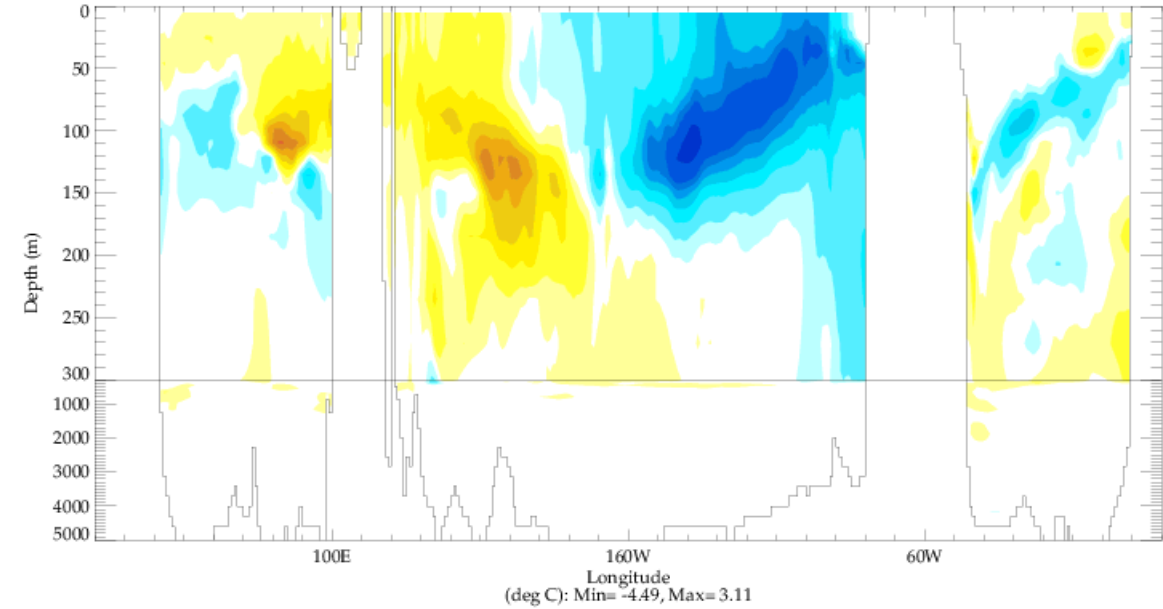
- Most of the forecast skill comes from initial conditions
 - Small amount from changes in forcing terms such as CO₂, volcanoes, solar variability
- Slowly-changing parts of the earth system are typically most important
 - Often these are hard to observe (e.g. sub-surface ocean or below the ground), so creating right initial conditions not easy
- Initial conditions use data assimilation: blending the (often very limited) observational data with model calculations to get the best estimate of the initial state
- A good set of initial conditions will include estimates of the uncertainty, so that this can be accounted for in the forecast.
 - We do this better for some components (e.g. the atmosphere) than we do for others (e.g. land surface).
- We need initial conditions in **real-time** (so that we can run today's forecast), but also for **many years in the past** to allow for calibration (see later).
 - Initial conditions need to be both accurate and consistent over time for calibration purposes. It is not easy to ensure both of these are true, when data is needed in real-time but also stretching several decades into the past. For example, the ocean is much better observed now than in the past.

Initial conditions – ocean analysis

Sea Surface Temperature
2017 November mean anomaly (1981-2009 climate)

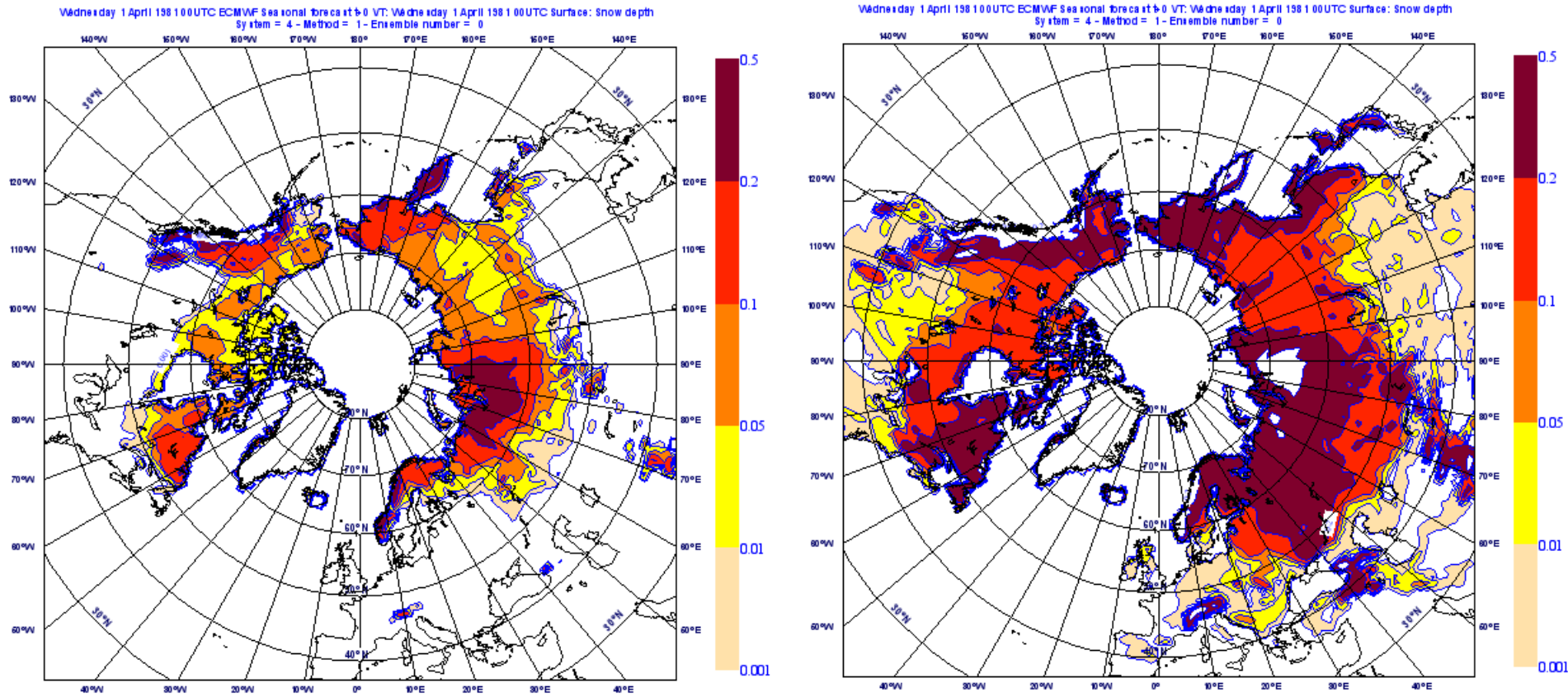


Ocean Potential Temperature Equatorial Section
2017 November mean anomaly (1981-2009 climate)



ECMWF Ocean Reanalysis ORA-S4 Dec 10 2017

Initial conditions - land surface



Snow depth limits, 1st April

Step 3: Run an ensemble forecast

- Need an ensemble to represent possible future evolutions of the atmosphere/ocean/earth system
 - After 10-15 days, details of day-to-day weather become decorrelated – all the ensemble members predict something different, and the exact future is unknowable
 - But we are looking for changes in the pdf of, for example, seasonal means. Ensembles will give us the information we need to evaluate such changes and probabilities.
- How to make the ensemble members different?
 - Explicitly generate an ensemble on the 1st of each month, with perturbations to represent the uncertainty *in the initial conditions*; run forecasts for 7 months
 - SST perturbations (symmetric pairs, based on sampling of past differences between different analyses)
 - 5 member ocean analysis, with perturbations to wind and data
 - Stochastic physics to represent *indeterminacy* of large scale (due to missing details of small scales) and also account for uncertainty due to our model being imperfect
- Now, simply run the ensemble of forecasts
 - This is where we need a very big computer
 - Worry about model biases later

ECMWF SEAS5 configuration

- Real time forecasts:
 - **51 member ensemble forecast to 7 months**
 - SST and atmosphere initial perturbations (SV, EDA) added to each member
 - **15 member ensemble forecast to 13 months**
 - Designed to give an 'outlook' for ENSO
 - Only runs once per quarter (Feb, May, Aug and Nov starts)
- Re-forecasts from 1981-2016 (36 years)
 - 25 member ensemble every month
 - 15 members extended to 13 months once per quarter

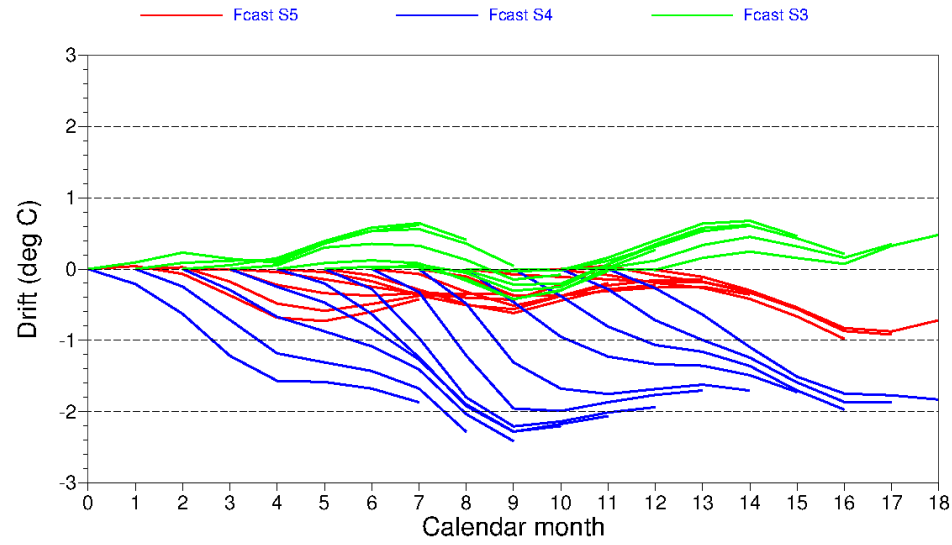
How many re-forecasts?

- Re-forecasts dominate total cost of system
 - SEAS5: 10800 re-forecasts (must be in first year)
612 real-time integrations (per year)
- Re-forecasts define model climate
 - Need both climate mean and the pdf, latter needs large sample
 - May prefer to use a “recent” period (SEAS5 has 36 years available, but uses only last 24 years for web products)
 - SEAS5 has 600 member climate (25 members * 24 years) for web products, so sampling is basically OK
- Re-forecasts provide information on skill
 - A forecast cannot be used unless we know (or assume) its level of skill
 - Observations have only 1 member, so large ensembles are less helpful than large numbers of cases.
 - Care needed e.g. to estimate skill of 51 member ensemble based on past performance of 25 member ensemble
 - For regions of high signal/noise, SEAS5 gives adequate skill estimates
 - For regions of low signal/noise (eg ≤ 0.5), need hundreds of years, 36 years available is not enough

Step 4: Calibrate the output

- Model biases are typically comparable in size to the signal we are predicting
 - True for both SST and atmosphere fields
 - This bias **MUST** be accounted for in some way – a fundamental requirement for seasonal forecasting
- Forecast calibration requires a corresponding set of **re-forecasts**:
 - Re-forecasts should use the same model and (where possible) the same method of initialization, to ensure that the biases are consistent
 - There are different ways of using the re-forecasts for calibration, but in general need a large number of re-forecasts. The full set of re-forecasts for a given calendar start date (e.g. 1 May) define the *model climate* for forecasts starting on that date.
- Compare the model forecast to the model climate:
 - SEAS5 forecasts are calibrated using 25-member re-forecasts for each of the years 1993-2016 (24 years), so the model climate has 600 members.
 - Model climate has both a mean and a distribution, allowing us to estimate e.g. tercile boundaries.
 - Model climate is a function of start date and forecast lead time – model biases often get larger the further into the future we calculate.

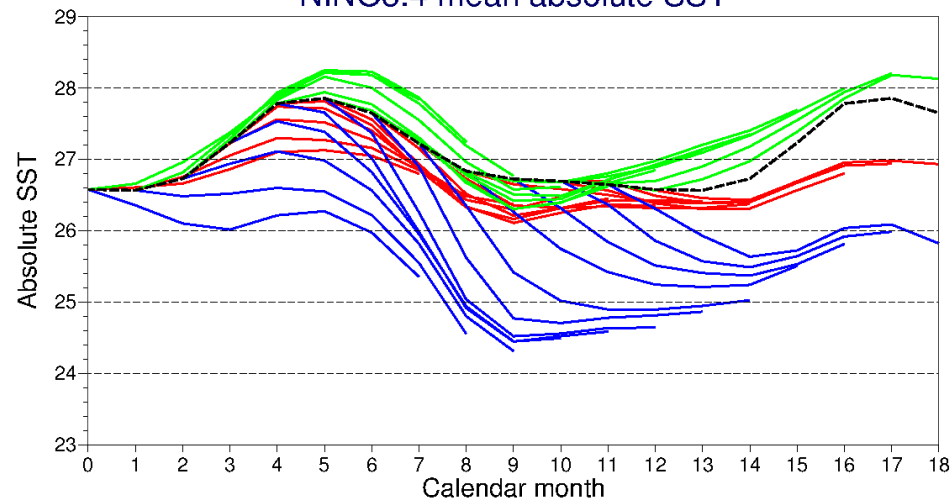
NINO3.4 mean SST drift

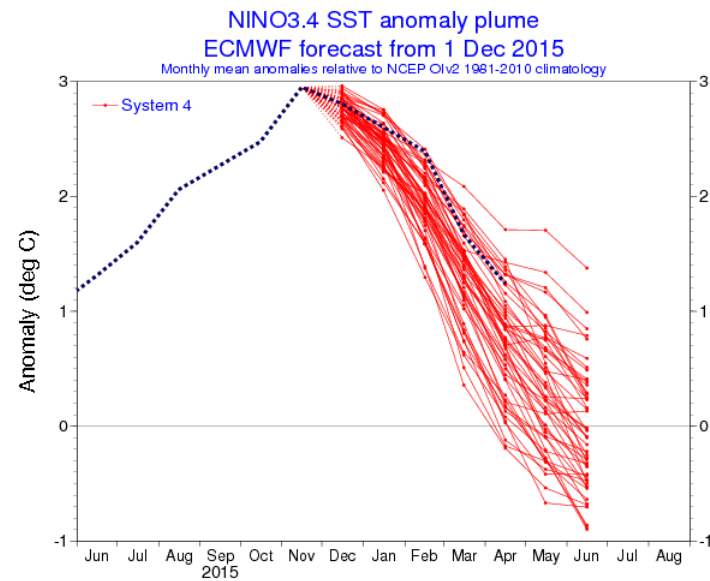
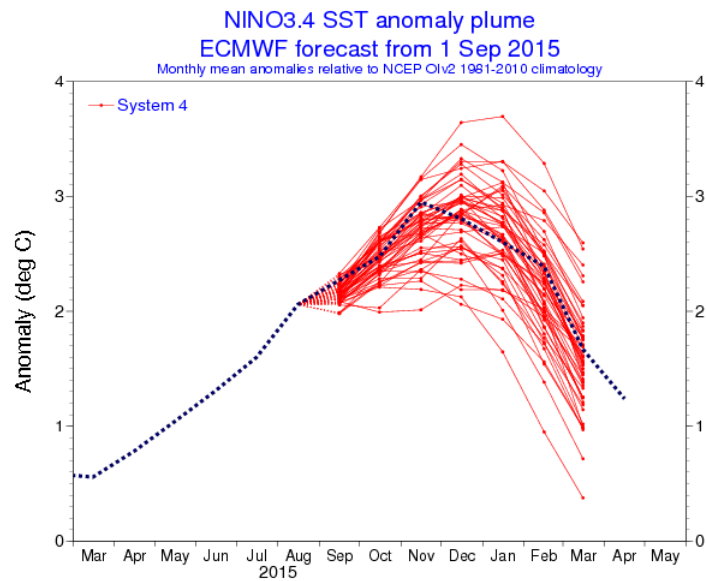
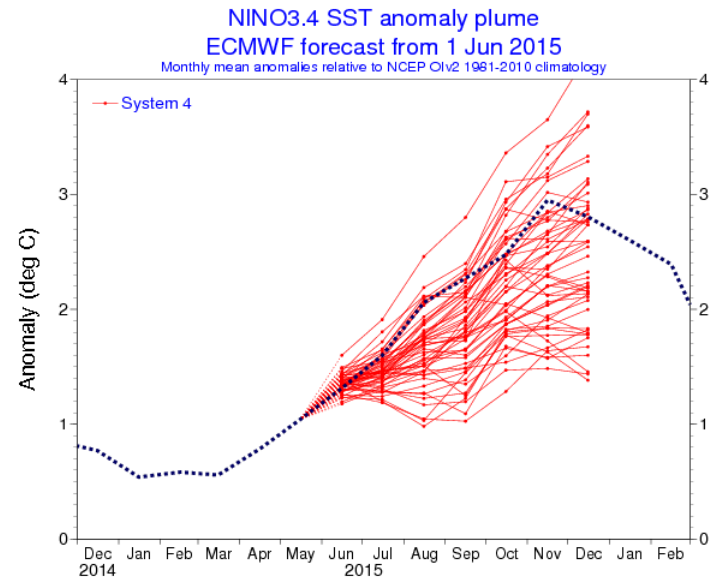
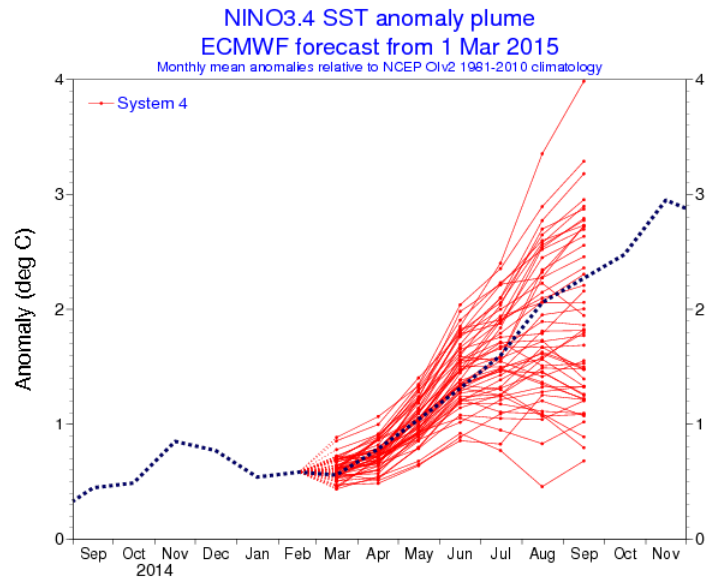


SST bias is a function of lead time and season.

