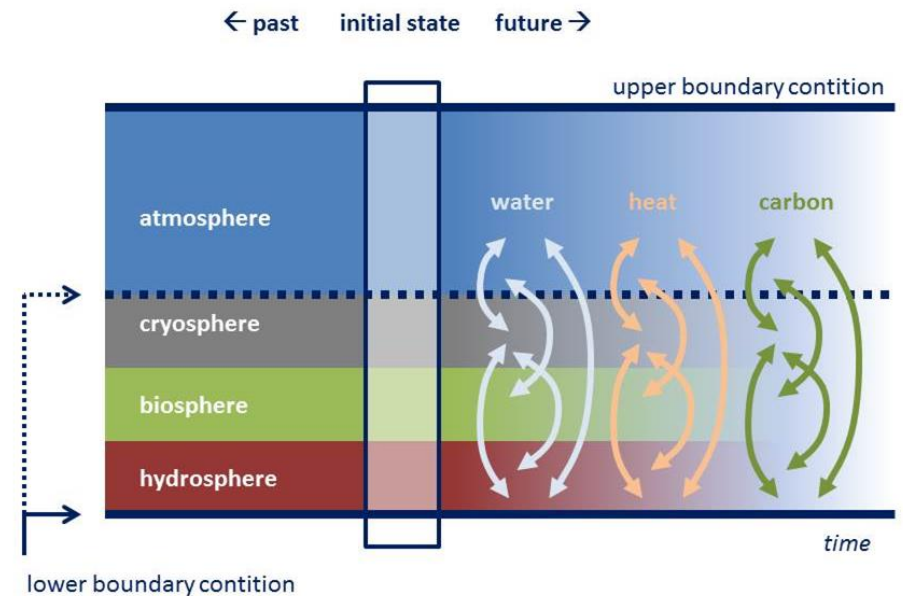


Coupled land-atmosphere variability: does land contribute to predictability?

Bart van den Hurk / Tim Stockdale

Tim.Stockale@ecmwf.int



Why do we care about land processes?

- Energy-budget
 - Albedo

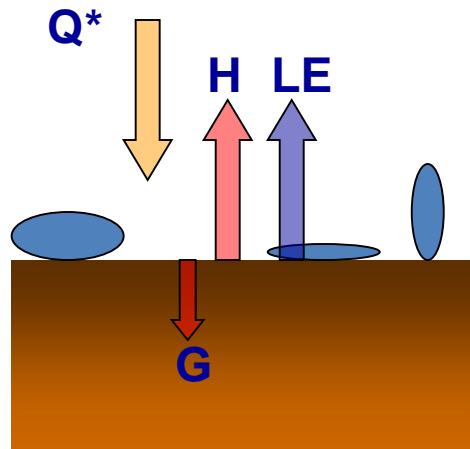
Surface	Albedo
Dark forest	9-12%
Grassland	15-20%
Bare soil	20-30%
Snow in forest	15-25%
Open snow	50-85%



BLD-U10642306 - © - Fotosearch

Why do we care about land processes?

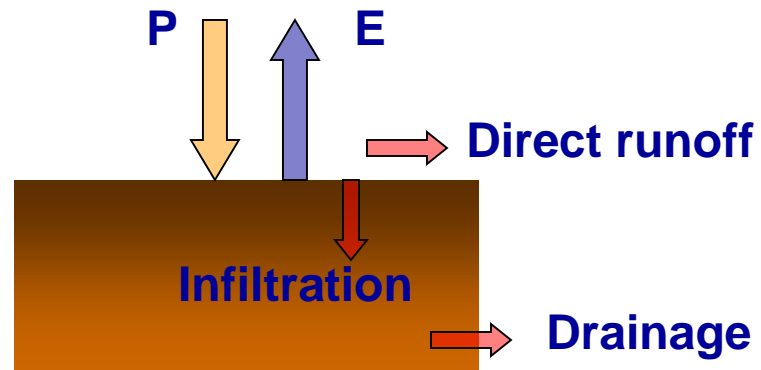
- Energy-budget
 - Albedo
 - Evaporative fraction



Surface	LE/Q^*
Boreal forest	25%
Forest in temperate climate	65%
Dry vineyard	20%
Irrigated field in dry area	100%

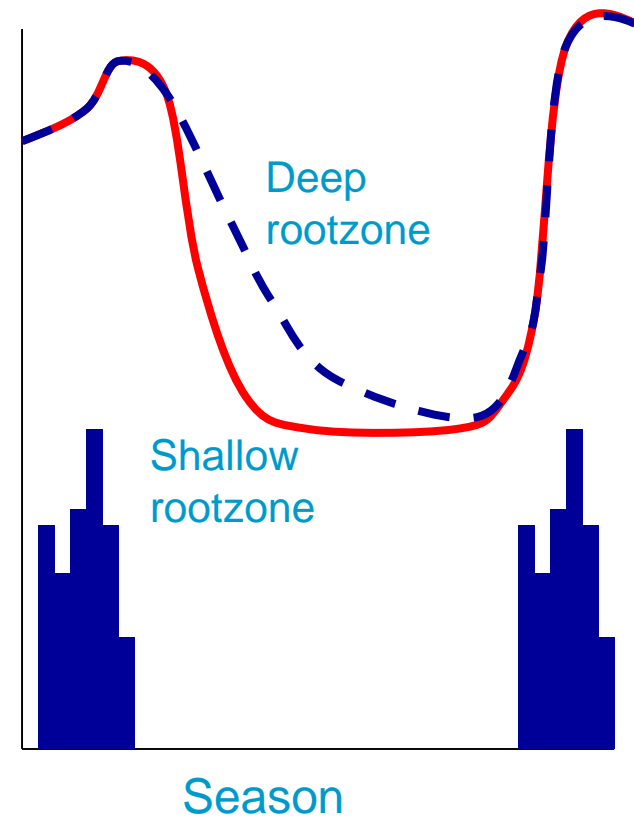
Why do we care about land processes?

- Energy-budget
 - Albedo
 - Evaporative fraction
- Water budget
 - Runoff-fraction



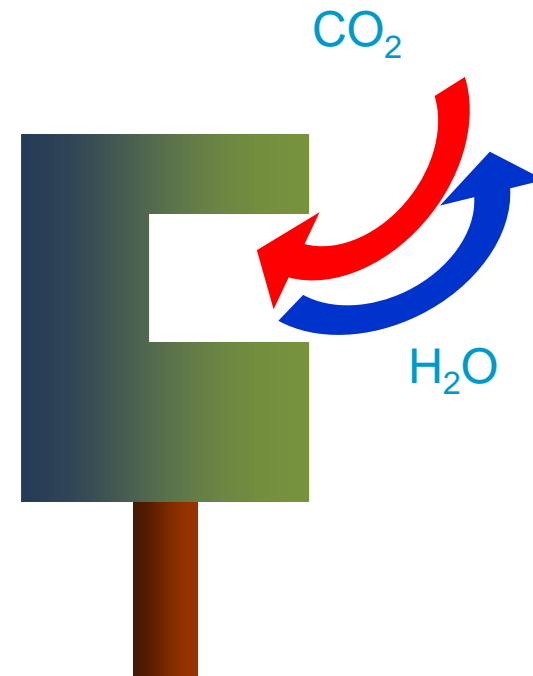
Land processes in atmospheric models

- Energy-budget
 - Albedo
 - Evaporative fraction
- Water budget
 - Runoff-fraction
 - Soil water reservoir



Land processes in atmospheric models

- Energy-budget
 - Albedo
 - Evaporative fraction
- Water budget
 - Runoff-fraction
 - Soil water reservoir
- Carbon budget

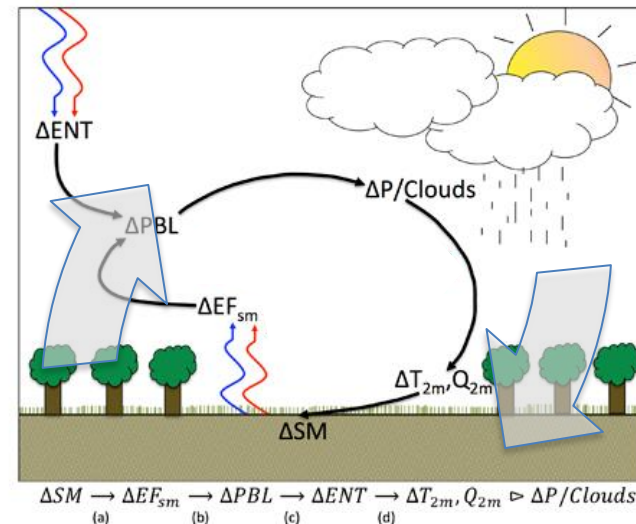


What is needed to contribute to predictability?

