

Weather regimes

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Also see: ICTP School on Weather Regimes and Weather Types (Oct 2013): http://indico.ictp.it/event/a12220/



Outline

- Introduction:
 - Dynamical concepts
 - Examples of recurrent flow patterns
- Historical overview
- Detection of regimes in atmospheric and model datasets
 - PDF estimation and statistical significance
 - Examples of cluster analysis for NH domains
- Sources of extended-range predictability
 - Impact of external/boundary forcing on atmospheric regimes
 - > MJO and Euro-Atlantic regimes
 - Non-linear impact of ENSO on regime properties

Weather regime:

A persistent and/or recurrent large-scale atmospheric circulation pattern which is associated with specific weather conditions on a regional scale

Flow regime:

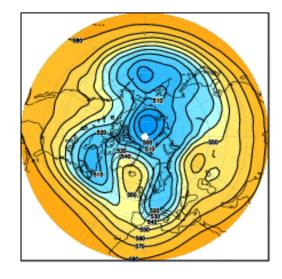
A persistent and/or recurrent large-scale flow pattern in a (geophysical) fluid-dynamical system

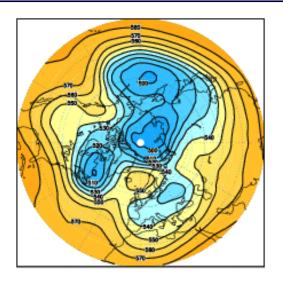
Multiple equilibria:

Multiple stationary solutions of a non-linear dynamical system

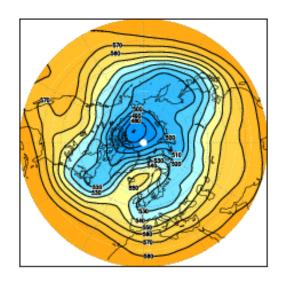


Recurrent flow patterns: examples



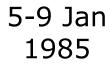


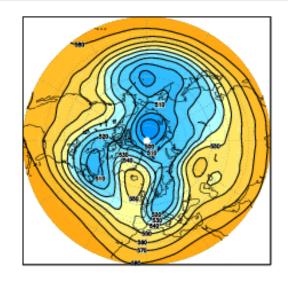
A sequence of 5-day mean fields of 500 hPa geopotential height during boreal winter ...

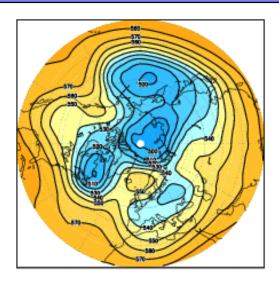




Recurrent flow patterns: examples

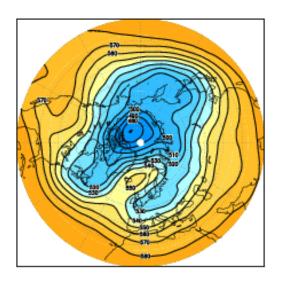






4-8 Feb 1986

... but each one occurred in a different winter !