Some systems have less bias, but it is still large enough to require correcting for.

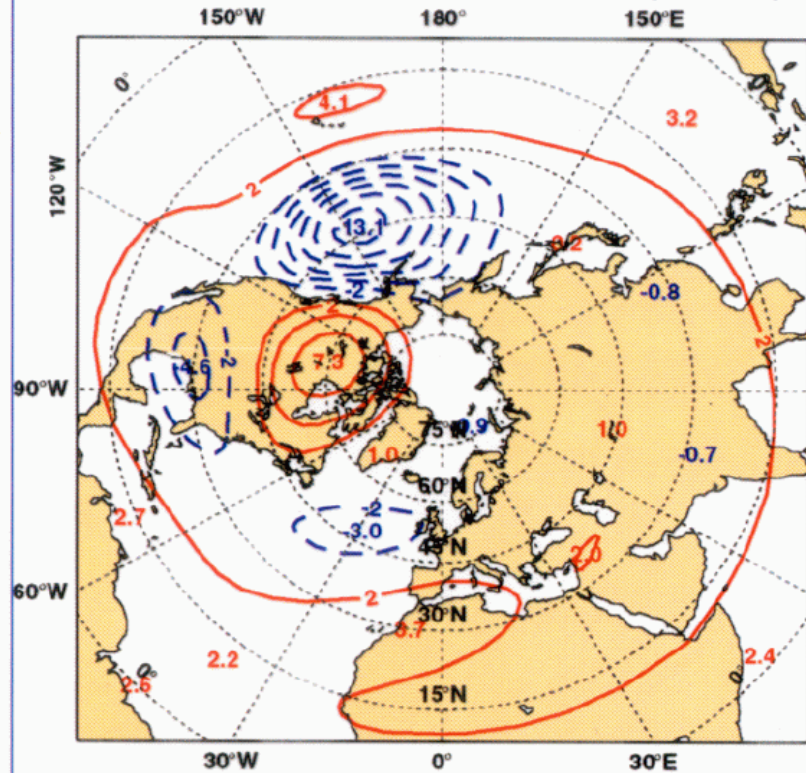
NINO3.4 mean absolute SST



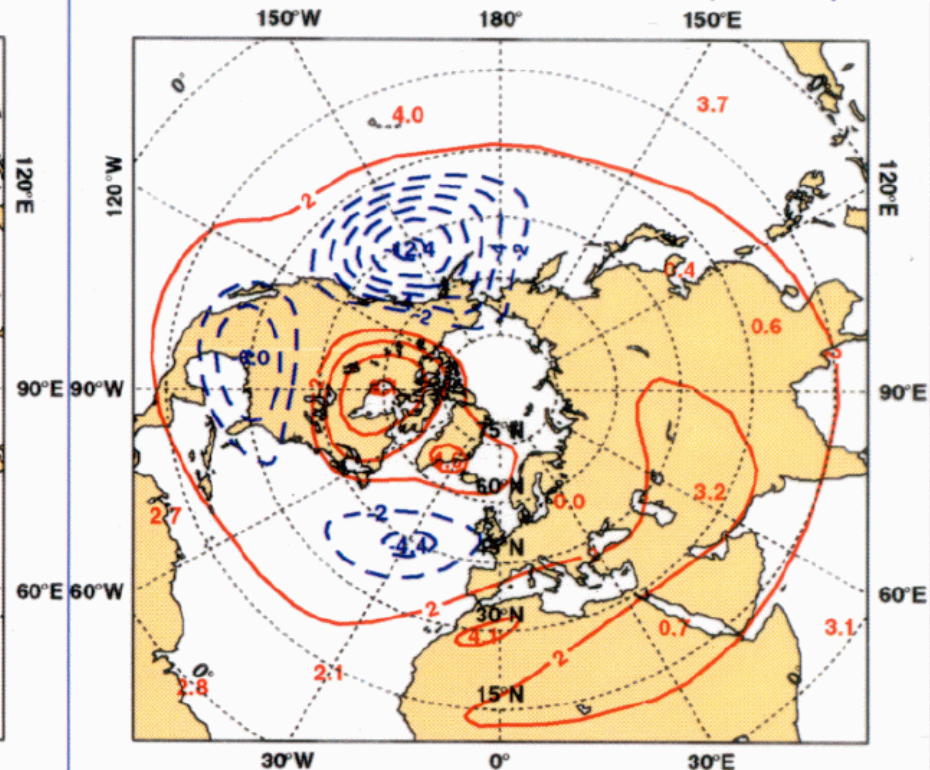


Despite SST bias and other errors, anomalies in the coupled system can be remarkably similar to those obtained using observed (unbiased) SSTs

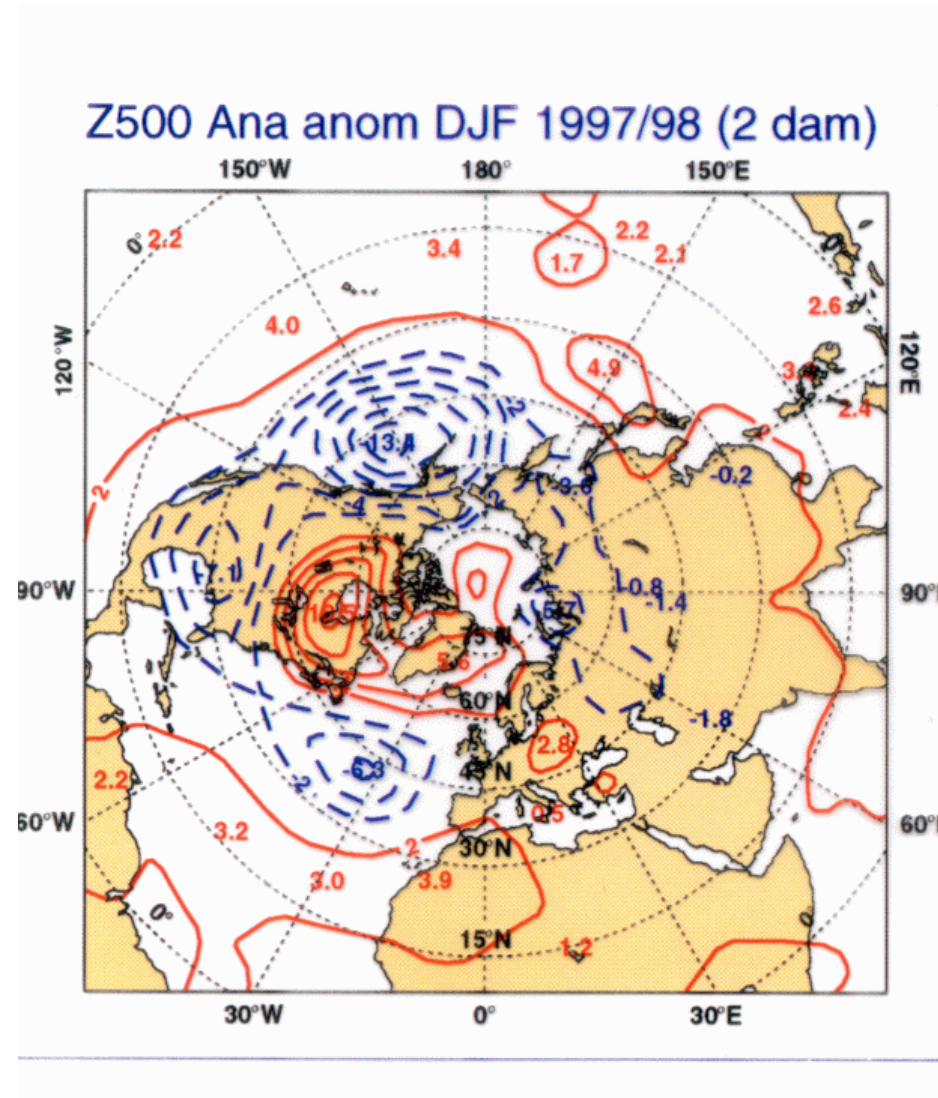
Z500 COA anom DJF 1997/98 (2 dam)



Z500 UNC anom DJF 1997/98 (2 dam)

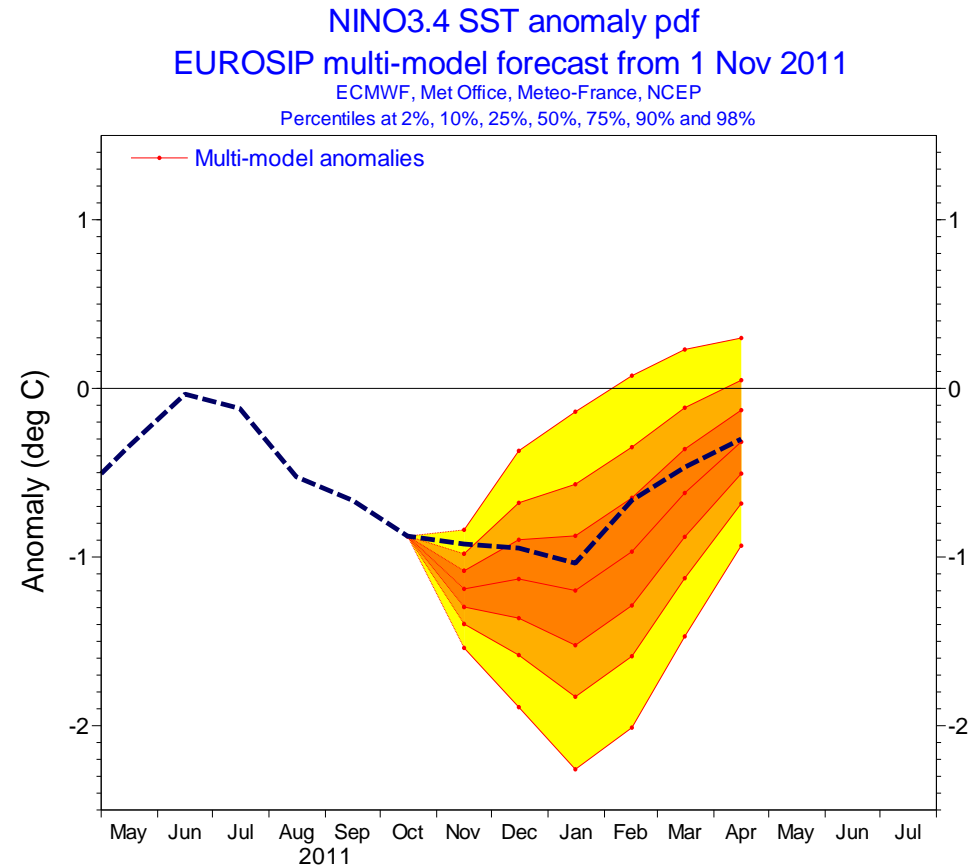
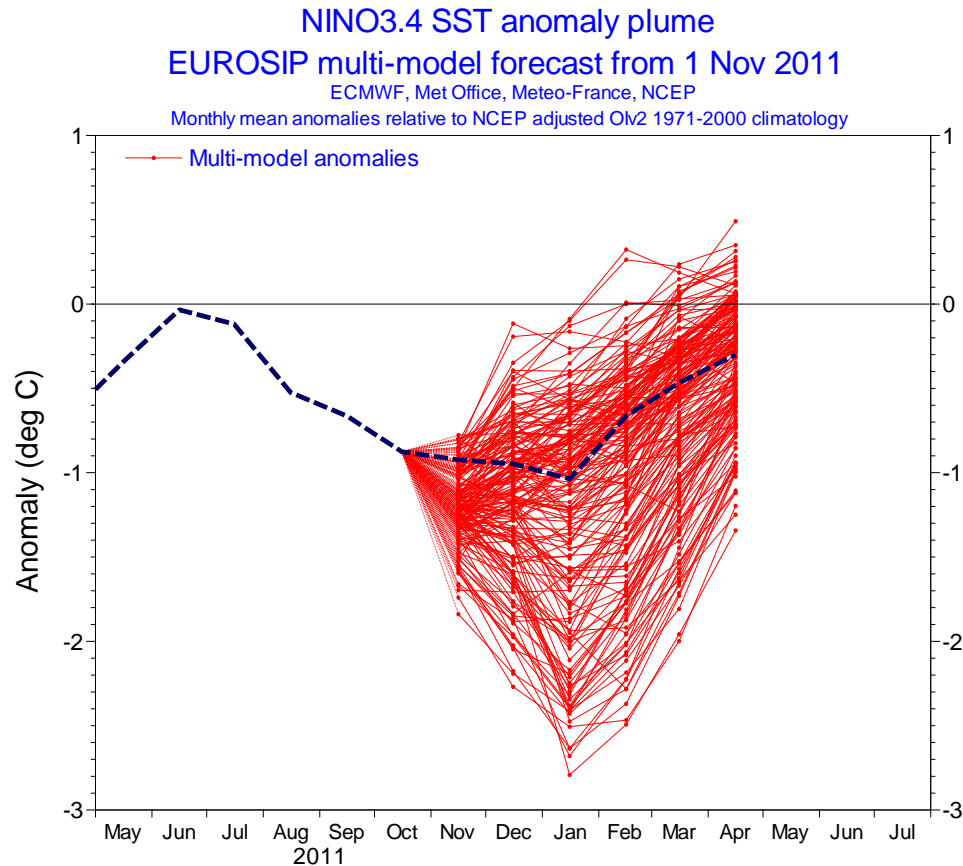



... and can also verify well against observations



- Implicit assumption of linearity
 - We implicitly assume that a shift in the model forecast relative to the model climate corresponds to the expected shift in a true forecast relative to the true climate, despite differences between model and true climate.
 - Most of the time, this assumption seems to work pretty well. But not always. This is one reason we are always seeking to improve the models and reduce the size of the biases. The smaller the bias, the less chance there is of non-linear effects causing errors in the seasonal forecast.
- More advanced calibration, based on forecast skill
 - The calibration mentioned so far is rather basic, and is designed to remove zeroth or at most first order errors from the forecast
 - Forecast interpretation has to take account of past skill – are there grounds for trusting the model forecast?
 - More advanced calibration methods can be used – we covered some in the previous lecture