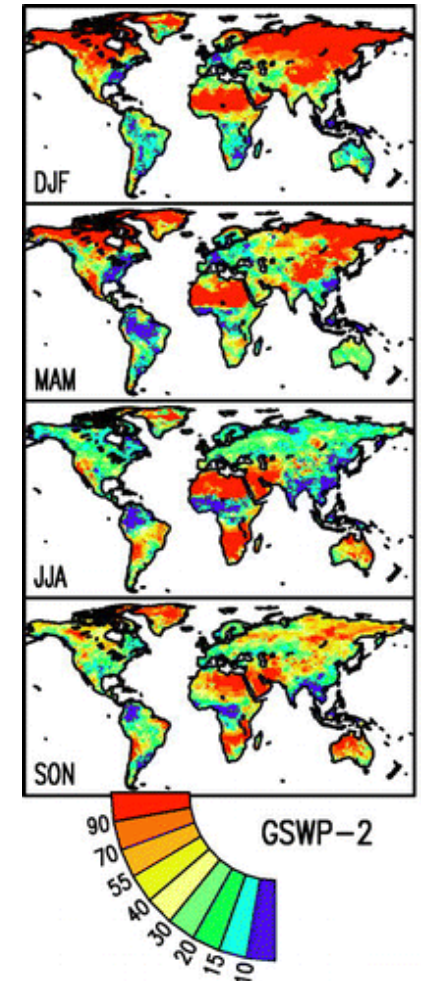
- In the climate system all processes are connected



- A systematic influence of land surface on atmosphere requires:
 - **Variability**
 - **Memory**
 - **Coupling to the atmosphere**

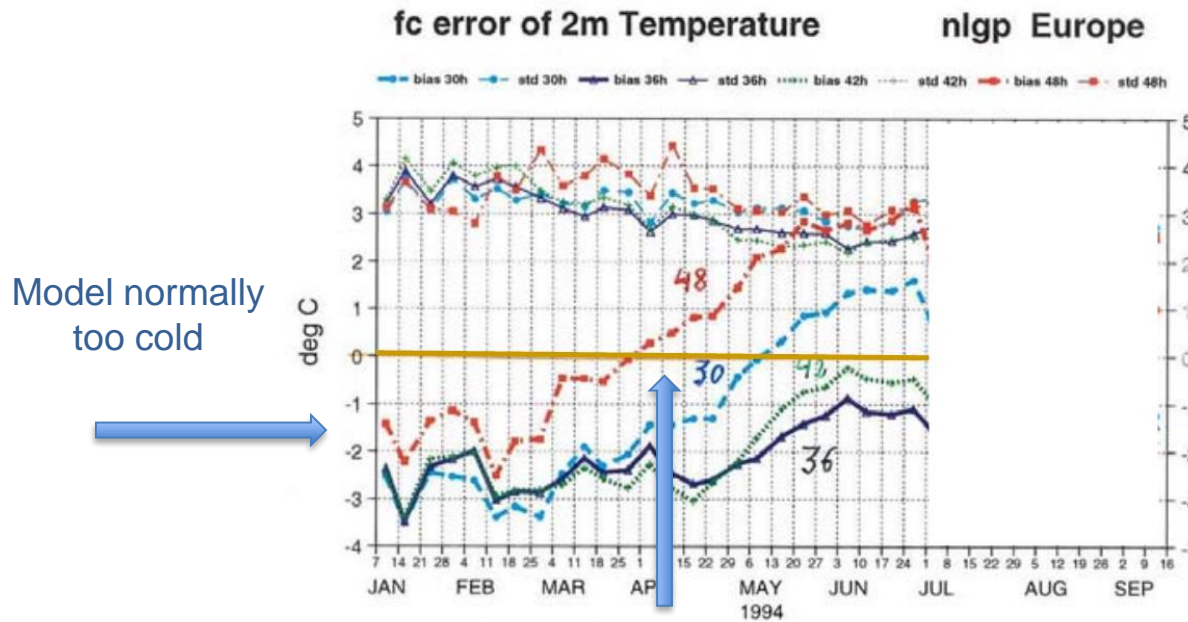


Dirmeyer et al, 2009



An anecdote demonstrating impact of soil moisture

- Mid '90's: introduction of **prognostic soil moisture** scheme



Model normally too cold



Suddenly extreme warm dry drift

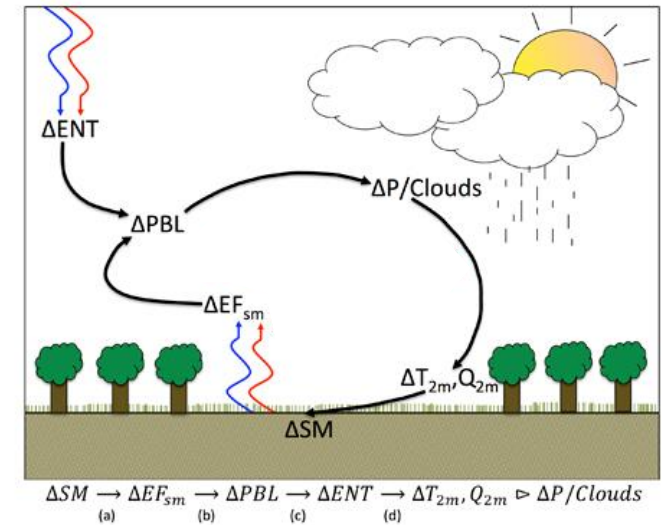
- Soil moisture **data assimilation** needed to control drift

(Root cause of drift was model bias, but once unphysical constraint was removed, model bias led to errors that grew over time)