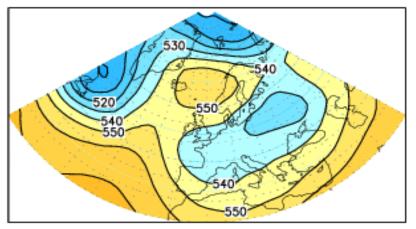


10-14 Jan 1987

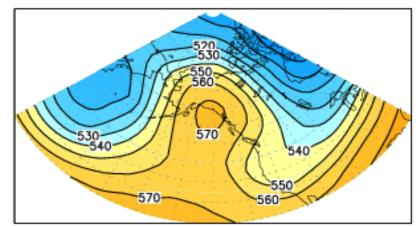


500 hPa geop. height

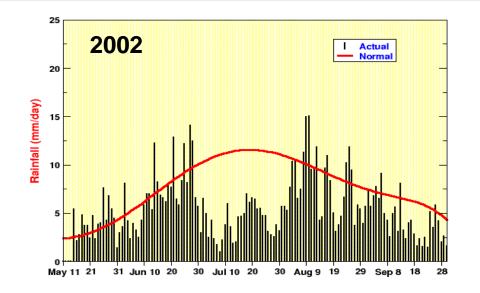
4-8 Feb. 1986



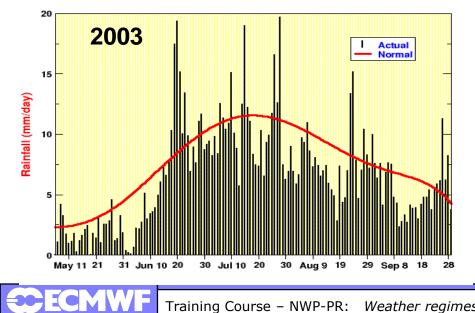
4-8 Feb. 1989

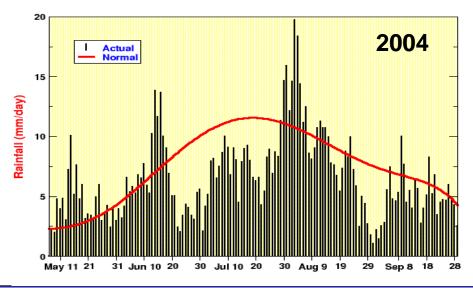


Regional regime behaviour: monsoon active/brake phases



All-India Rainfall time-series (May-September)





Training Course - NWP-PR: Weather regimes

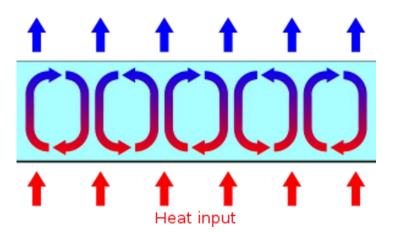


3-variable model of Rayleigh-Benard convection (Lorenz 1963)

- $dX/dt = \sigma (Y X)$
- dY/dt = -XZ + rX Y
- dZ/dt = X Y b Z

Unstable stationary states

- X = Y = Z = 0
- $X = Y = \pm [b(r-1)] \frac{1}{2}, Z = r-1$







 ${\bf q}$: barotropic or quasi-geostrophic potential vorticity

$$\partial_t q = -V_{\Psi} \cdot grad q - D(q - q^*)$$

steady state for instantaneous flow:

$$0 = -V_{\Psi} \cdot grad q - D (q - q^*)$$

steady state for time-averaged flow:

$$0 = - \langle V_{\psi} \rangle \cdot grad \langle q \rangle - D (\langle q \rangle - q^*) \\ - \langle V'_{\psi} \cdot grad q' \rangle$$

Multiple steady states of low-order barotropic model with wave-shaped bottom topography

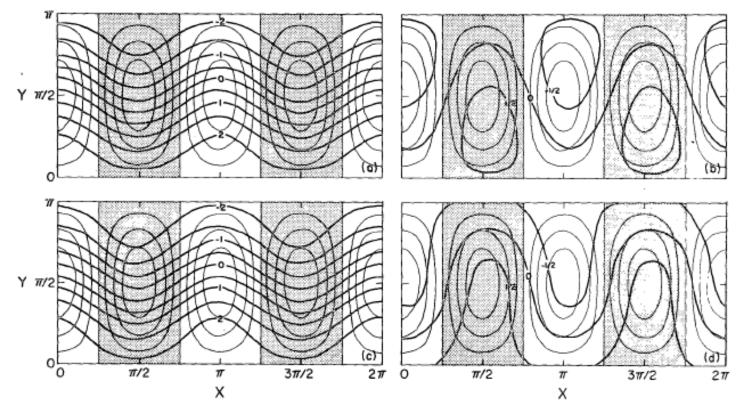
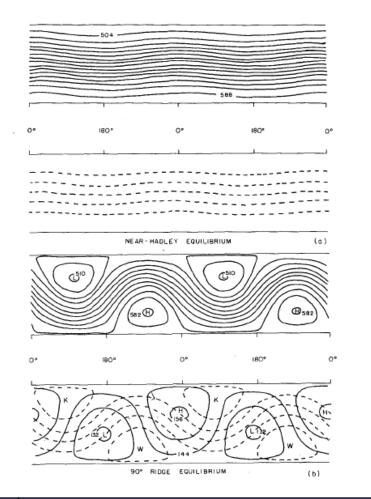
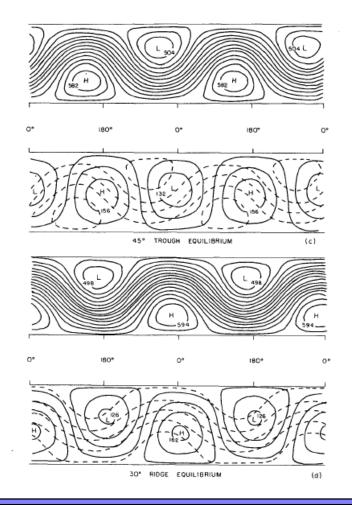


FIG. 4. Streamfunction fields of the stable first mode equilibria of a topographically forced flow for $k = 10^{-2}$, $L/a = \frac{1}{4}$, n = 2, $h_0/H = 0.2$ and $\psi_A^* = 0.2$; for the spectral model above resonance (a) and slightly below resonance (b); and for the grid-point model above resonance (c) and slightly below resonance (d). The nondimensional topographic heights are shown with light lines; the contour spacing is 0.05 units, with negative regions shaded.