Nino 3.4 plume and pdf – calibrated multi-model forecast



Each model bias and variance corrected, then multi-model  ensemble formed, then calibrated to give pdf with correct average width (“t” distribution, variance is half climatological, half varying with model distribution – a variant of NGR)

3. Operational forecast products from SEAS5

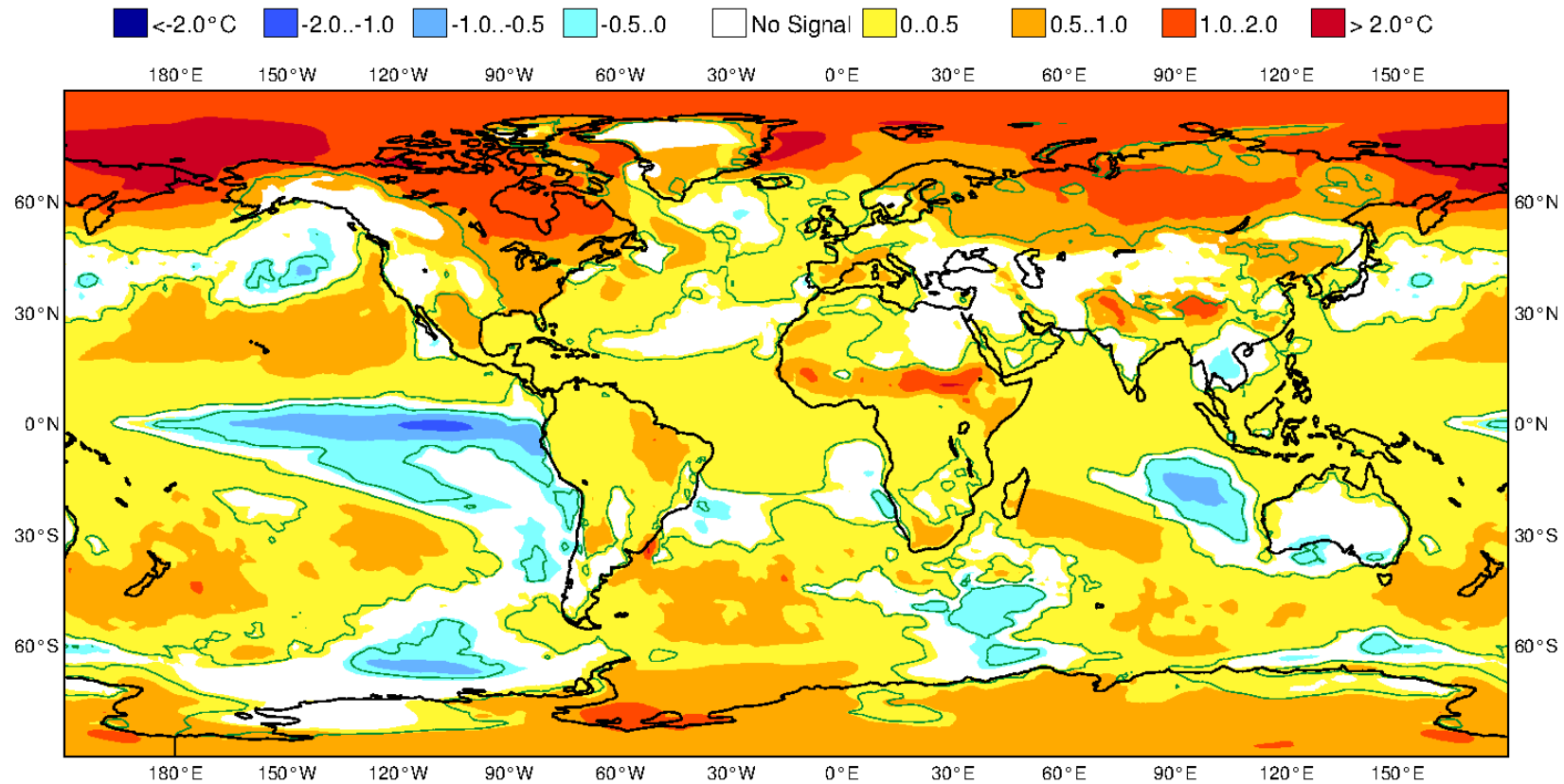
- A few examples only – see ECMWF or C3S web pages for full details and assessment of skill
- All graphical products are supplied with corresponding skill estimates, as we illustrate in our examples. It is always important to look at the skill, to reduce the risk of over-interpreting the forecasts!
- Note: Significance values on plots
 - Ensembles are large enough to test whether any apparent signals are real shifts in the model pdf, or are due to the limited ensemble size giving a false signal by chance
 - We use the Wilcoxon-Mann-Whitney rank-sum test, which is non-parametric and is both robust and efficient at detecting shifts in the mean
 - The significance levels on the plots are a test as to *whether the model has a signal*, and are **NOT** related in any way as to whether the model signal should be trusted. The past skill estimates should be looked at to get a sense of the reliability of the model forecasts.

ECMWF Seasonal Forecast Mean 2m temperature anomaly

Forecast start is 01/09/17, climate period is 1993-2016
Ensemble size = 51, climate size = 600

System 5
OND 2017

Shaded areas significant at 10% level
Solid contour at 1% level

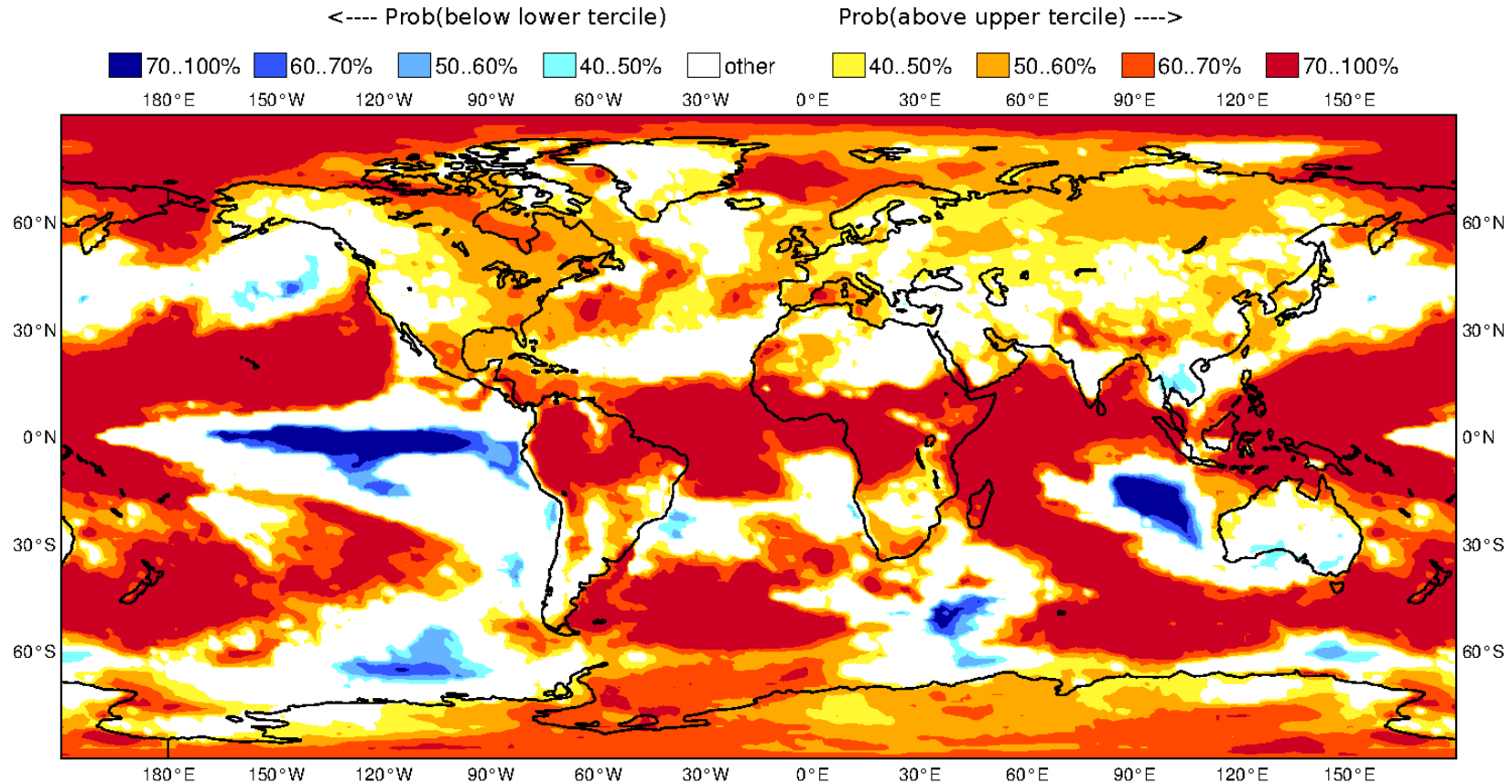


ECMWF Seasonal Forecast

Prob(most likely category of 2m temperature)

Forecast start is 01/09/17, climate period is 1993-2016
Ensemble size = 51, climate size = 600