Dry atmosphere, too little clouds

START HERE

Positive radiation bias

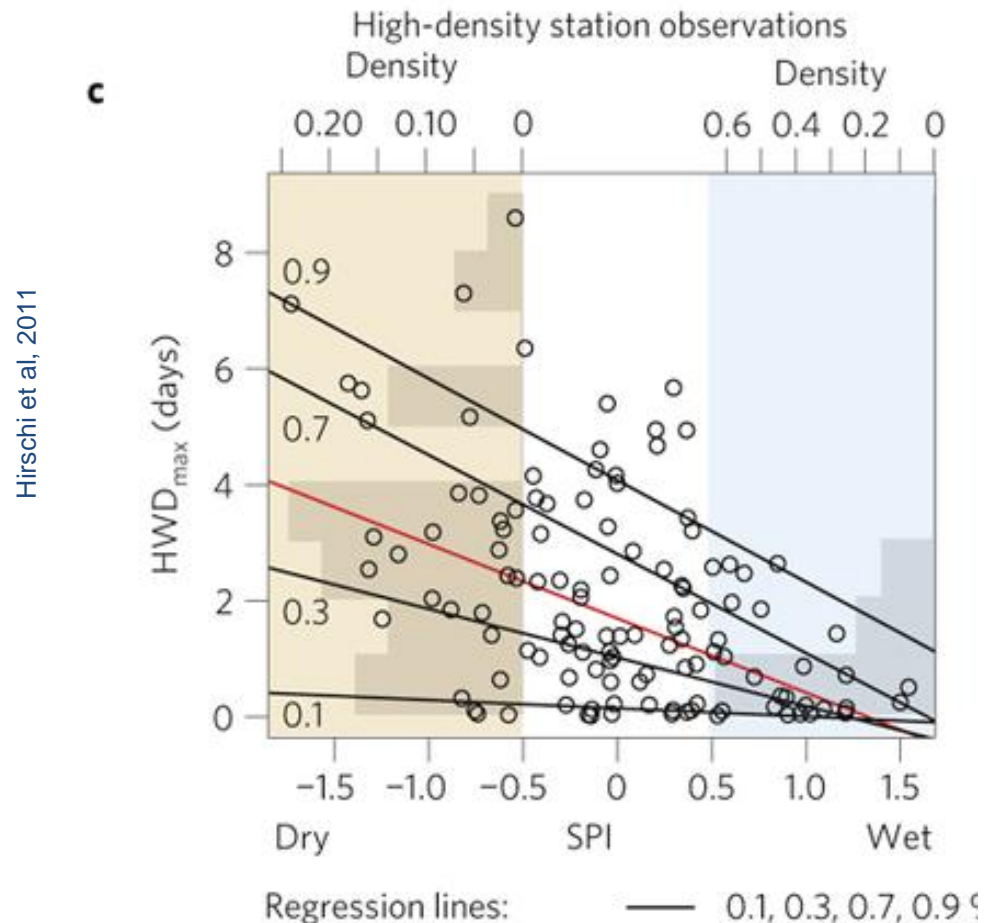


Evaporation stops, Less land cooling

Soil drying due to overestimated evaporation

Measures to quantify land-atmosphere coupling

- From observations:
 - relation between (soil) wetness and extreme temperatures

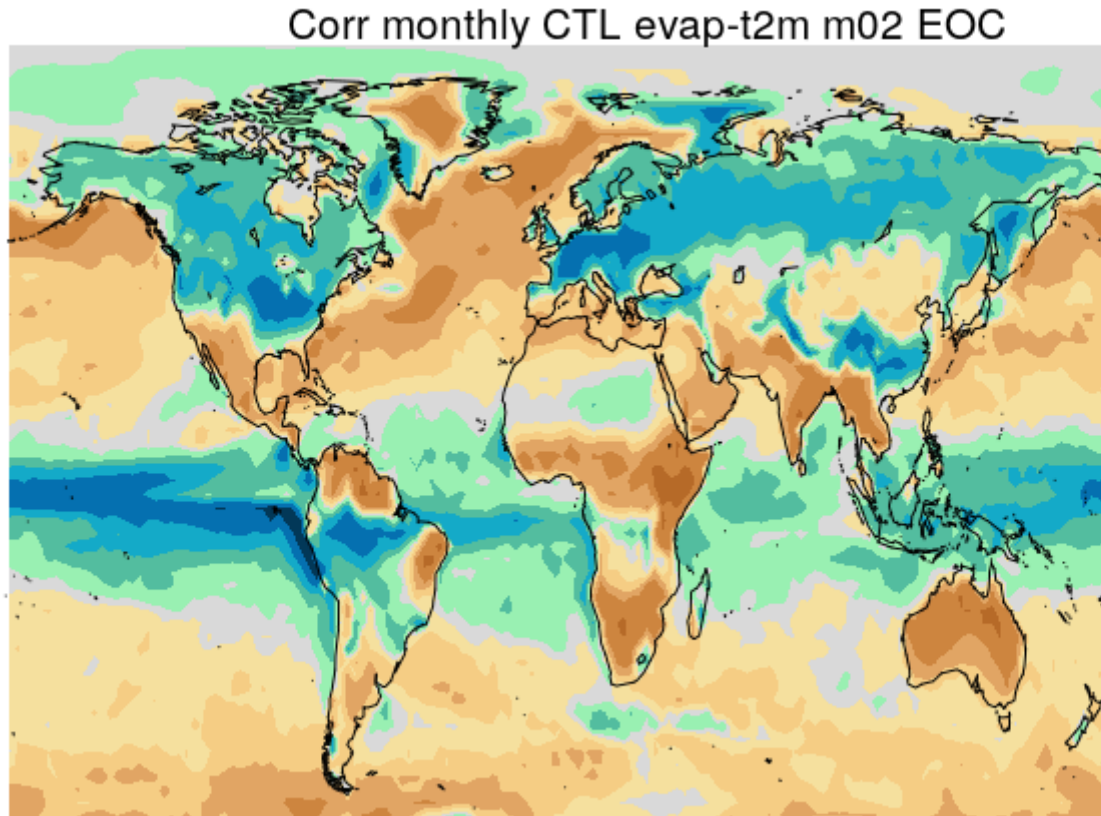


Predictability over wet conditions better than over dry conditions

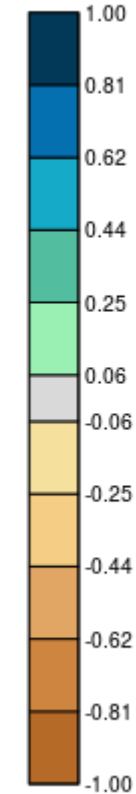
Measures to quantify land-atmosphere coupling

- From (pseudo)observations:
 - Correlation between evaporation and temperature

Feb-Apr

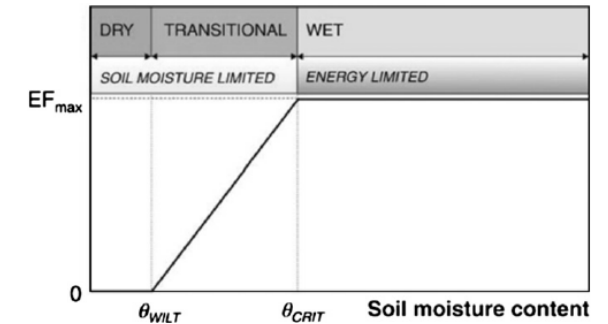


Energy limited



Soil water limited

Seneviratne et al, 2010

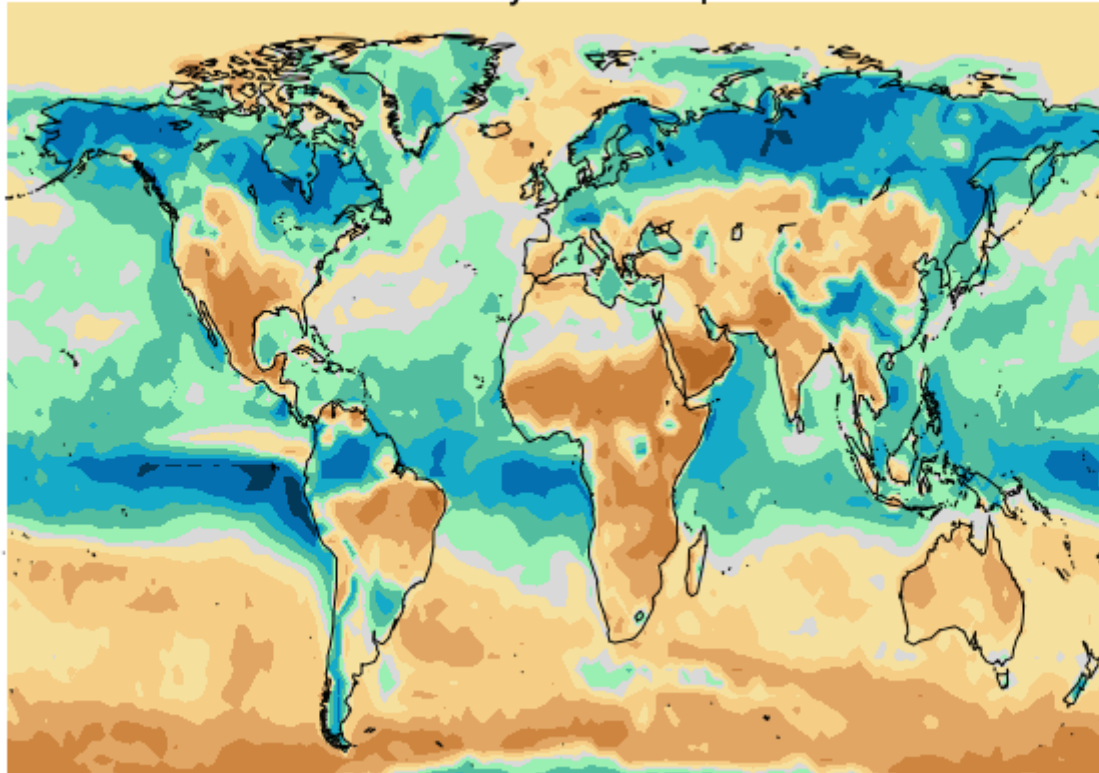


Measures to quantify land-atmosphere coupling

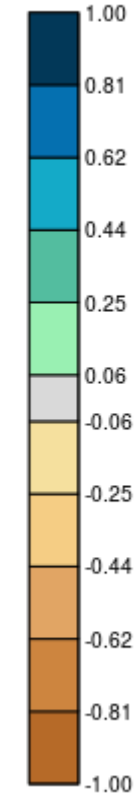
- From (pseudo)observations:
 - Correlation between evaporation and temperature