Seminal papers: Reinhold and Pierrehumbert 1982

Hemispheric weather regimes arising from equilibration of large-scale dynamical tendencies and "forcing" from transient baroclinic eddies



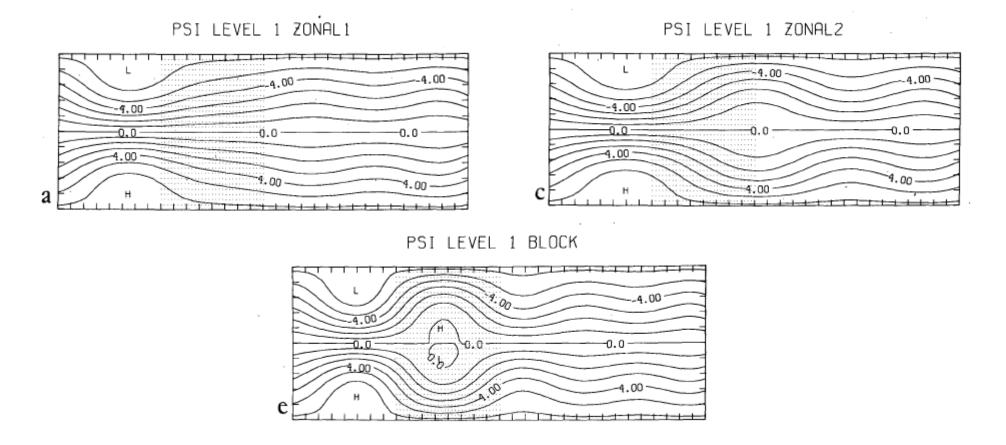




- **Green 1977**: The weather during July 1976: some dynamical consideration of the drought
- **Illari and Marshall 1983**: *On the interpretation of eddy fluxes during a blocking episode*
- **Shutts 1986**: *A case study of eddy forcing during an Atlantic blocking episode*
- Haines and Marshall 1987: Eddy-forced coherent structures as a propotype of atmospheric blocking



Regional weather regimes arising from equilibration of large-scale dynamical tendencies and PV fluxes from transient baroclinic eddies



Bimodality in the probability density function (PDF) of an index of N. Hem. planetary wave amplitude due to near-resonant wave-numbers (m=2-4)

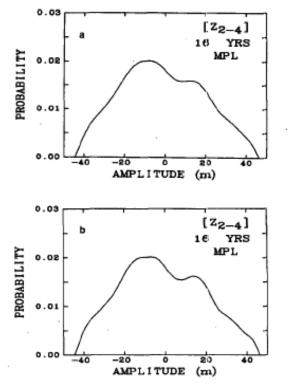
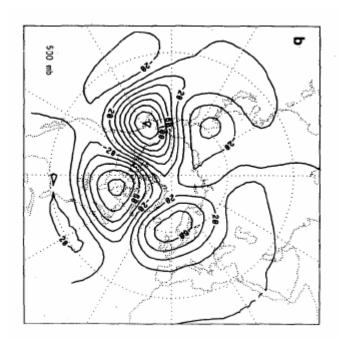
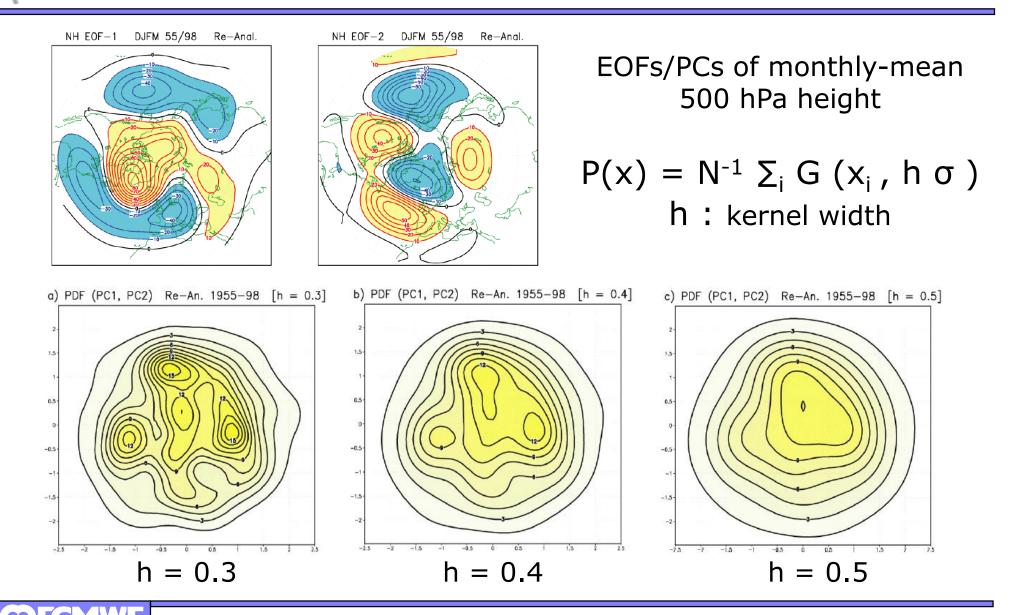