System 5
OND 2017

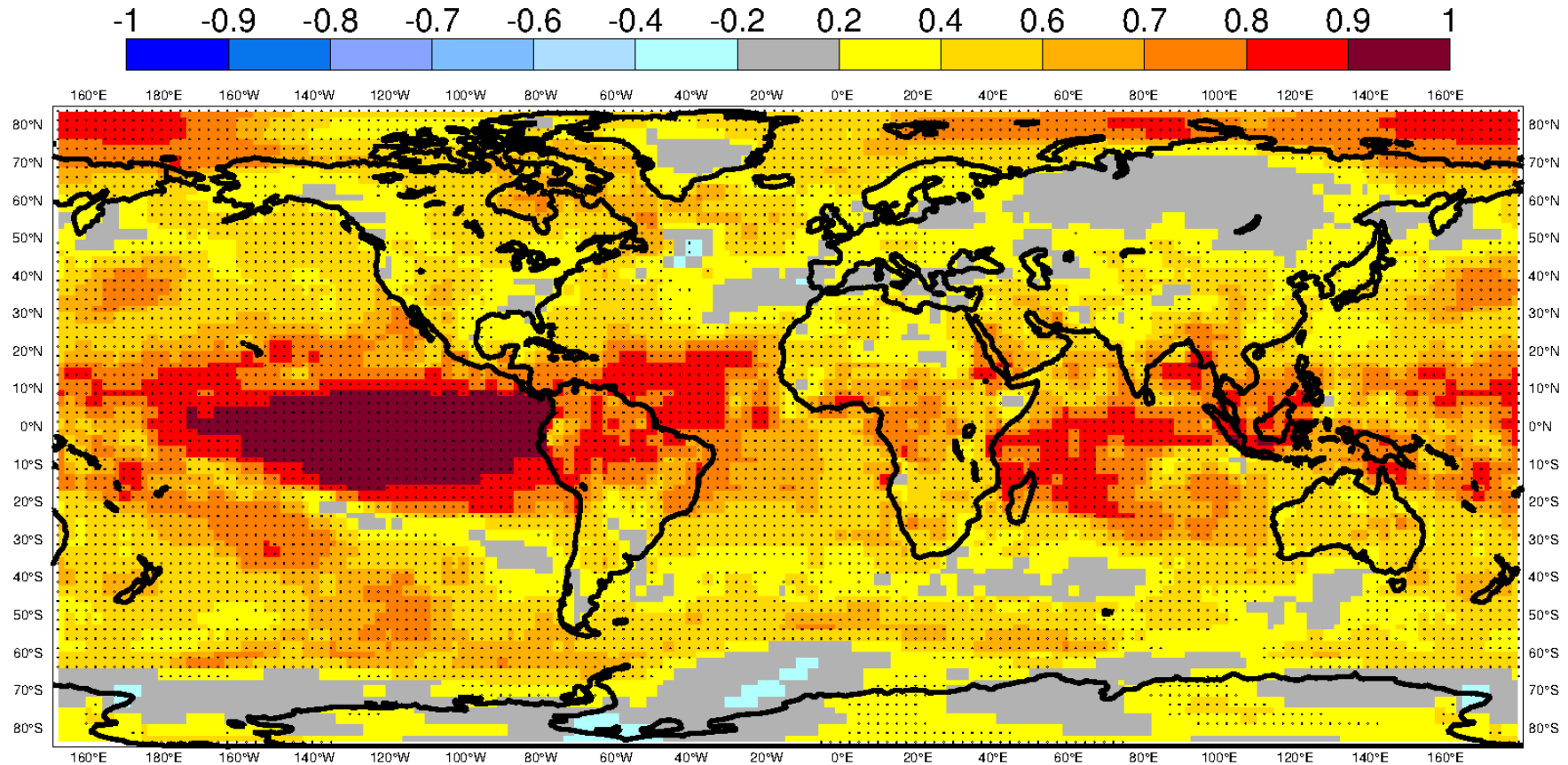


Anomaly Correlation Coefficient for 0001 with 25 ensemble members

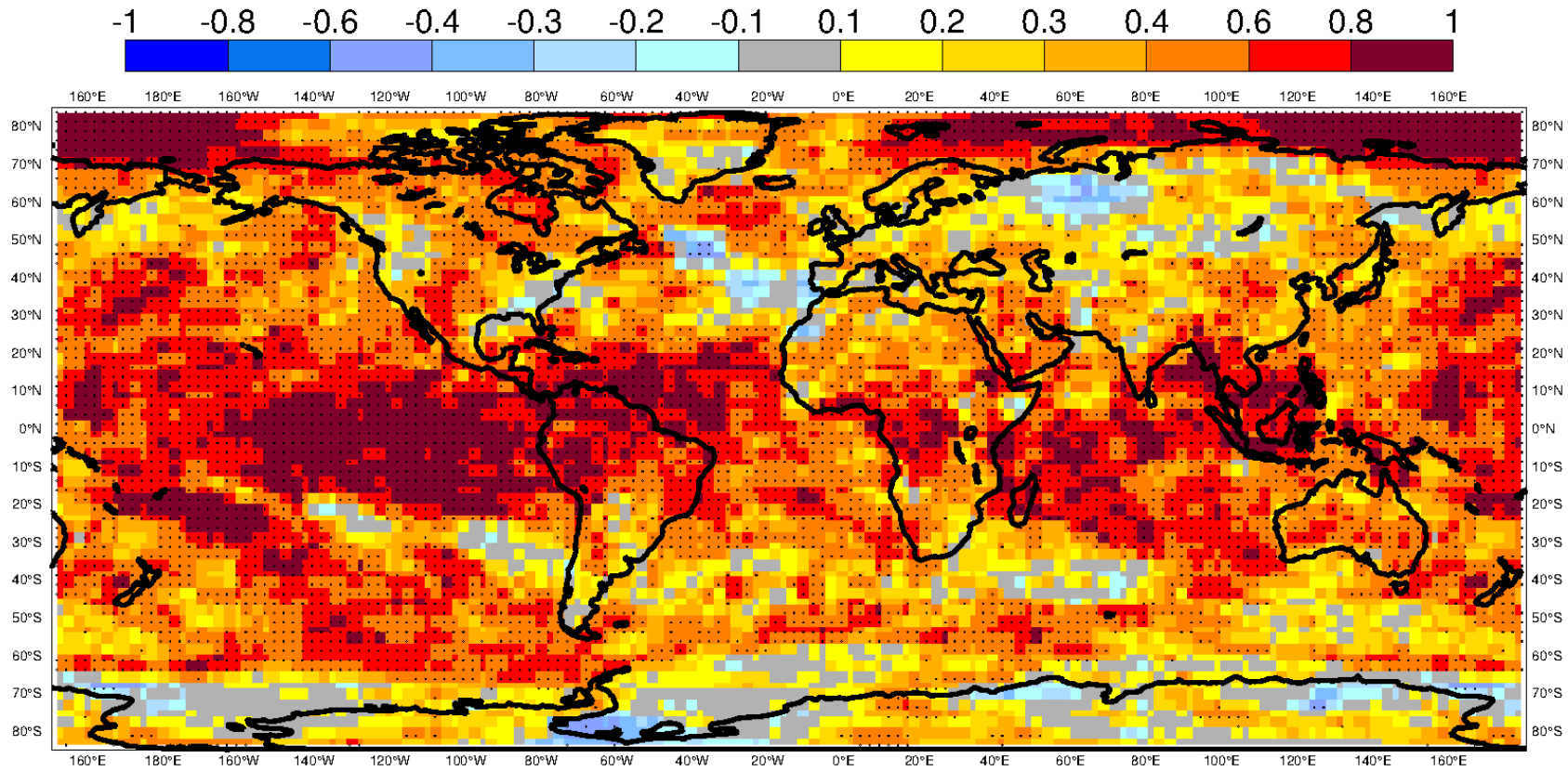
Near-surface air temperature

Hindcast period 1981-2016 with start in September average over months 2 to 4

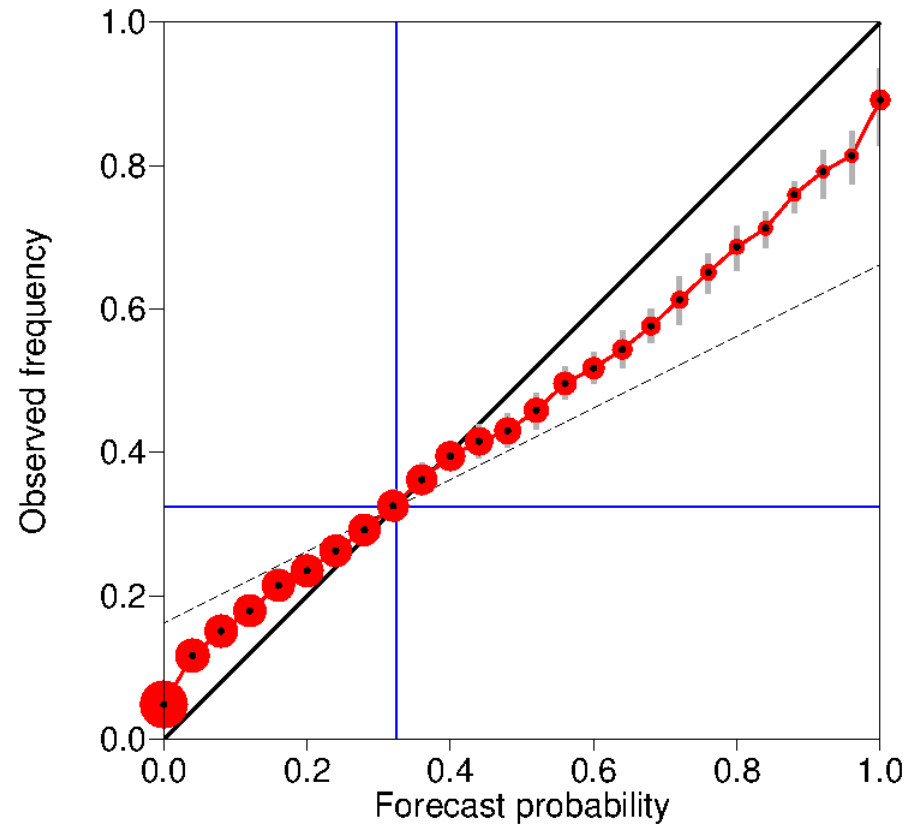
Black dots for values significantly different from zero with 95% confidence (1000 samples)



ROC Skill Score for ORecmfEX0001SY05M1 with 25 ensemble members and 26 bins
Near-surface air temperature anomalies above the upper tercile
Hindcast period 1981-2016 with start in September and averaging period 2 to 4
Threshold estimated with a kernel method for the PDF



Reliability diagram for 0001 with 25 ensemble members
Near-surface air temperature anomalies above the upper tercile
Accumulated over global (land and sea points)
Hindcast period 1981-2016 with start in September average over months 2 to 4
Skill scores and 95% conf. intervals (1000 samples)
Brier skill score: 0.175 (0.134, 0.215)
Reliability skill score: 0.982 (0.976, 0.987)
Resolution skill score: 0.193 (0.155, 0.231)

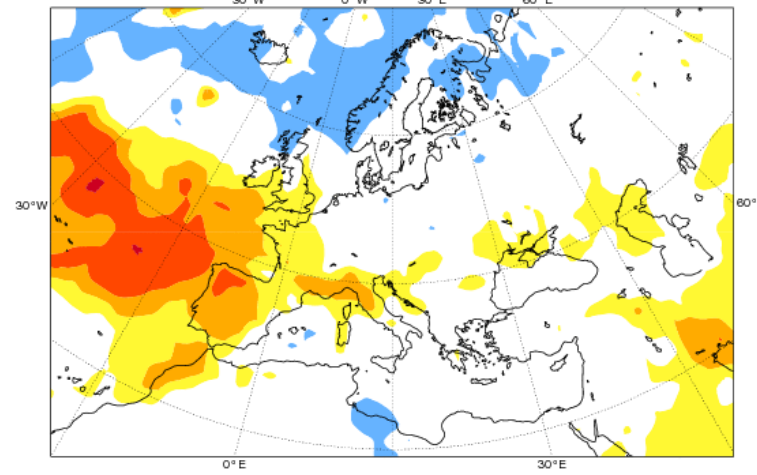


Other operational plot examples

ECMWF Seasonal Forecast
 Prob(lowest 20% of climatology) - precipitation
 Forecast start reference is 01/12/11
 Ensemble size - 51, climate size - 450

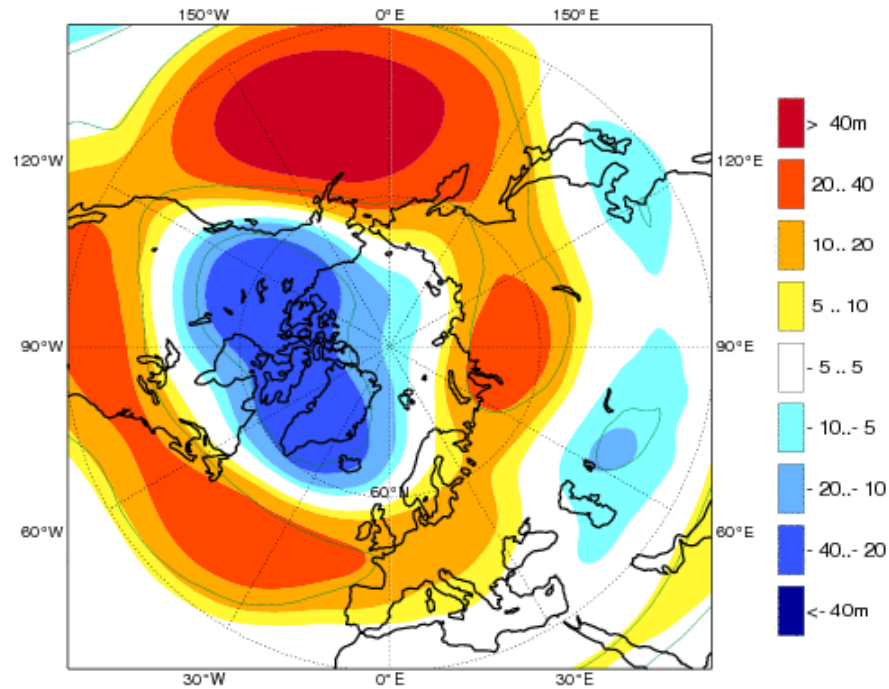
System 4
 JFM 2012

0..10% 10..30% 30..40% 40..50% 50..70% 70..100%



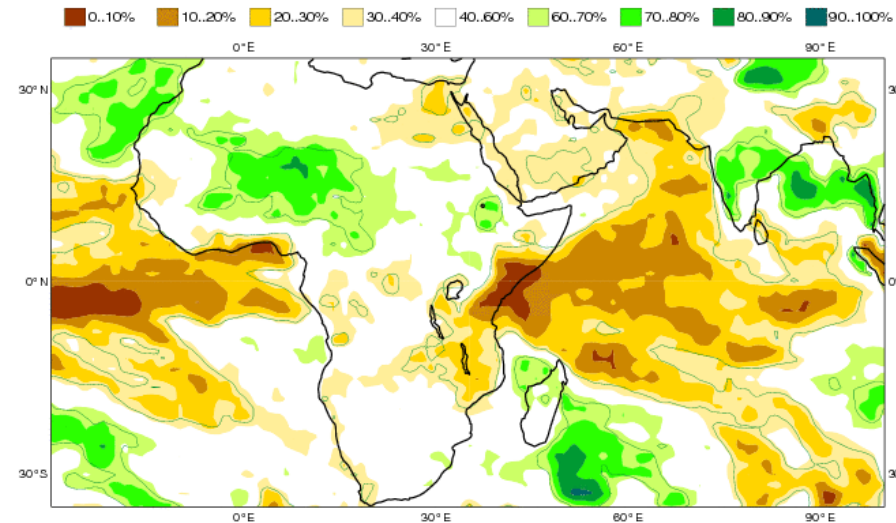
ECMWF Seasonal Forecast
 Mean Z500 anomaly
 Forecast start reference is 01/11/11
 Ensemble size - 51, climate size - 450

System 4
 DJF 2011/12
 Solid contour at 1% significance level



ECMWF Seasonal Forecast
 Prob(precipitation > median)
 Forecast start reference is 01/05/12
 Ensemble size - 51, climate size - 450

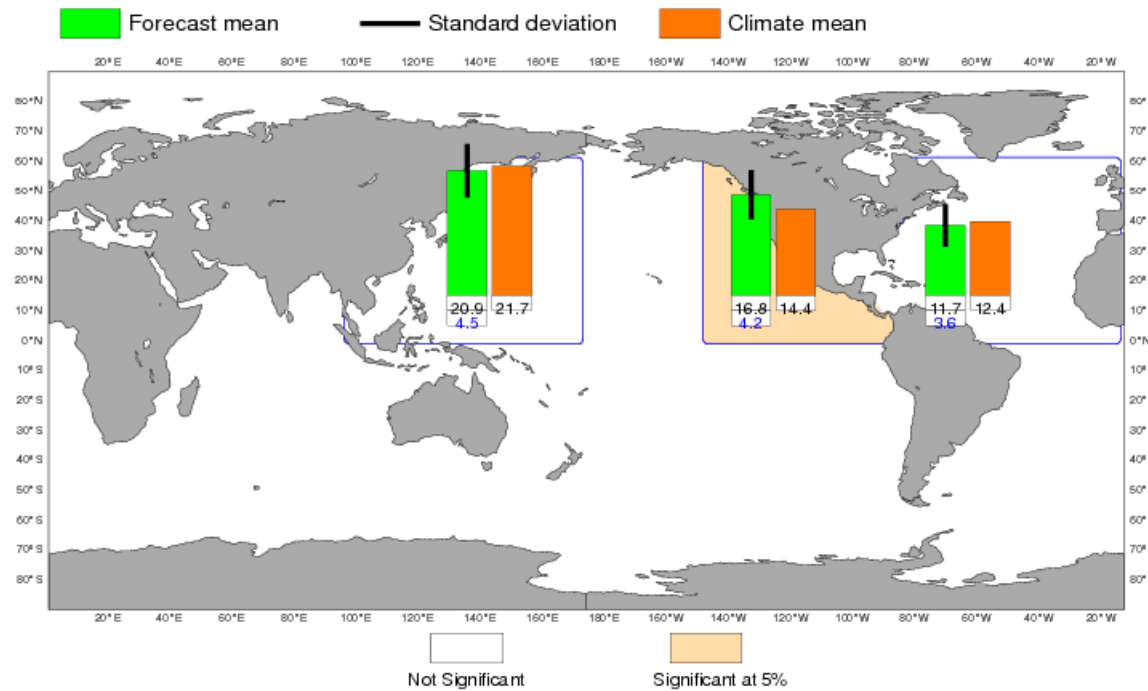
System 4
 JJA 2012
 Solid contour at 1% significance level



Tropical storm forecasts

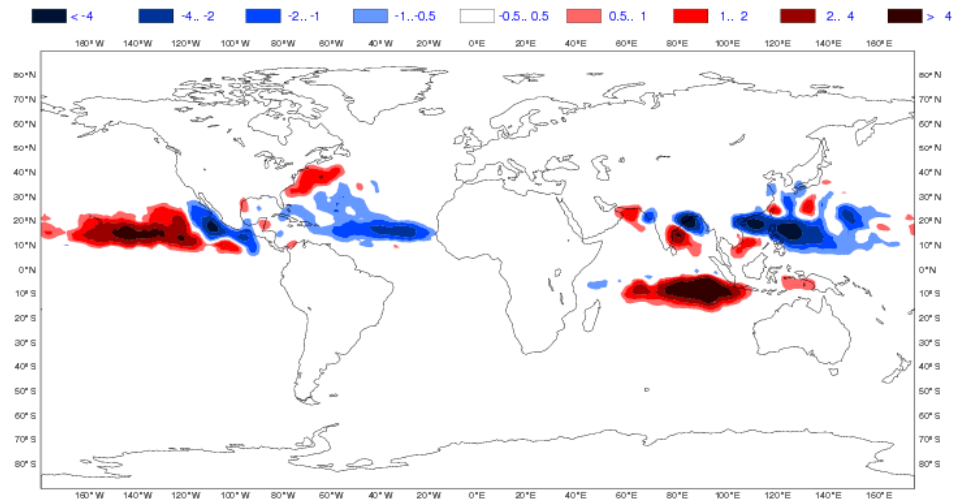
ECMWF Seasonal Forecast
Tropical Storm Frequency
Forecast start reference is 01/05/2016
Ensemble size = 51, climate size = 300

System 4
JJASON 2016
Climate (initial dates) = 1990-2009



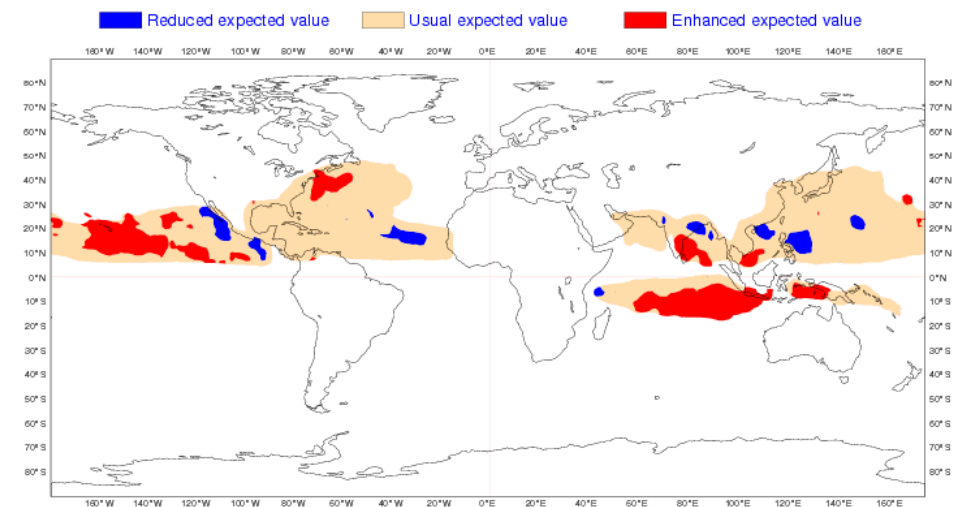
ECMWF Seasonal Forecast
Tropical Storm Density Anomaly
Forecast start reference is 01/05/2016
Ensemble size = 51, climate size = 300