May-Jul

Corr monthly CTL evap-t2m m05 EOC

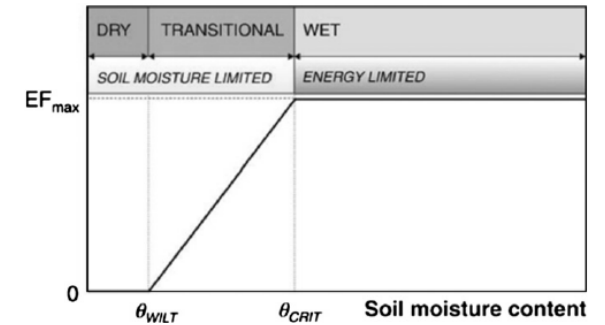


Energy limited



Soil water limited

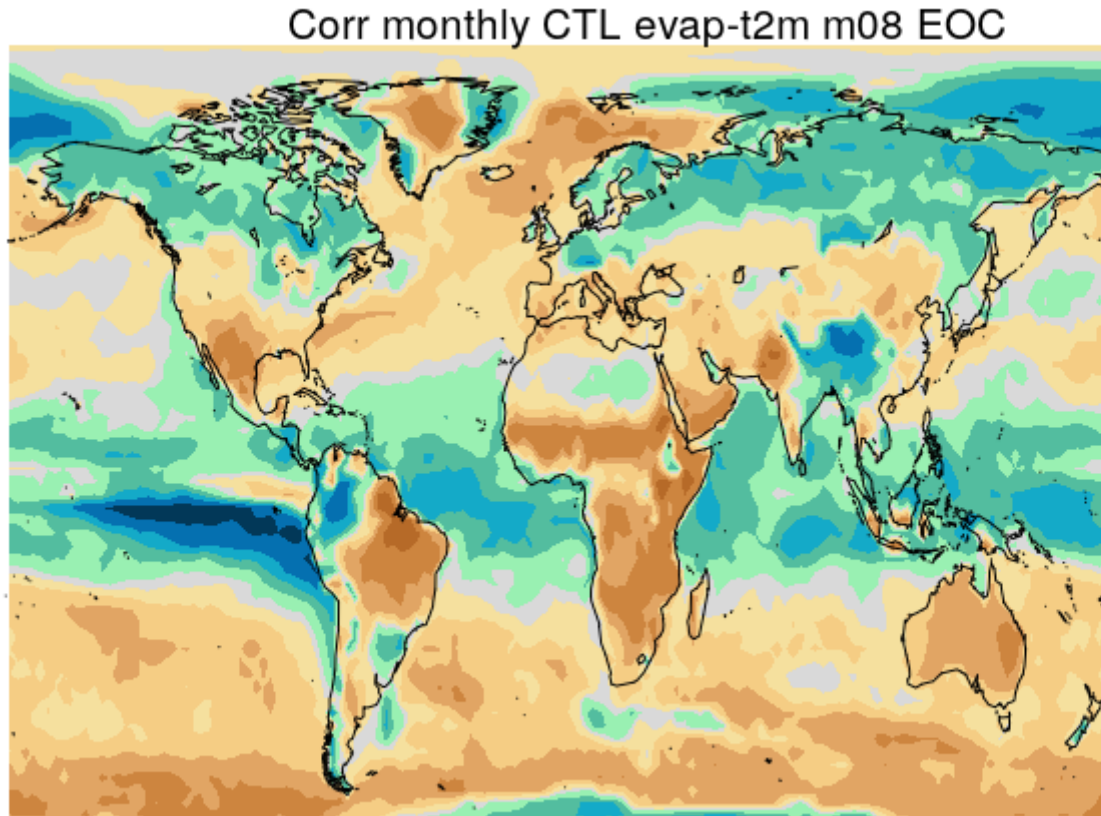
Seneviratne et al, 2010



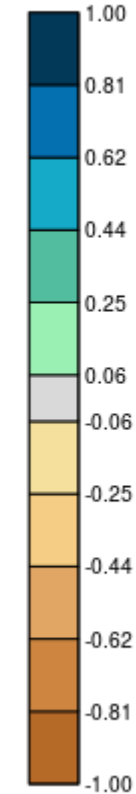
Measures to quantify land-atmosphere coupling

- From (pseudo)observations:
 - Correlation between evaporation and temperature

Aug-Oct

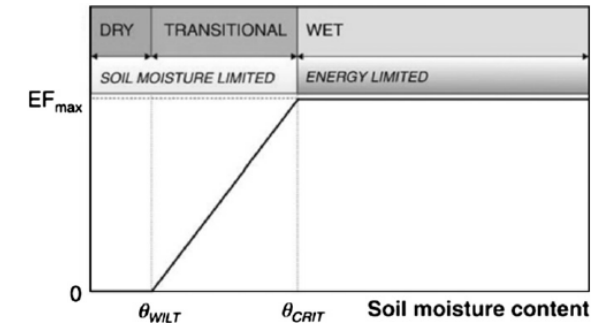


Energy limited



Soil water limited

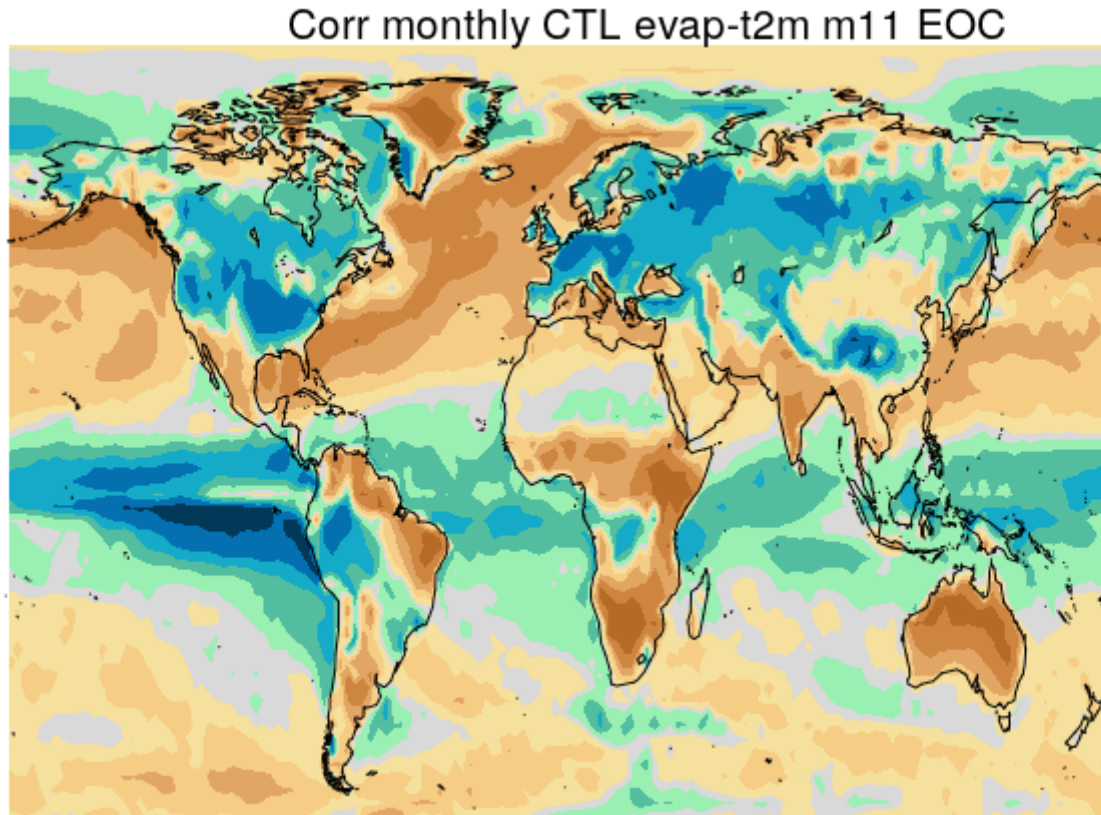
Seneviratne et al, 2010



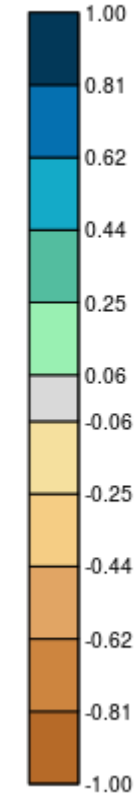
Measures to quantify land-atmosphere coupling