FIG. 4. MPL probability density estimates of $[Z_{2-4}]$ formed from the 16 winter composite filtered data for (a) $\alpha = 10^{7}$ and (b) $\alpha = 5 \times 10^{6}$.

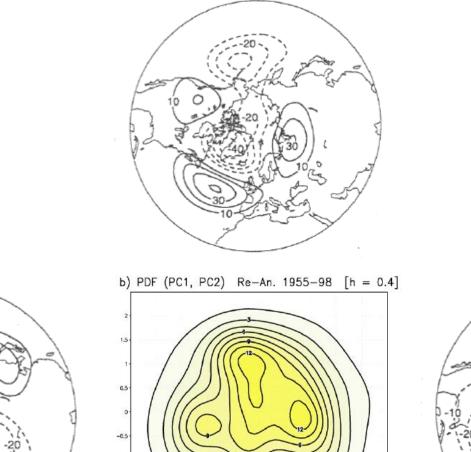


PDF estimation with the Gaussian kernel method



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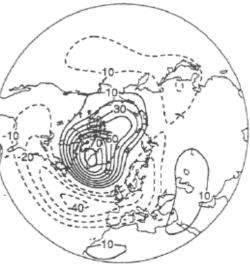
Regimes from PDF estimation (Corti et al. 1999)

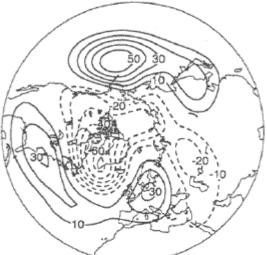


-0.5

0.5

1.5





-1.5

-2.5 -2 -1.5



Regimes defined by multi-dim PDF and cluster analysis

- Mo and Ghil 1988
- Molteni et al. 1990
- Cheng and Wallace 1993
- Kimoto and Ghil 1993a, b
 - Michelangeli et al. 1995

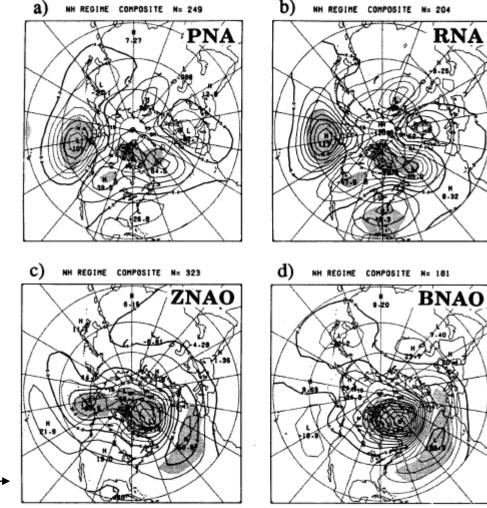
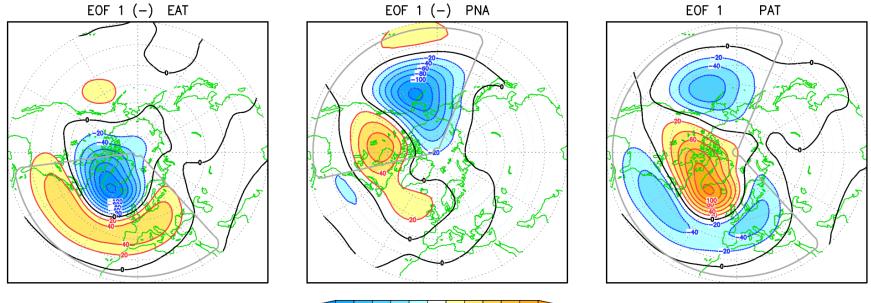