System 4
JJASON 2016
Climate (initial dates) = 1990-2009

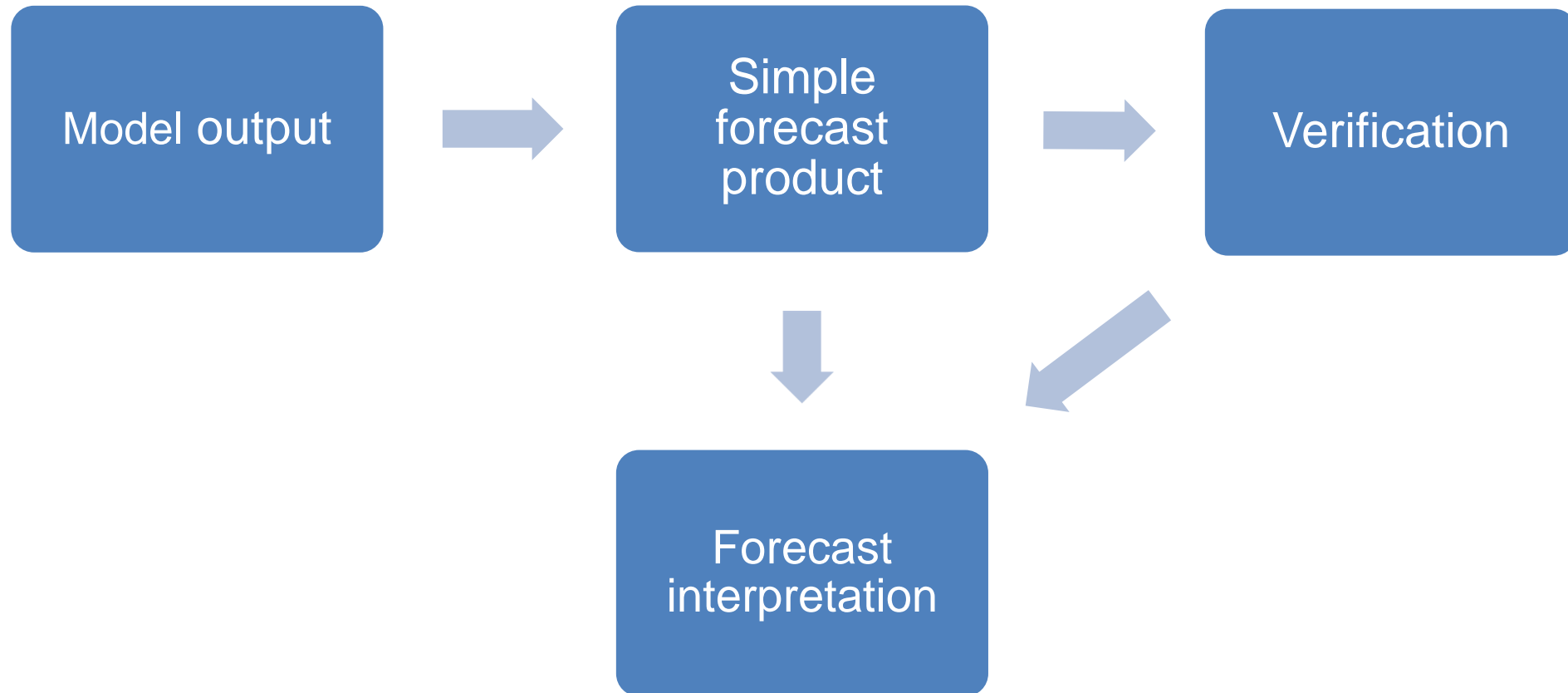


ECMWF Seasonal Forecast
Standardized Tropical Storm Density
Forecast start reference is 01/05/2016
Ensemble size = 51, climate size = 300

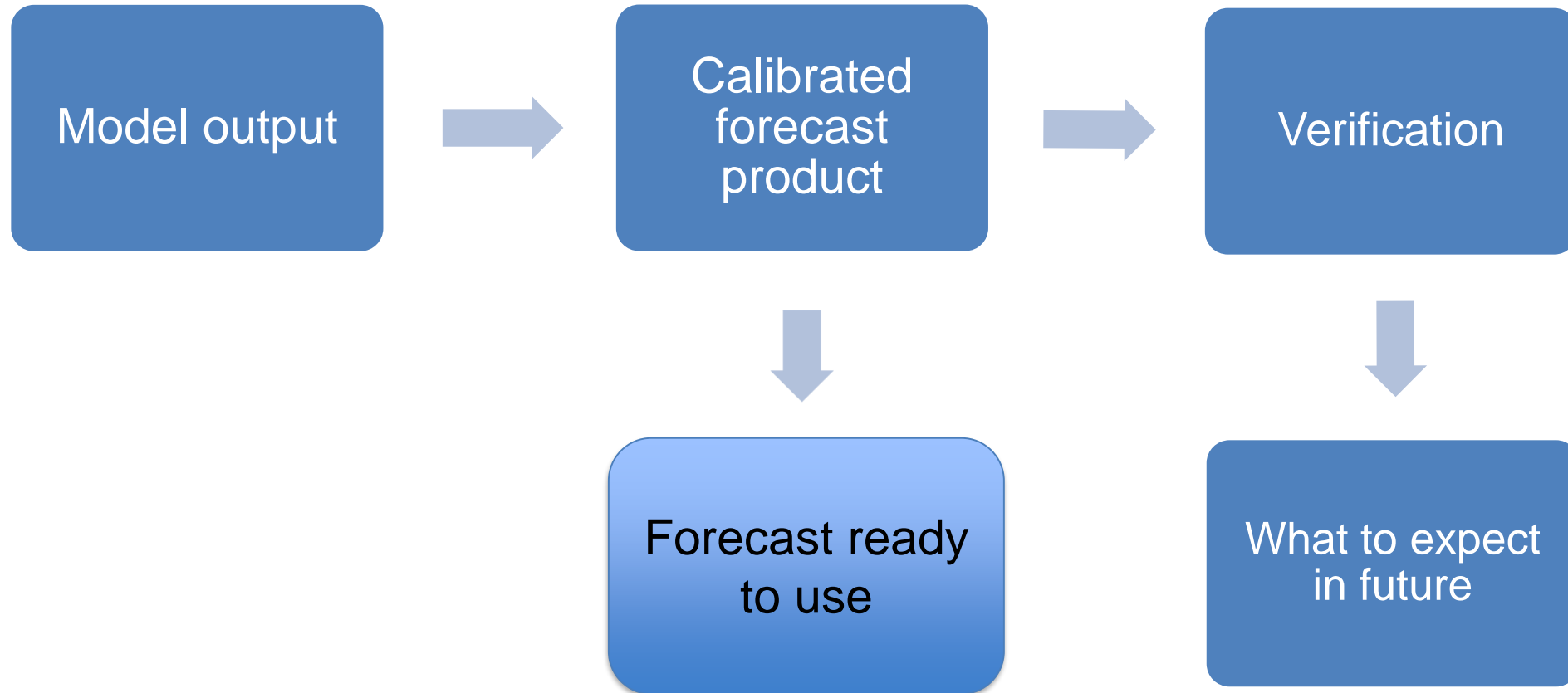
System 4
JJASON 2016
Climate (initial dates) = 1990-2009



4. Forecast process I: How it is often done (including at ECMWF)



Forecast process II: How it should be done



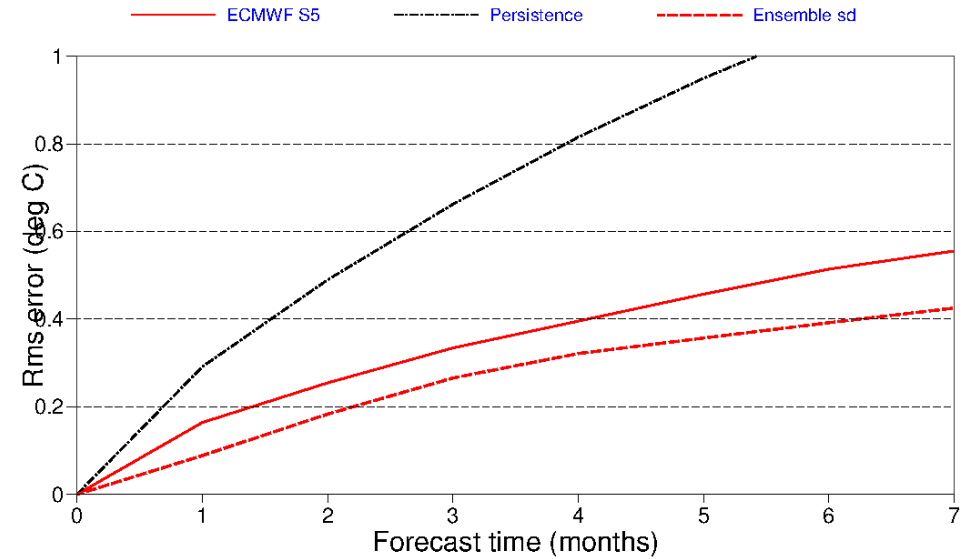
SST forecast performance

Actual r.m.s. errors > model estimate of “perfect model” errors

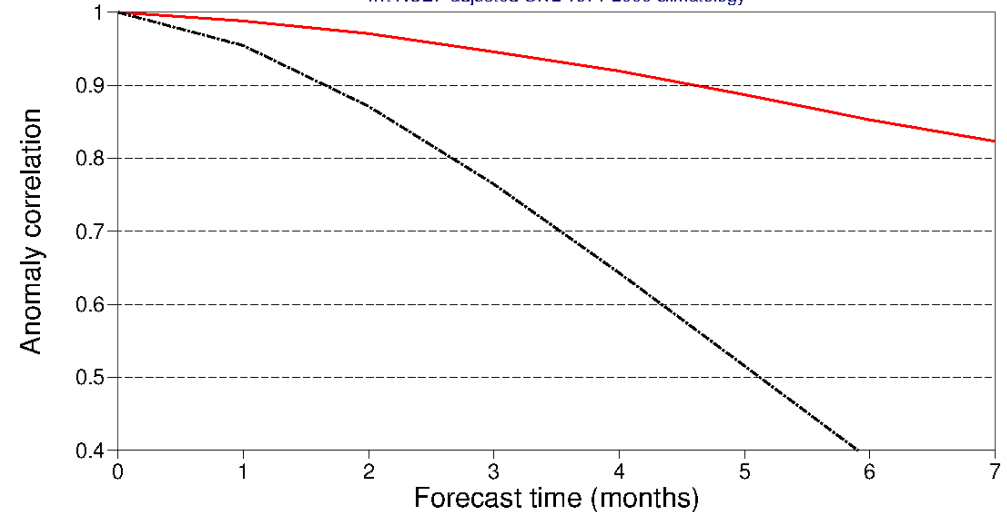
NOTE: Stochastic physics gives increased spread to Nino SSTs, due to its representation of low-frequency model error.

This gives better probabilistic scores, but means the ensemble spread is not a predictability limit: if in future systems we reduce the model error, we can reduce the amplitude of the stochastic “noise” to match, and the ensemble spread will reduce.

NINO3.4 SST rms errors
432 start dates from 19810101 to 20161201, bias corrected
Ensemble size is 25



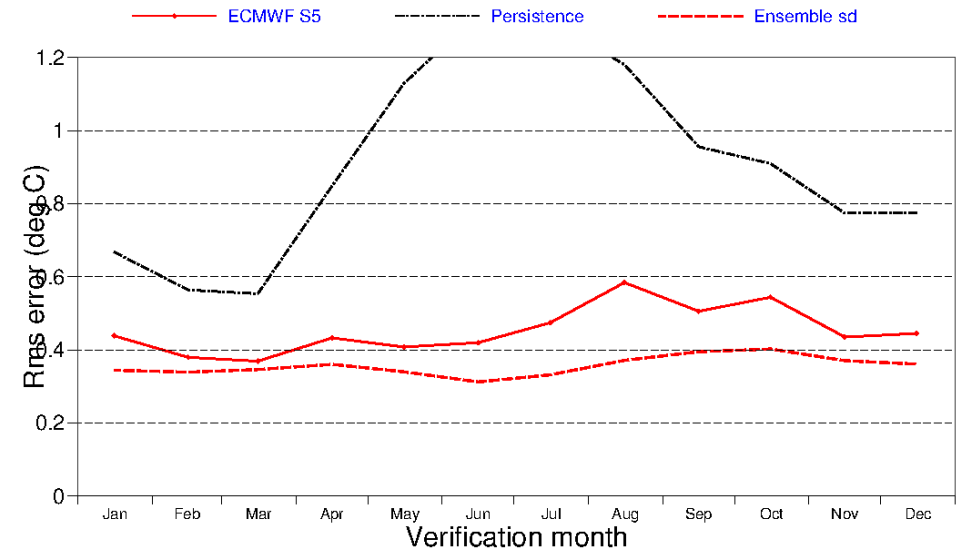
NINO3.4 SST anomaly correlation
wrt NCEP adjusted OIv2 1971-2000 climatology



Seasonal dependence

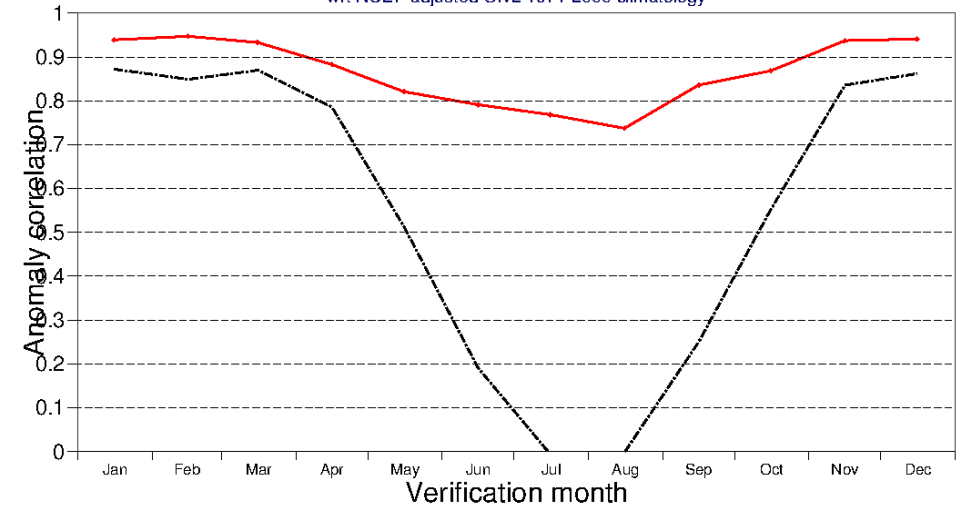
NINO3.4 SST rms errors at 5 months

432 start dates from 19810101 to 20161201, bias corrected
Ensemble size is 25



NINO3.4 SST anomaly correlation at 5 months

wrt NCEP adjusted OIv2 1971-2000 climatology



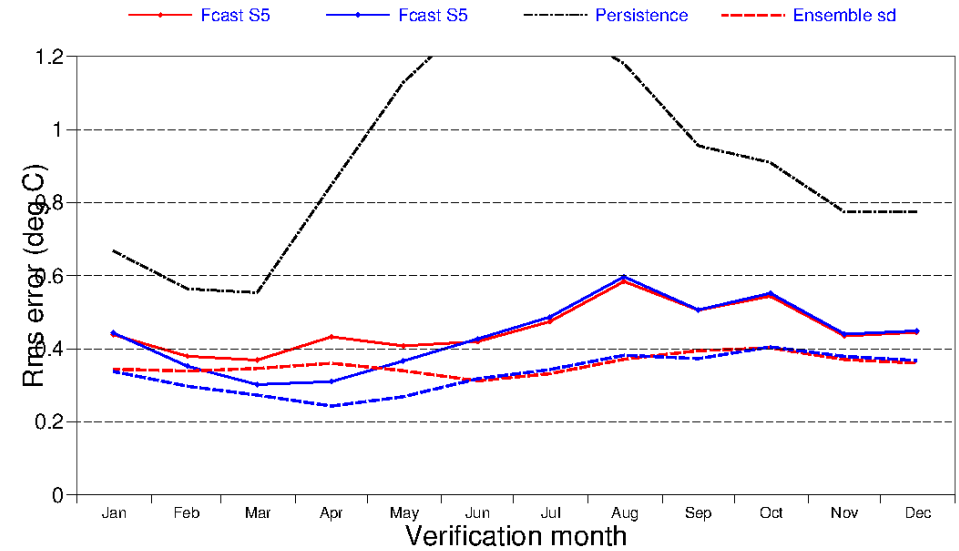
Variance adjustment

This very simple calibration scales the forecast climatological variance to match the observed climatological variance. The scaling is seasonally dependent. This calibration can substantially improve forecast products (and their verification scores). This calibration was used for our previous system, but was turned off in SEAS5.

SEAS5 verification includes the amplitude ratio, which should be used a *posteriori* to interpret the Nino plumes. This is important for forecasts of March, April and May.