- From (pseudo)observations:
 - Correlation between evaporation and temperature

Nov-Jan

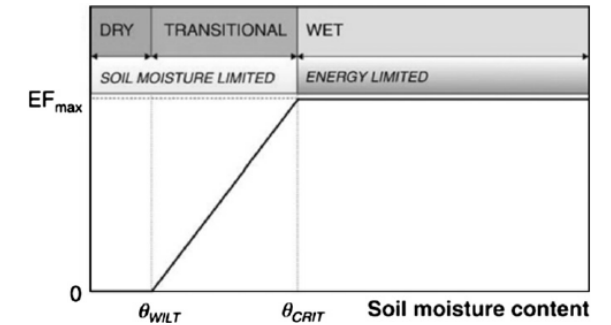


Energy limited



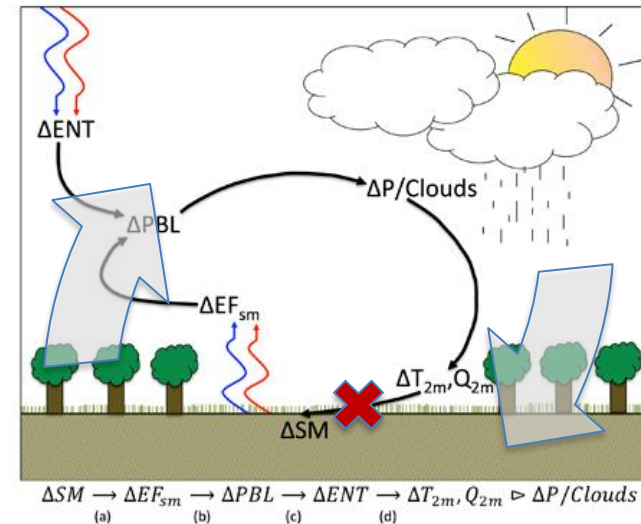
Soil water limited

Seneviratne et al, 2010

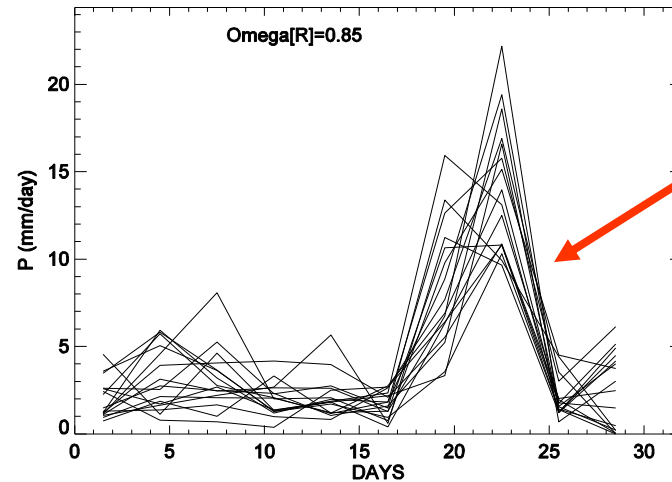


Measures to quantify land-atmosphere coupling

- From a model experiment (GLACE = Global Land Atmosphere Coupling Experiment)
- How?
 - Simulate the hydrological cycle **with** and **without** interactive land-atmosphere coupling and compare.
- How to remove coupling?
 - In second ensemble, replace soil moisture by values from one of the integrations in the first (interactive) ensemble.
- How to measure the effect?
 - Ensemble simulations
 - Compare within-ensemble spread

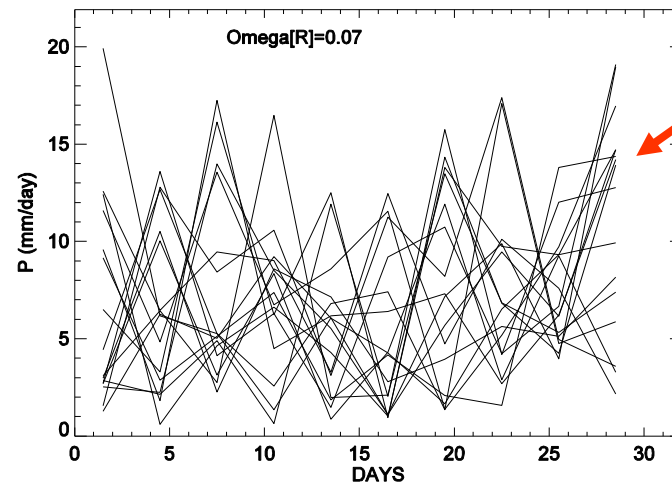


Comparison between ensembles



All simulations in ensemble respond to the land surface boundary condition in the same way

→ strong coupling



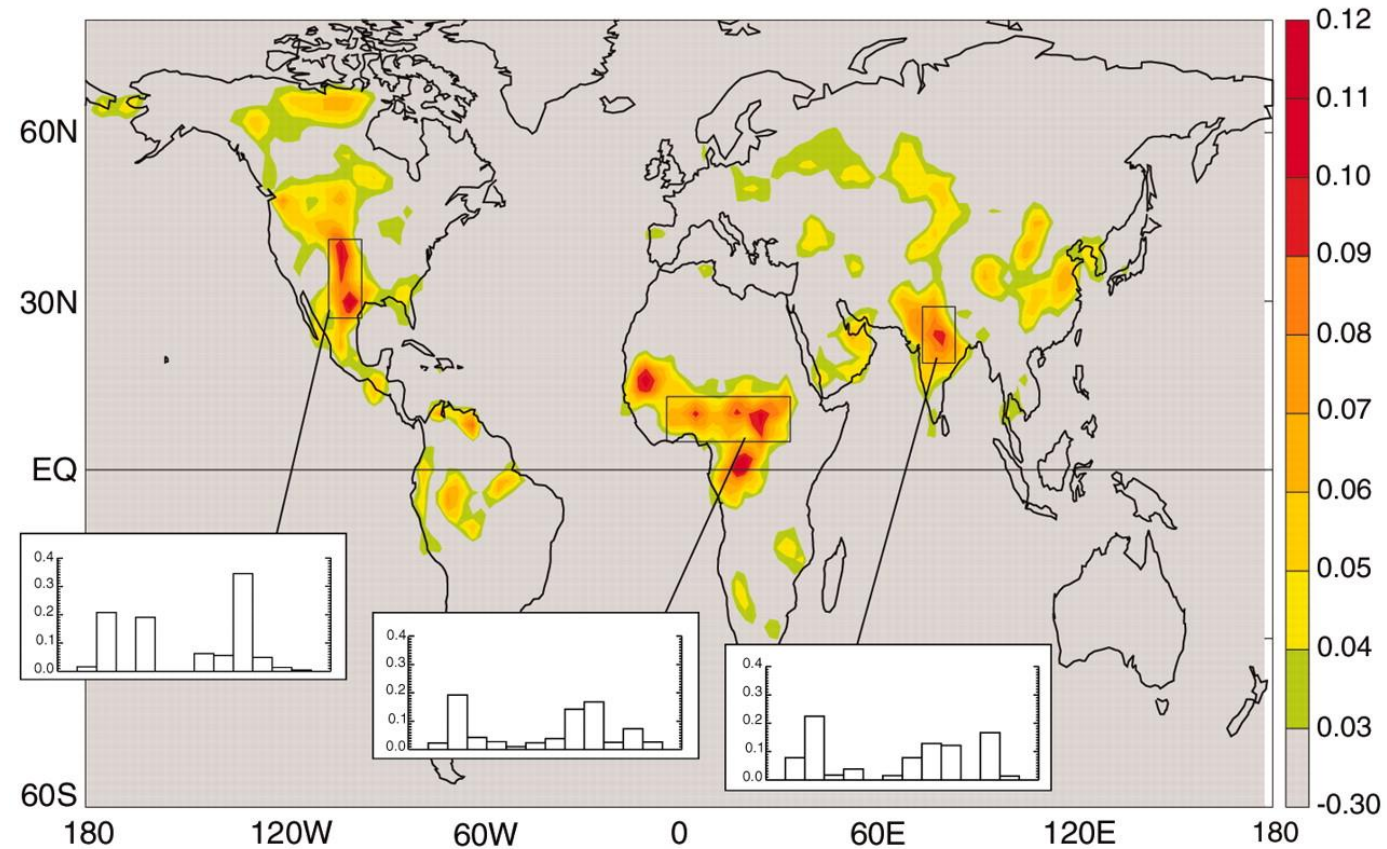
Simulations in ensemble have no coherent response to the land surface boundary condition