Fig. 14. Composite maps of unfiltered anomalies for the four NH regimes. Those samples falling in either of the four rectangles in Fig. 11a are collected for (a) PNA, (b) RNA, (c) ZNAO, and (d) BNAO. Numbers of collected daily maps are (a) 249, (b) 204, (c) 323, and (d) 181, respectively. Contour interval is 15 meters; shaded regions are significantly different from zero at a 99% level judged by a pointwise *t*-test.

Kimoto and Ghil 1993a

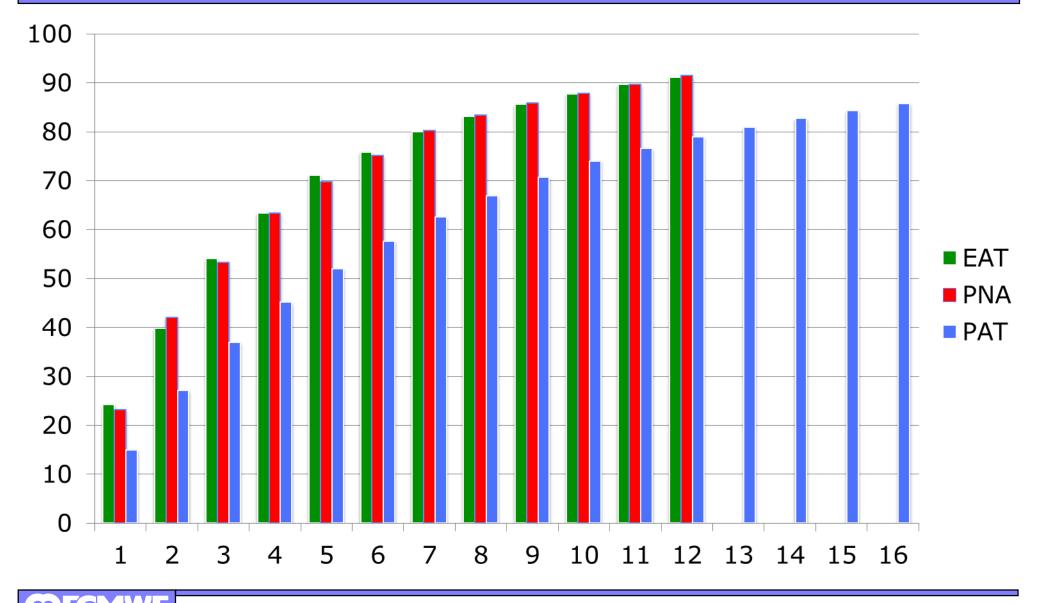
Data: 5-day means of Z 500 hPa in DJF 1979/80 to 2012/13 (from ERA-interim)



-120-100-80 -60 -40 -20 20 40 60 80 100 120

Cluster analysis method: k-means (Michelangeli et al. 1995, Straus et al. 2007)

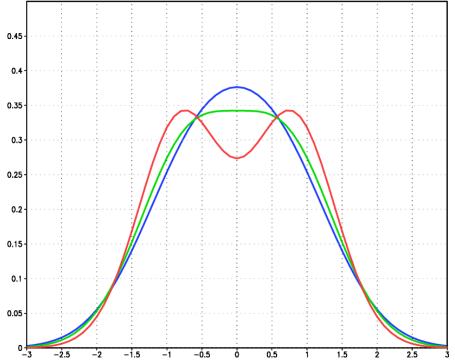
variance of N EOFs (%) in the three domains





a) 2 regimes in 1 dimension:

P(x) = 0.5 [G(μ , σ) + G(- μ , σ)] Total variance = $\mu^2 + \sigma^2$ S/N variance ratio = μ^2 / σ^2 <u>P(x) is bimodal if S/N > 1</u> $\mu = 0.8, \sigma_x = 0.6: S/N = 1.78$ $\mu = 0.71, \sigma_x = 0.71: S/N = 1.00$ $\mu = 0.6, \sigma_x = 0.8: S/N = 0.56$ b) 2 regimes in 2 dimensions