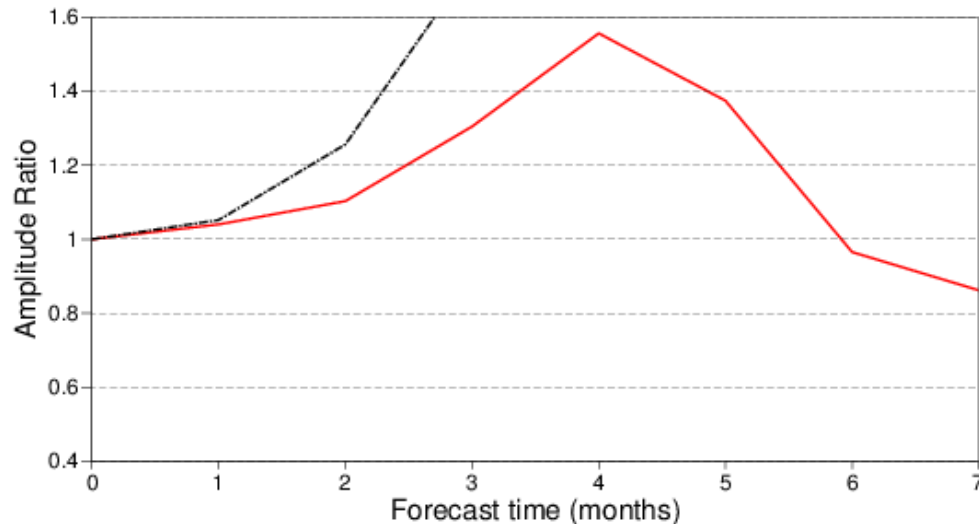
NINO3.4 SST rms errors at 5 months

432 start dates from 19810101 to 20161201, various corrections
Ensemble sizes are 25 (0001) and 25 (0001)



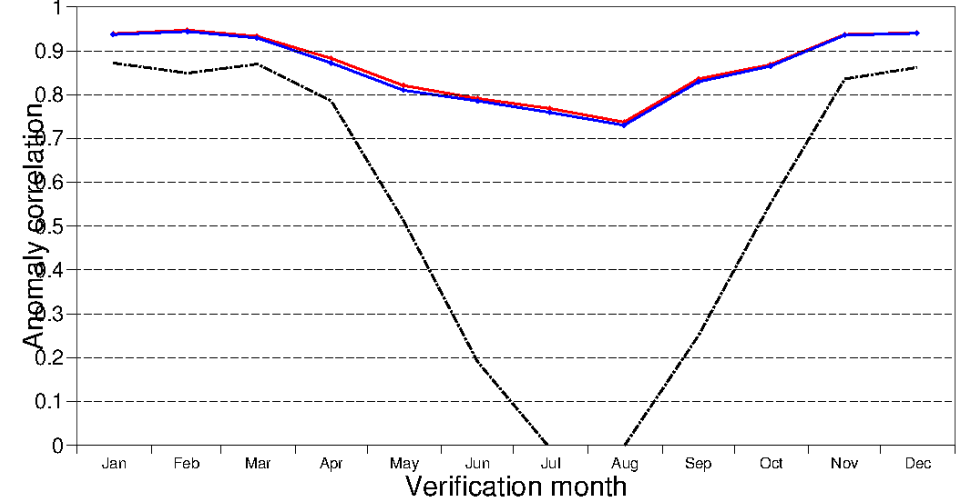
NINO3.4 SST anomaly amplitude ratio

ECMWF S5 (Jan starts)



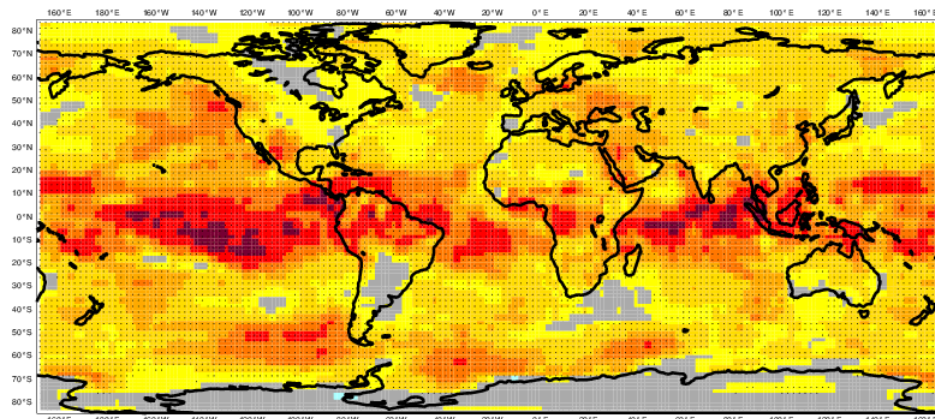
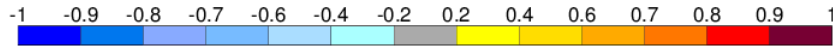
NINO3.4 SST anomaly correlation at 5 months

wrt NCEP adjusted Olv2 1971-2000 climatology



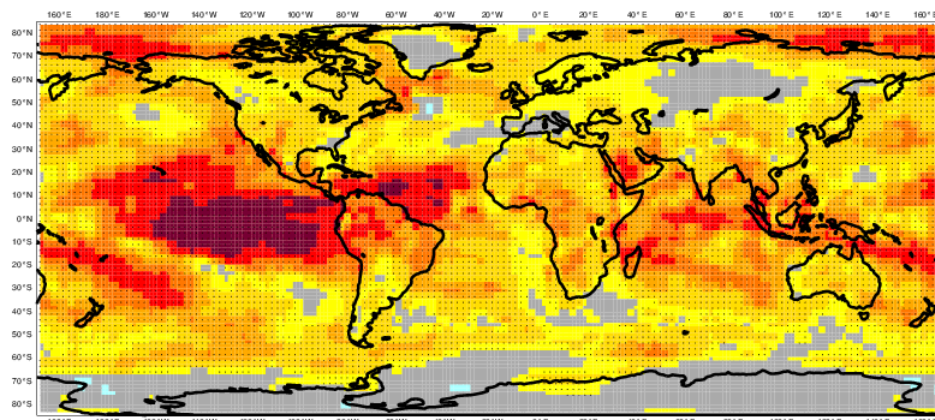
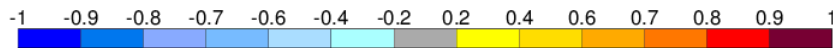
Seasonal dependence of surface parameters

Anomaly Correlation Coefficient for 0001 with 25 ensemble members
 Near-surface air temperature
 Hindcast period 1981-2016 with start in February average over months 2 to 4
 Black dots for values significantly different from zero with 95% confidence (1000 samples)



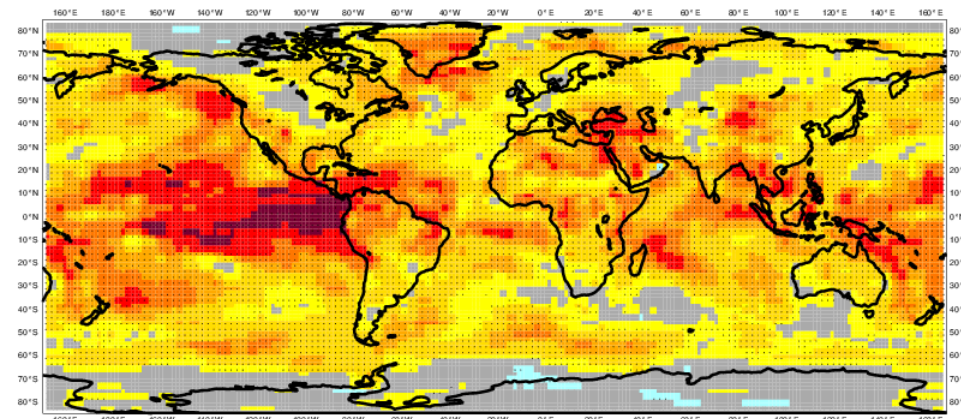
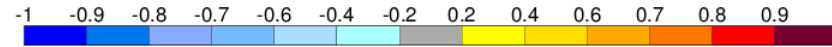
MAM

Hindcast period 1981-2016 with start in August average over months 2 to 4
 Black dots for values significantly different from zero with 95% confidence (1000 samples)



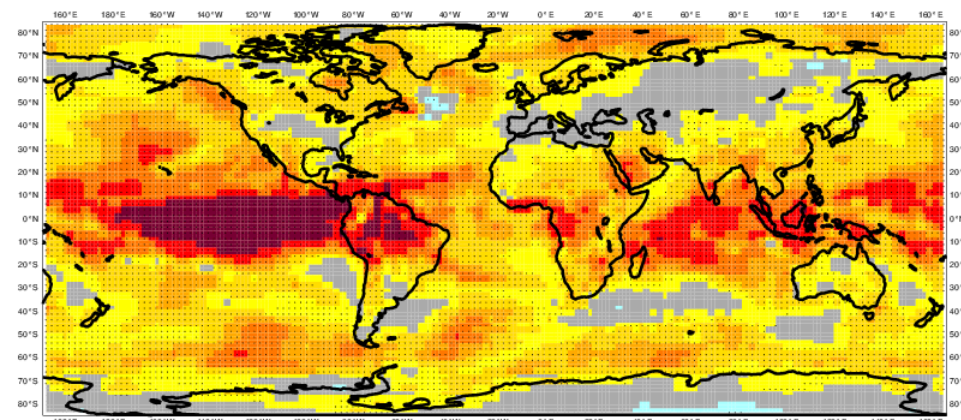
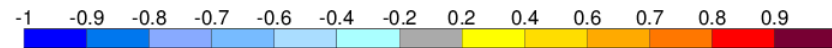
SON

Anomaly Correlation Coefficient for 0001 with 25 ensemble members
 Near-surface air temperature
 Hindcast period 1981-2016 with start in May average over months 2 to 4
 Black dots for values significantly different from zero with 95% confidence (1000 samples)



JJA

Hindcast period 1981-2016 with start in November average over months 2 to 4
 Black dots for values significantly different from zero with 95% confidence (1000 samples)

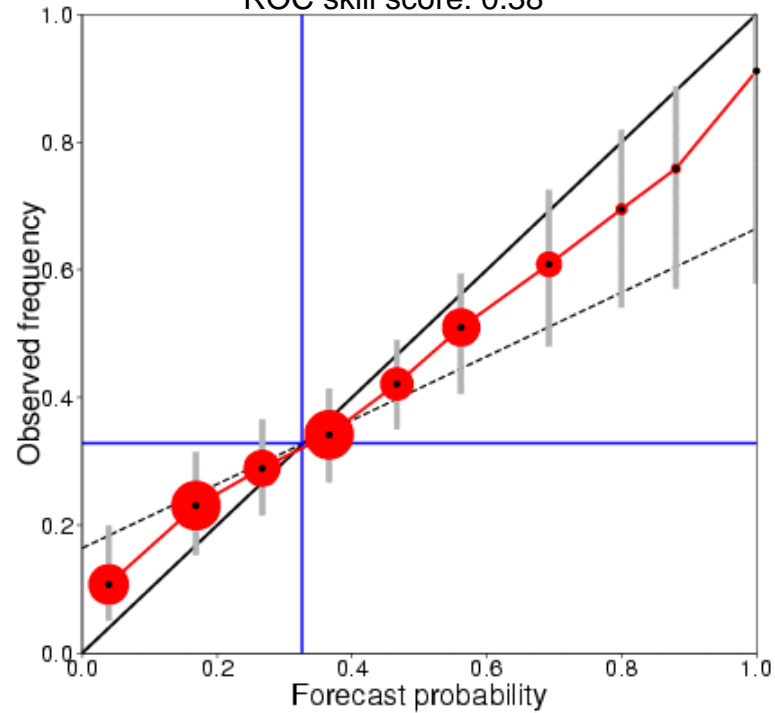


DJF

Role of ensemble size: Scores for Europe in JJA

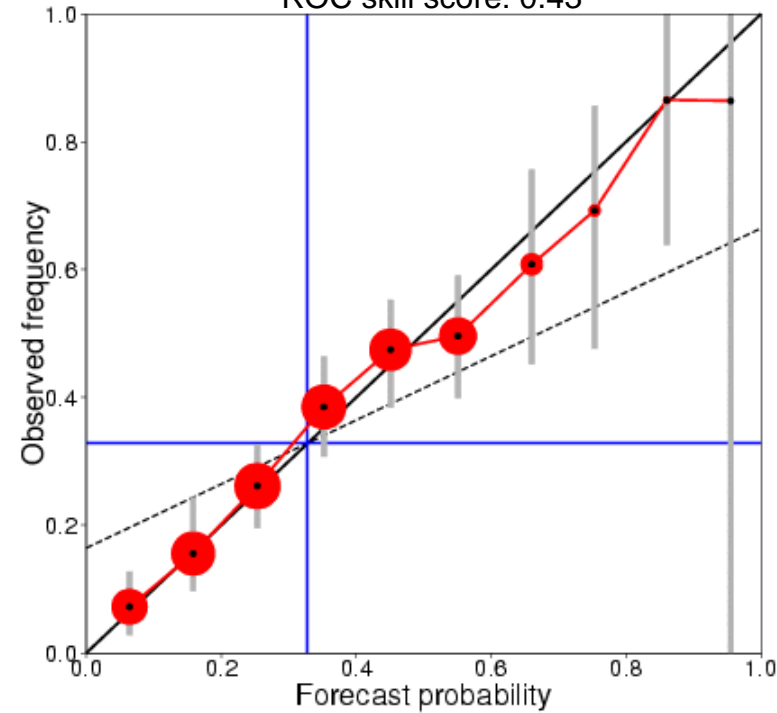
15 members

JJA Europe T2m>upper tercile
Re-forecasts from 1 May, 1981-2010
Reliability score: 0.987
ROC skill score: 0.38



51 members

JJA Europe T2m>upper tercile
Re-forecasts from 1 May, 1981-2010
Reliability score: 0.996
ROC skill score: 0.43

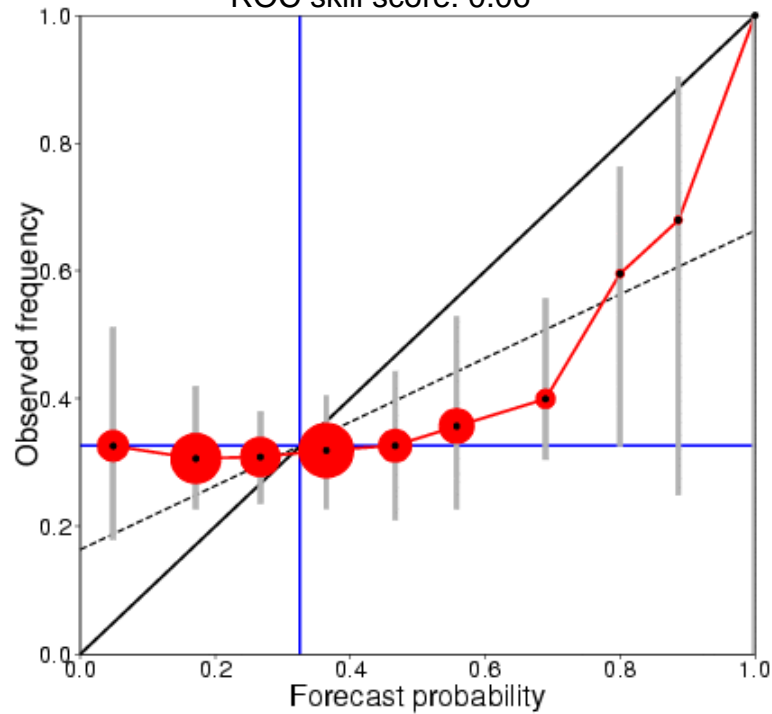


(System 4)

Scores for Europe: DJF

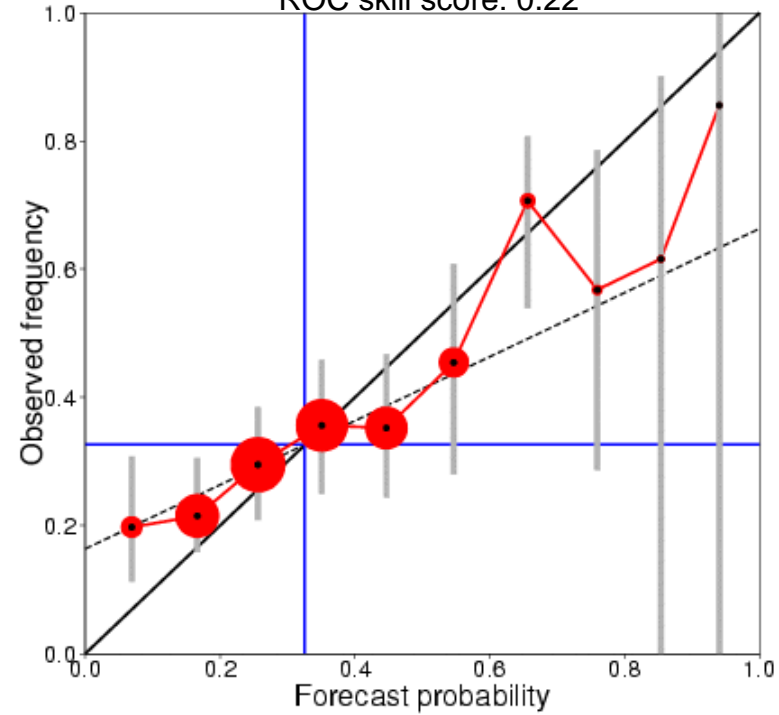
15 members

DJF Europe T2m>upper tercile
Re-forecasts from 1 Nov, 1981-2010
Reliability score: 0.902
ROC skill score: 0.06



51 members

DJF Europe T2m>upper tercile
Re-forecasts from 1 Nov, 1981-2010
Reliability score: 0.981
ROC skill score: 0.22



Overall assessment: how good are the forecasts?