→ weak coupling

$$\Omega = \frac{\sigma_P^2(W) - \sigma_P^2(S)}{\sigma_P^2(W)}$$

Areas with strong feedback

Land-atmosphere coupling strength (JJA), averaged across AGCMs

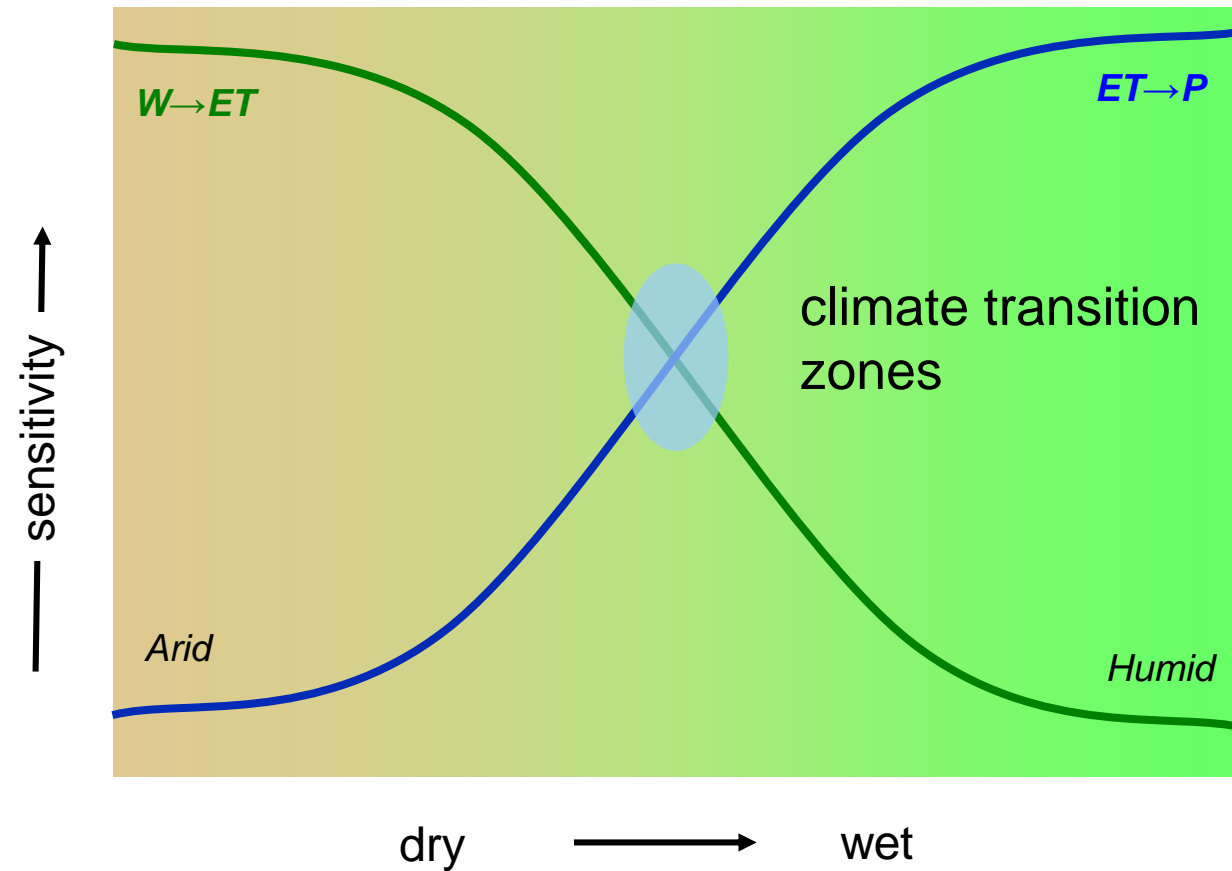


$$\Omega = \frac{\sigma_P^2(W) - \sigma_P^2(S)}{\sigma_P^2(W)}$$

This is a famous figure, and looks very nice. But note that different models gave substantially different results. Model representation of land surface processes is improving, but still has some way to go.

Koster et al, 2004, Science

Strong coupling needs combination of sensitivities

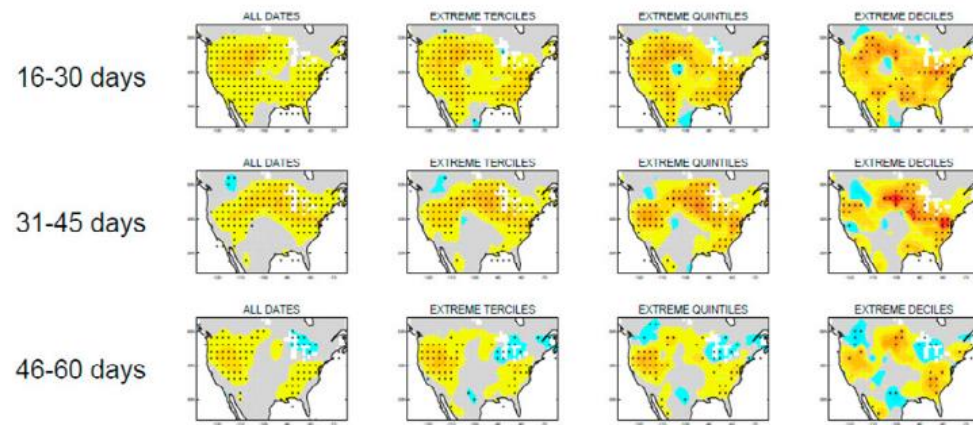


Some “real” land-surface predictability experiments

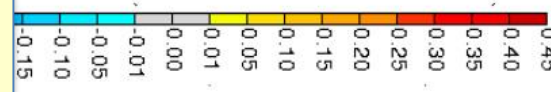
- Global Land Atmosphere Coupling Experiment – 2
 - Compare 2 ensembles of seasonal forecasts (8 weeks ahead)
 - Ensemble 1: all members use the same realistic initial conditions
 - Ensemble 2: every member gets a randomly selected initial condition
 - Measure R^2 difference using real observations

Van den Hurk et al, 2012

1b. AIR TEMPERATURE FORECAST SKILL (r^2 with land ICs minus r^2 w/o land ICs)



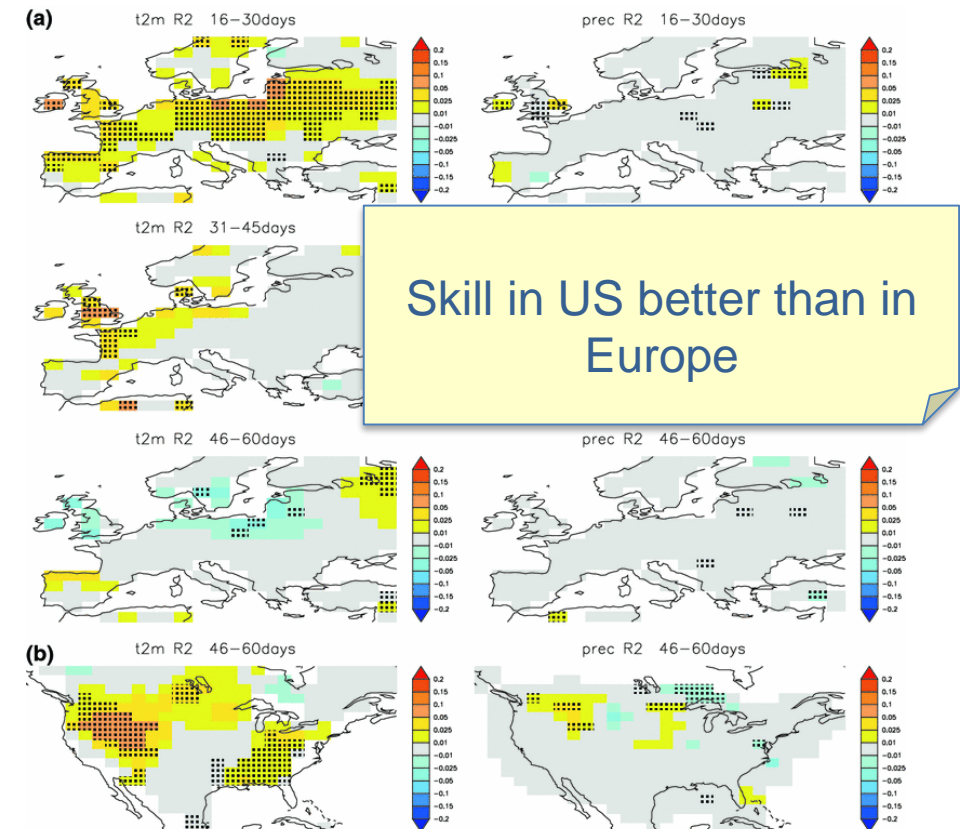
Dates for conditioning vary w/location



Skill improves for more extreme conditions



Koster et al, 2010



Skill in US better than in Europe

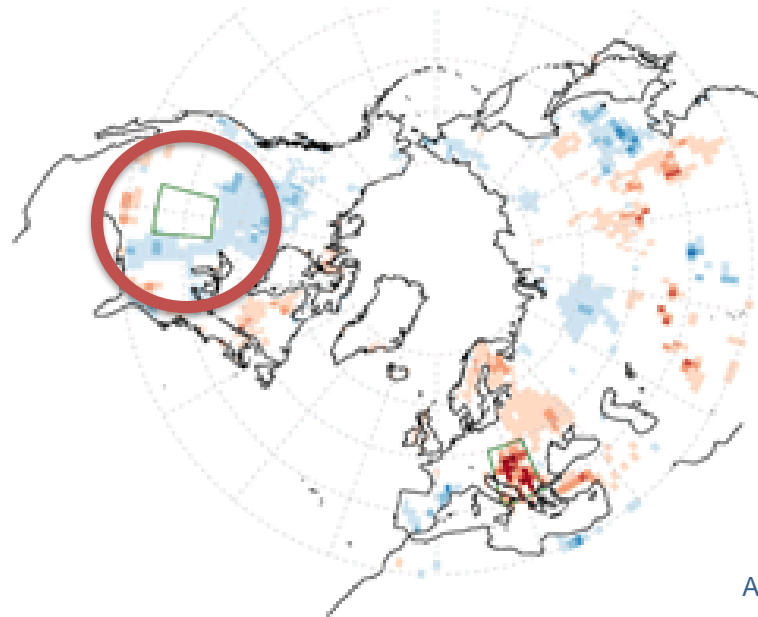
Another experiment, similar set-up, different results!