P(x, y) = P(x) P(y) $P(x) = 0.5 [G(\mu, \sigma_x) + G(-\mu, \sigma_x)], P(y) = G(0, \sigma_y)$ If $\mu = \sigma_x = \sigma_y = 0.71$: $S/N = \mu^2 / (\sigma_x^2 + \sigma_y^2) = 0.5$

For N regimes, S/N should be > 1 in a subspace of N-1 dimensions

(a lower limit applies to regimes with different population)

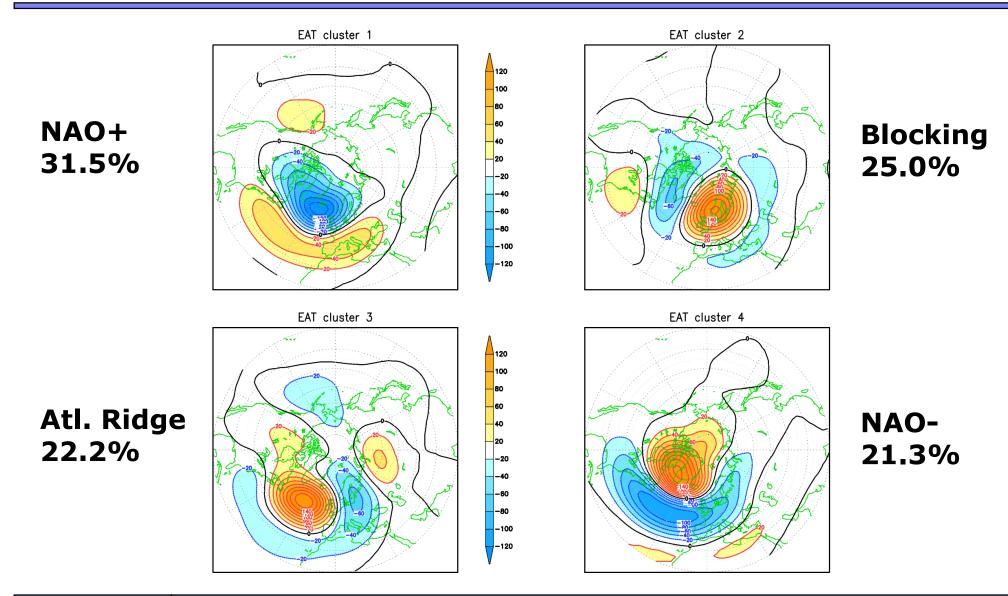


Statistics for N-cluster partitions (%)

	2 cl	3 cl	4 cl	5 cl	6 cl	7 cl	8cl
E-AT var s/n	24.7	42.3	59.3	71.4	81.5		
E-AT conf.lev	52.7	86.8	99.8	99.6	99.8		
P-NA var s/n	24.2	43.8	57.9	69.4	79.1		
P-NA conf.lev	76.0	87.6	98.6	98.8	99.0		
P-AT var s/n	15.6	27.3	36.3	43.6	50.0	55.7	61.2
P-AT conf.lev	57.0	76.2	90.4	93.0	97.4	98.0	98.8