The best way to answer this question is to browse the various skill scores and maps made available on the ECMWF (or other provider) website. But a highly simplified summary is:

- Skill (ACC, BSS, ROC, ...) relative to climate is typically moderate to high in the tropics, moderately low to sometimes very low in mid-latitudes.
- Reliability is on average moderately good; large ensemble sizes are needed to measure this in low-predictability areas.
- Even with large ensemble sizes, the limited number of years means that skill assessments have large uncertainties in mid-latitude regions.
- We can average skill over many gridpoints, seasons etc. to try to reduce uncertainty, but we inevitably trade spatial resolution to gain a bit more accuracy.

- As we are about to see, there are indications that in some cases the forecast spread is too large.

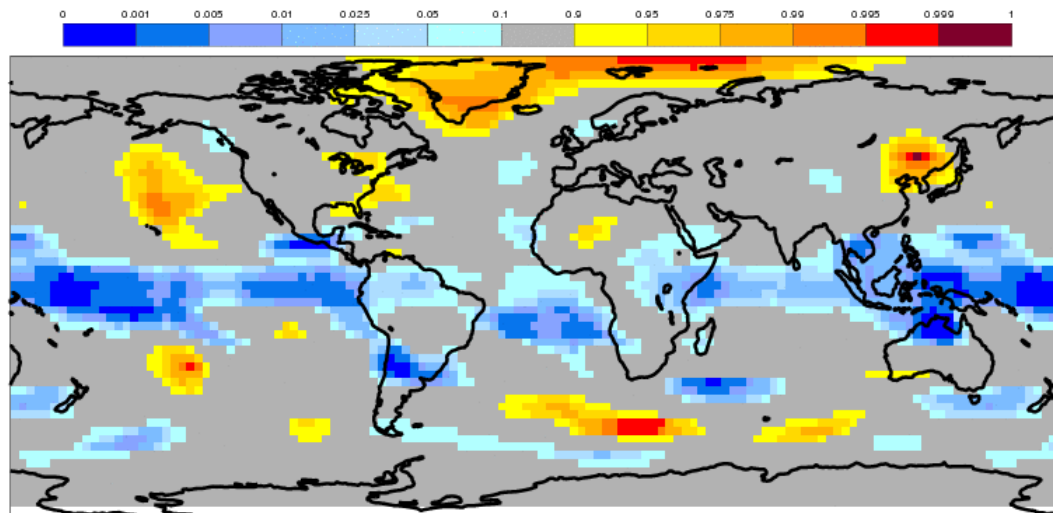
5. Future Challenges and prospects

- NH winter predictability
- QBO teleconnections

Predictive skill vs. Predictability limit

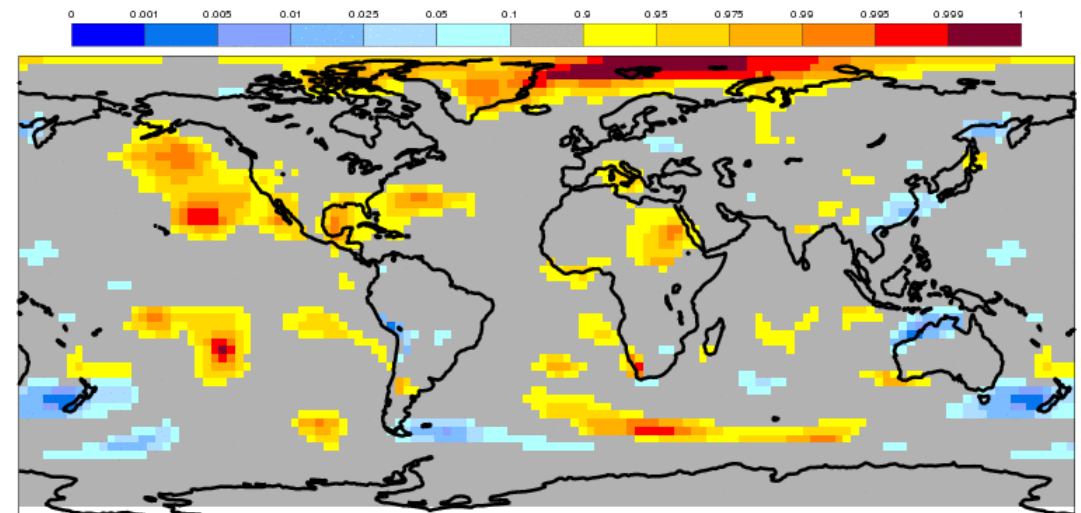
DJF Z500

p-value for observed ACC, assuming perfect model for ECMWF S4 with 51 ensemble members
500 hPa geopotential height
Hindcast period 1981-2010 with start in November average over months 2 to 4



DJF MSLP

p-value for observed ACC, assuming perfect model for ECMWF S4 with 51 ensemble members
Mean sea level pressure
Hindcast period 1981-2010 with start in November average over months 2 to 4



Indistinguishable from perfect

Worse than perfect

Better than perfect

Predictability of the Arctic Oscillation

Predictability can be under-estimated if we miss or under-represent important processes

Scope for improvement

Unbiased variance estimates: Obs/Tot/Int/Ext: 1.0000 0.8390 0.8316

Model/obs stddev ratio: 0.9159

Model/obs stddev ratio interval: 0.693 1.129

Bootstrap over nens, pval for ratio=1: 0.7960

← model variability consistent

=====
SNR actual : 0.0941

SNR jackknife over nens : 0.0202 0.1029 0.1857
=====

=====
ACC actual : 0.6085

ACC basic bootstrap over nens : 0.5568 0.7121 0.8144

ACC basic bootstrap over nyears: 0.2052 0.6069 0.8326
=====

← 95% interval due to ensemble size

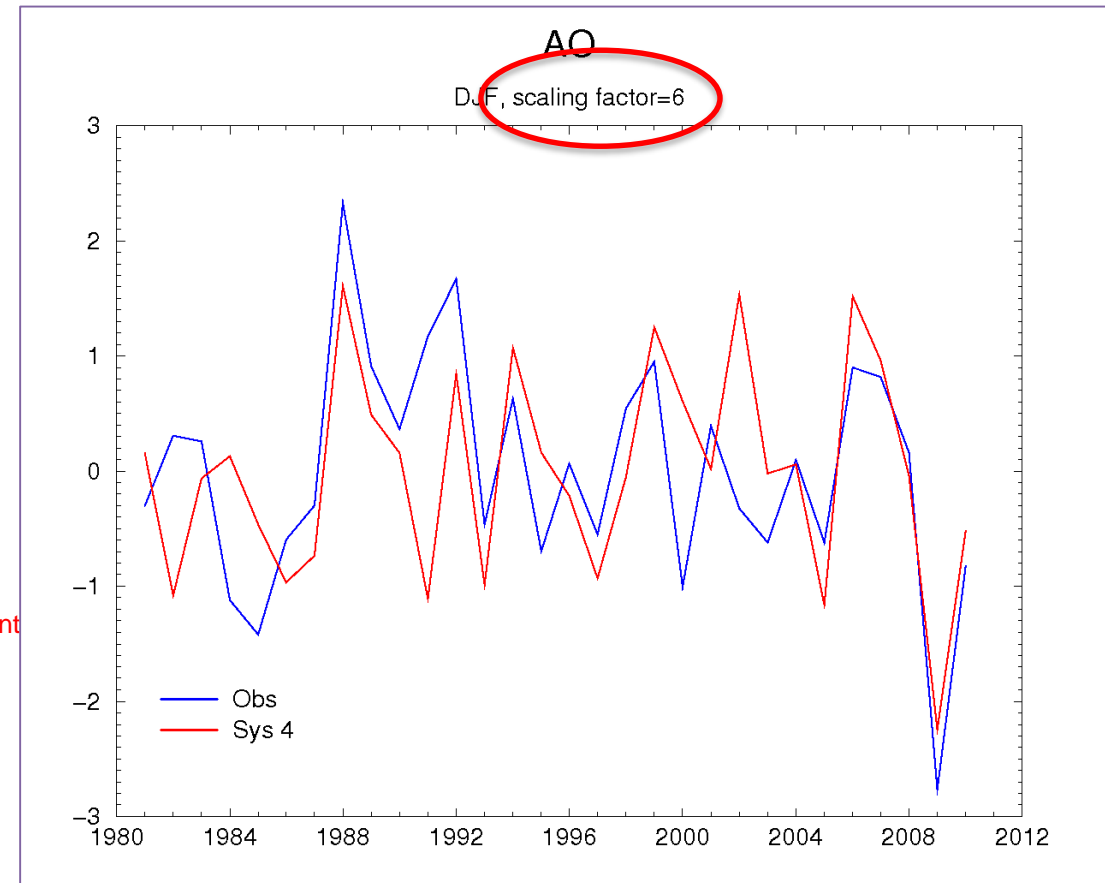
← bigger uncertainty range here

ACP from internal sampling: -0.2947 0.0583 0.4010

Mean ACC for nens-1: 0.6049

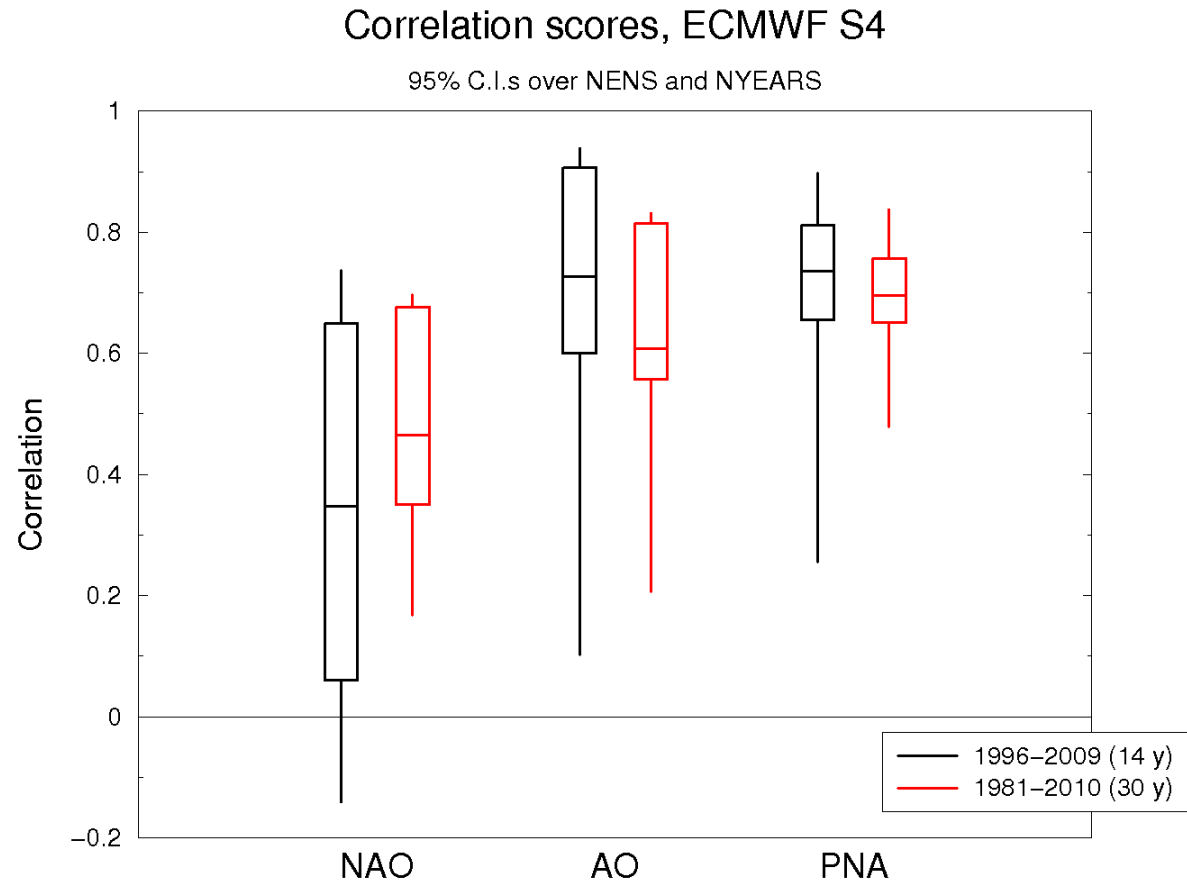
p val of measured acc if model perfect: 0.9996

← only a 0.0004 chance we could get this correlation



Correlation (30y) =0.608

Challenge: sampling errors are large!



Box = 95% interval, bootstrapping
on ensemble size

Whiskers = 95% interval,
bootstrapping on years included

QBO teleconnections – NH winter polar vortex (50 hPa)

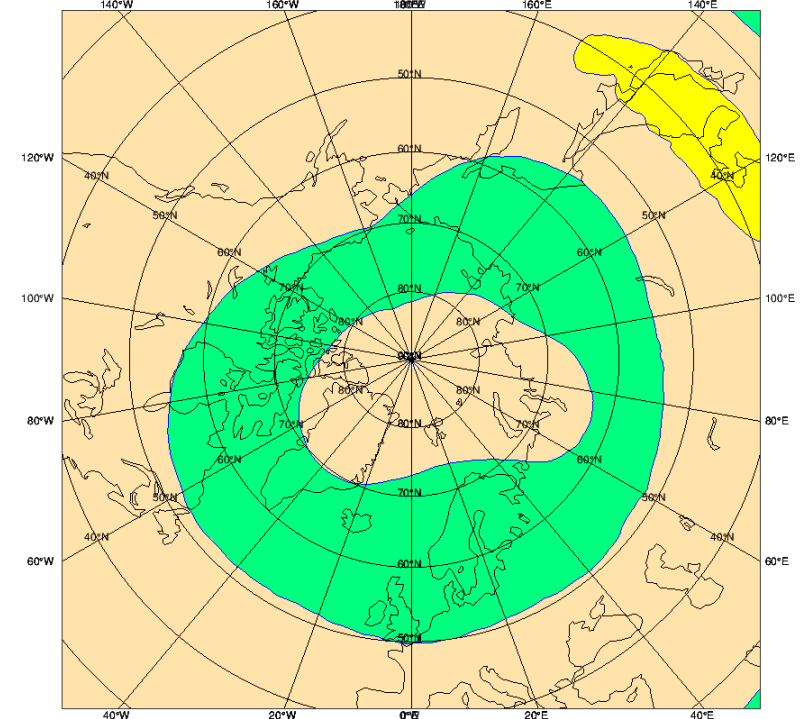
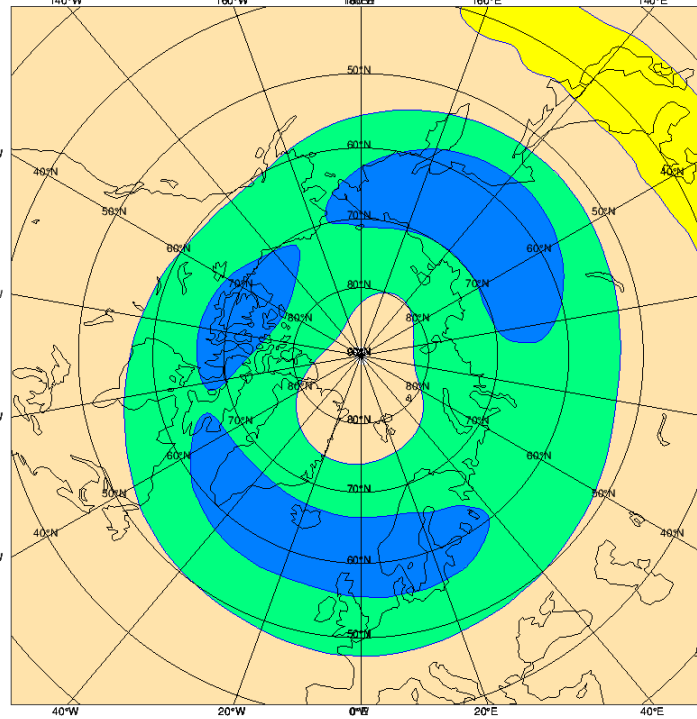
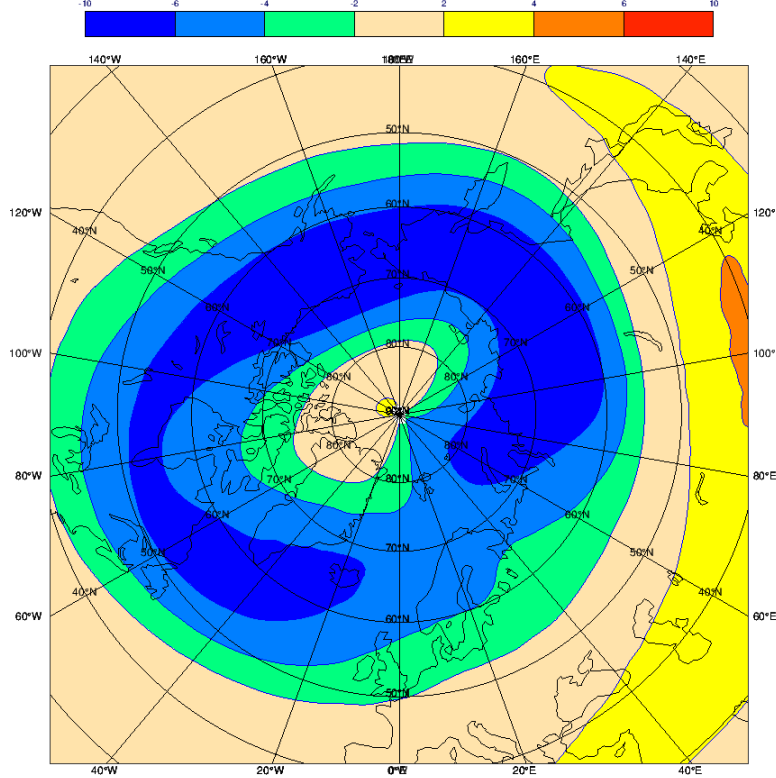
System 4