- Similar to GLACE-2, multi-model study (5 models), but
 - comparing realistic versus climatological initial conditions
 - coupled ocean model instead of prescribed SSTs
 - Longer period (19 yrs instead of 10 yrs)

Model bias in correlation between soil moisture and temperature gives poor results in US

(Models have dry bias, which results in a too-strong sensitivity of T2m to initial soil moisture).

RMS skill INIT – CLIM



Ardilouze et al, 2017



Prediction of an individual event

- European heat wave 2003
- Different set-ups of ECMWF forecasting system

Combination of land surface and atmosphere is needed to improve forecasts

Old model

New model

New model
(old land surface)

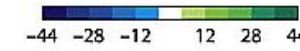
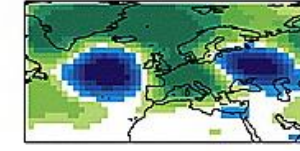
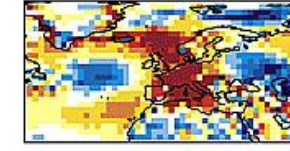
New model
(old radiation)

New model
(old convection)

Temperature anomaly

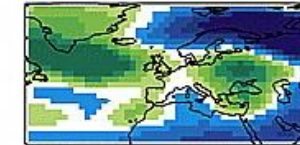
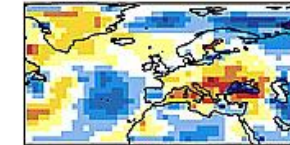
Z500 anomaly

a Analysis

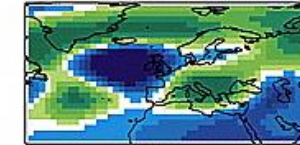
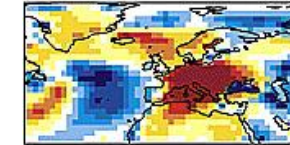


OBS

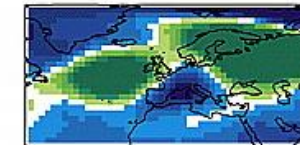
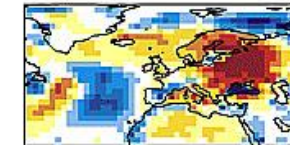
b System 3



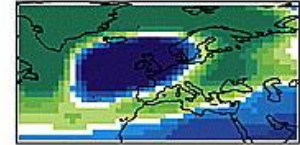
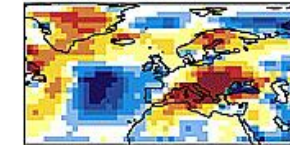
c Cy33r1



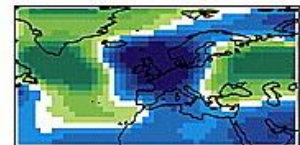
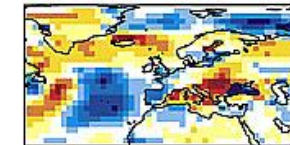
d Cy33r1 with TESSEL