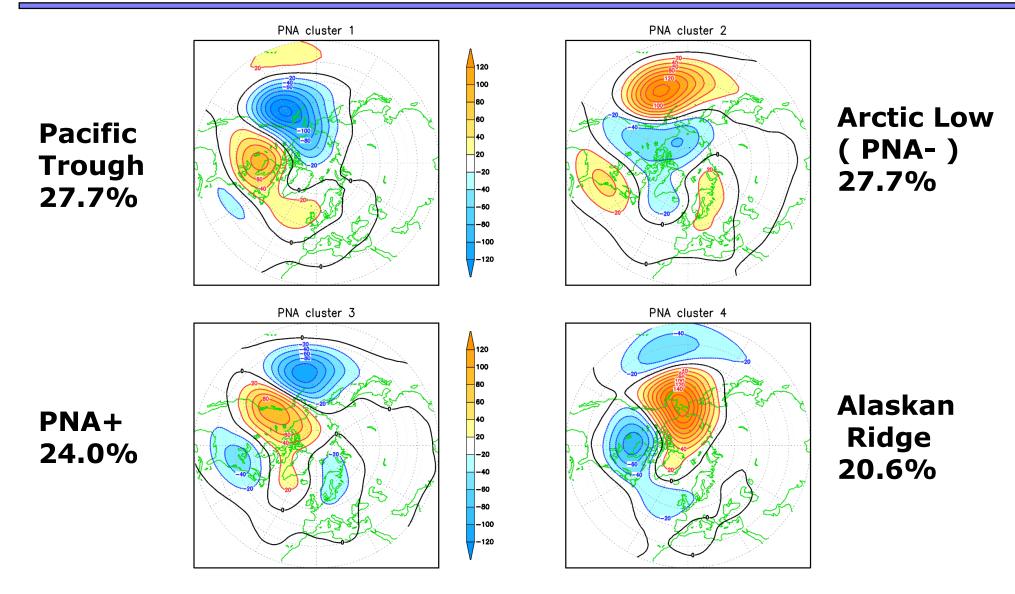


Euro-Atlantic 4-cluster centroids





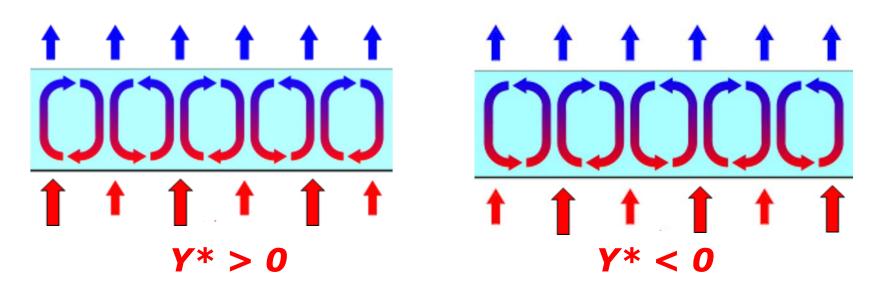
Pacific-North American 4-cluster centroids





Lorenz (1963) truncated convection model with additional forcing (Molteni et al. 1993; Palmer 1993)

- $dX/dt = \sigma (Y X)$
- $dY/dt = -XZ + rX (Y Y^*)$
- dZ/dt = X Y b Z

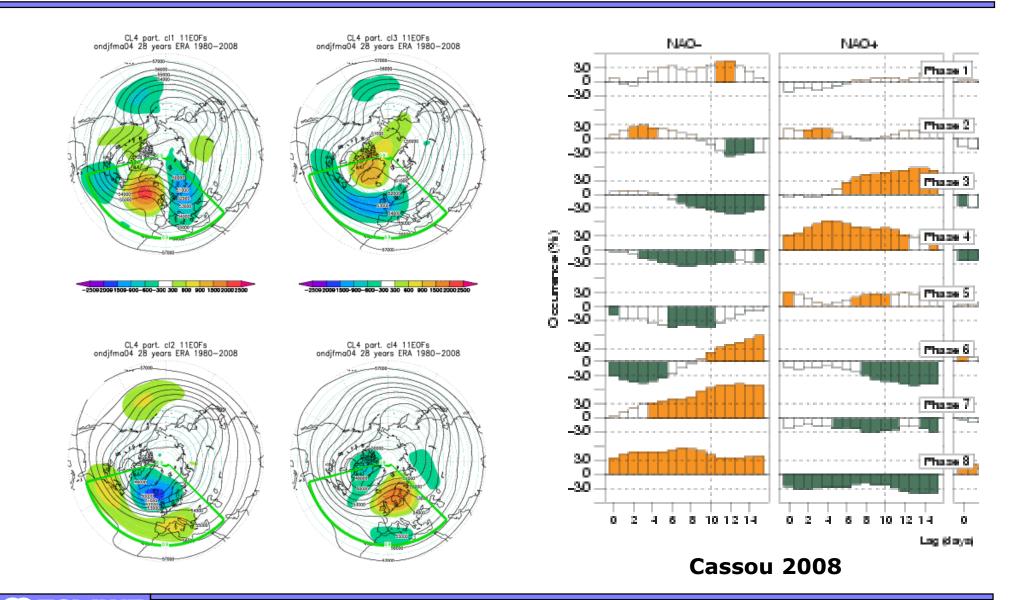


The properties of flow regimes may be affected by anomalous forcing in two different ways:

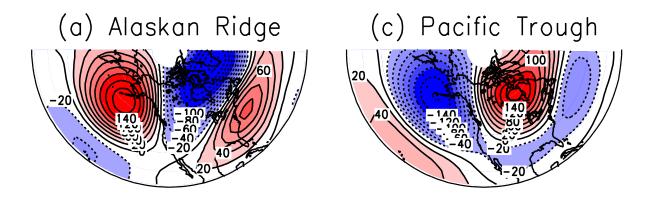
- Weak forcing anomaly: the number and spatial patterns of regimes remain the same, but their frequency of occurrence is changed
- Strong forcing anomaly: the number and patterns of regimes are modified as the atmospheric system goes through bifurcation points