SEAS5

ERA-Interim DJF mean U at 50 hPa: QBOE - QBOW (10 E years, 12 W years)

S4 DJF mean U at 50 hPa: QBOE - QBOW (10 E years, 12 W years)

S5 DJF mean U at 50 hPa: QBOE - QBOW (10 E years, 12 W years)



Factor 3 weaker than obs

Factor 5 weaker than obs

QBO composite years for 1981-2005, following Boer and Hamilton (2008). Contour interval is 2 m/s for ERAI, 1 m/s for models. Model composites based on 25 member ensemble.

QBO teleconnections – NH winter MSLP

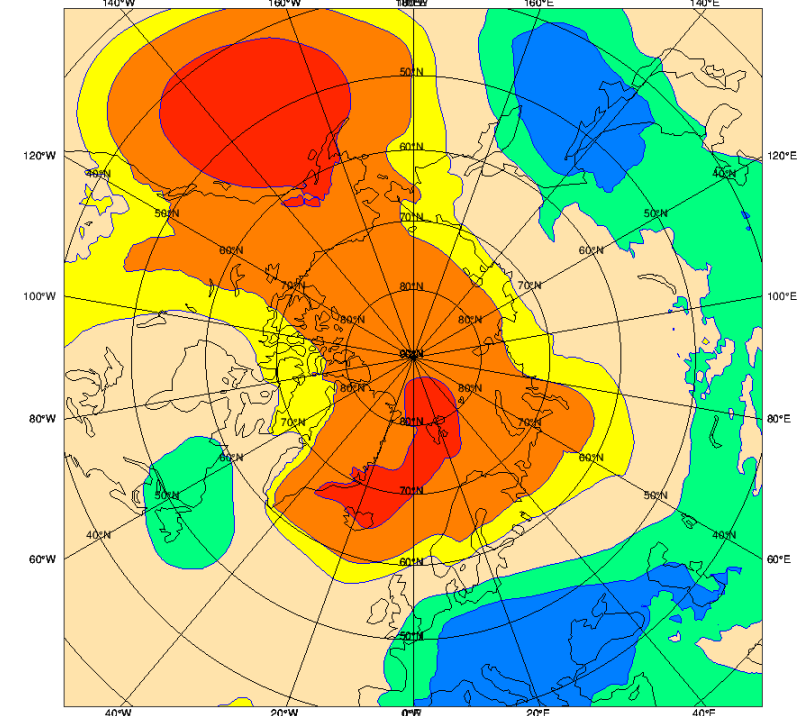
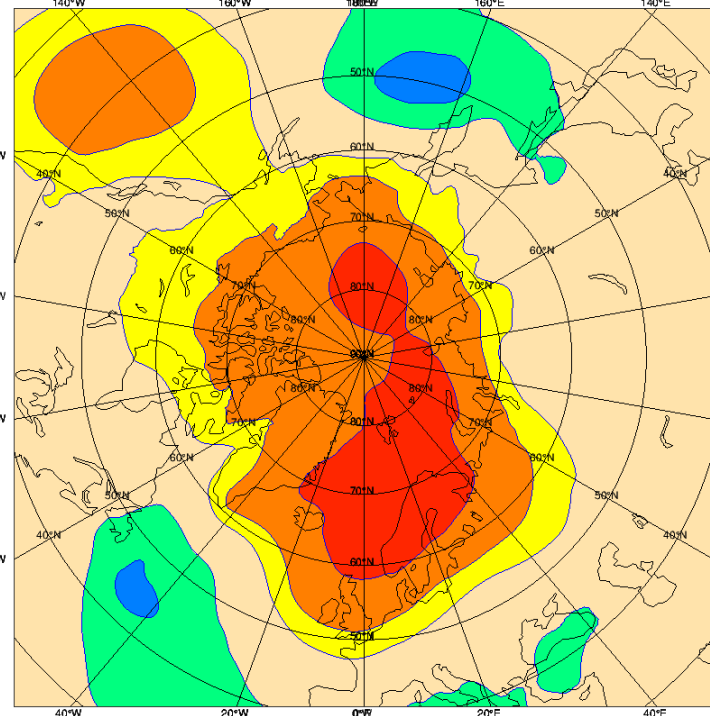
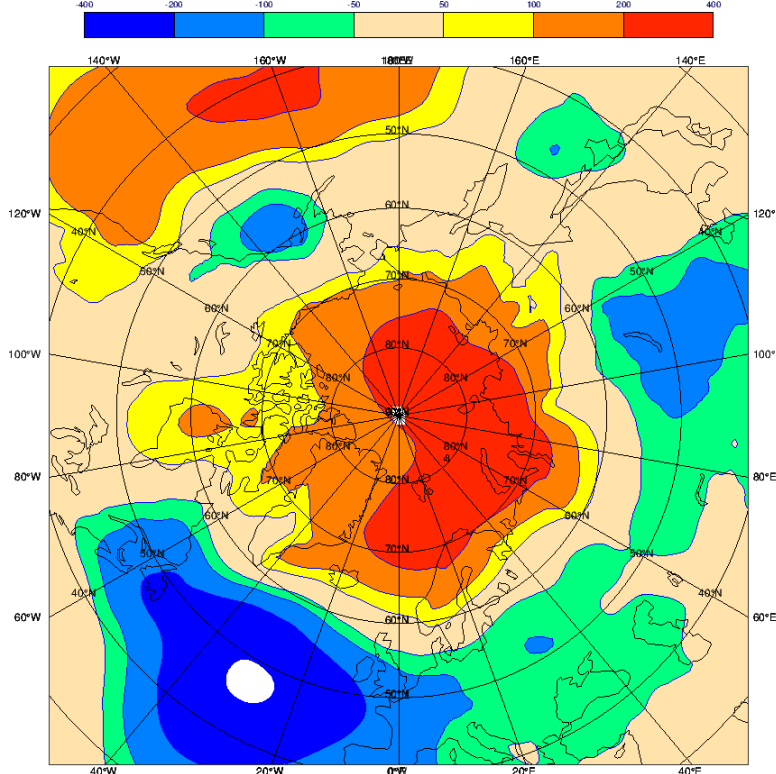
System 4

SEAS5

ERA-Interim (ERA-I) DJF mean mslp: QBOE - QBOW (10 E years, 12 W years)

System 4 (S4) DJF mean mslp: QBOE - QBOW (10 E years, 12 W years)

System 5 (S5) DJF mean mslp: QBOE - QBOW (10 E years, 12 W years)



QBO composite years for 1981-2005, following Boer and Hamilton (2008). Contour interval is 0.5 hPa for ERAI, 0.25 hPa for model. Model composites based on 25 member ensemble.

6. Final perspectives

- Seasonal prediction skill is variable and reliability is moderate, but both skill and reliability are at a level where many applications are possible.
- Care is needed to ensure forecast information is properly interpreted and used sensibly
- Forecasting models are fairly realistic in many ways, but remaining errors are enough to substantially impact forecast skill and reliability, even after calibration
- Creating consistent initial conditions for past and present is a challenge, due in particular to the lack of observational data in the past. Observing systems are better now, but still need some improvements.
- Limited predictability and limited past data prevent us being sure about the skill levels of today's forecast systems, and calibration is therefore subject to uncertainty.
- Although multi-model ensembles are helpful, they only partially span the space of model errors.
- In the end, the only way to achieve high reliability is to build trustworthy models

References and further reading

SEAS5 forecasts on www.ecmwf.int/en/forecasts/charts and <https://climate.copernicus.eu/seasonal-forecasts>

ECMWF Seasonal Forecast User Guide

SPECS fact sheets <http://www.specs-fp7.eu/Fact%20sheets> on seasonal forecasting

Boer, G.J. and K. Hamilton, 2008: QBO influence on extratropical predictive skill. *Clim. Dyn.* 31:987–1000. doi: 10.1007/s00382-008-0379-5

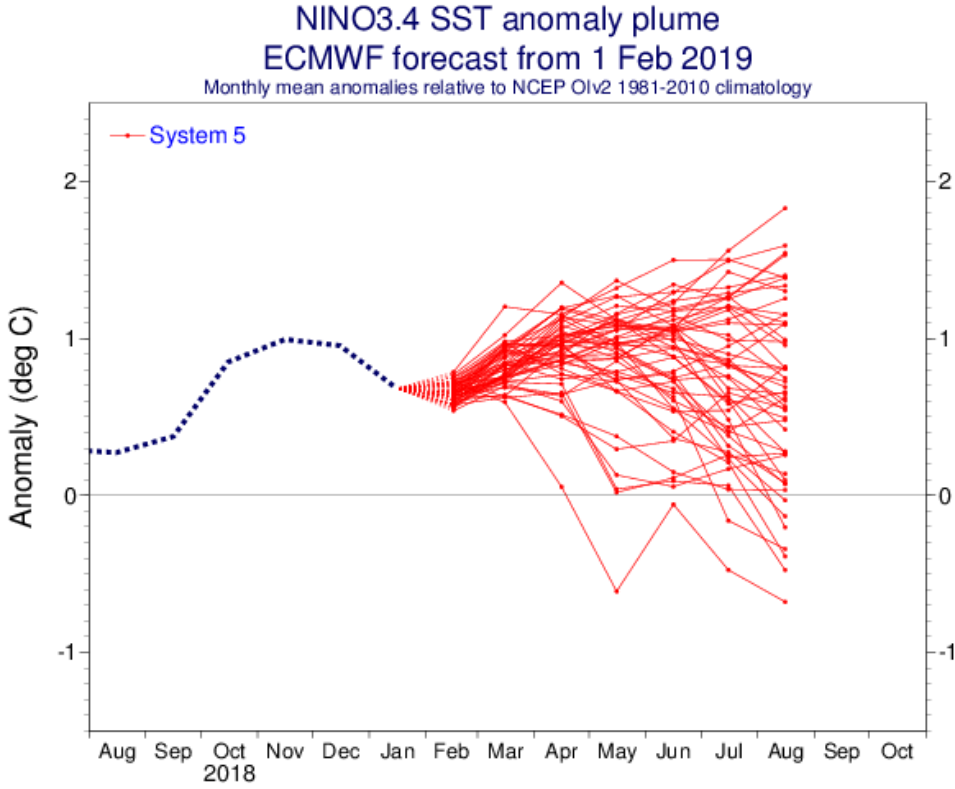
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Stockdale T.N., F. Molteni and L. Ferranti, 2015: Atmospheric initial conditions and the predictability of the Arctic Oscillation. *Geophys. Res. Lett.* doi: 10.1002/2014GL062681

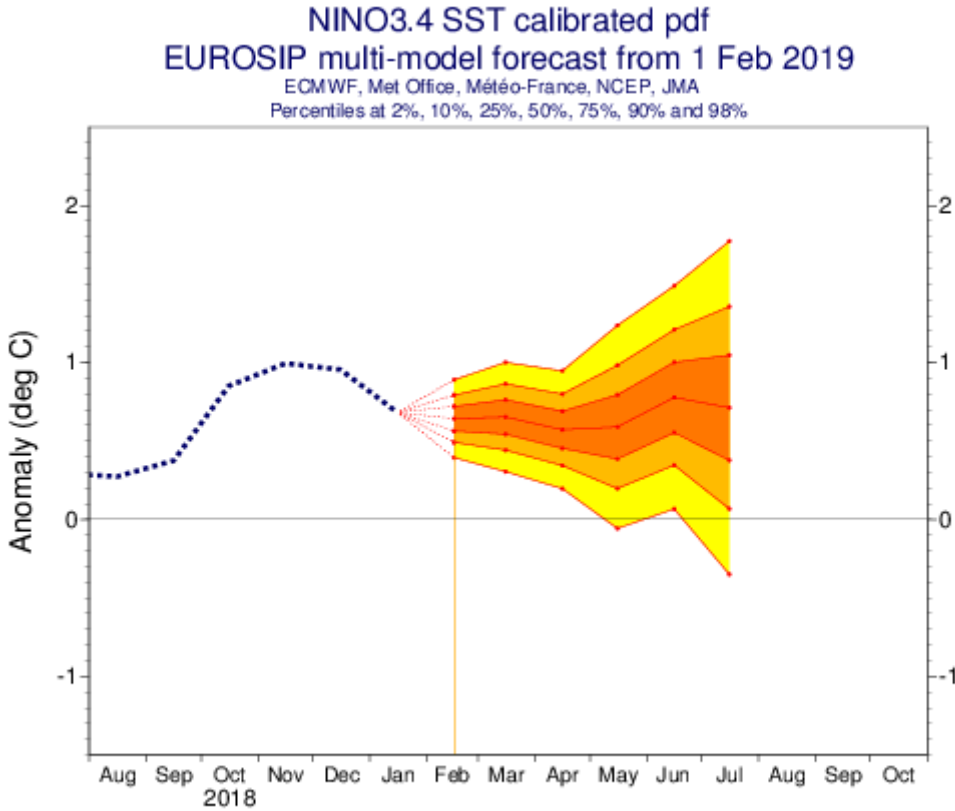
Johnson, S. J., Stockdale, T. N., Ferranti, L., Balmaseda, M. A., Molteni, F., Magnusson, L., Tietsche, S., Decremer, D., Weisheimer, A., Balsamo, G., Keeley, S., Mogensen, K., Zuo, H., and Monge-Sanz, B., 2018: SEAS5: The new ECMWF seasonal forecast system, *Geosci. Model Dev.* DOI: 10.5194/gmd-2018-228

Stockdale, T., Alonso-Balmaseda, M., Johnson, S, Ferranti, L, Molteni, F, Magnusson, L, Tietsche, S, Vitart, F, Decremer, D, Weisheimer, A, Roberts, CD, Balsamo, G, Keeley, S, Mogensen, K, Zuo, H, Mayer, M, and Monge-Sanz, BM, 2018: SEAS5 and the future evolution of the long-range forecast system. ECMWF Tech Memo 835, DOI: 10.21957/z3e92di7y

Latest ENSO forecast



ECMWF



ECMWF

Predicted MAM 2018 (left) and 2019 (right)

ECMWF Seasonal Forecast
 Mean 2m temperature anomaly
 Forecast start is 01/02/18, climate period is 1993-2016
 Ensemble size = 51, climate size = 600

System 5
 MAM 2018

Shaded areas significant at 10% level
 Solid contour at 1% level

ECMWF Seasonal Forecast
 Mean 2m temperature anomaly
 Forecast start is 01/02/19, climate period is 1993-2016
 Ensemble size = 51, climate size = 600

System 5
 MAM 2019

Shaded areas significant at 10% level
 Solid contour at 1% level