e Cy33r1 - radiation



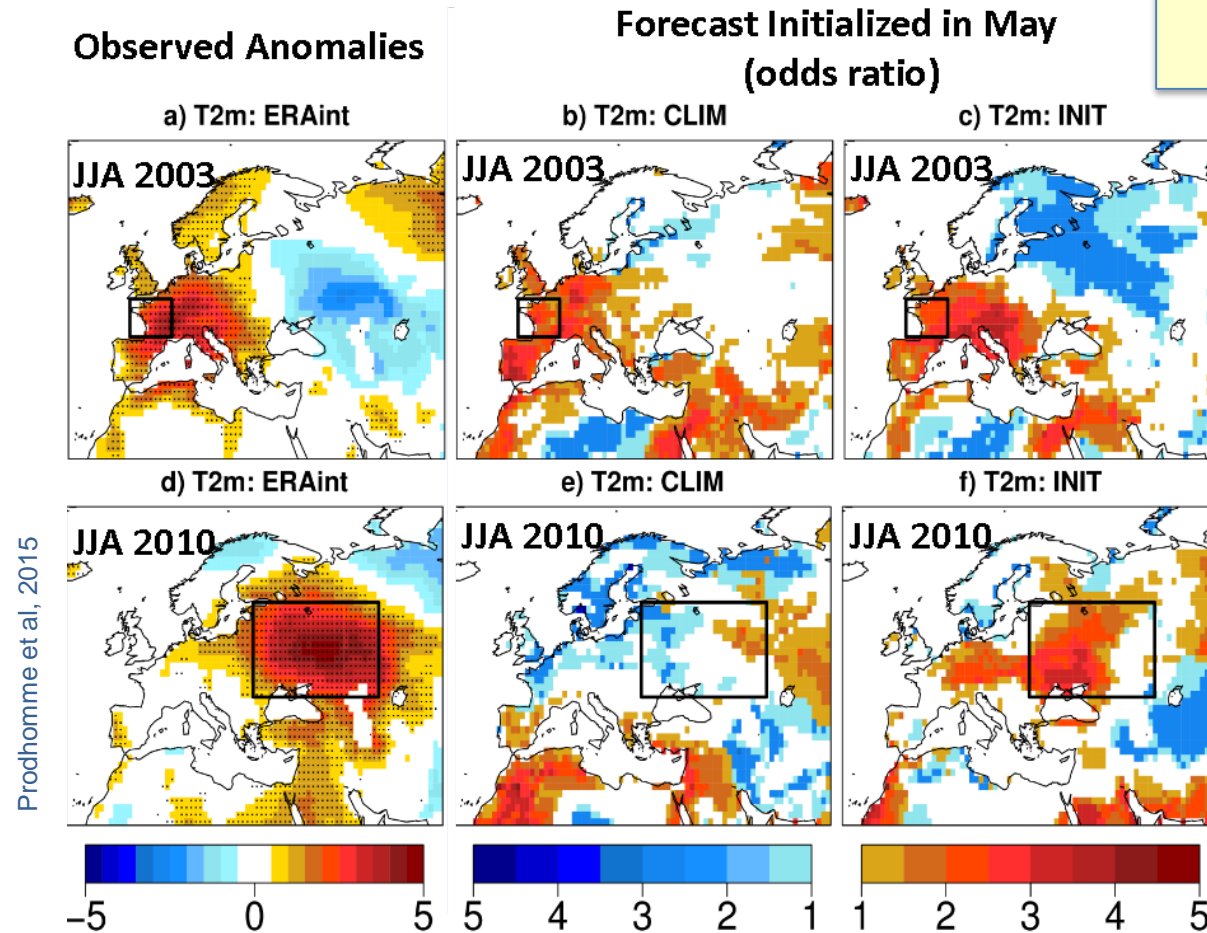
f Cy33r1 - convection



New study, somewhat different results

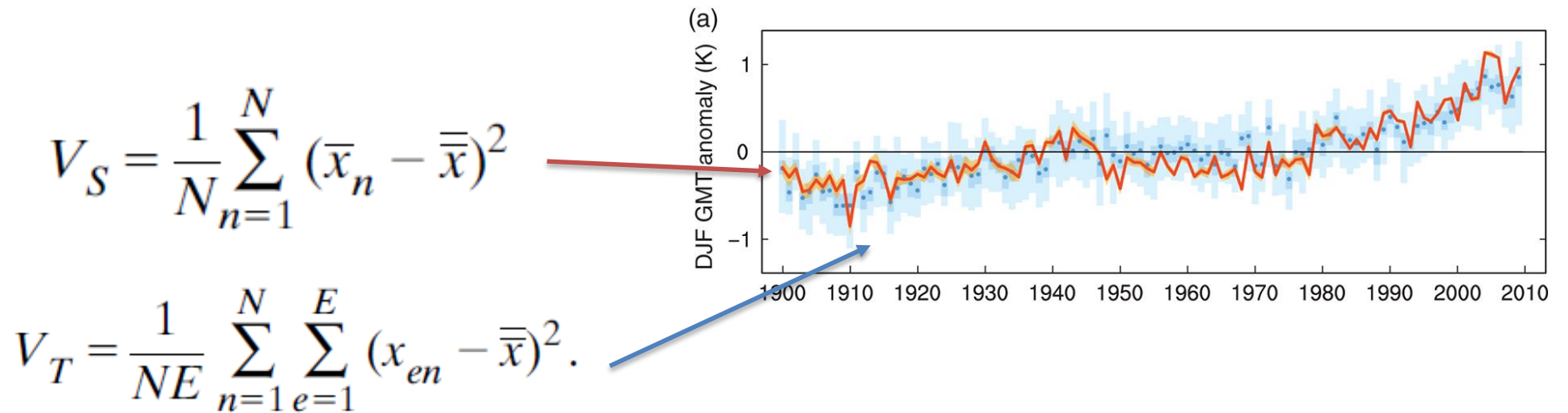
- 5 models, comparing INIT with CLIM initialization
- Start date 1 May, evaluation JJA

European heatwave 2003
is less affected by soil
than Russian heatwave
2010

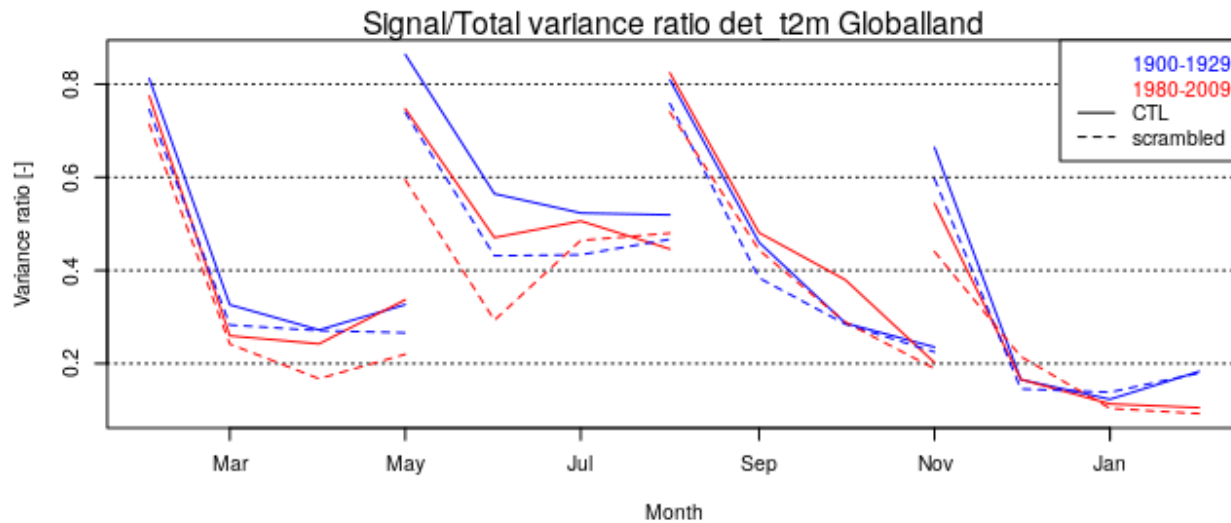
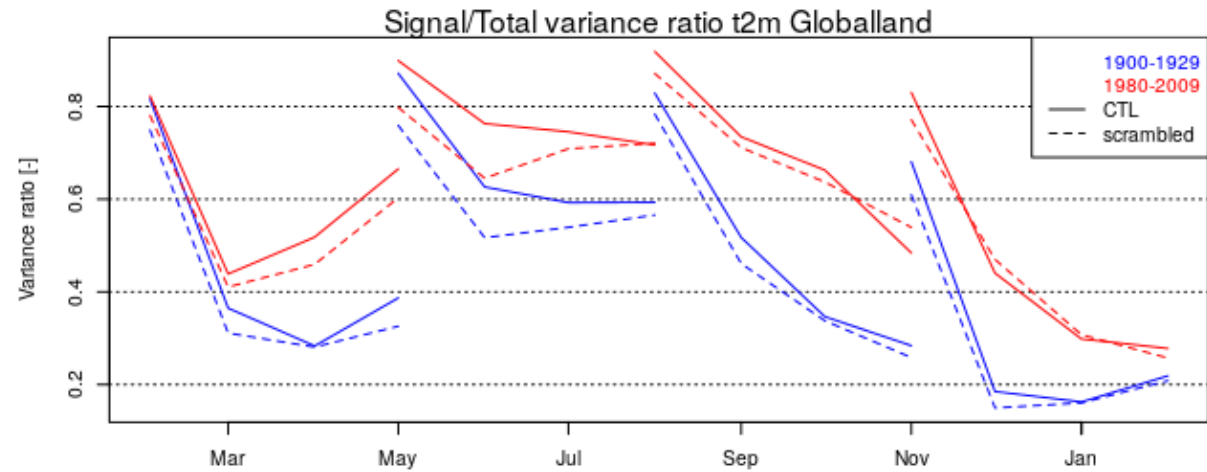


How about trends in predictability?

- Can we see climate trends in predictability?
 - Model experiment: compare ensemble seasonal forecasts 1900-1929 to 1980-2009
- Can we see trend in land surface contribution to this predictability?
 - Model experiment: same forecasts but with random initial land conditions
- Metric: ratio between **signal** and **total** variance

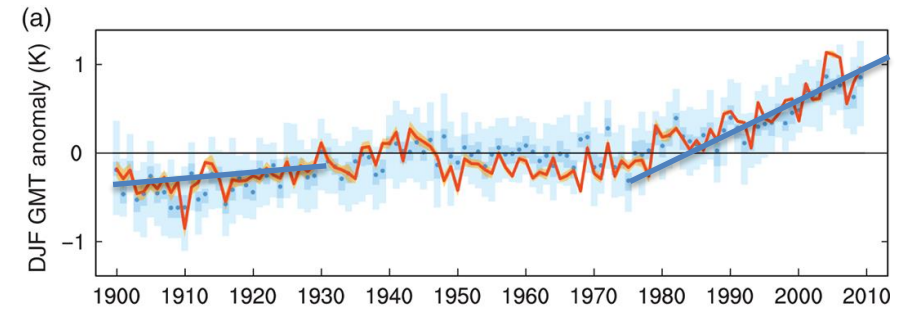


Trend contributes to predictability



Note: initialized land surface (solid line) gives additional signal in T2m, especially in early summer. Note these plots do not show skill – extra skill would require the additional signal to be correct.

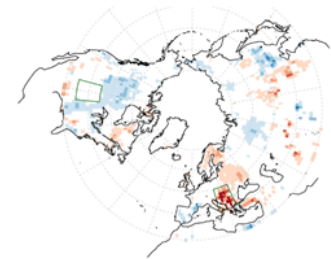
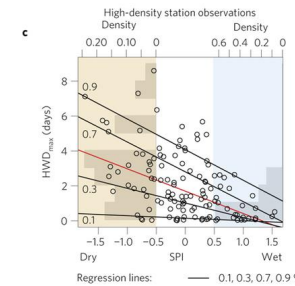
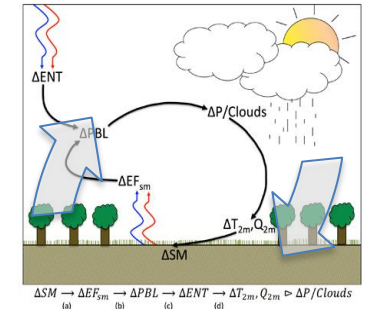
Before detrending



After detrending

Conclusions

- For land-related predictability we need
 - Variability
 - Memory
 - Coupling
- Predictability affects multiple time scales which can interact
 - Predictions of heatwaves → short time scales
 - Predictions of long warm/cool spells → seasonal time scales
- Land surface signal is moderately small in a noisy climate system
 - We need unbiased model systems...
 - ... and pretty large ensembles and long periods
- Land surface initialization *is* improving over time
 - Seasonal t2m skill is improving faster than other fields
 - We believe further progress is possible



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- S. I. Seneviratne et al., 'Investigating soil moisture–climate interactions in a changing climate: A review', *Earth-Science Reviews*, vol. 99, no. 3–4, pp. 125–161, mei 2010.
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