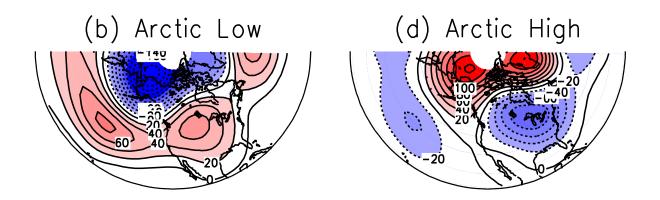


Impact of MJO on Euro-Atlantic regimes (see lecture on MJO)

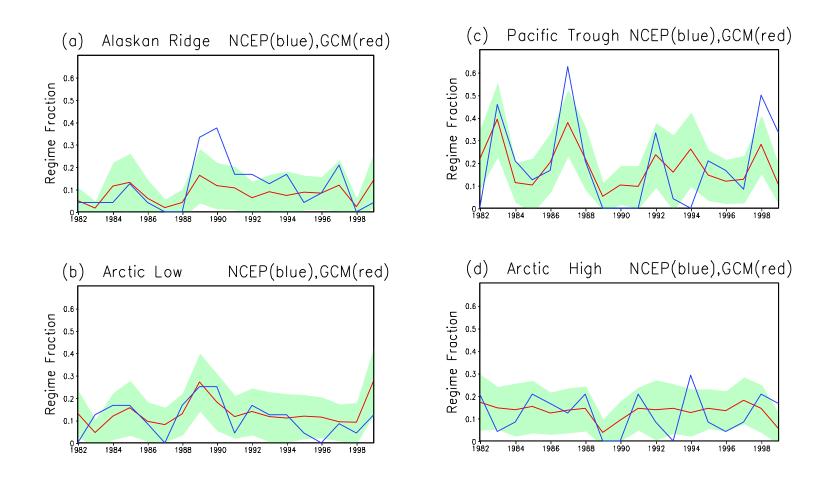


Cluster analysis of low-freq. (T>10 d) Z 200 in NCEP re-analysis and COLA AGCM ensembles (Straus, Corti, Molteni 2007)





Predictability of cluster frequencies (SCM 2007)





 Ratio of inter-cluster to intra-cluster variance as a function of ENSO indices (Straus and Molteni 2004)

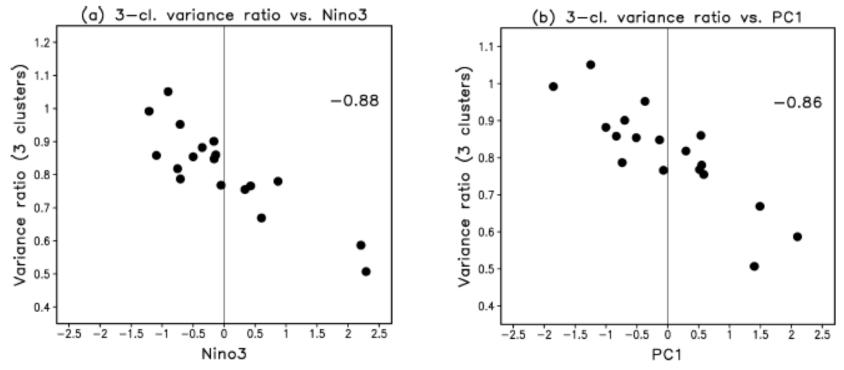


FIG. 4. Scatterplots of (a) the 3-cluster (k = 3) variance ratio vs Niño-3, and (b) the 3-cluster variance ratio vs the leading PC of ensemble/seasonal means. The leading PC and SST index time series are standardized.



- Flow regime behaviour can be reproduced in a variety of dynamical models of different complexity.
- Atmospheric flow regimes may be defined on a hemispheric or regional domain.
- Detection of regimes in atmospheric and model datasets is usually performed by PDF estimation or cluster analysis; results are dependent on adequate time-filtering and proper use/interpretation of statistical significance tests.
- The impact of forcing anomalies on regime properties may occur through changes in regime frequencies or bifurcation effects.
- Predictability of regime frequencies and variations in the number of regimes as a function of the ENSO and MJO phases have been detected in ensembles of GCM simulations, and offer an alternative approach to long-range prediction.



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